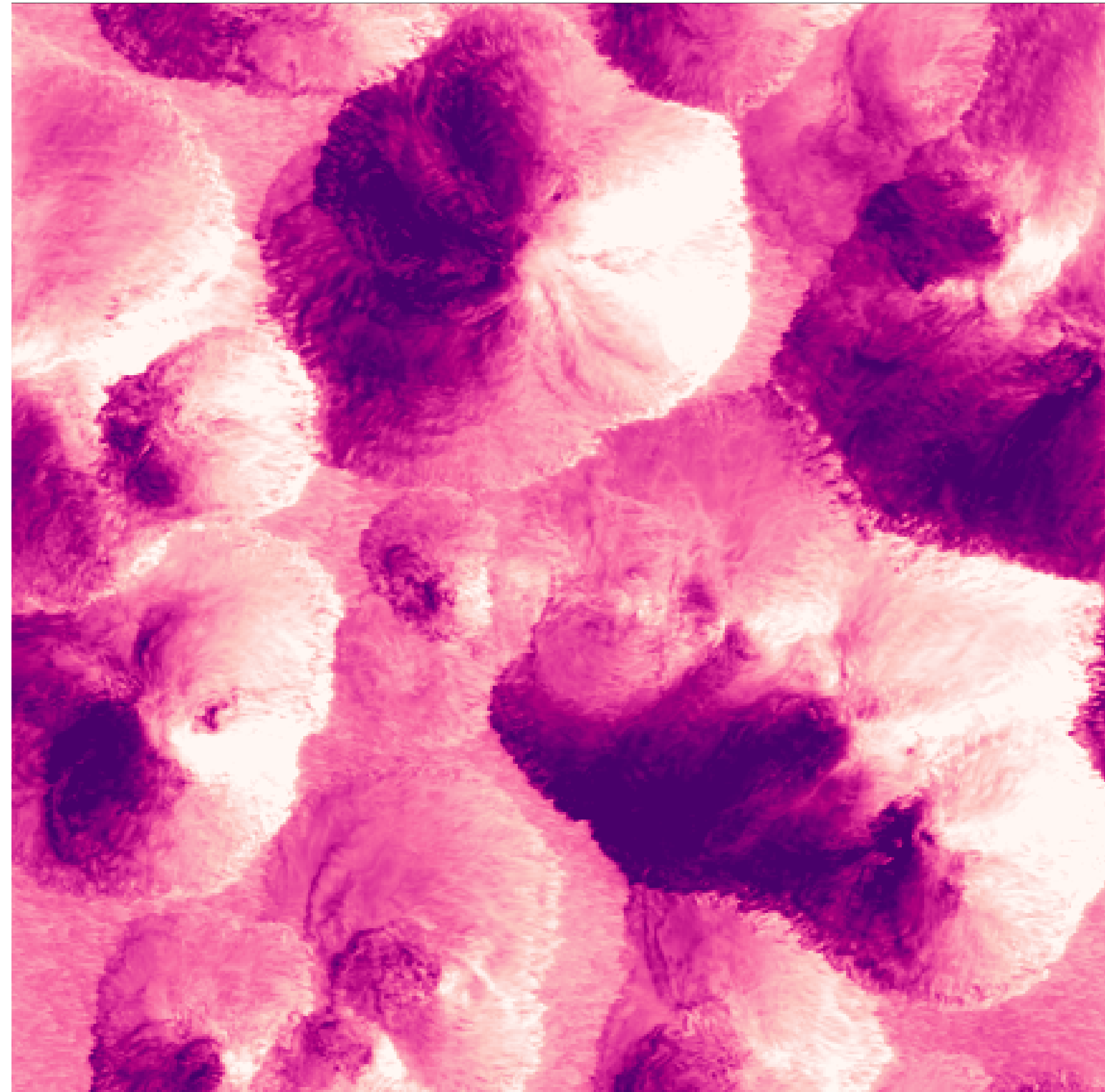


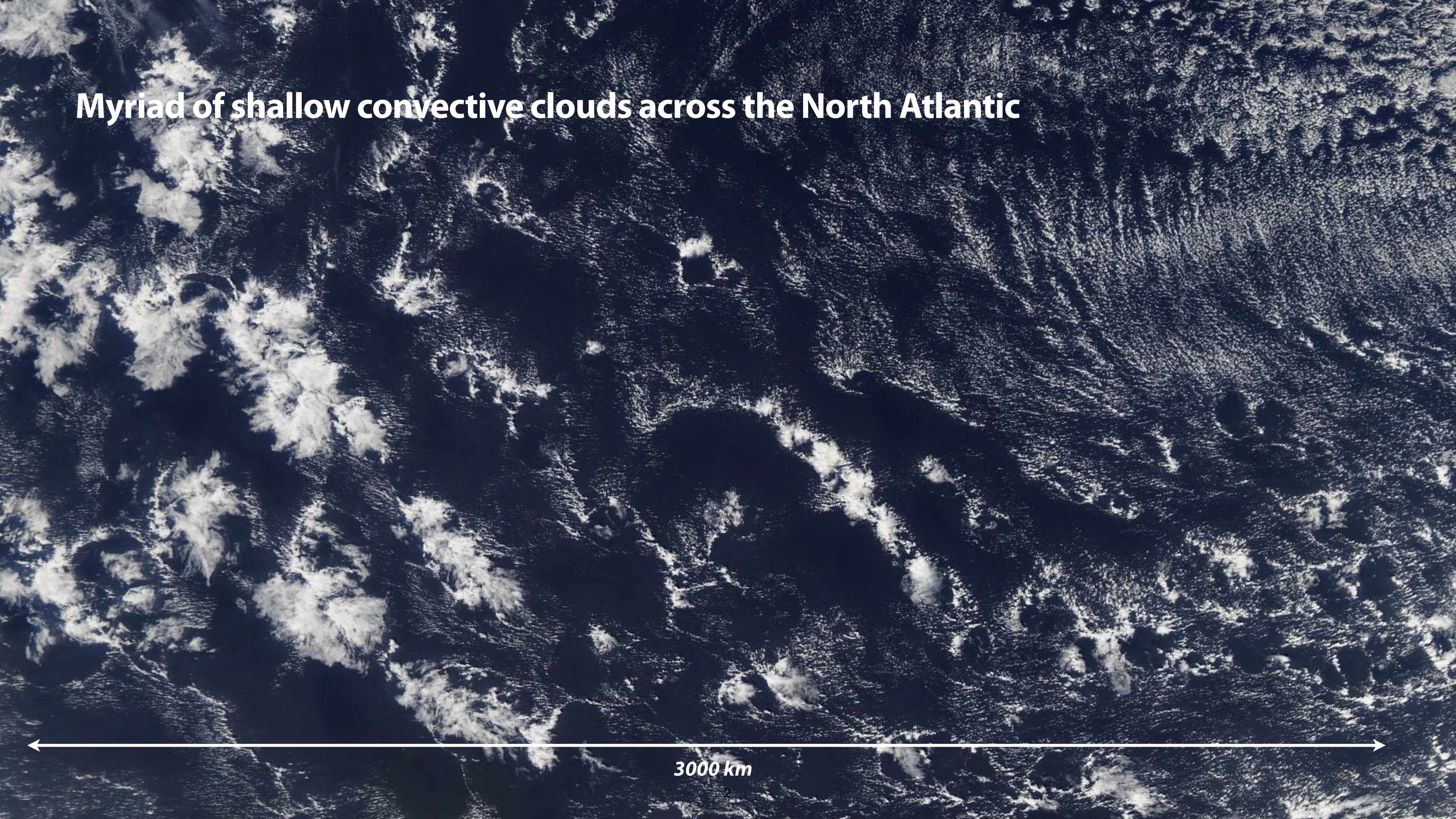
Momentum transport across scales in complex cloudy atmospheres

Louise Nuijens, Alessandro Savazzi, Vishal Dixit, Pier Siebesma, Wim de Rooy

ECMWF Annual Seminar
September 12, 2022

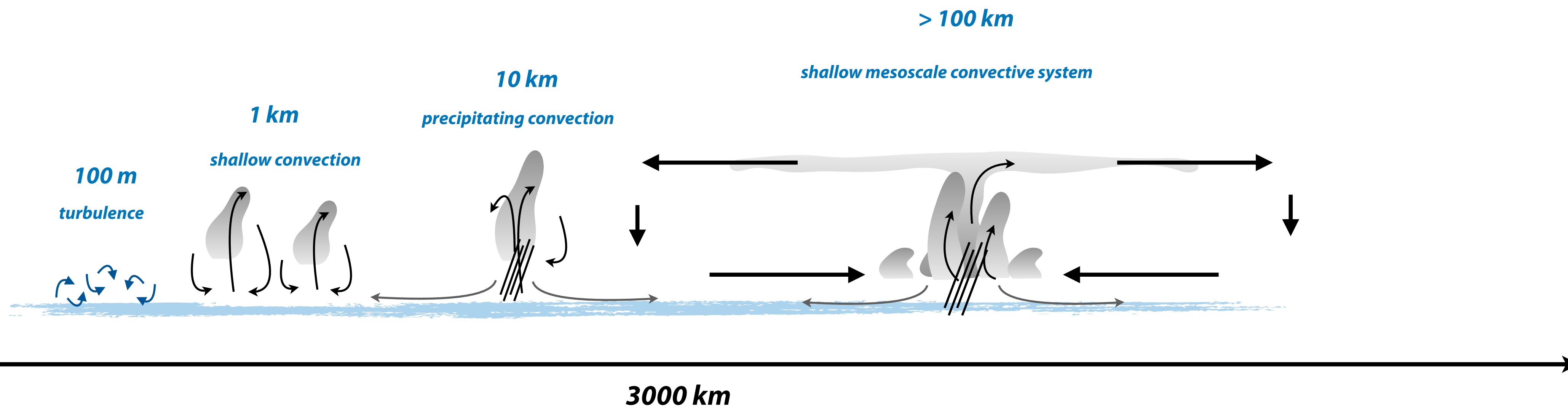


Myriad of shallow convective clouds across the North Atlantic



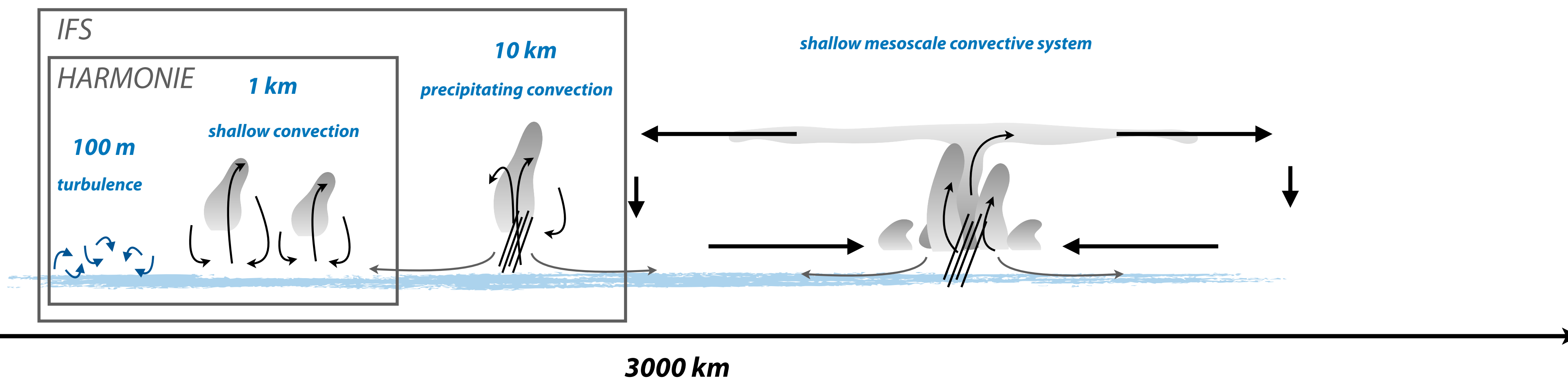
3000 km

Multi-scale flows in the presence of shallow cloud organization patterns

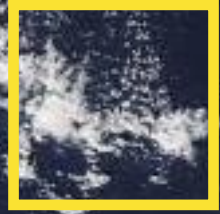


Turbulence and shallow convection parameterized in weather models

parameterized processes

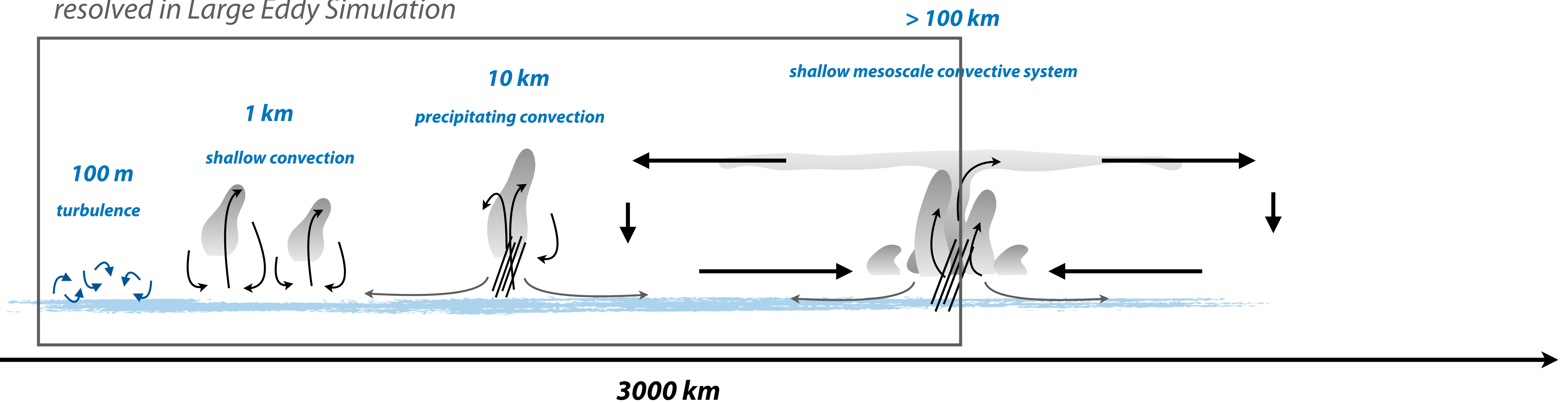


From idealised to realistic large-eddy simulation run in 'weather' mode

150 km  LES

20 km  LES

resolved in Large Eddy Simulation



What have we learned about shallow convective momentum transport?

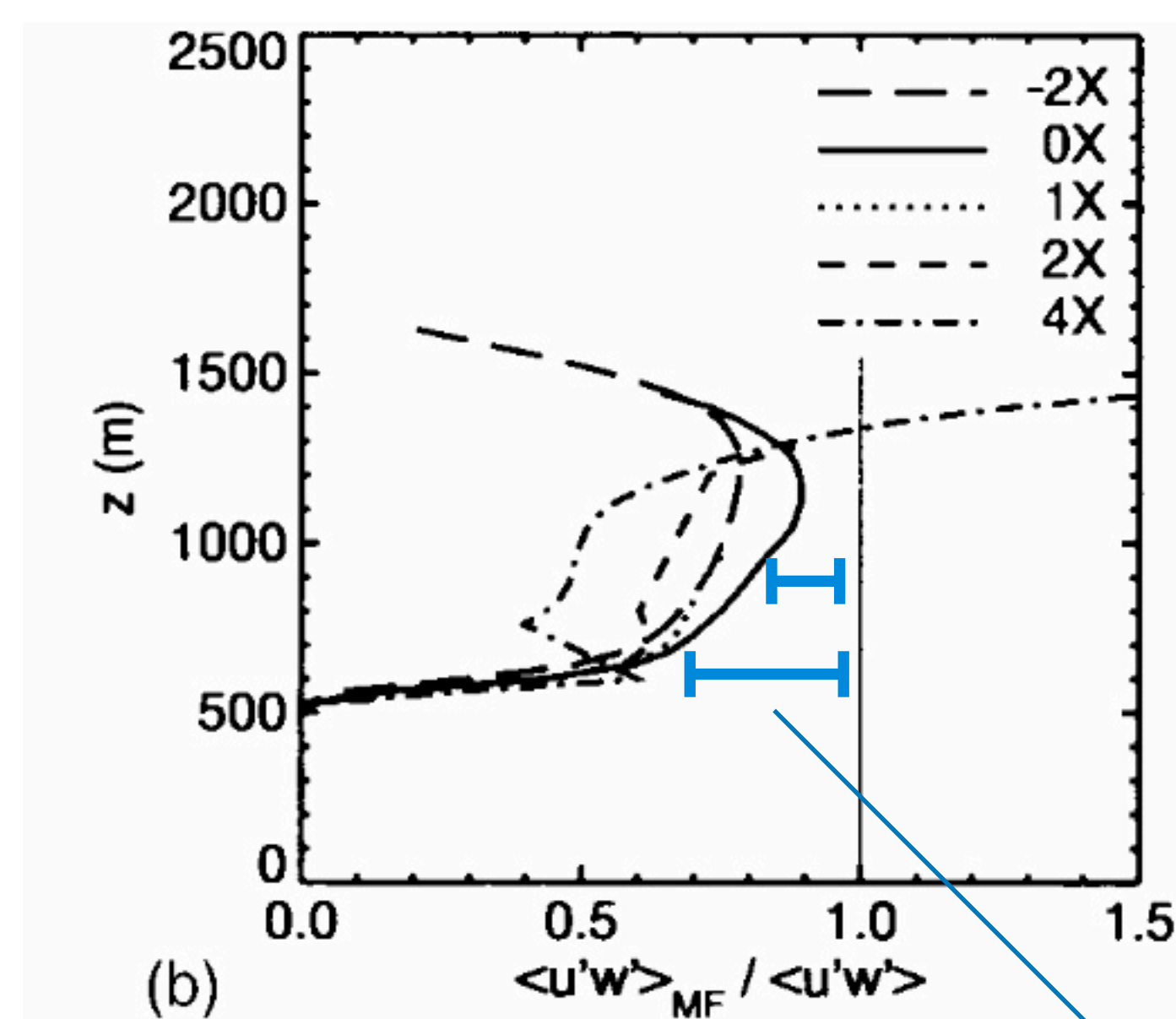
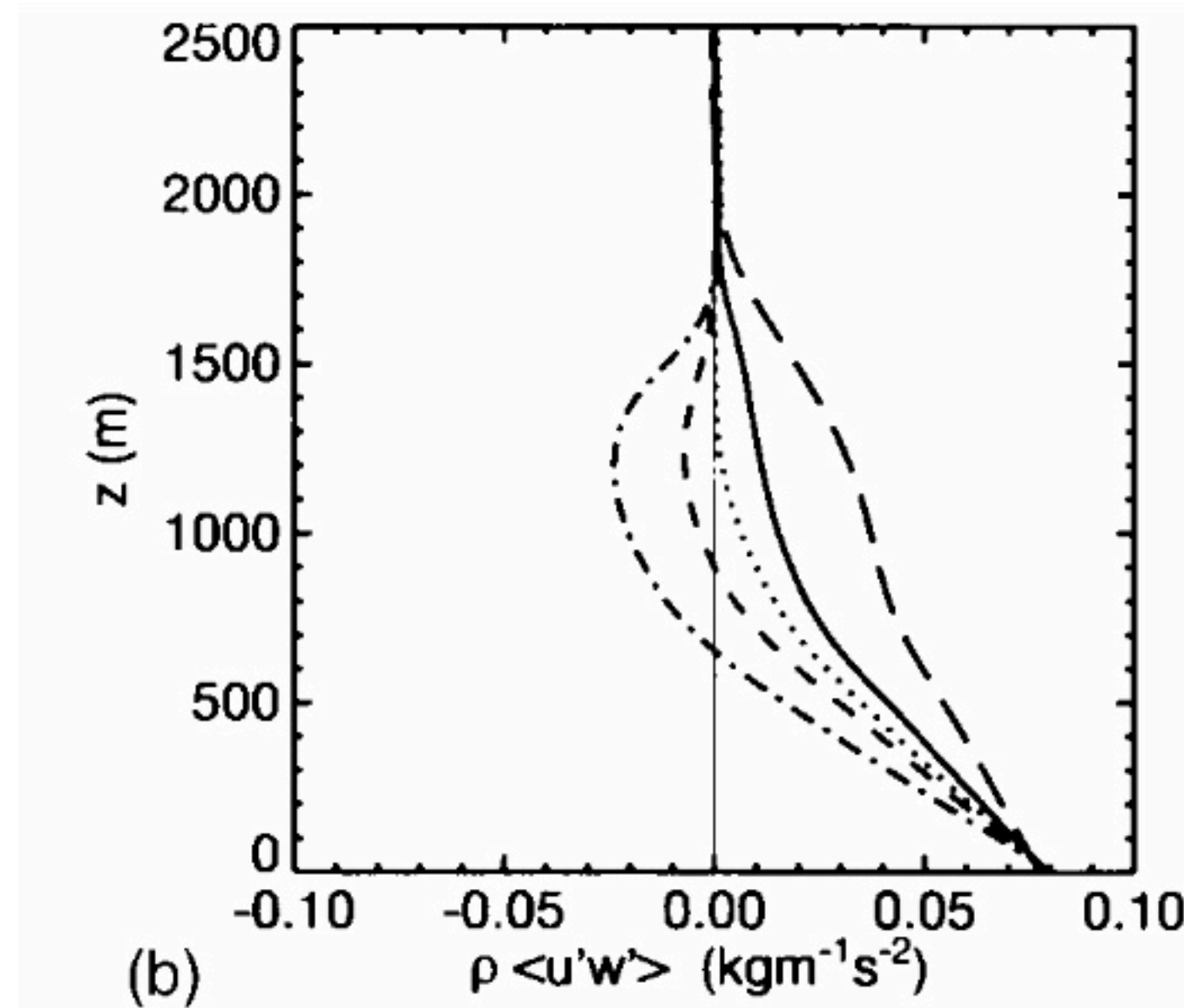
1. Down-gradient and up-gradient momentum transport - since the 90's
2. Convective momentum transport in Large Eddy and Storm Resolving Models: relevant scales
3. Link to wind biases in the IFS in the trades

What have we learned about shallow convective momentum transport?

