

Eidgenössische Technische Hochschule Zürich Swiss Federal Institute of Technology Zurich

Embedded convection in the warm conveyor belt of a North Atlantic cyclone and its relevance for large-scale dynamics

Workshop on warm conveyor belts – a challenge to forecasting

Annika Oertel, Maxi Boettcher, Hanna Joos, Michael Sprenger, and Heini Wernli

March 10, 2020

23 Sep 2016 – Cyclone Vladiana (IOP 3 - NAWDEX)

slow slantwise ascent





Satellite observations

Binder 2016, PhD thesis; Flaounas et al. 2016, QJRMS; Crespo and Posselt 2016, MWR; Flaounas et al. 2018, ClimDyn; Oertel et al. 2019, QJRMS

23 Sep 2016 – Cyclone Vladiana (IOP 3 - NAWDEX) Slow slantwise ascent 40°N 40°N Certel et al. 2019, QIRMS

- convective clouds
- cirrus clouds

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23 Sep 2016 – Cyclone Vladiana (IOP 3 - NAWDEX) Slow slantwise ascent fast convective ascent fast convective ascent for N dor N d

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23 Sep 2016 – Cyclone Vladiana (IOP 3 - NAWDEX) slow slantwise ascent "escalator" fast convective ascent "elevator" 60°N 40°N cf. Neiman et al. 1993, MWR 40°W 20°W 0° "escalator-elevator concept" Oertel et al. 2019, QJRMS

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Convection-permitting simulations

Rasp et al. 2016, MWR; Oertel et al. 2019, QJRMS; Oertel et al. 2019, WCDD



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cloud diabatic processes

$$\frac{D}{Dt}PV = \frac{1}{\rho}\boldsymbol{\omega} \cdot \nabla \dot{\boldsymbol{\theta}}$$

 $PV = \frac{1}{\rho} \frac{\mathbf{vorticity}}{\mathbf{vorticity}}$ potential temperature gradient





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WCB ascent e.g., Wernli and Davies 1997 QJRMS; Joos and Wernli 2012, QJRMS

cf. presentation by Ben Harvey

Revisiting the isentropic view of PV modification in warm conveyor belts



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modification of the waveguide Wernli and Davies 1997, QJRMS; Joos and Wernli 2012, QJRMS; Joos and Forbes 2016, QJRMS



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gradient

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smaller-scale diabatic heating

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>> embedded convection in WCBs



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>> embedded convection in WCBs



Key questions

I. How does embedded convection influence the cloud and precipitation structure?



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II. How does embedded convection modify the PV distribution?





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II. How does embedded convection modify the PV distribution?

III. How does embedded convection influence the mesoand large-scale circulation?





Cyclone Vladiana (IOP 3 NAWDEX)

• **Convection-permitting simulation** with the non-hydrostatic model COSMO

 $\Delta \text{lon} = \Delta \text{lat} = 0.02^{\circ} (\sim 2 \text{ km})$

limited-area model COSMO





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Miltenberger et al. 2013, 2014

- calculated from **3D wind field at every model timestep** ($\Delta t = 20 \text{ s}$)

- explicitly capture rapid convective ascent





surface height [m]

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• WCB trajectories 600 hPa in 48 h e.g., Madonna et al. 2014, JCLI



500

1000



1500

2000

surface height [m]

2500

3000

3500

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 - **Slantwise 'escalator'**-like WCB trajectories 'typical' WCB





time [h]

10°F

3000

2500

20°F

30°F

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 - Convective 'elevator'-like WCB trajectories
 >600 hPa in 3 h