1. Down-gradient and up-gradient momentum transport - since the 90's
2. Convective momentum transport in Large Eddy and Storm Resolving Models: relevant scales
3. Link to wind biases in the IFS in the trades

There is momentum to be gained at (sub)mesoscales

Mass flux does not suffice (environmental flux is significant) and pressure gradients matter for in-cloud momentum



Brown (1999): BOMEX on 6.4 km x 6.4 km x 3.0 km, $\Delta x = 100\text{m}$

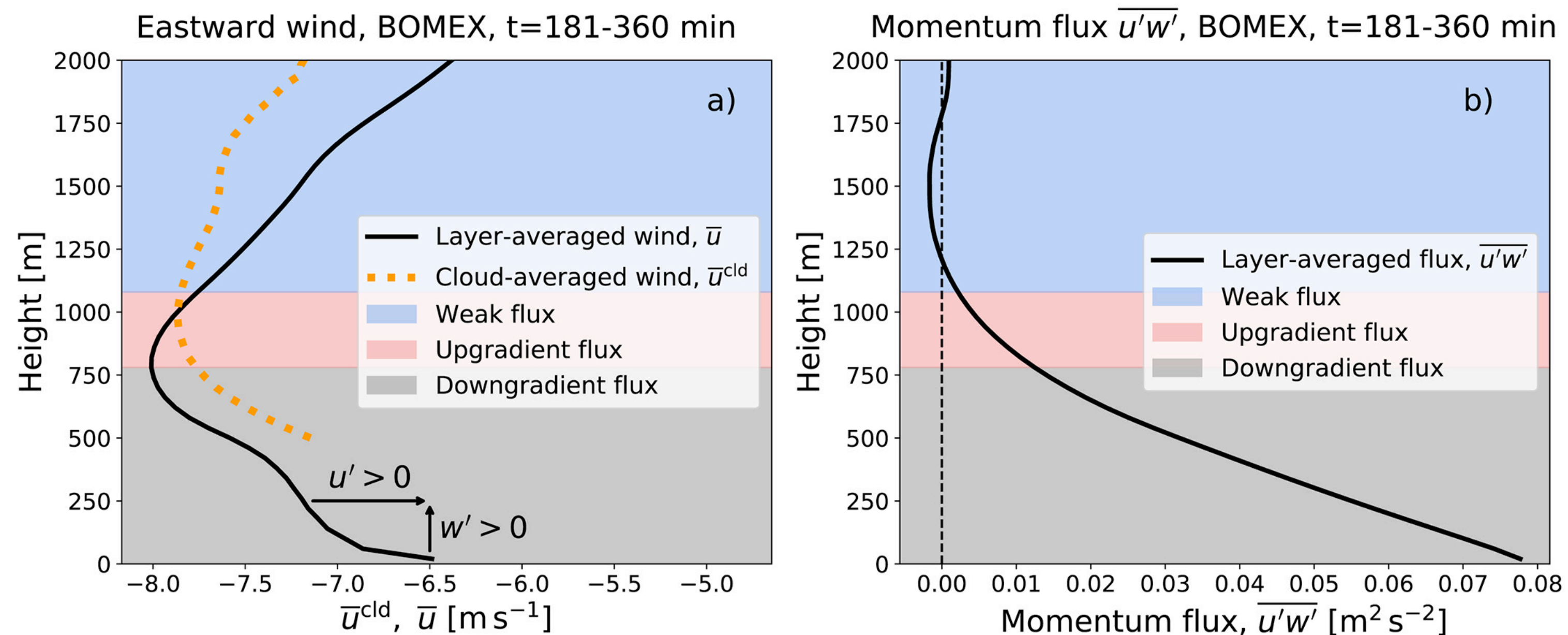
Zhu (2005): BOMEX/RICO on 16 x 16 km x 3 km, $\Delta x = 25\text{m}$

Schlemmer et al (2017): RICO on 51.2 km x 51.2 km x 5 km, $\Delta x = 25\text{m}$

Larson et al (2019): RICO on 51.2 km x 51.2 km x 5 km, $\Delta x = 100\text{m}$

Momentum flux not captured by mass flux

A layer of upgradient momentum flux is created by non-local transport through a zonal wind jet



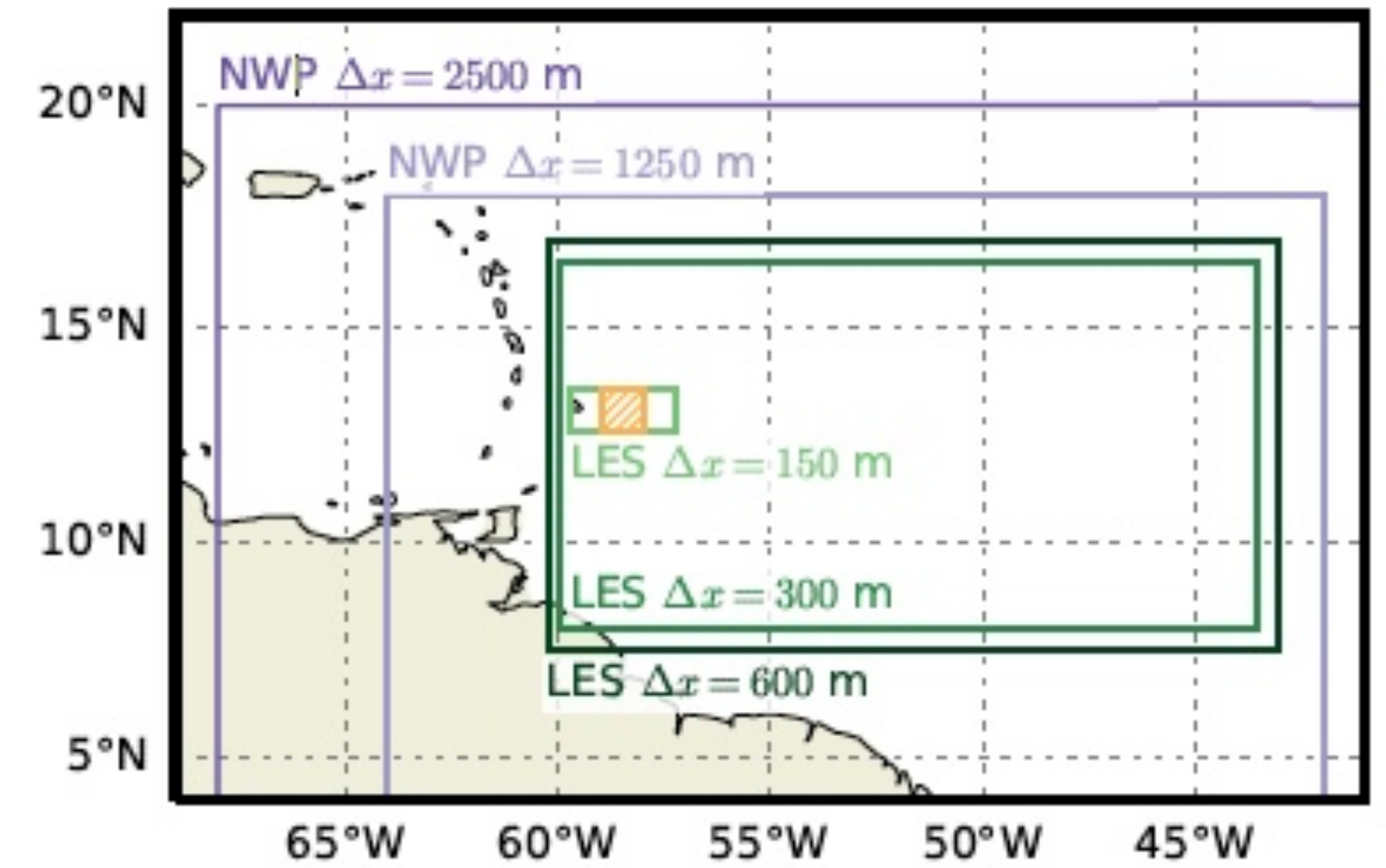
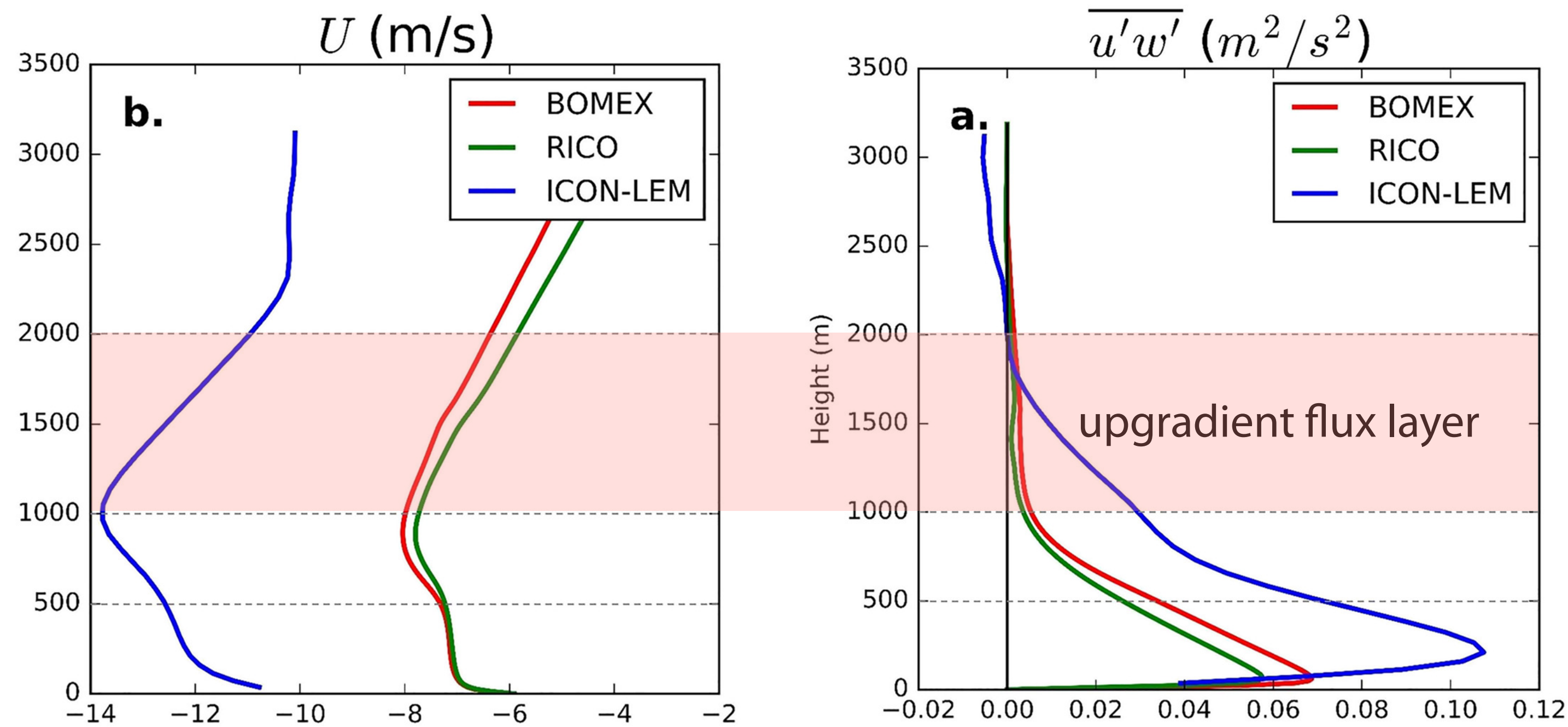
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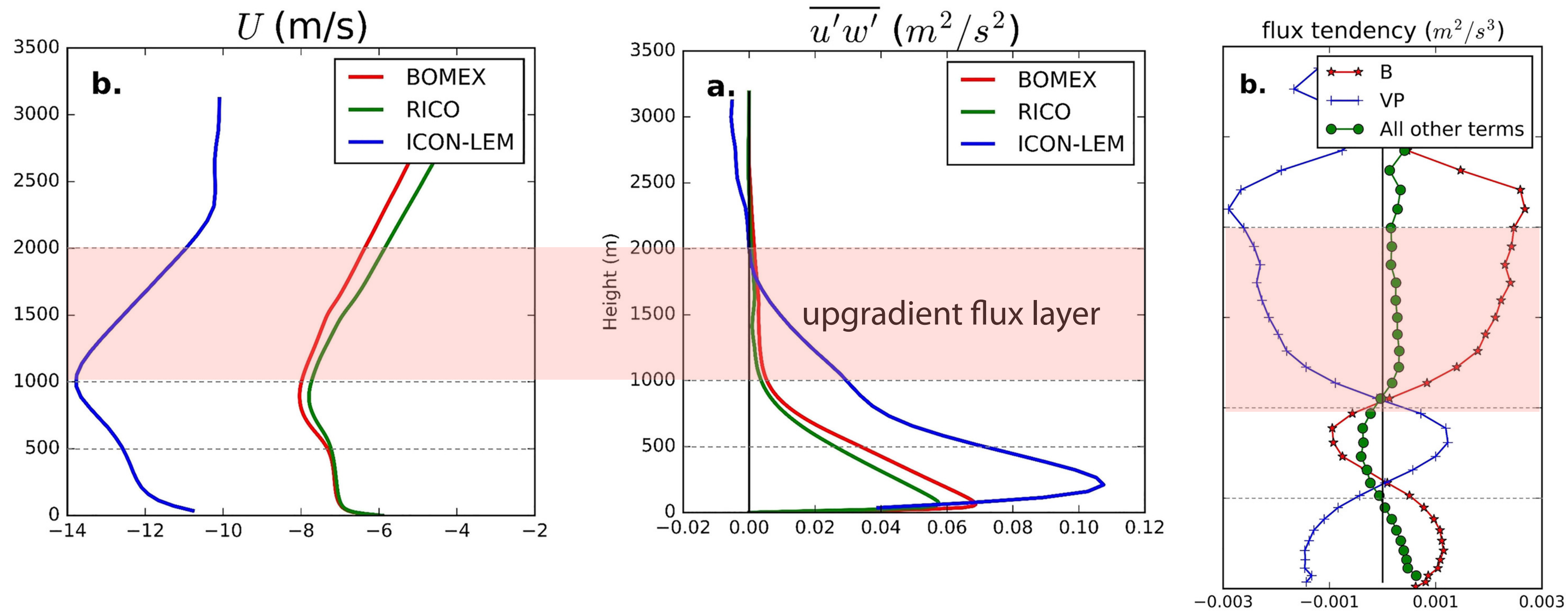
Larson et al (2019): BOMEX on 25.6 km x 25.6 km, $\Delta x = 100\text{m}$

Deeper layer of upgradient flux in open-boundary ICON-LEM hindcasts



Dixit et al (2020): NARVAL with ICON-LEM on 100 km x 100 km x > 5 km, $\Delta x = 150$ m

Buoyancy generation of flux largely compensated by vertical pressure perturbations

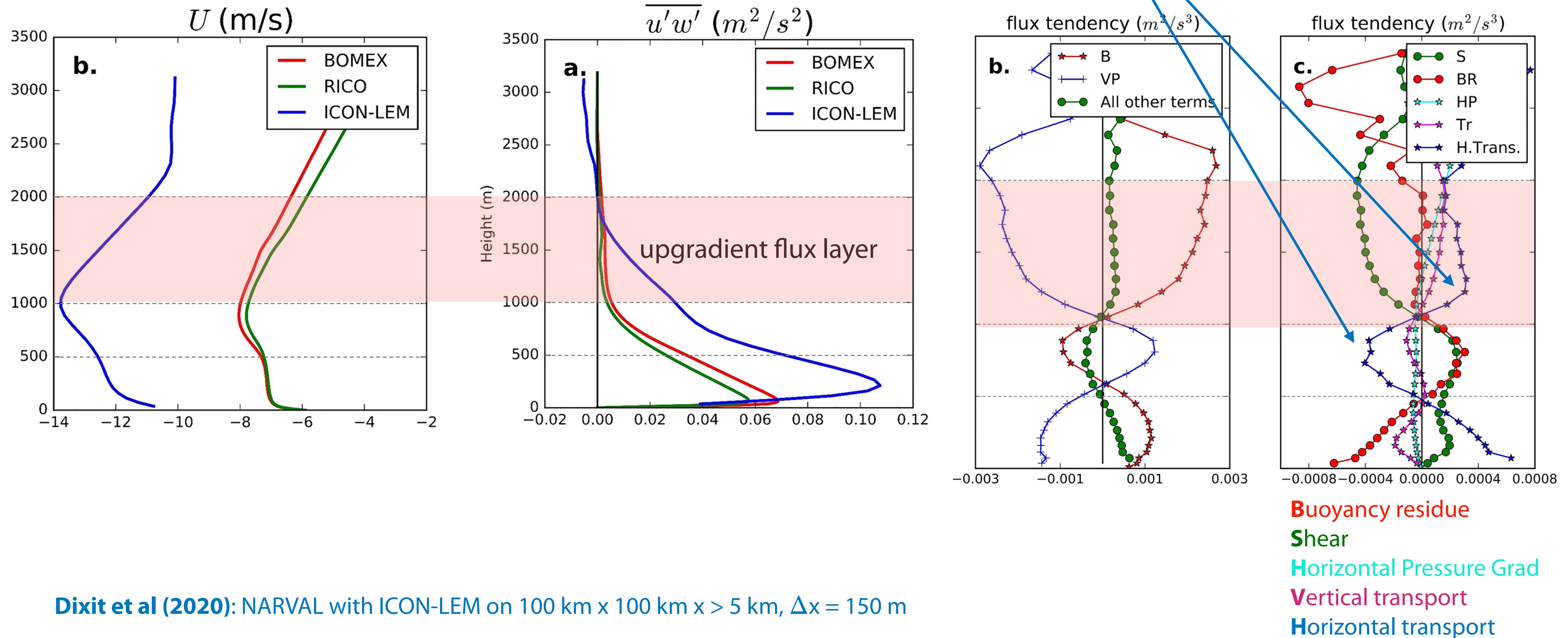


Dixit et al (2020): NARVAL with ICON-LEM on 100 km x 100 km x > 5 km, $\Delta x = 150$ m

$$BR = \frac{g}{T_v} \overline{u'T'_v} - \frac{\overline{u' \partial p'}}{\bar{\rho} \partial z}$$

Buoyancy **Vertical Pressure Gradient**

Upgradient flux generated by horizontal convergence of flux by wind (perturbations)

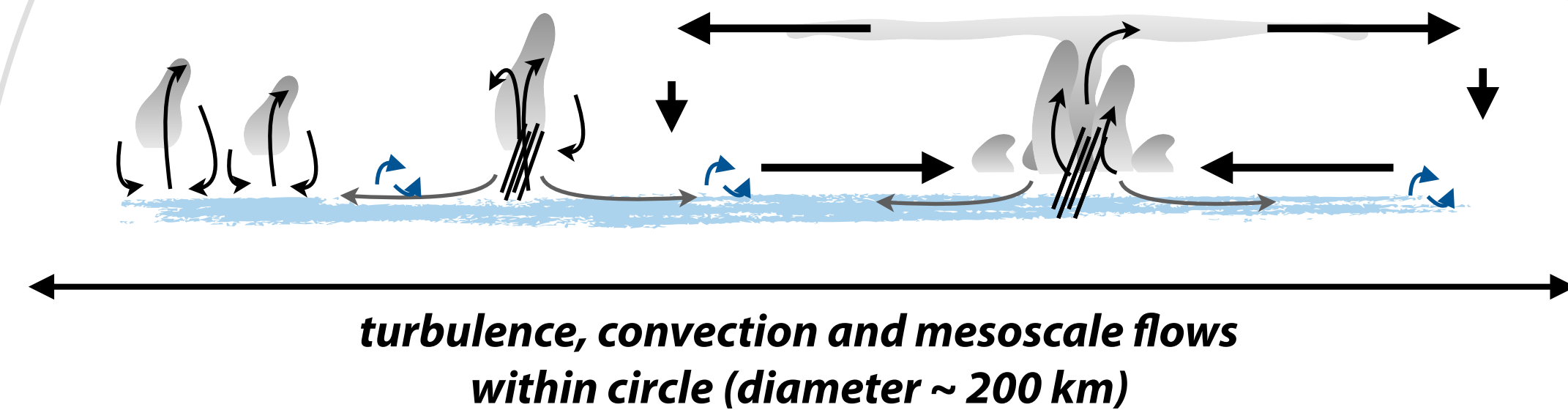
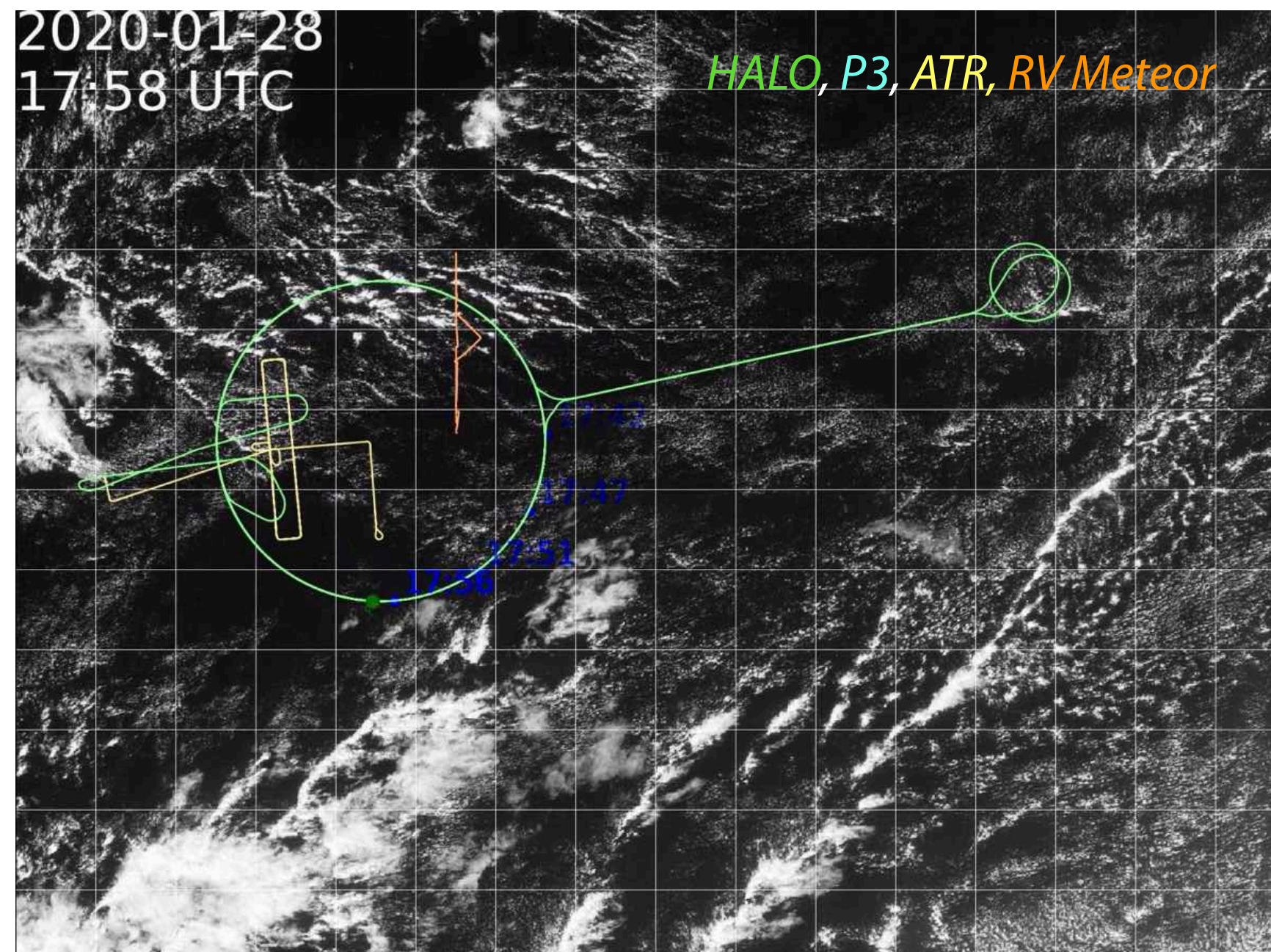


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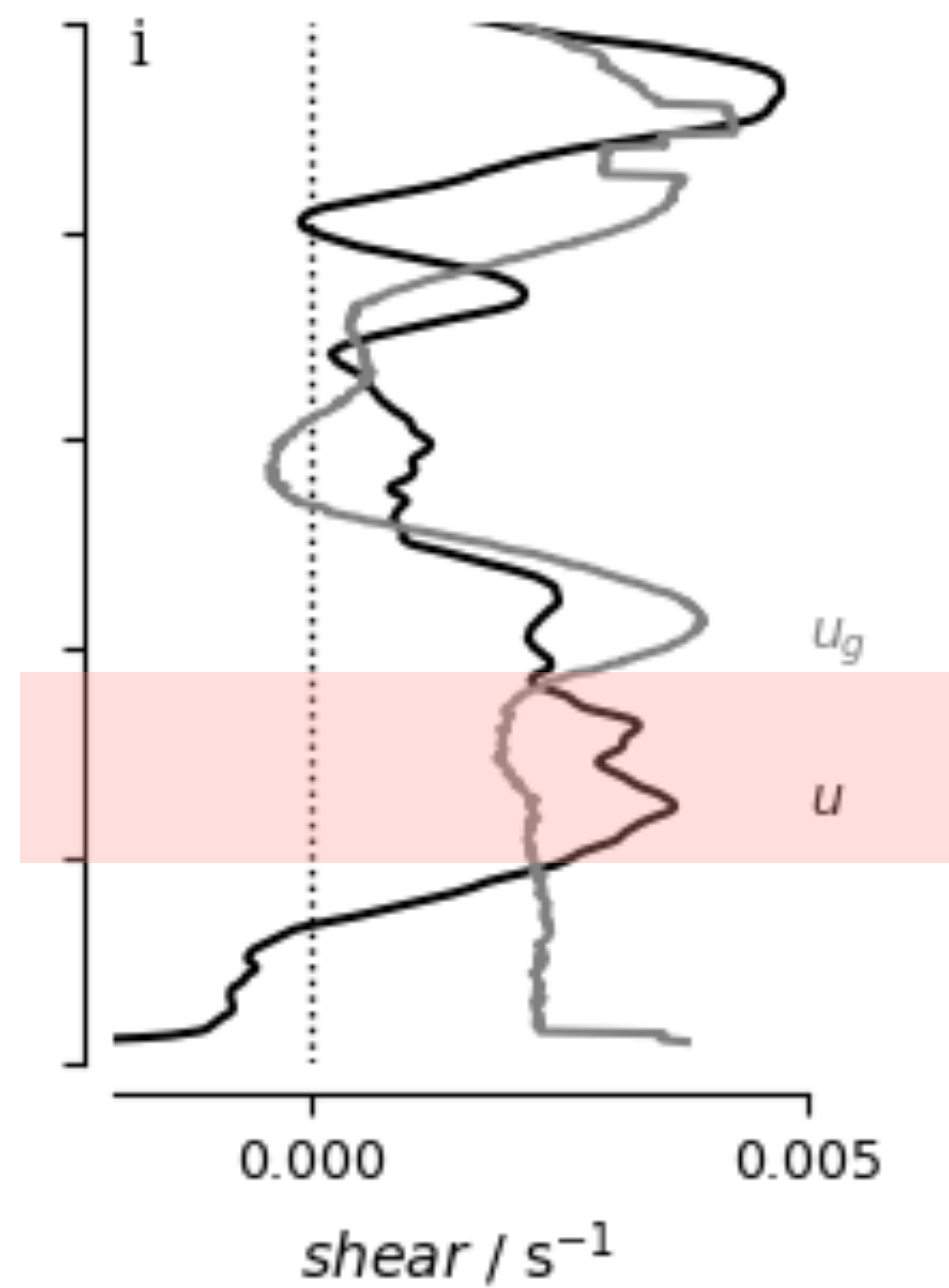
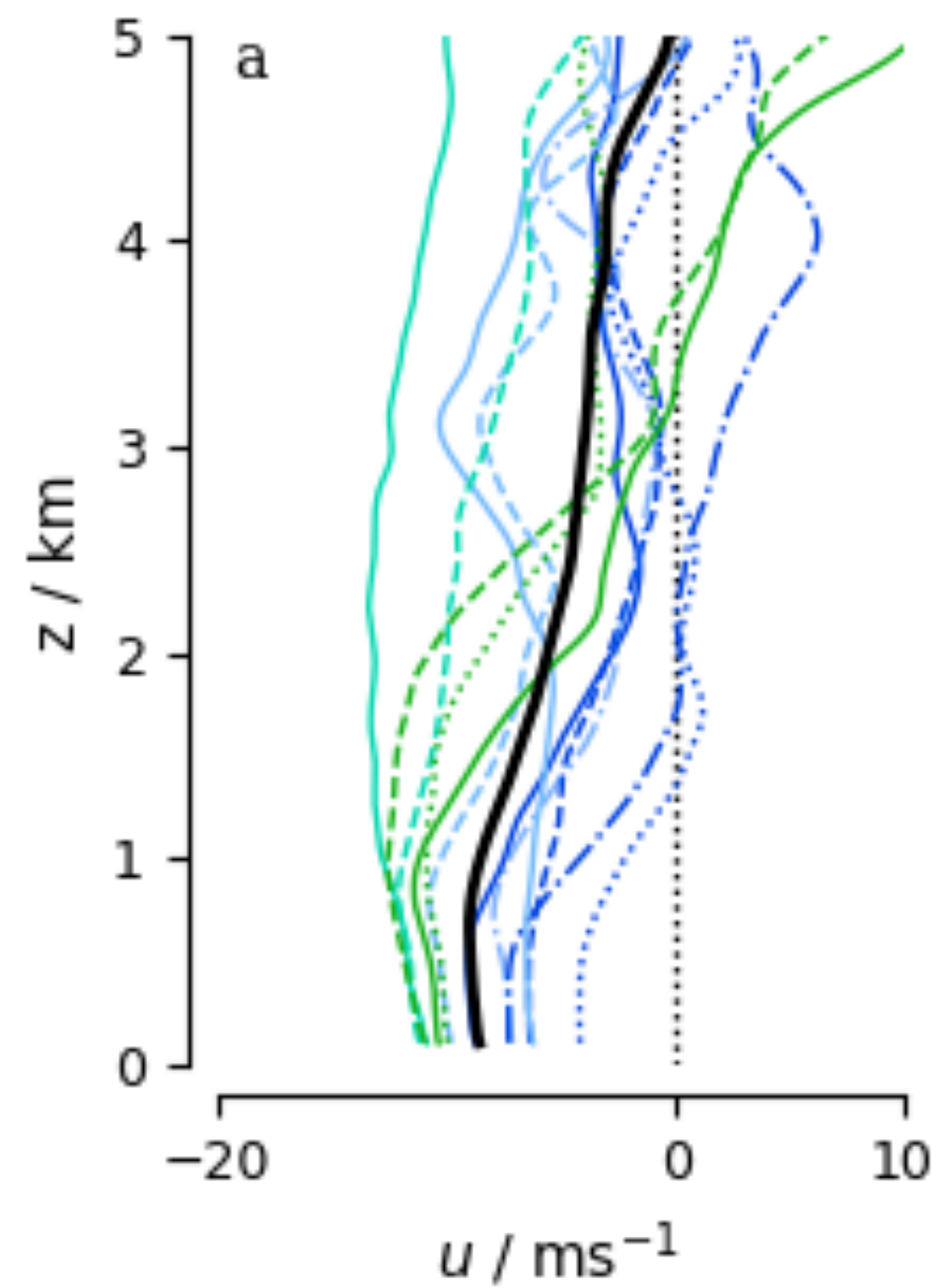
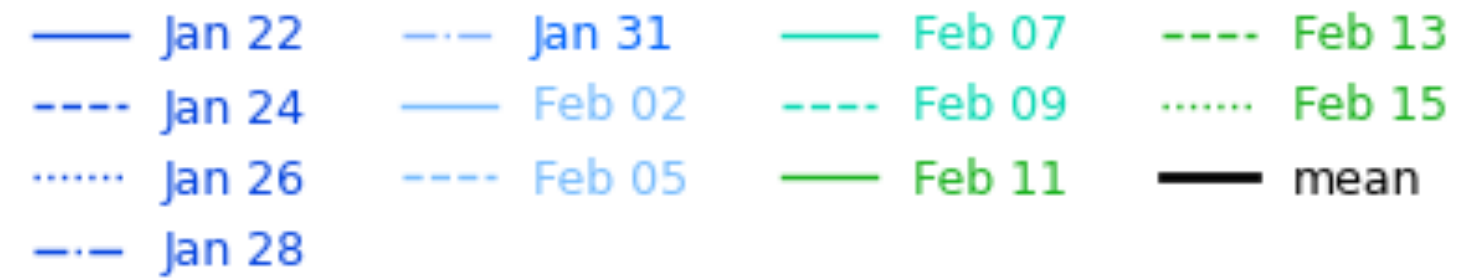
EUREC4A 's rich data set and hierarchy of model simulations are used to explore momentum transport by different flows



- * **JOANNE: circular dropsonde arrays** (85 circles, 13 flight days): meso-scale divergence, pressure gradients and geowind
- * **French ATR Safire aircraft / UAVs:** profiles of in-situ turbulence

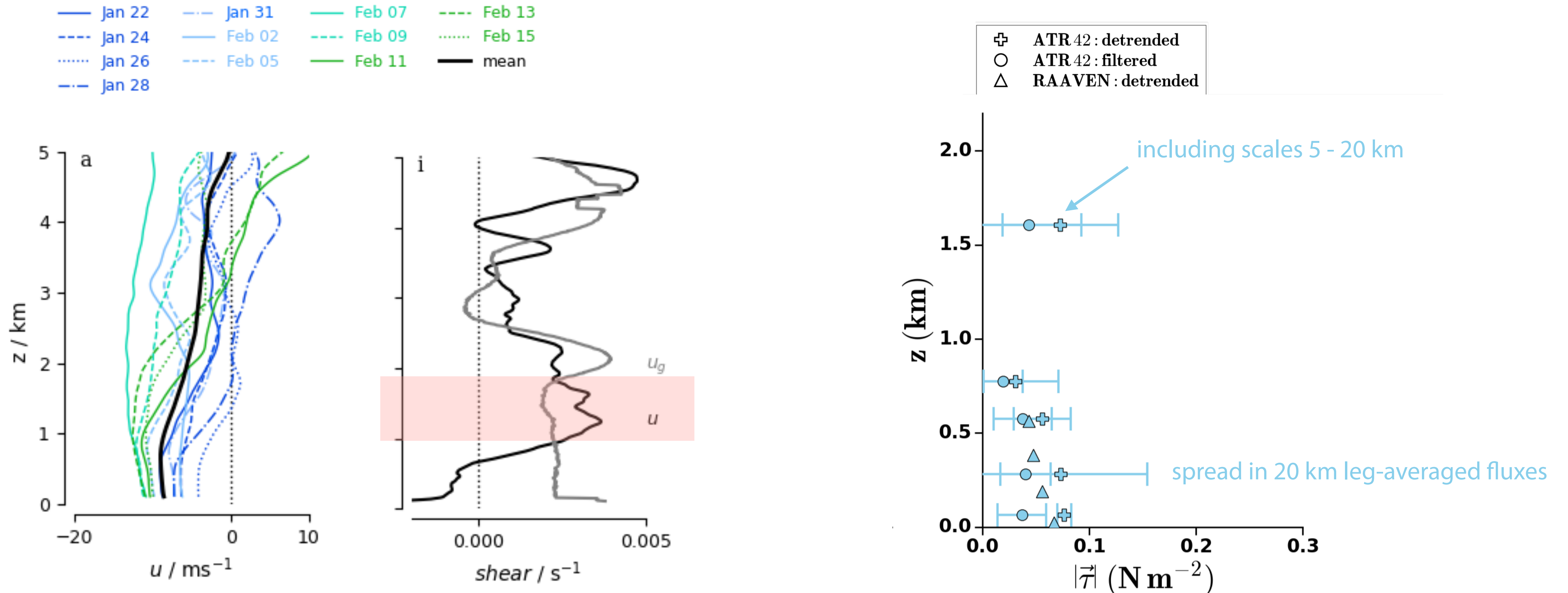


Supportive evidence of upgradient momentum transport and heterogeneity in flux along flight tracks, including near cloud tops

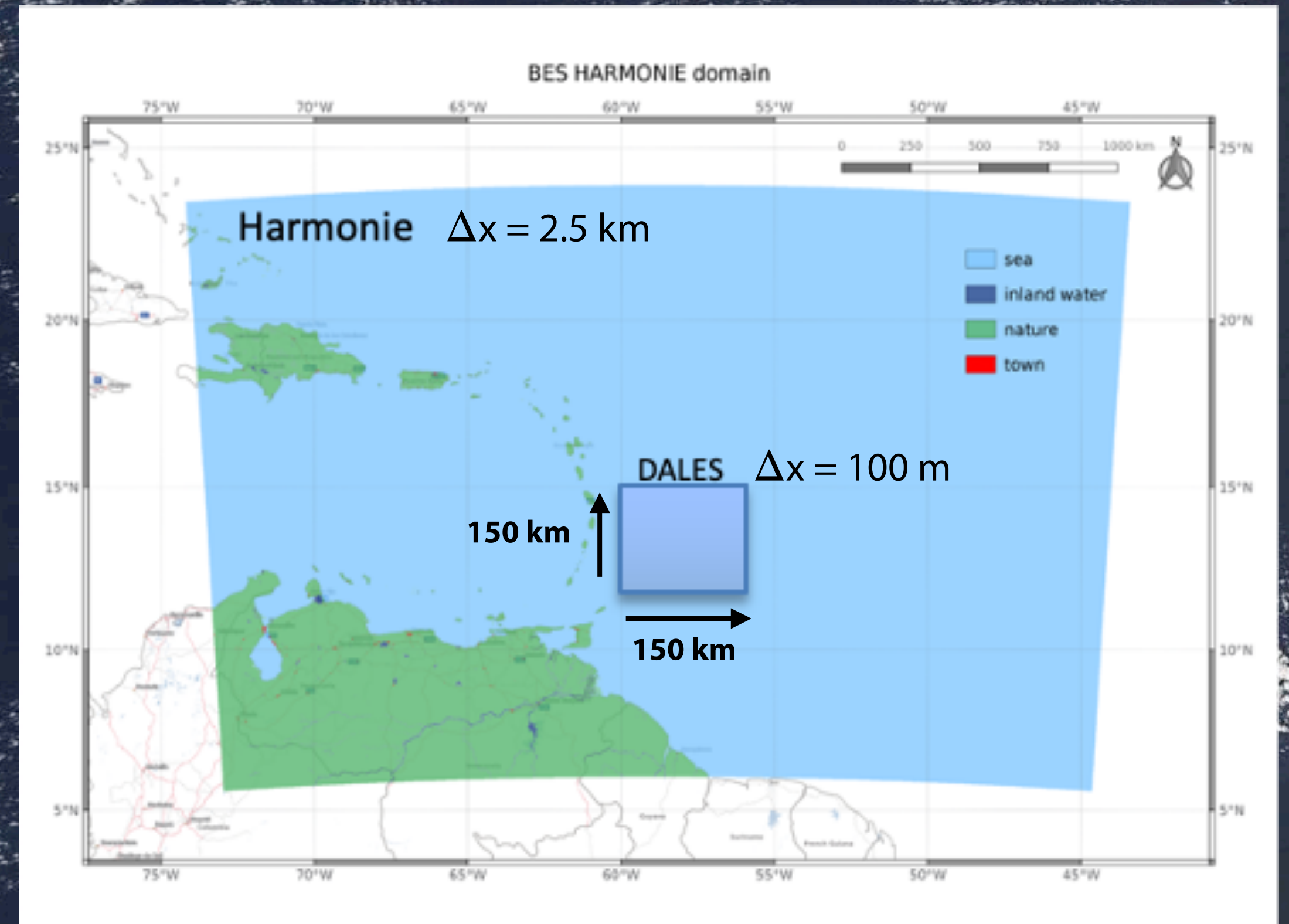


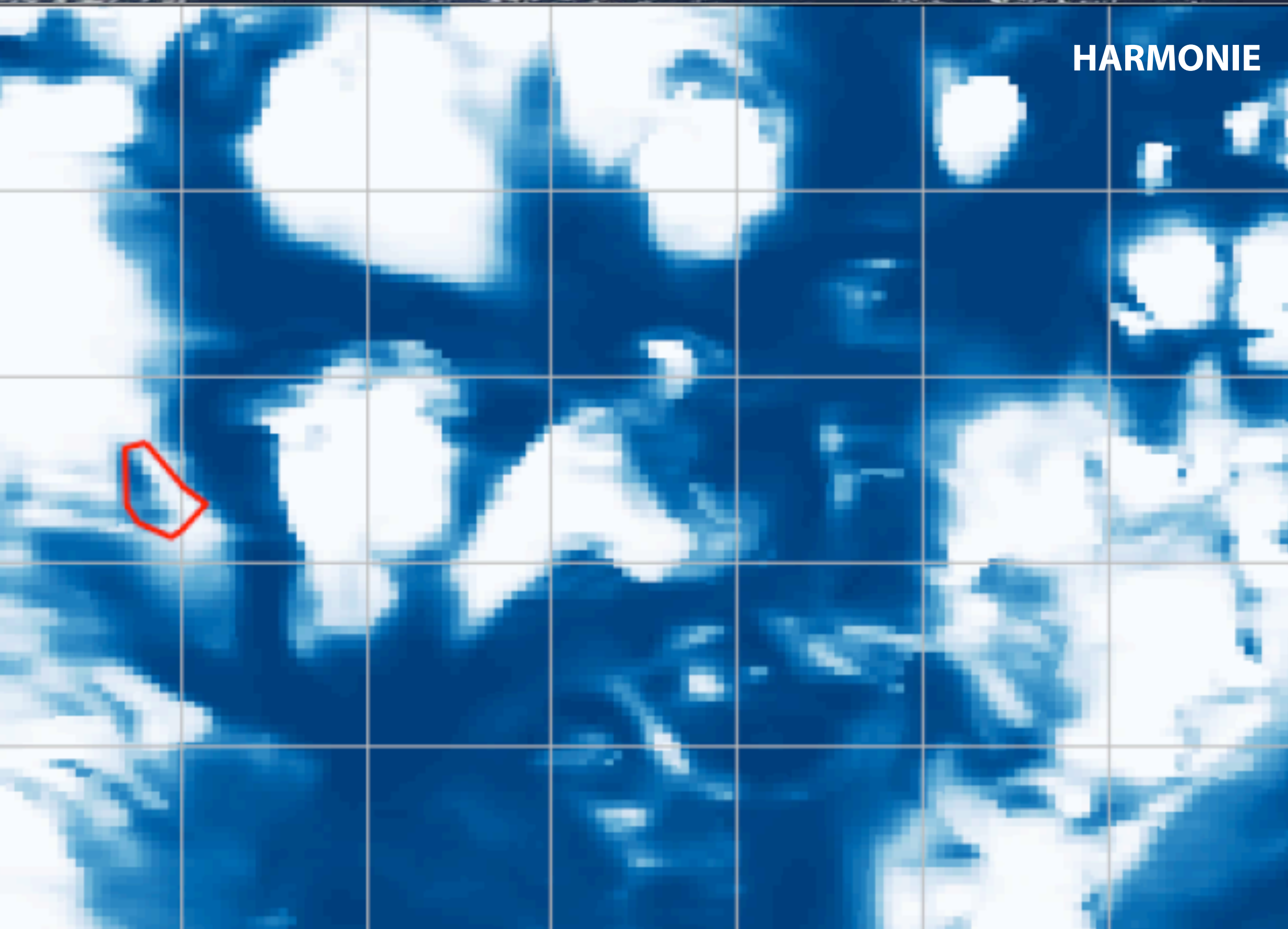
actual wind shear > shear in geo-wind

Supportive evidence of upgradient momentum transport and heterogeneity in flux along flight tracks, including near cloud tops