0.05



0.25

0.05











Online trajectories from 2 km COSMO simulation

convective WCB ascent - 'elevator' > 400 hPa / 2 h



PV@320K

Online trajectories from 2 km COSMO simulation

convective WCB ascent - 'elevator' > 400 hPa / 2 h



PV@320K

Online trajectories from 2 km COSMO simulation

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1000 km


Composite low-level PV structure (at 800 m)

slantwise WCB^{1 m s⁻¹}



600 km



Composite low-level PV structure (at 800 m)

slantwise WCB^{1 m s⁻¹} convective WCB^{1 m s⁻¹} **x** WCB trajectory position **x** convective updraft - 13 13 8 8 4 4 600 km 120 km PV [pvu] PV [pvu] 2 1 0 0 $^{-1}$ -1-2-2 600 km 120 km 11111



600 km





mmmm







mmmm







 $\frac{D}{Dt}PV = \frac{1}{\rho} \left[(f+\zeta)\frac{\partial\dot{\theta}}{\partial z} + \omega_h \cdot \nabla_h \dot{\theta} \right]$



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Composite upper-level vertical vorticity structure (at 320 K)



$$\frac{D}{Dt}PV = \frac{1}{\rho} \left[(f+\zeta)\frac{\partial\dot{\theta}}{\partial z} + \omega_h \cdot \nabla_h \dot{\theta} \right]$$

II. How does embedded convection influence the PV structure?









- 1) existence of PV anomaly on a larger-scale >> coarse-graining to 60 km
- 2) Temporal evolution of negative PV





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1) existence of PV anomaly on a larger-scale >> coarse-graining to 60 km

2) Temporal evolution of negative PV

>> neg. PV has a relatively long lifetime, can interact with the upper-level waveguide and form jet streaks



Conclusions

 Embedded convective ascent in WCBs forms a denser cloud with higher hydrometeor content and more intense surface precipitation



Conclusions

- Embedded convective ascent in WCBs forms a denser cloud with higher hydrometeor content and more intense surface precipitation
- II. Embedded convection in WCBs leads to
 - stronger pos. low-level PV anomalies
 - mesoscale upper-level horizontal PV dipoles





Conclusions

- Embedded convective ascent in WCBs forms a denser Ι. cloud with higher hydrometeor content and more intense surface precipitation
- Embedded convection in WCBs leads to П.
 - stronger pos. low-level PV anomalies
 - mesoscale upper-level horizontal PV dipoles
- On the larger-scale, convectively produced PV anomalies III. aggregate to **elongated PV dipole bands** around elongated convective updrafts, which can
 - interact with the upper-level waveguide
 - strengthen the isentropic PV gradient
 - result in jet streaks







Thank you



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PV structure

20 km





upper-level PV dipole



120 km

800 m

4

2

0

-1

-2

PV [pvu]

120 km

II. How does embedded convection influence the PV structure?



Potential (in-)stability ($d\Theta_e/dz$)



slantwise WCB
>> typical ascent


Equivalent potential temperature



"average" WCB
>> typical ascent

convective WCB

Vertical vorticity



 $\frac{120 \text{ km}^{1 \text{ ms}^{-1}}}{120 \text{ km}}$

Different stages of convectively generated PV



Comparison of cyclones *Vladiana* and *Sanchez*

embedded convection of *Vladiana* and *Sanchez*



convective WCB

Composite upper-level PV structure (at 320 K)



PV structure and CAPE



convective WCB – 1h prior to ascent



Flow attribution



PV modification

$$\frac{D}{Dt}PV \approx \frac{1}{\rho} \left[\zeta \frac{\partial \dot{\theta}}{\partial z} + \omega_h \cdot \nabla_h \dot{\theta} \right]$$





Harvey et al.2020, QJR

$$\rho \frac{\tilde{D}P}{Dt} = \frac{P \nabla \cdot (\rho \mathbf{u}_D)}{P \nabla \cdot (\rho \mathbf{u}_D)} + \frac{\nabla \cdot (\boldsymbol{\zeta}_{//} \dot{\theta})}{\nabla \cdot (\boldsymbol{\zeta}_{//} \dot{\theta})} - \nabla \cdot (\mathbf{F} \times \nabla \theta)$$



Jet acceleration

18 UTC 23 Sep 2016







Vertical cross-section





Vertical cross-section