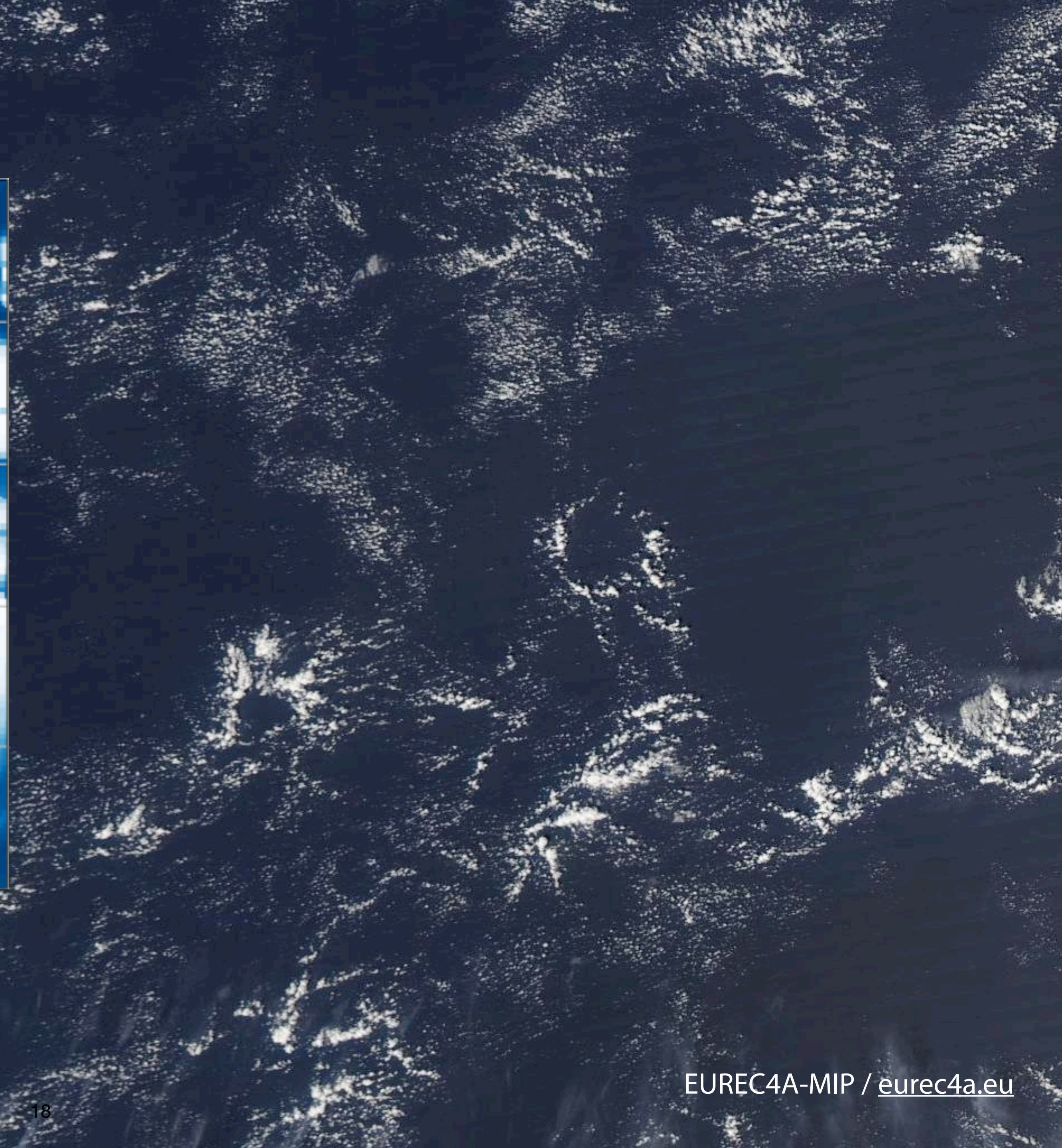
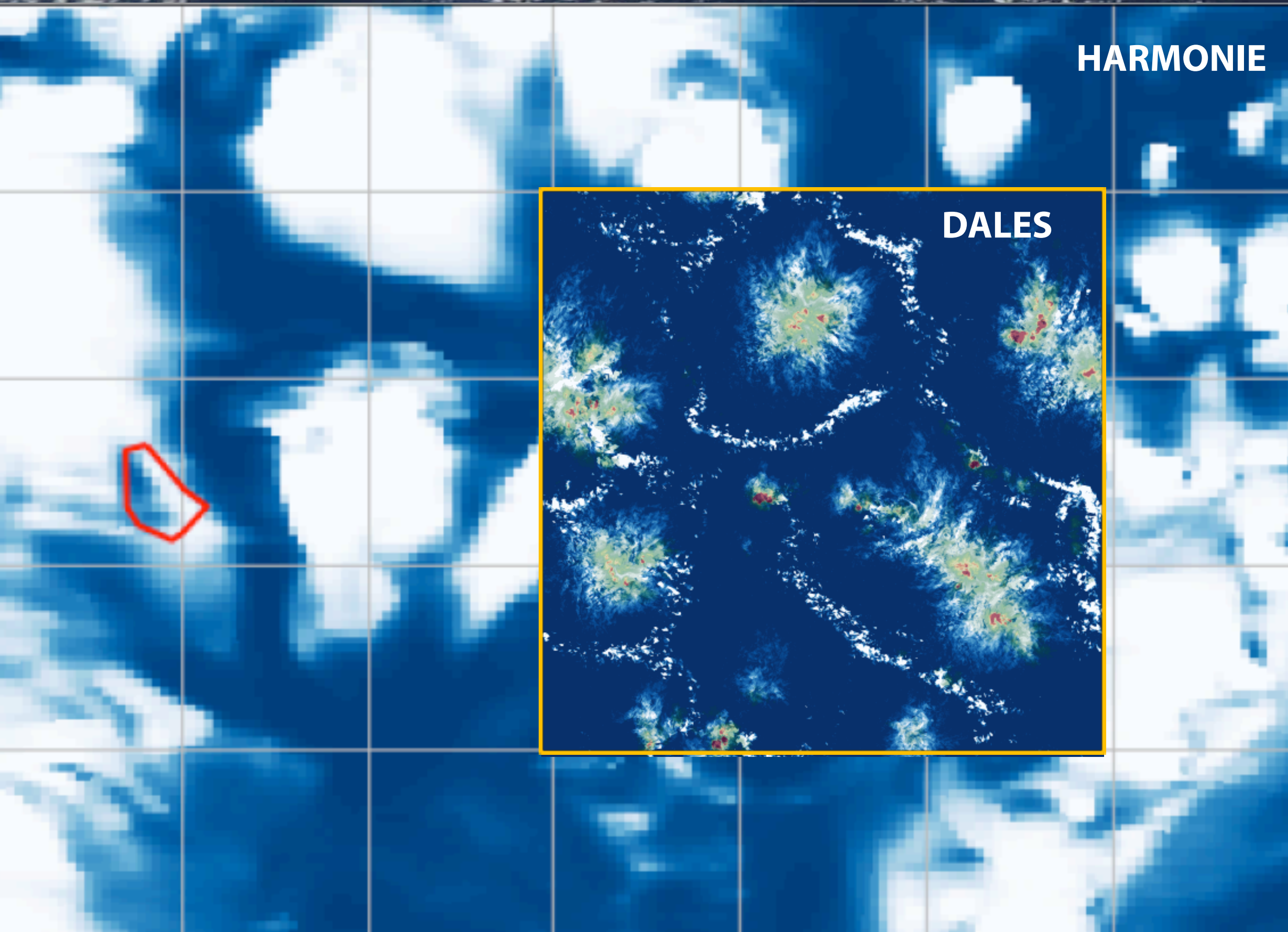


LEM and SRM hindcasts/climatological simulations of a ten-day EUREC4A period

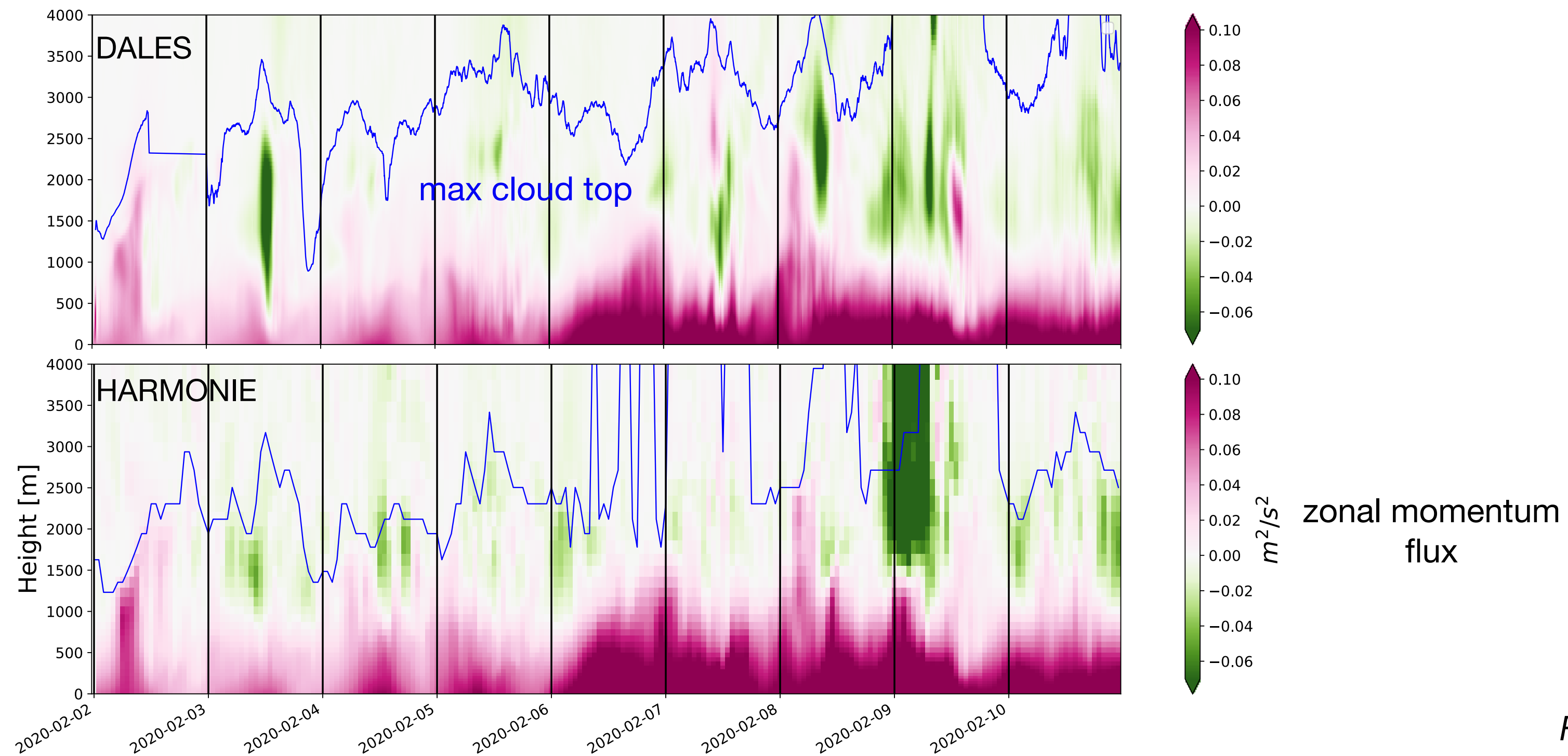
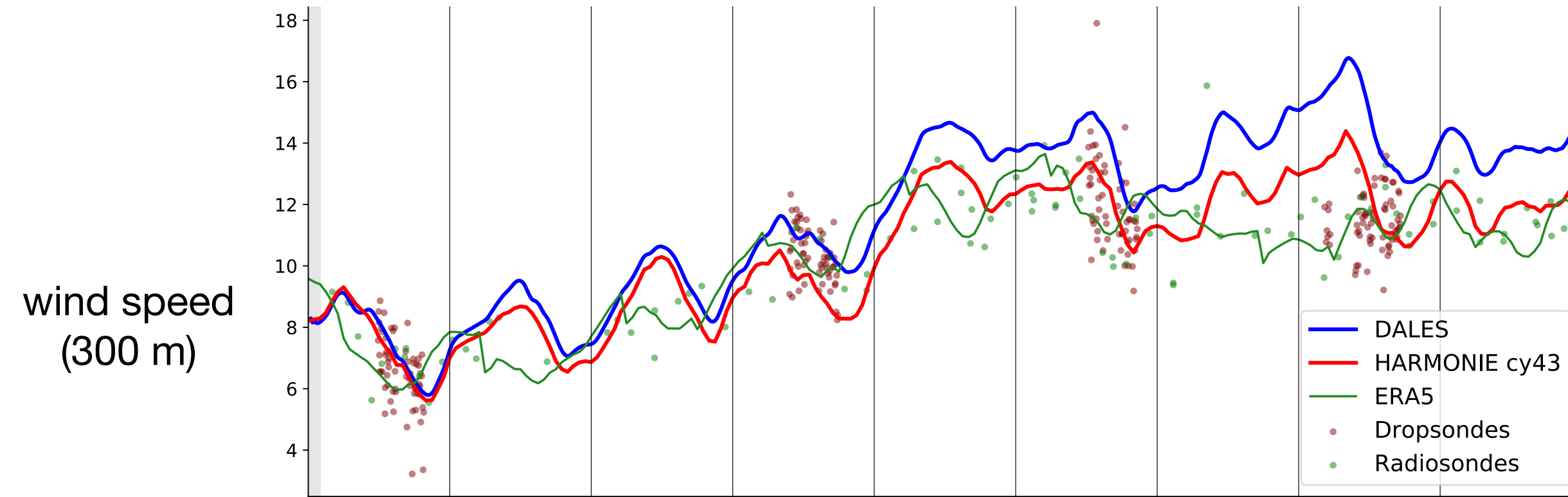




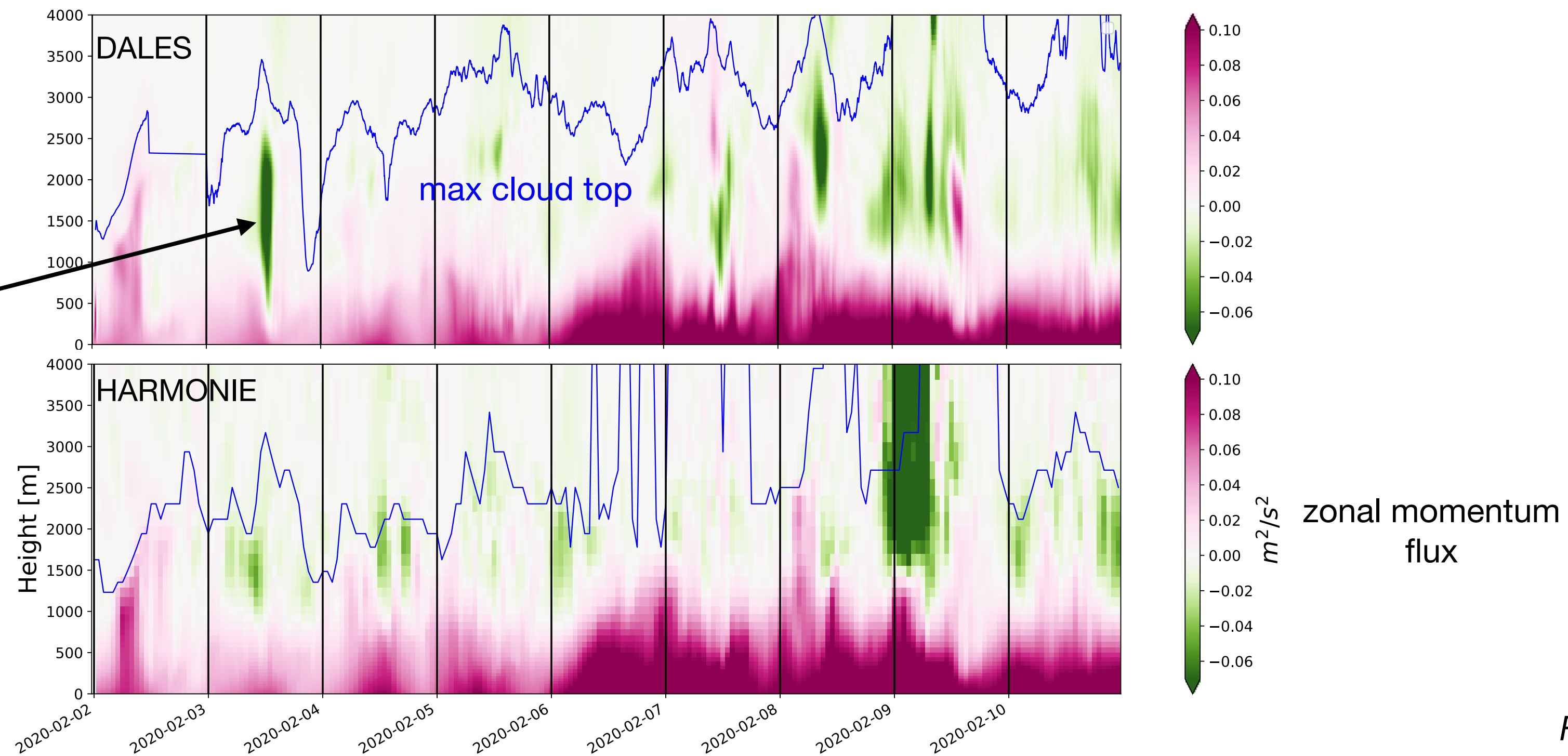
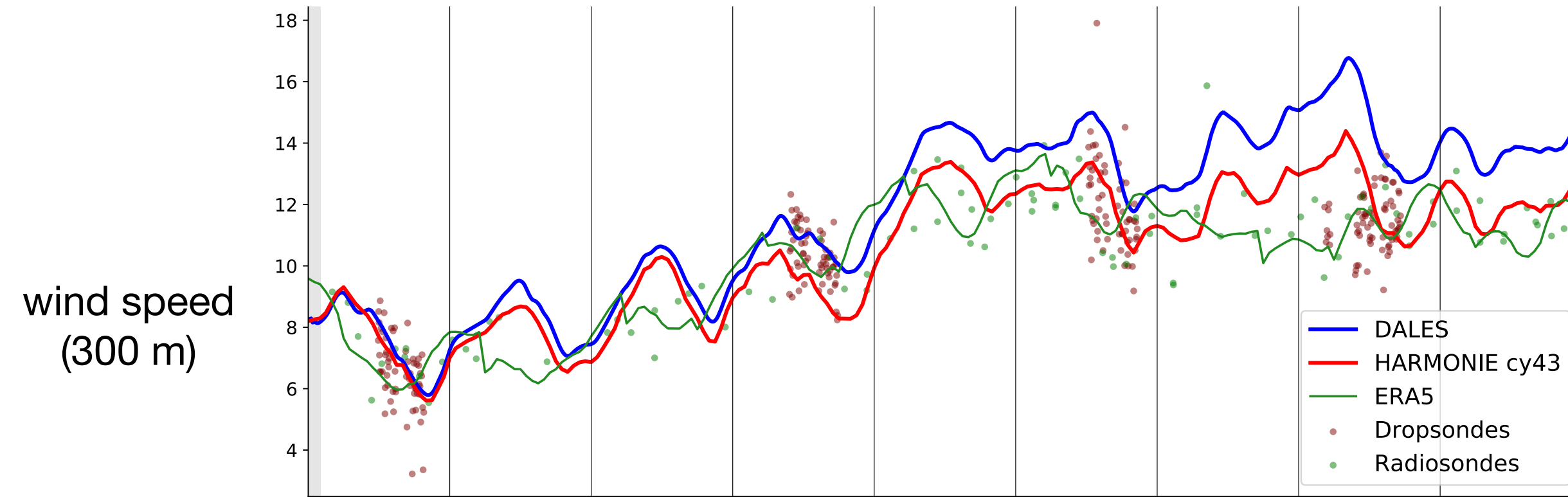
February 4, 2020 - 14 UTC



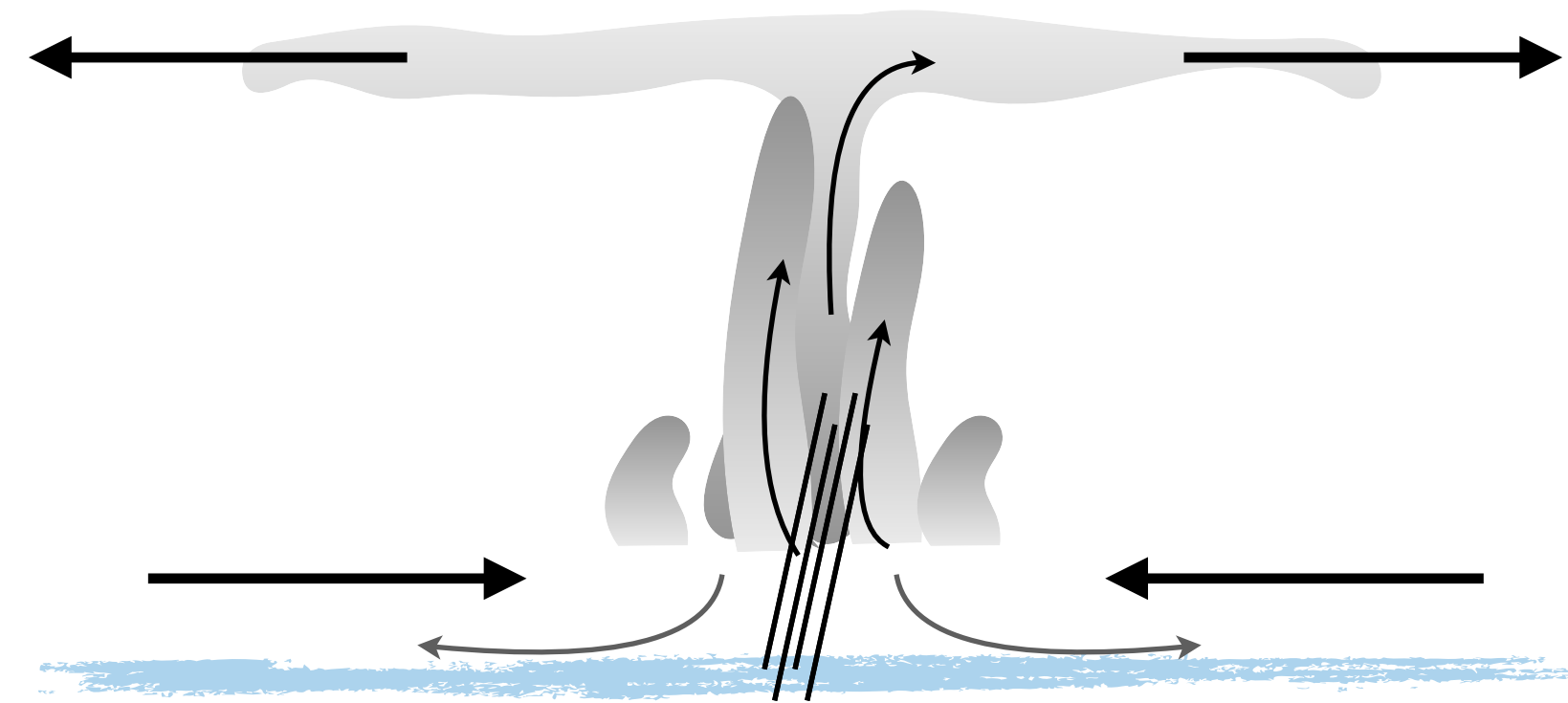
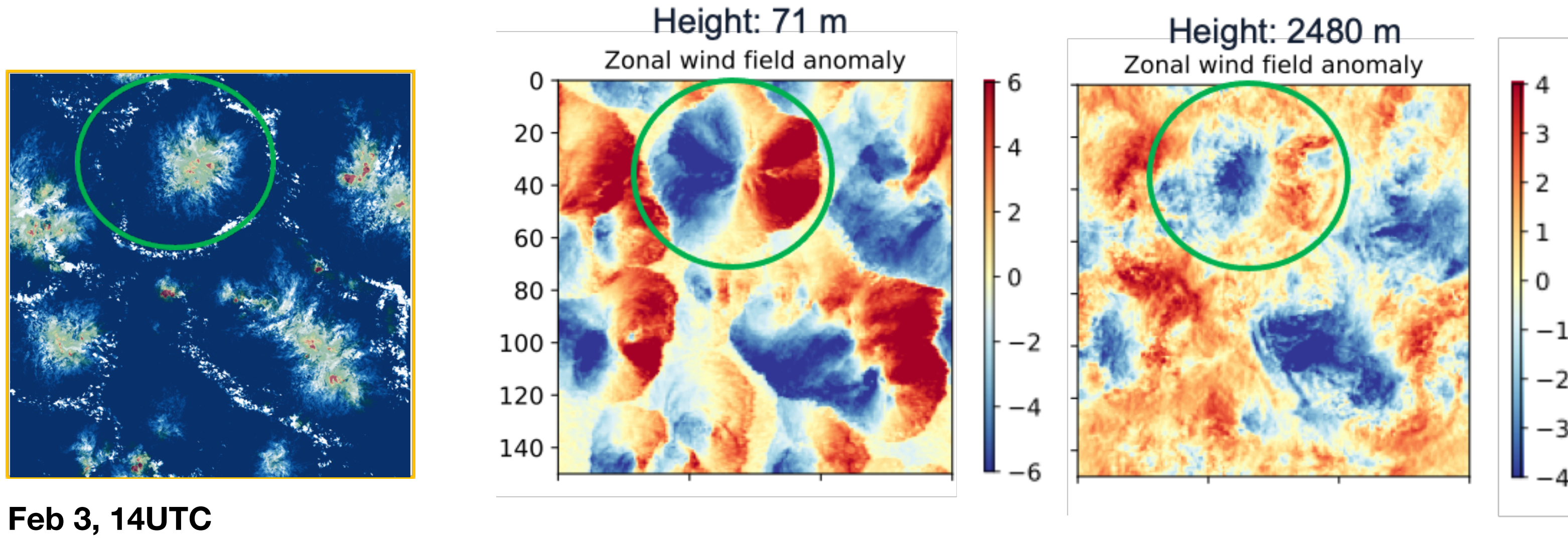
As trade-winds strengthen, convection drives strong (upgradient) momentum flux



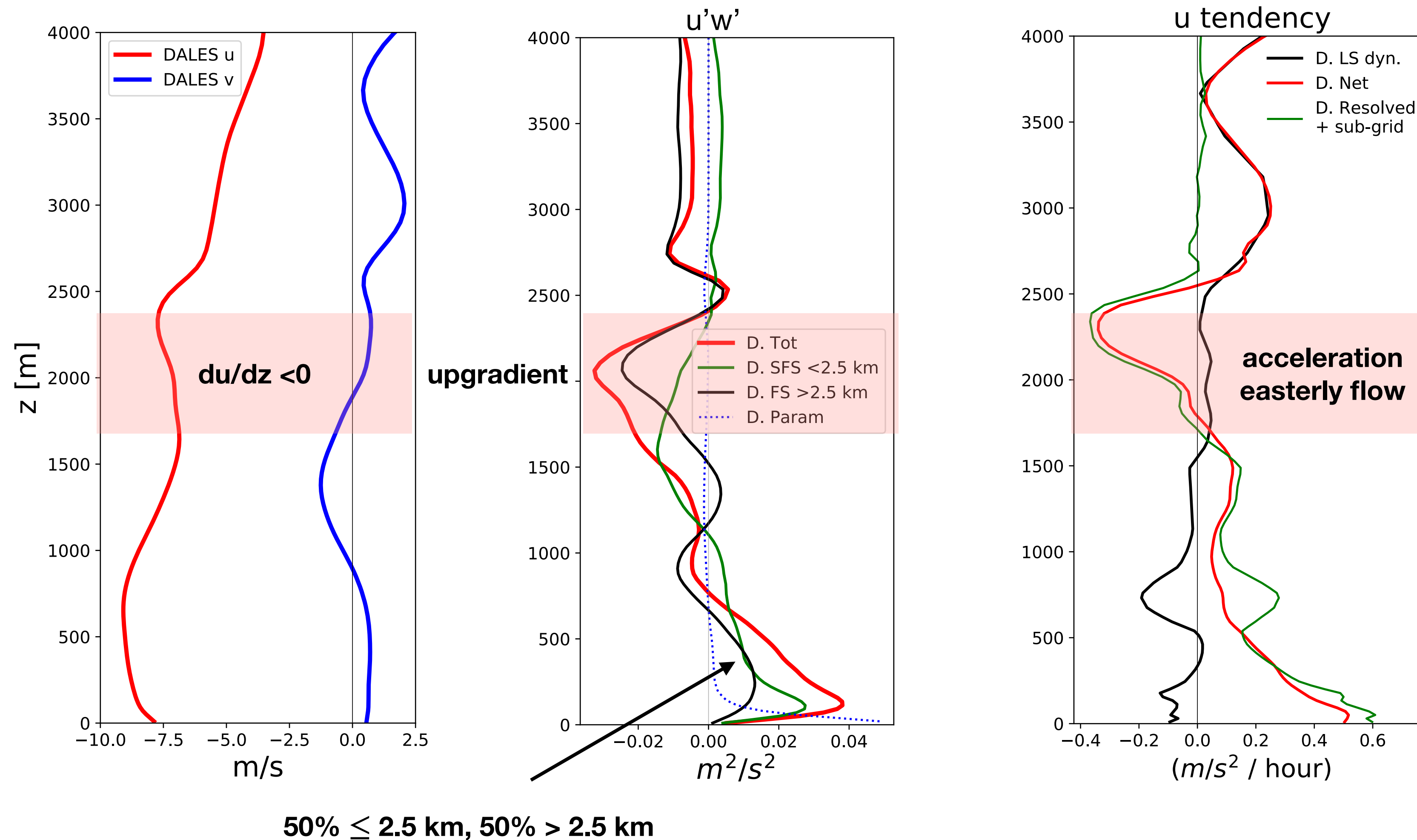
As trade-winds strengthen, convection drives strong (upgradient) momentum flux



Large anomalies in wind speed in surface layer and near cloud tops

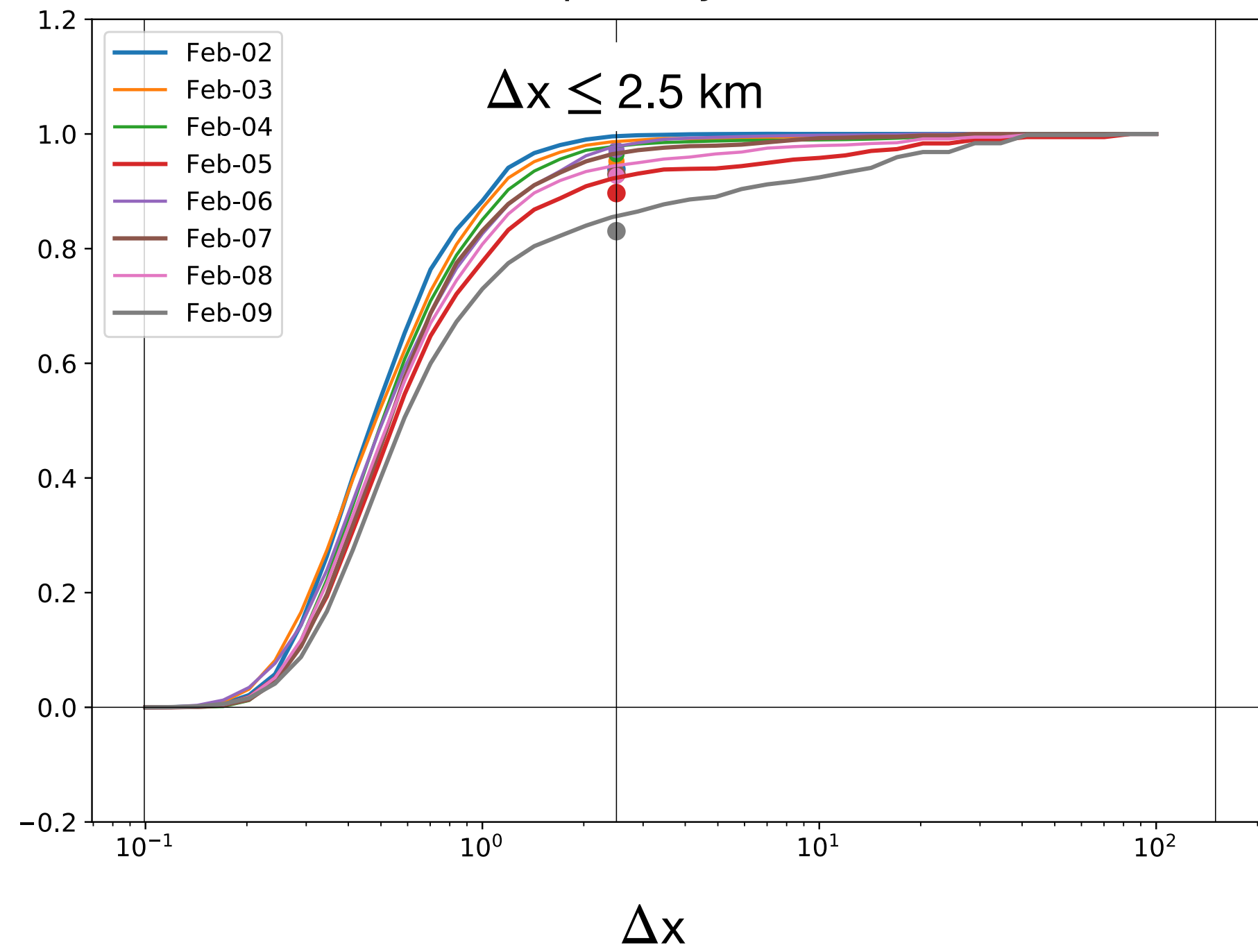


Mesoscale flows important for upgradient transport and acceleration near cloud tops, but also for generating flux in the mixed-layer

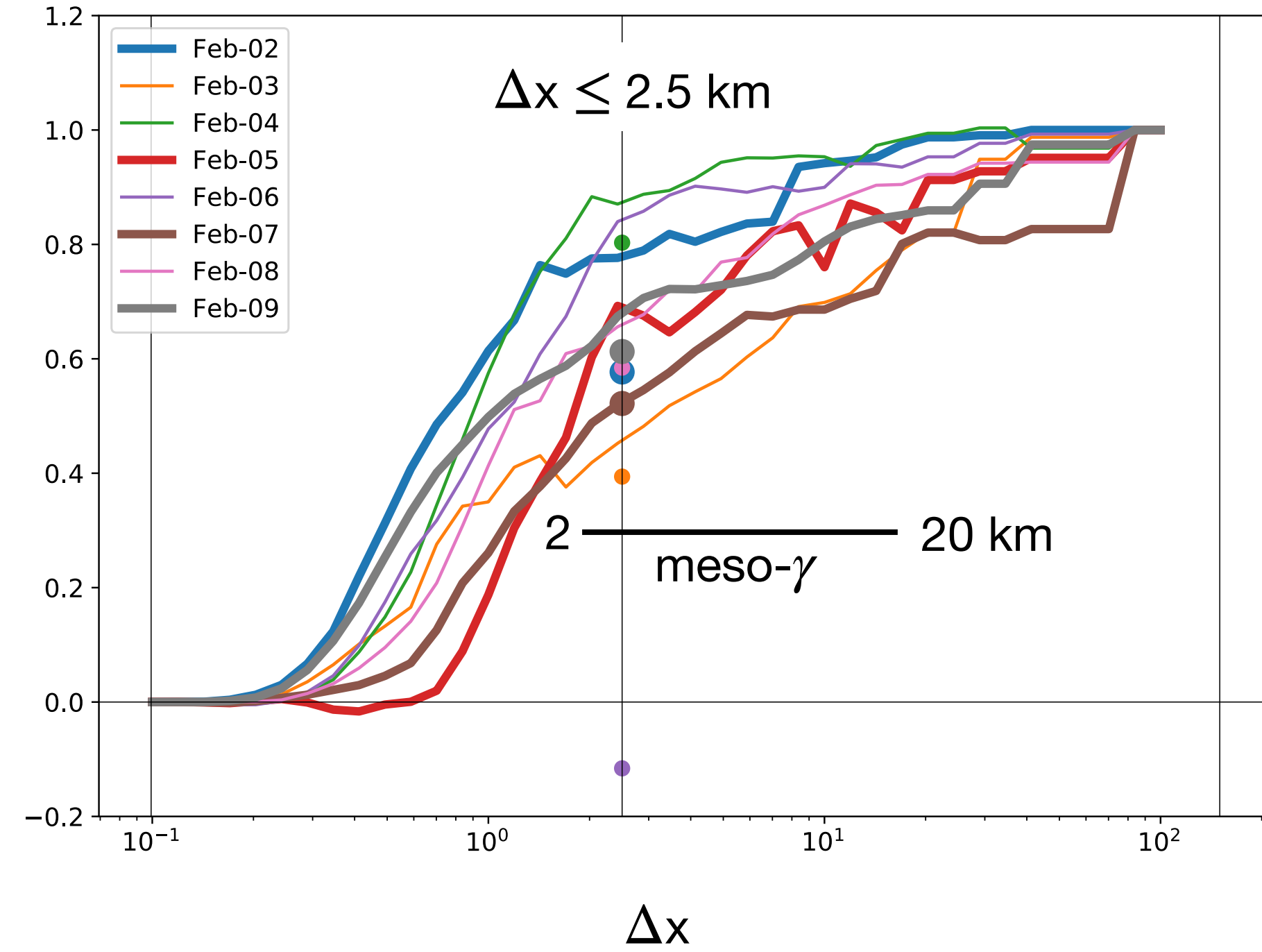


(Sub)mesoscales can contribute up to 50% of the momentum flux in the cloud layer

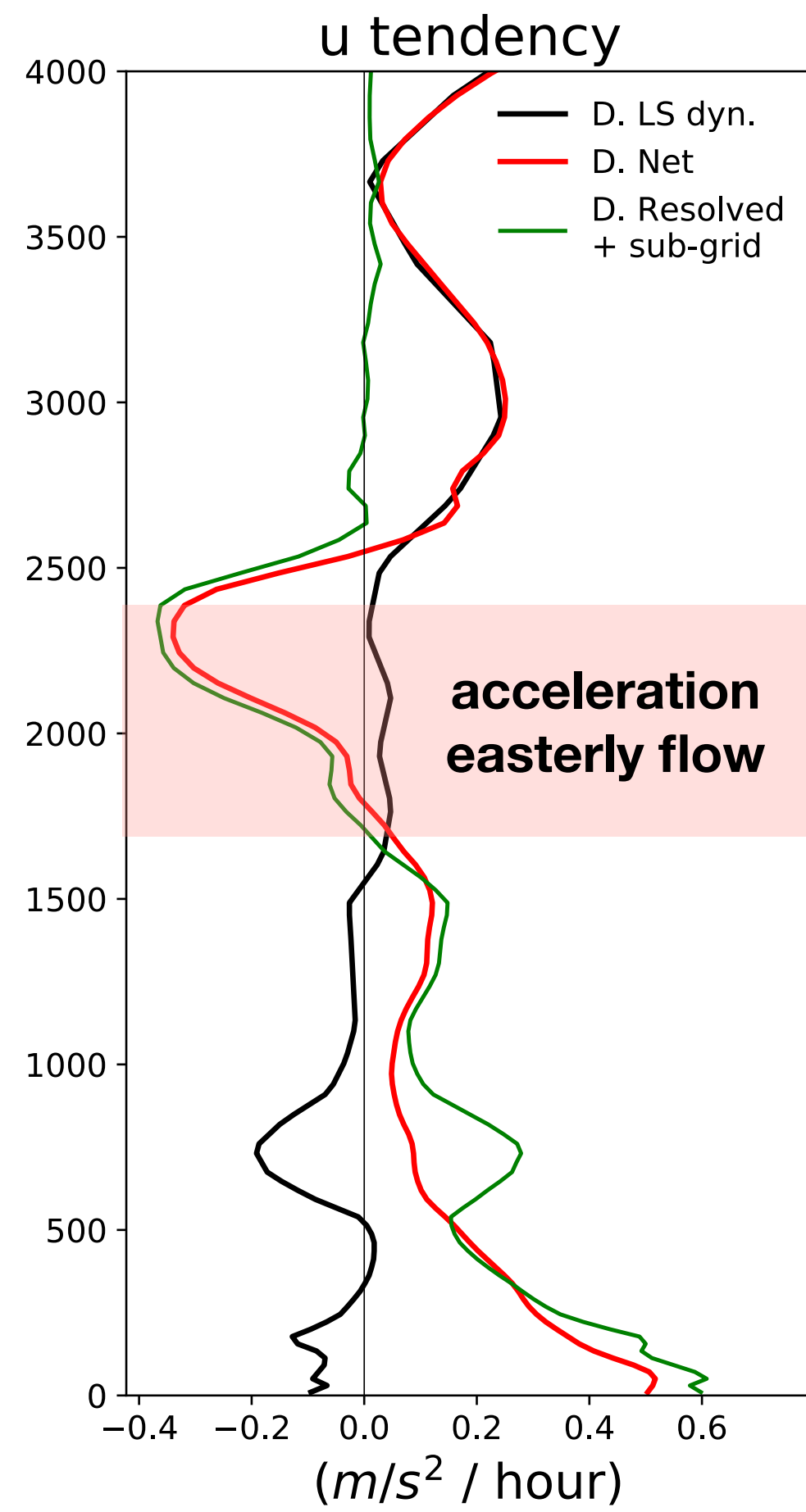
Contribution of grid scale to total momentum flux
At 200 m



At 1.5 km



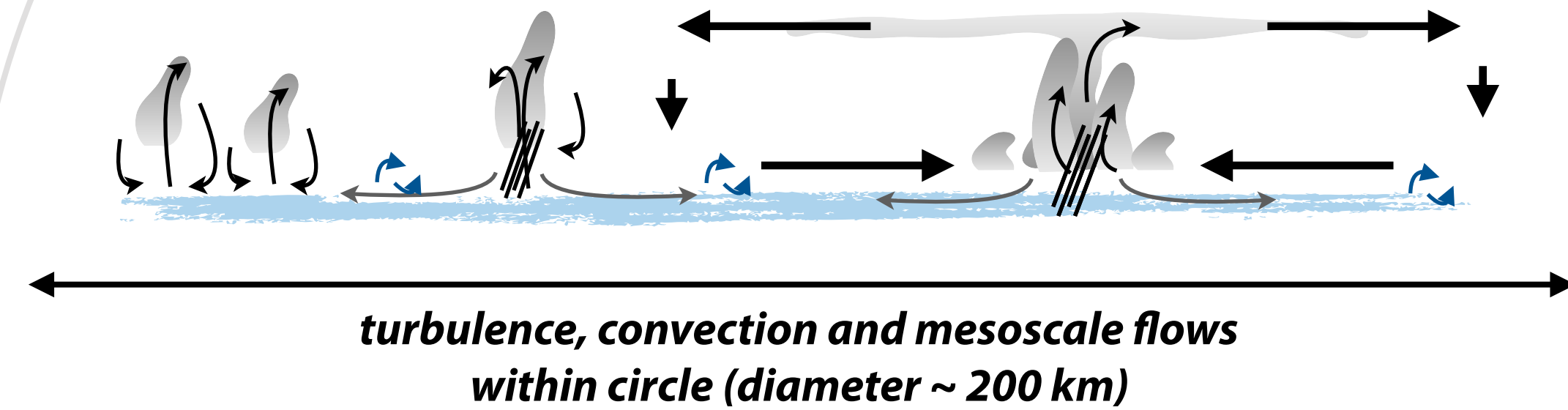
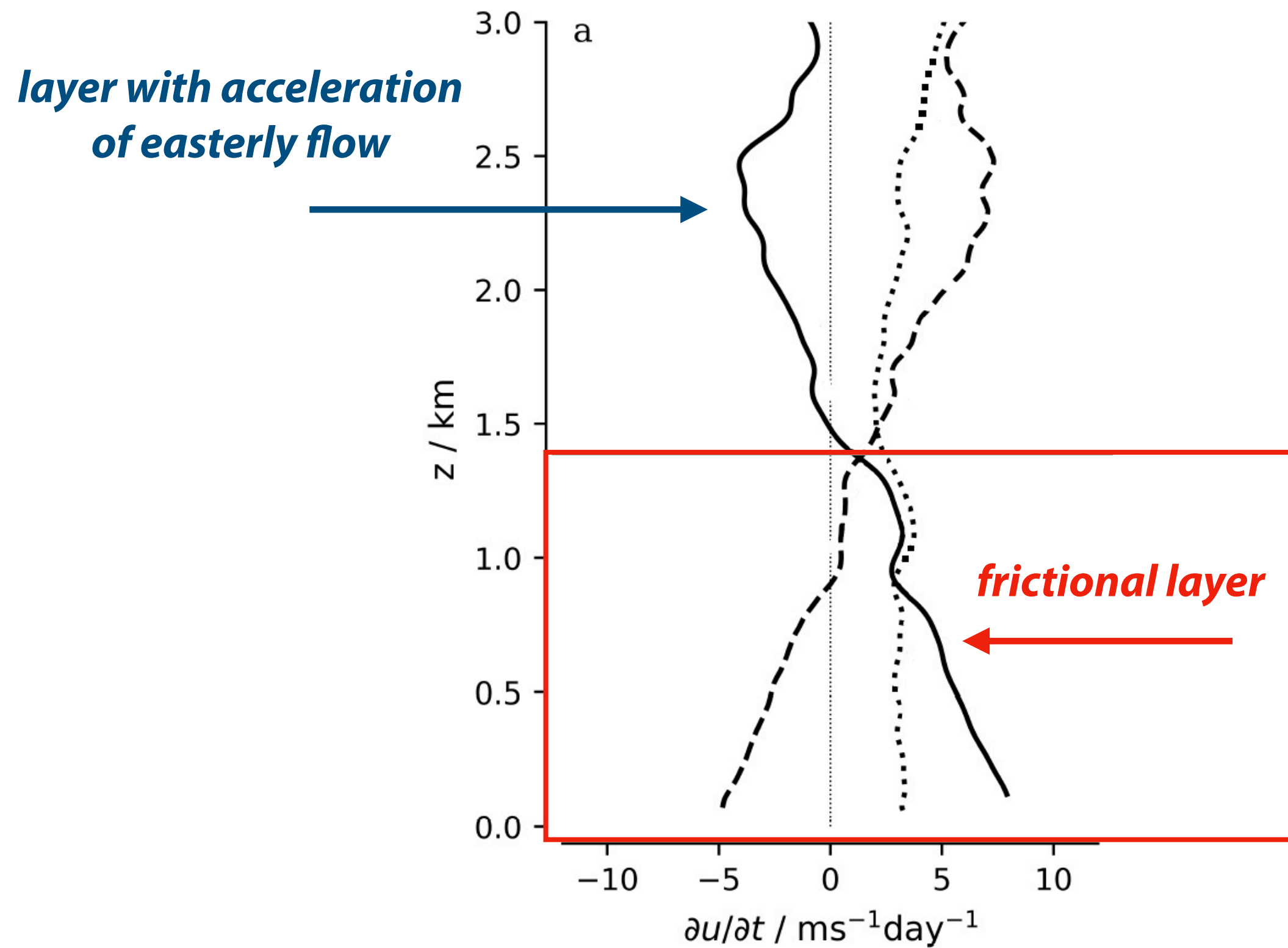
Link to bias in zonal winds?



Convective, mesoscale flows create a deeper layer of easterly flow



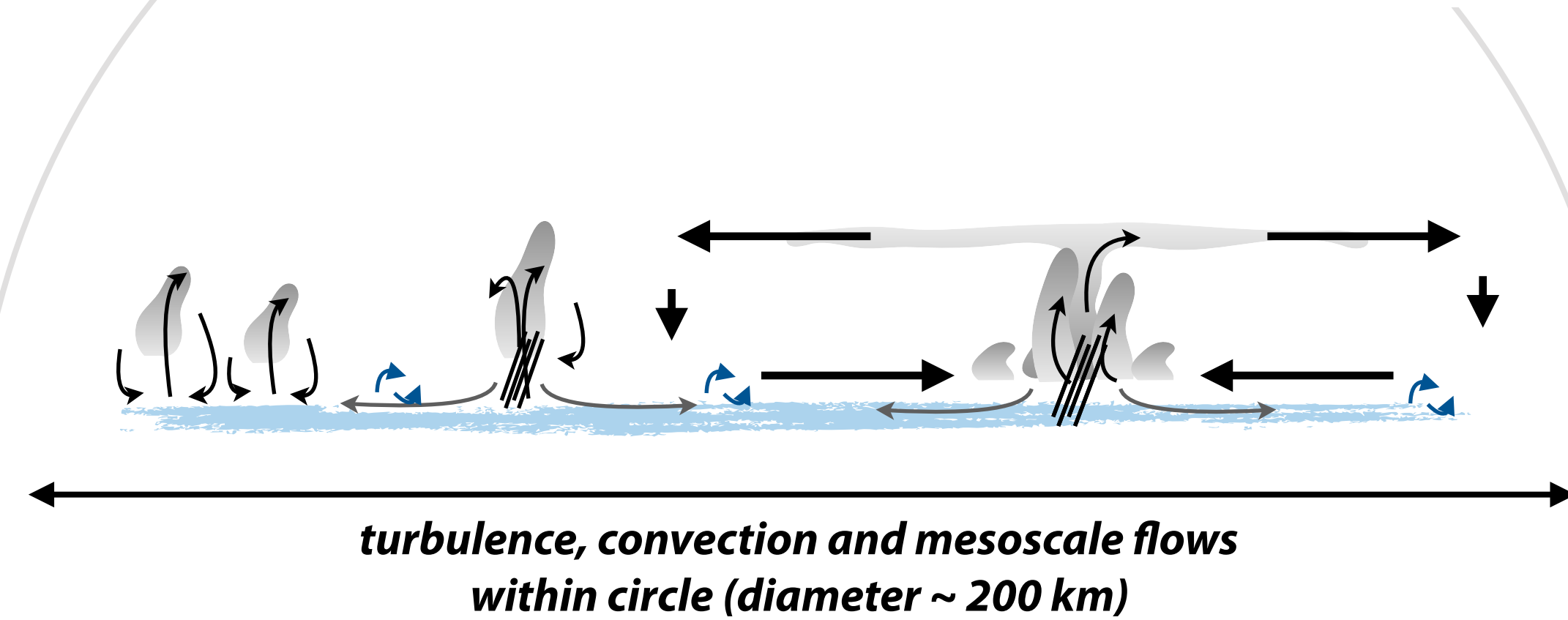
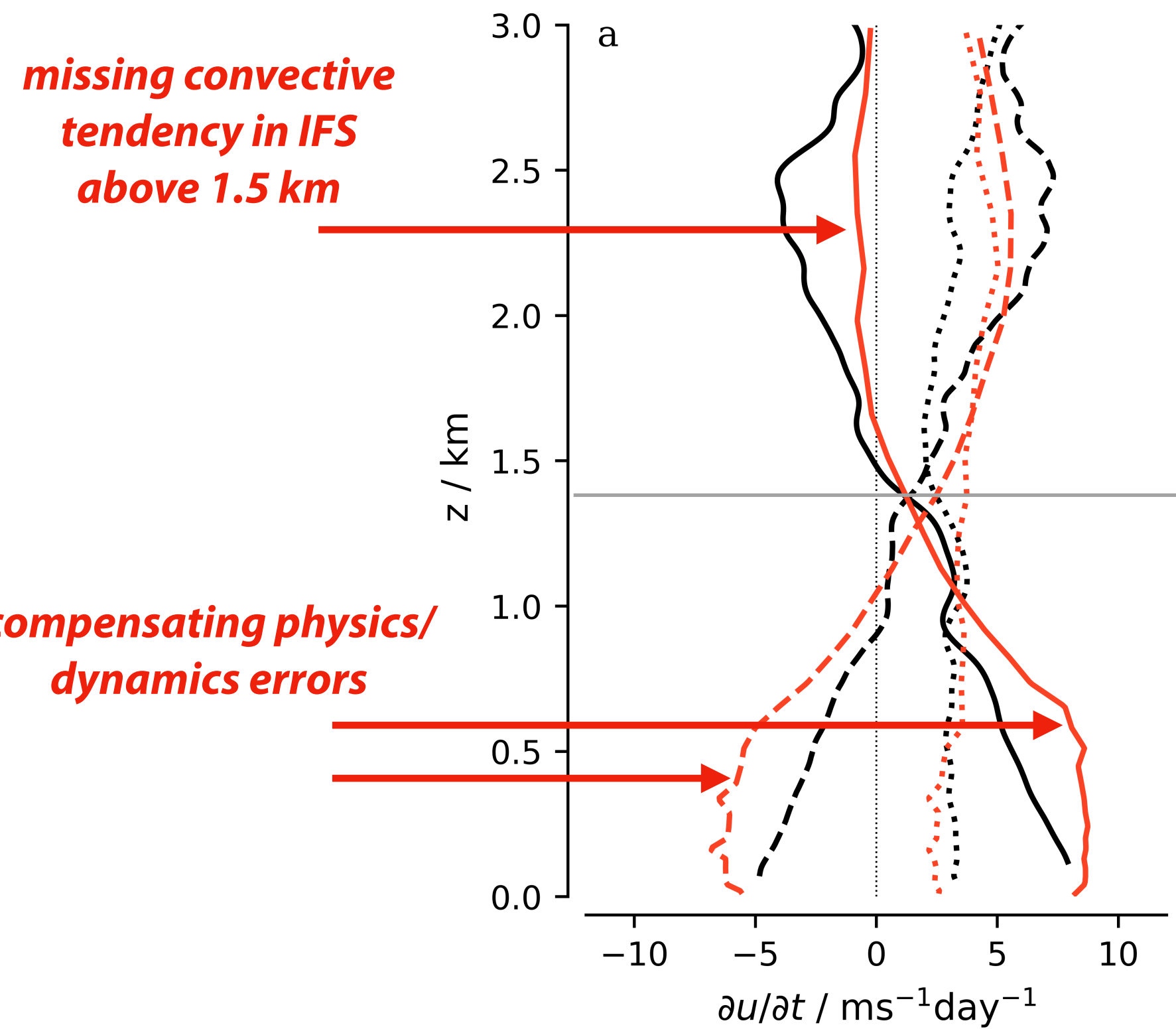
- JOANNE / temporal change
- - - JOANNE / large-scale dynamics
- JOANNE / residual (F_u, F_v) ~ friction



Does organised convection help explain the weak zonal wind bias ?

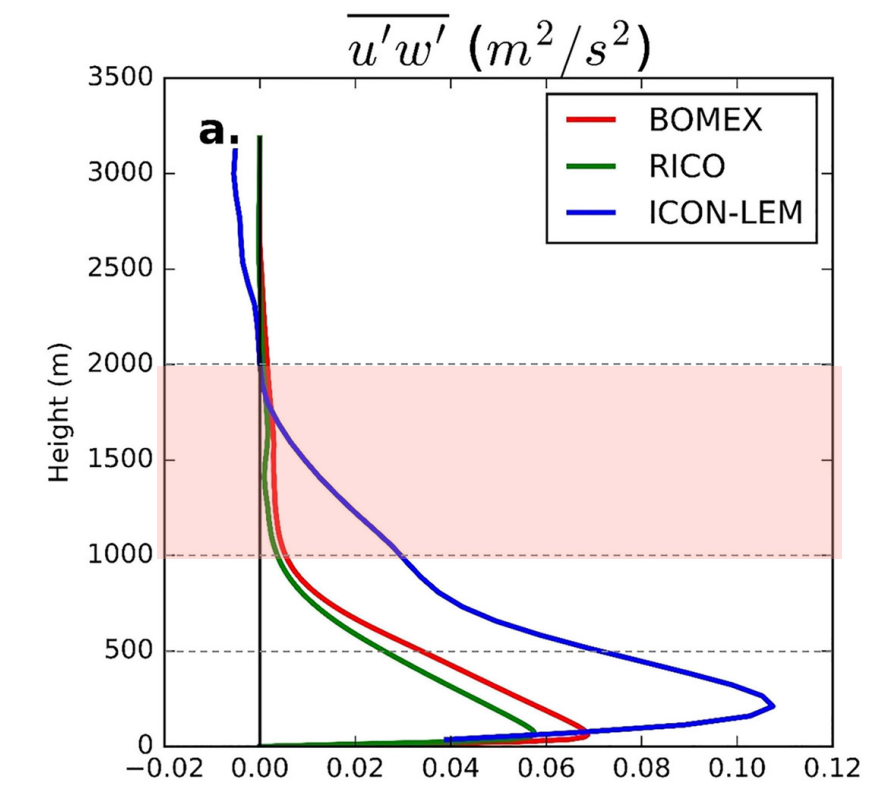


- JOANNE / temporal change
- - - JOANNE / large-scale dynamics
- JOANNE / residual (F_u, F_v)
- IFS / temporal change
- - - IFS / large-scale dynamics
- IFS / turbulence + convection



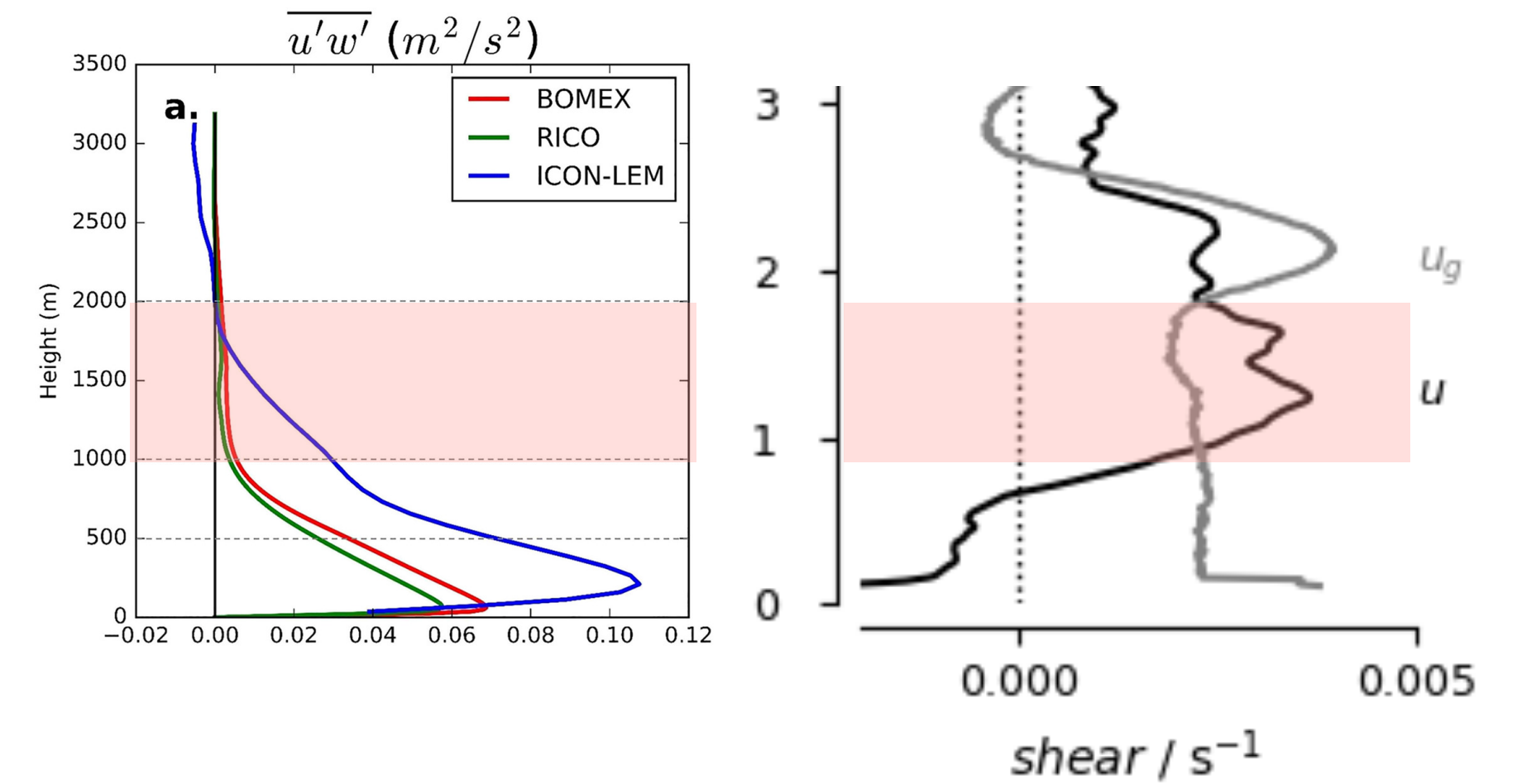
Summary

- ✦ In larger domain LEM with open boundary conditions and varying large-scale tendencies, horizontal transport and generation of momentum flux gains importance: more upgradient transport than perhaps appreciated



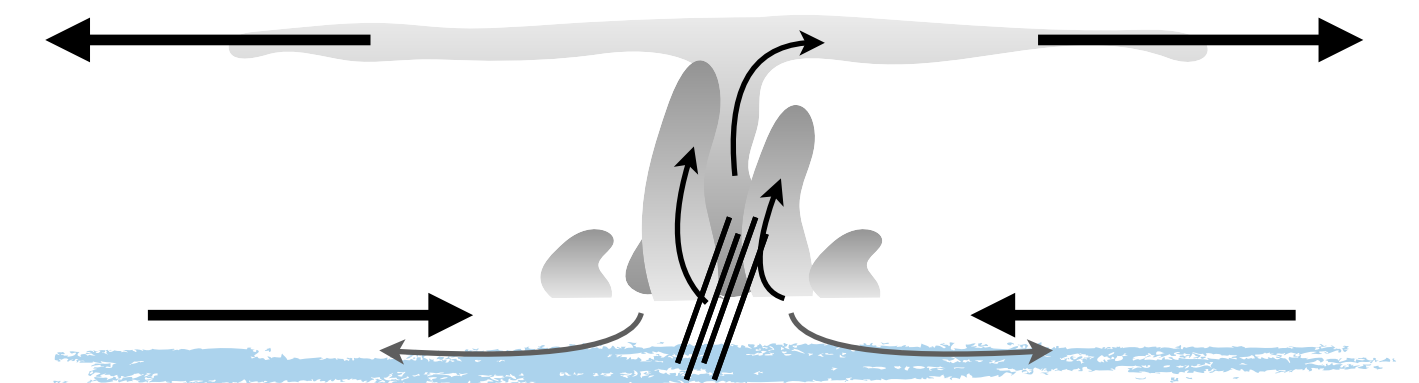
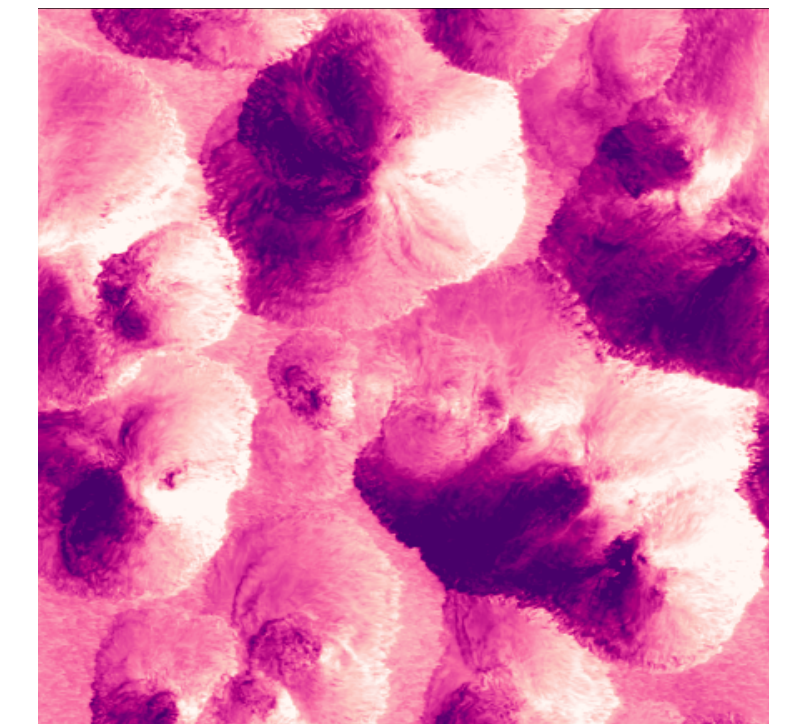
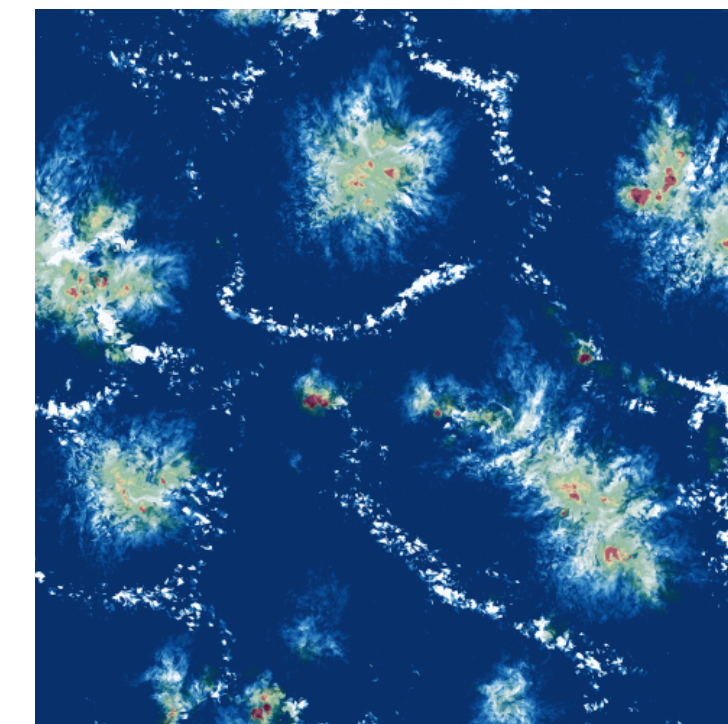
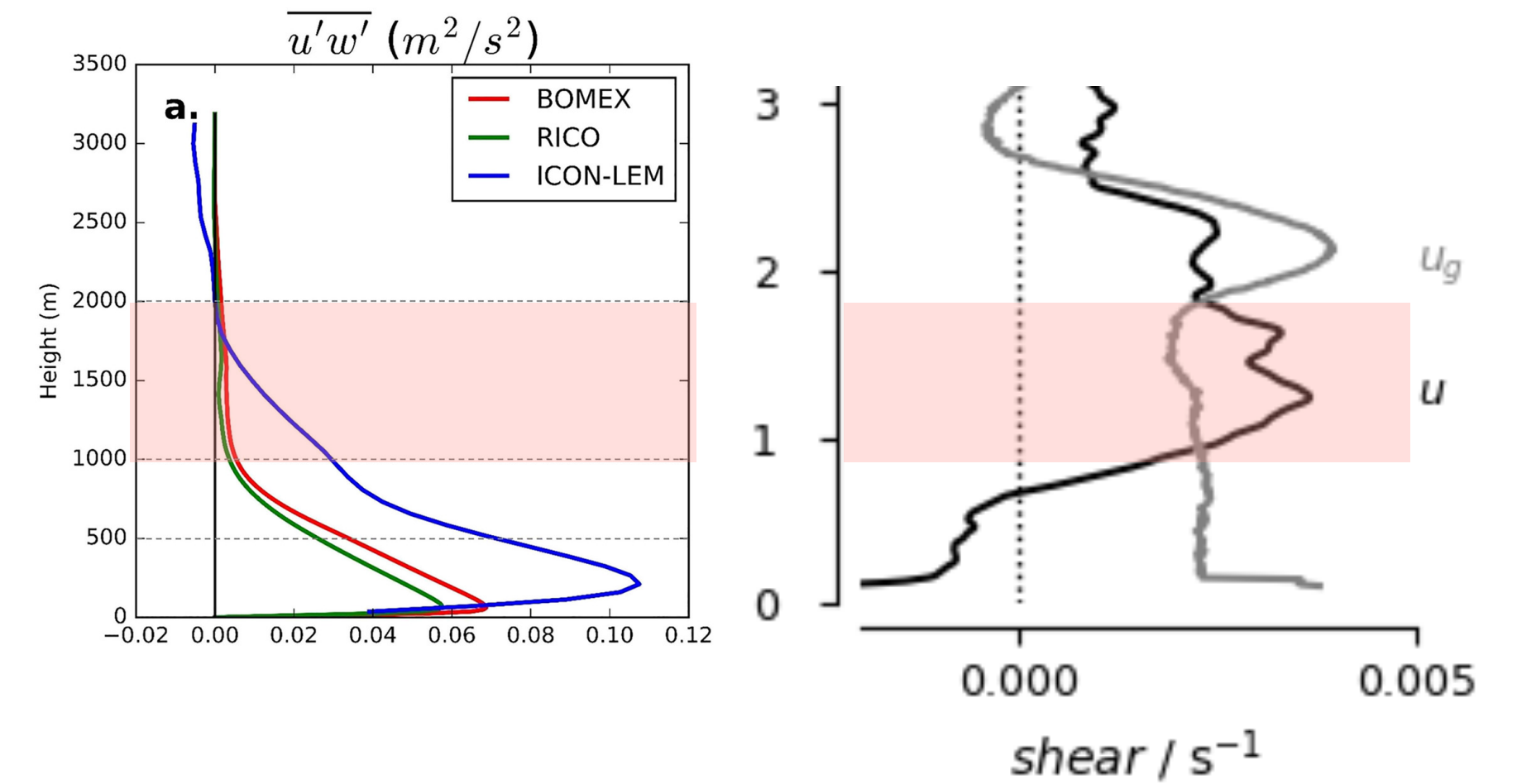
Summary

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- ✦ Observations suggest convection enhances wind shear in the cloud layer, significant momentum flux near cloud top and importance of wind variance on scales 5 - 20 km



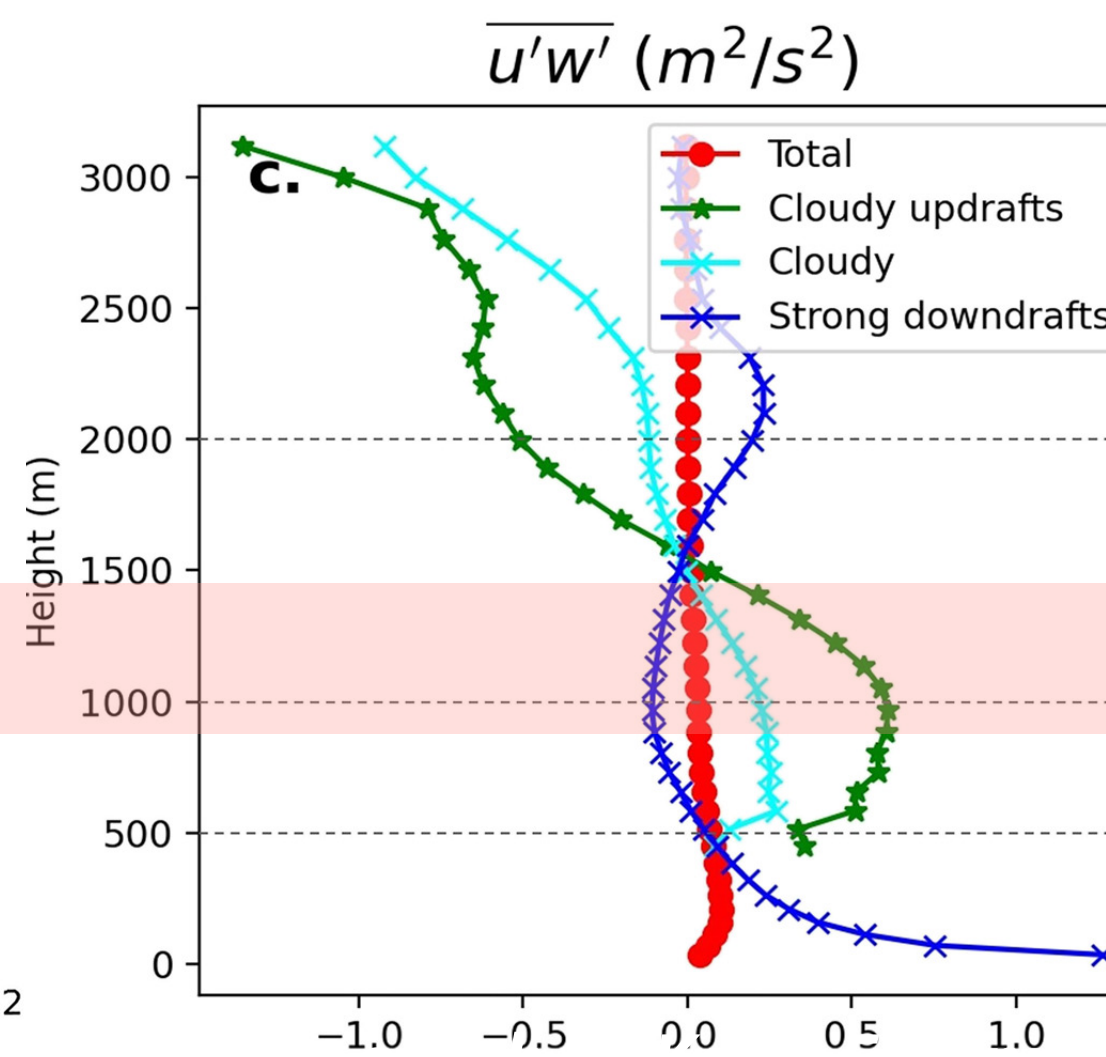
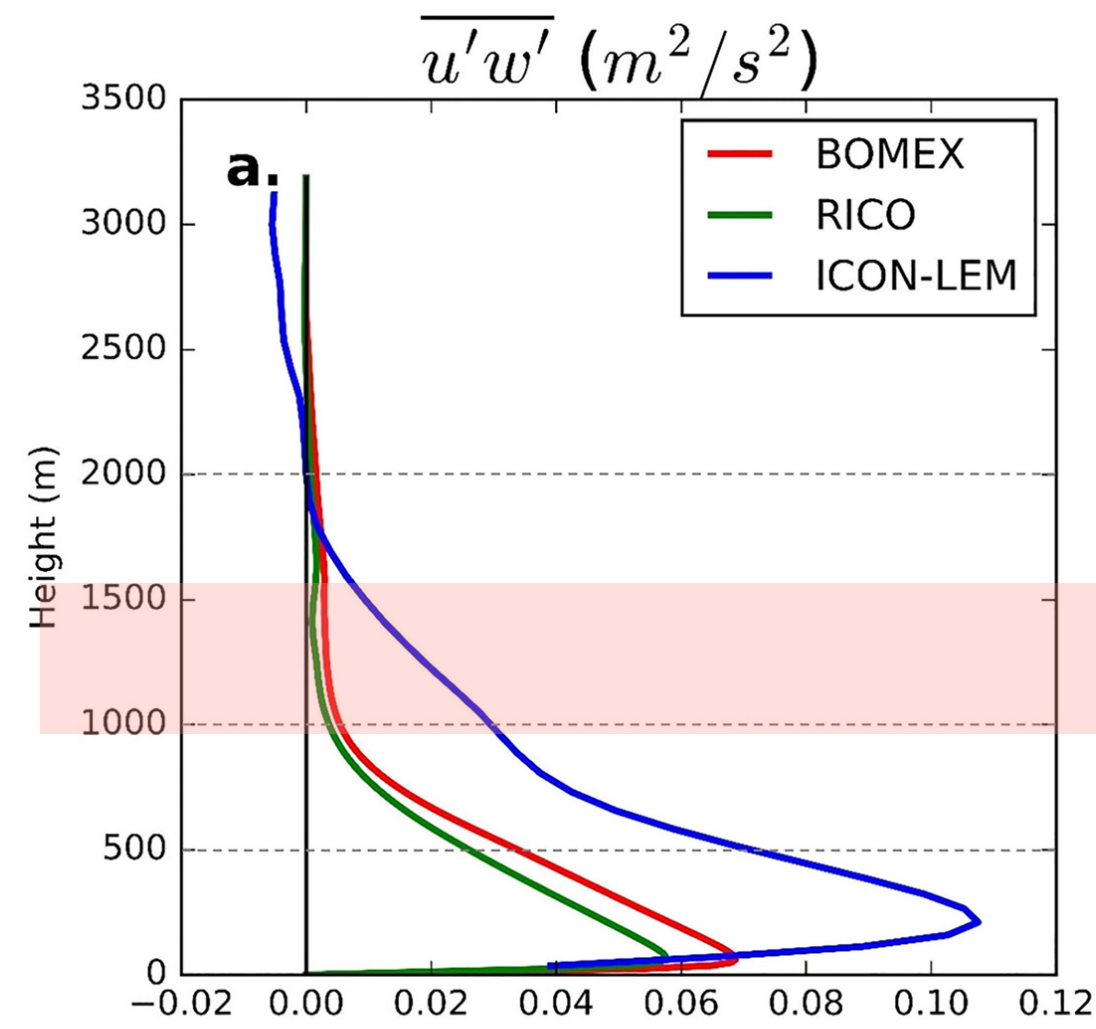
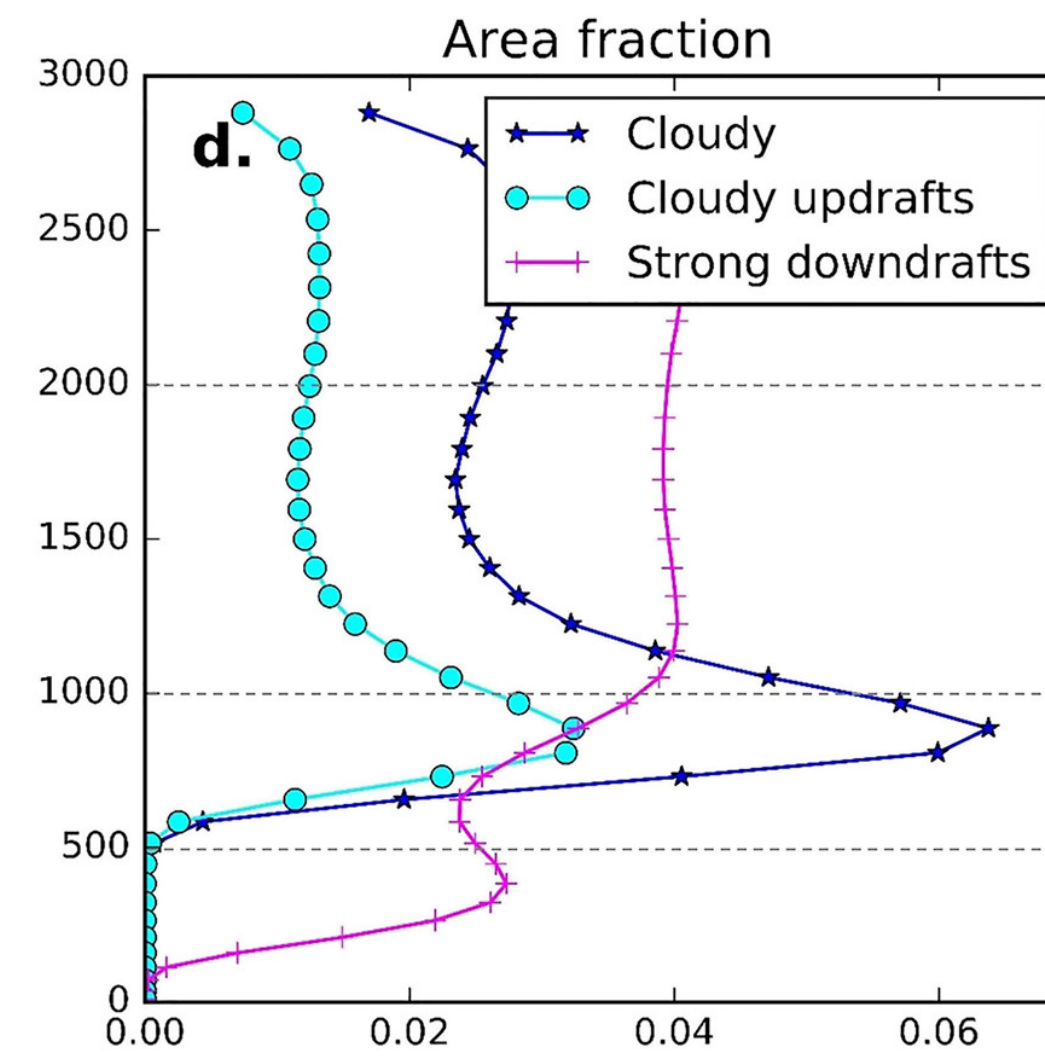
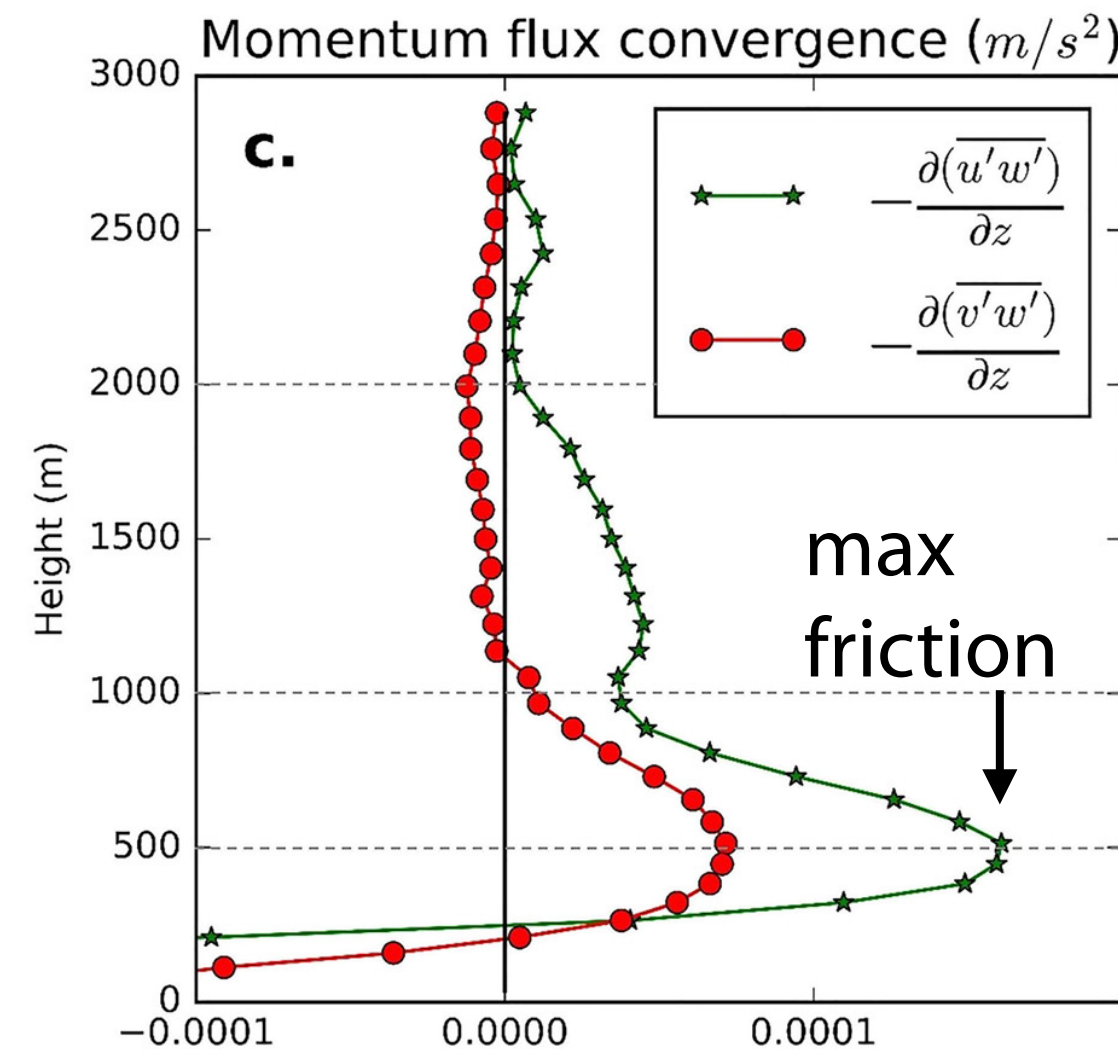
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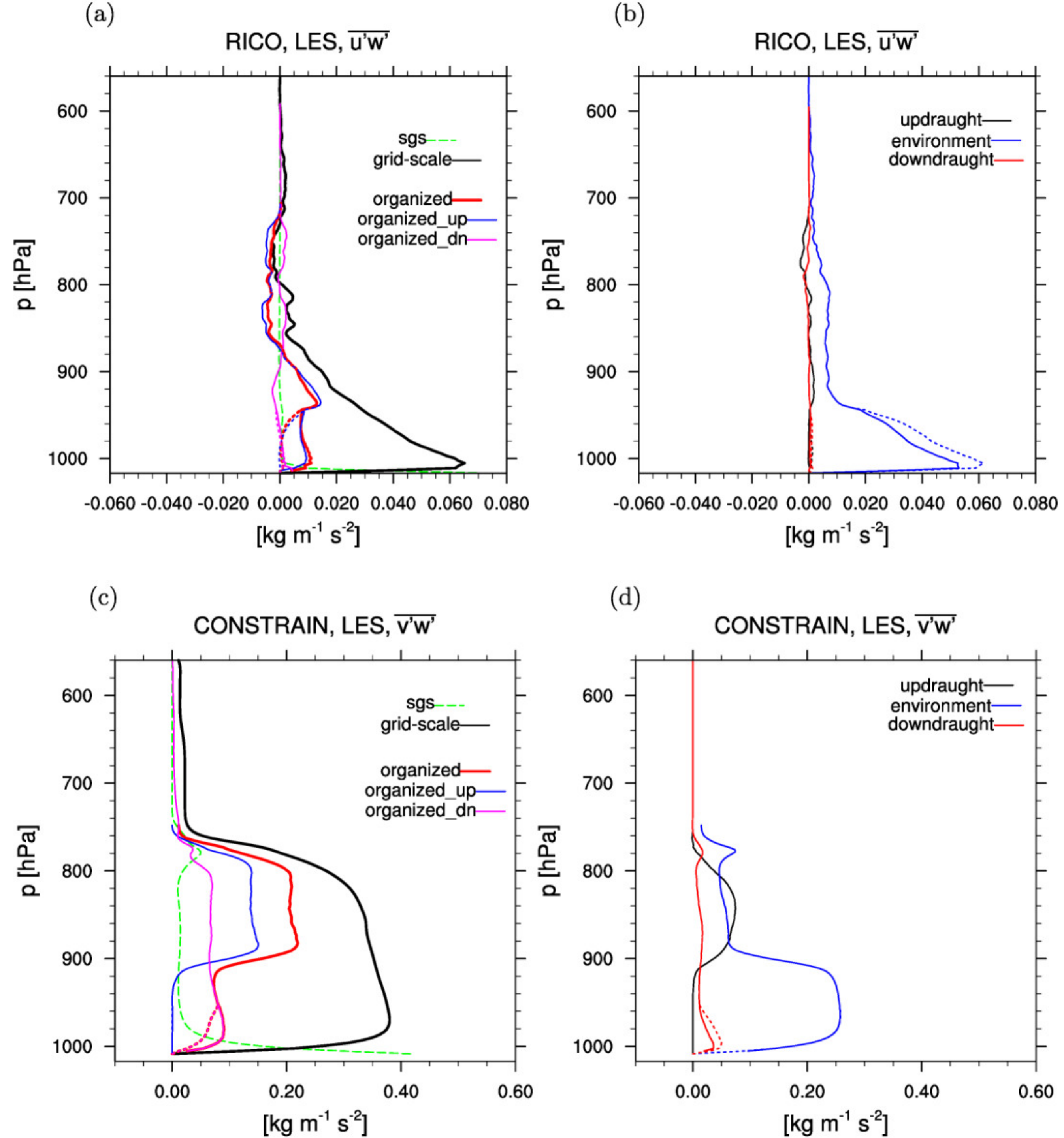
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- ✦ Observations suggest convection enhances wind shear in the cloud layer, significant momentum flux near cloud top and importance of wind variance on scales 5 - 20 km
- ✦ At times of vigorous shallow convection (gravel, flowers - congestus) up to half the momentum flux is carried by scales > 2.5 km (in the cloud layer, but also the sub-cloud layer)
- ✦ Convection accelerates winds near the surface, but also near cloud tops.

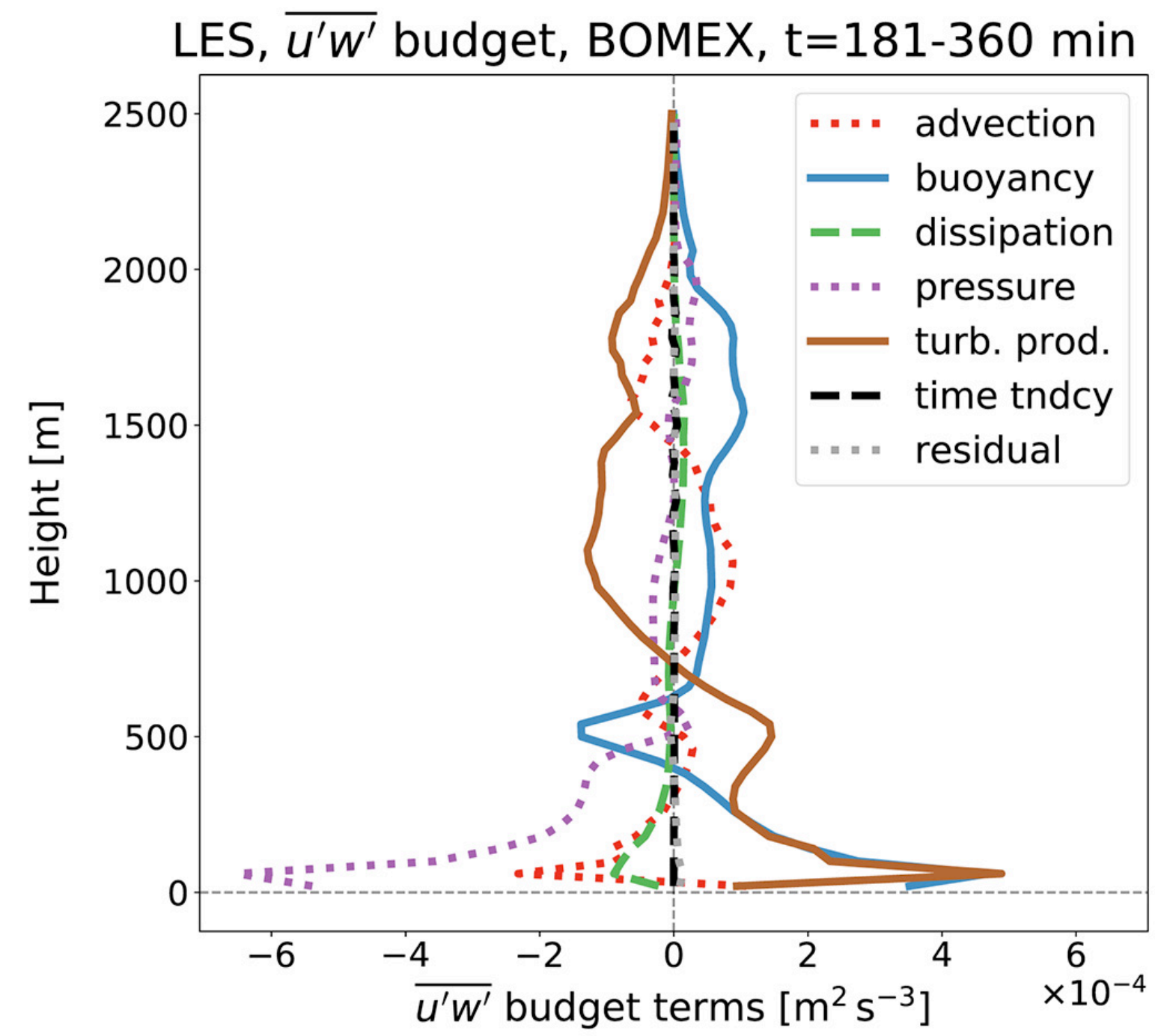
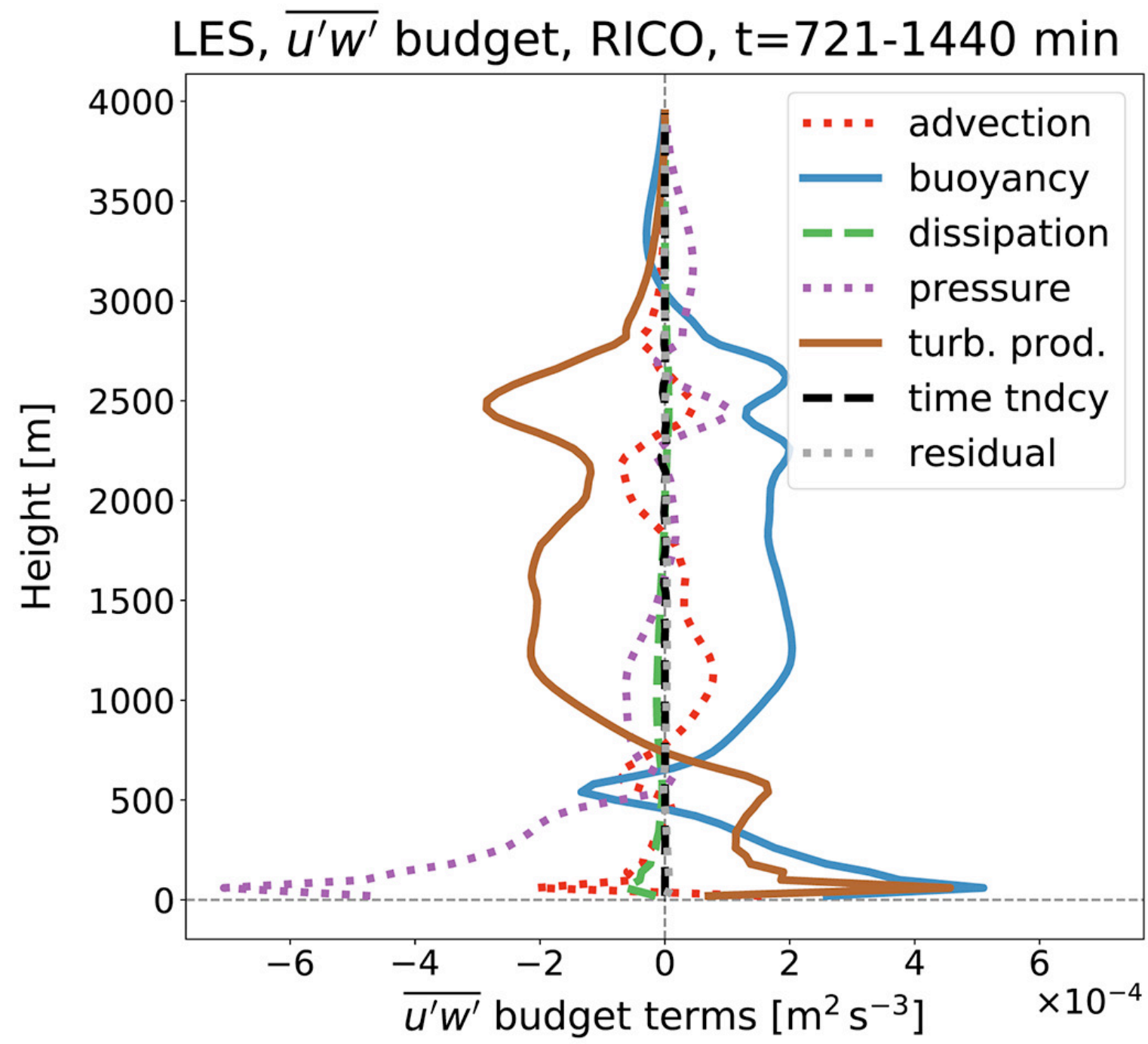


Backup slides

The maximum friction is introduced below cloud base

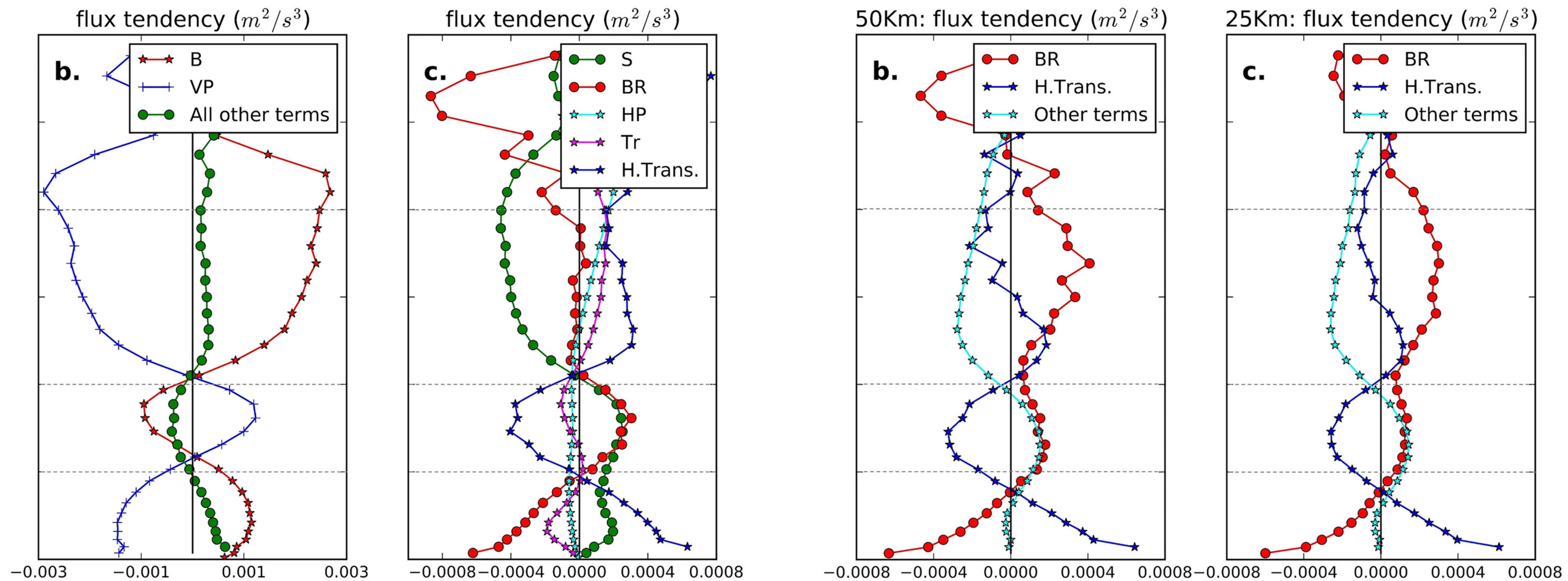






$$\frac{\partial \overline{u'w'}}{\partial t} = \underbrace{-\overline{w'^2} \frac{\partial \bar{u}}{\partial z}}_{\text{Turb Prod}} + \underbrace{\frac{g}{\theta_{1s}} \overline{u'\theta'_v}}_{\text{Buoy Prod}} - \underbrace{C_7 \frac{g}{\theta_{1s}} \overline{u'\theta'_v}}_{\text{Press-Buoy}} - \underbrace{\frac{C_6}{\tau} \overline{u'w'}}_{\text{Return to Isotropy}} - \underbrace{\frac{1}{\rho} \frac{\partial \overline{\rho w'^2 u'}}{\partial z}}_{\text{Turb Advection}}.$$

Buoyancy generation of flux largely compensated by vertical pressure perturbations



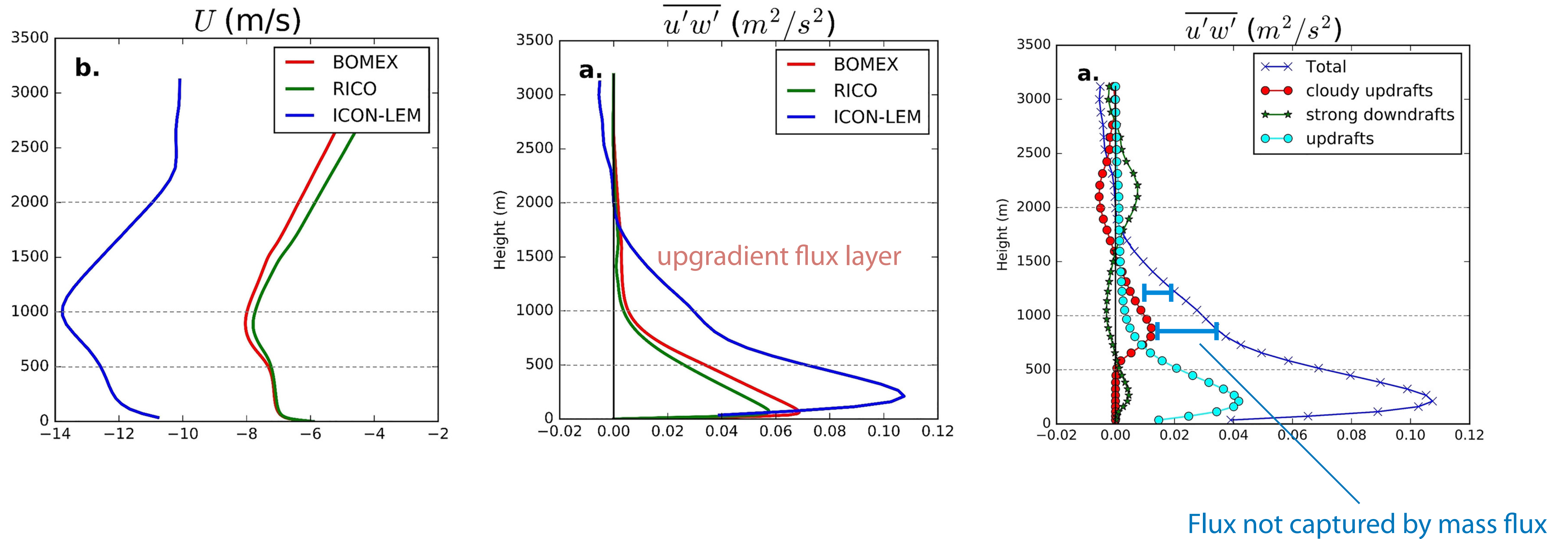
Dixit et al (2020): NARVAL with ICON-LEM on 150 km x 150 km x xx km, $\Delta x = 125m$

$$BR = \frac{g}{T_v} \overline{u'T'_v} - \frac{\overline{u' \partial p'}}{\bar{\rho} \partial z}$$

B

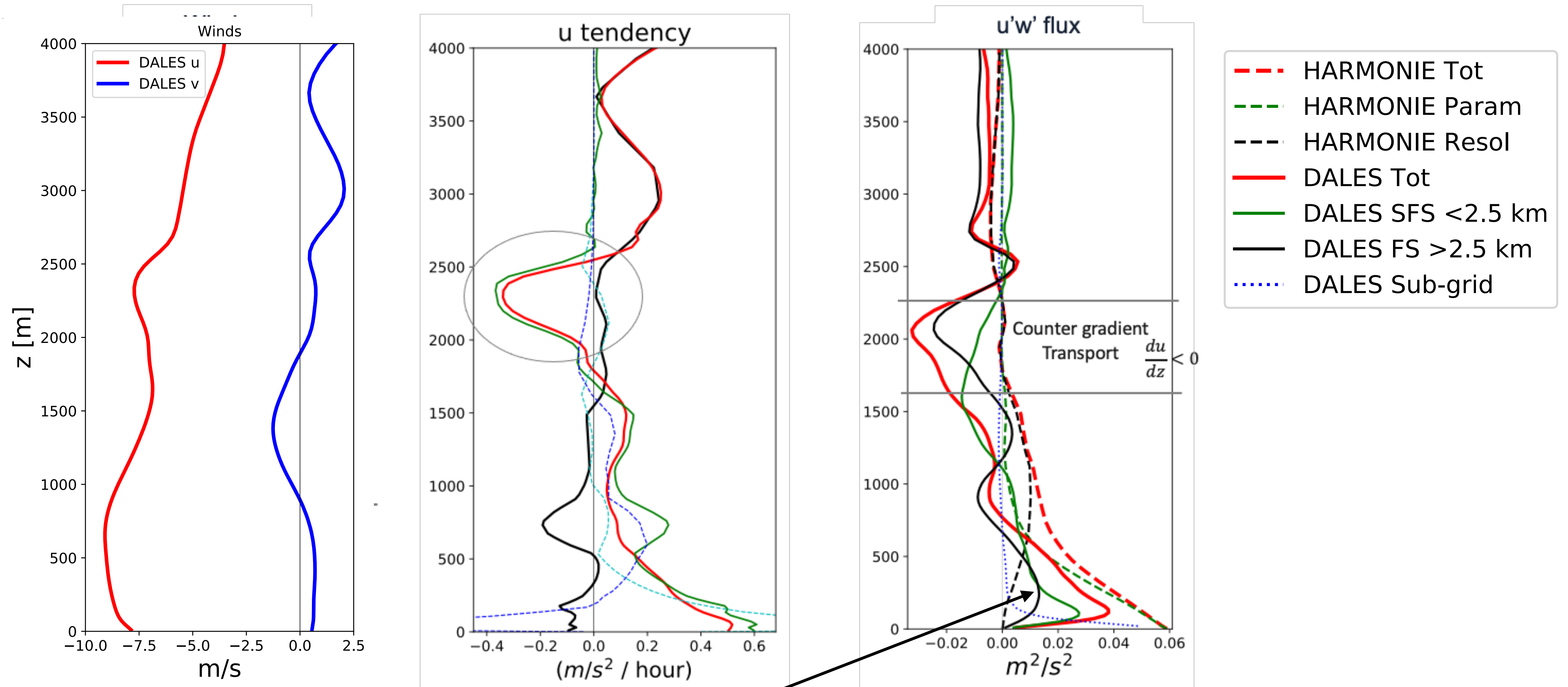
VP

Mass flux approach does not suffice (environmental flux is significant) and pressure gradients matter for in-cloud momentum



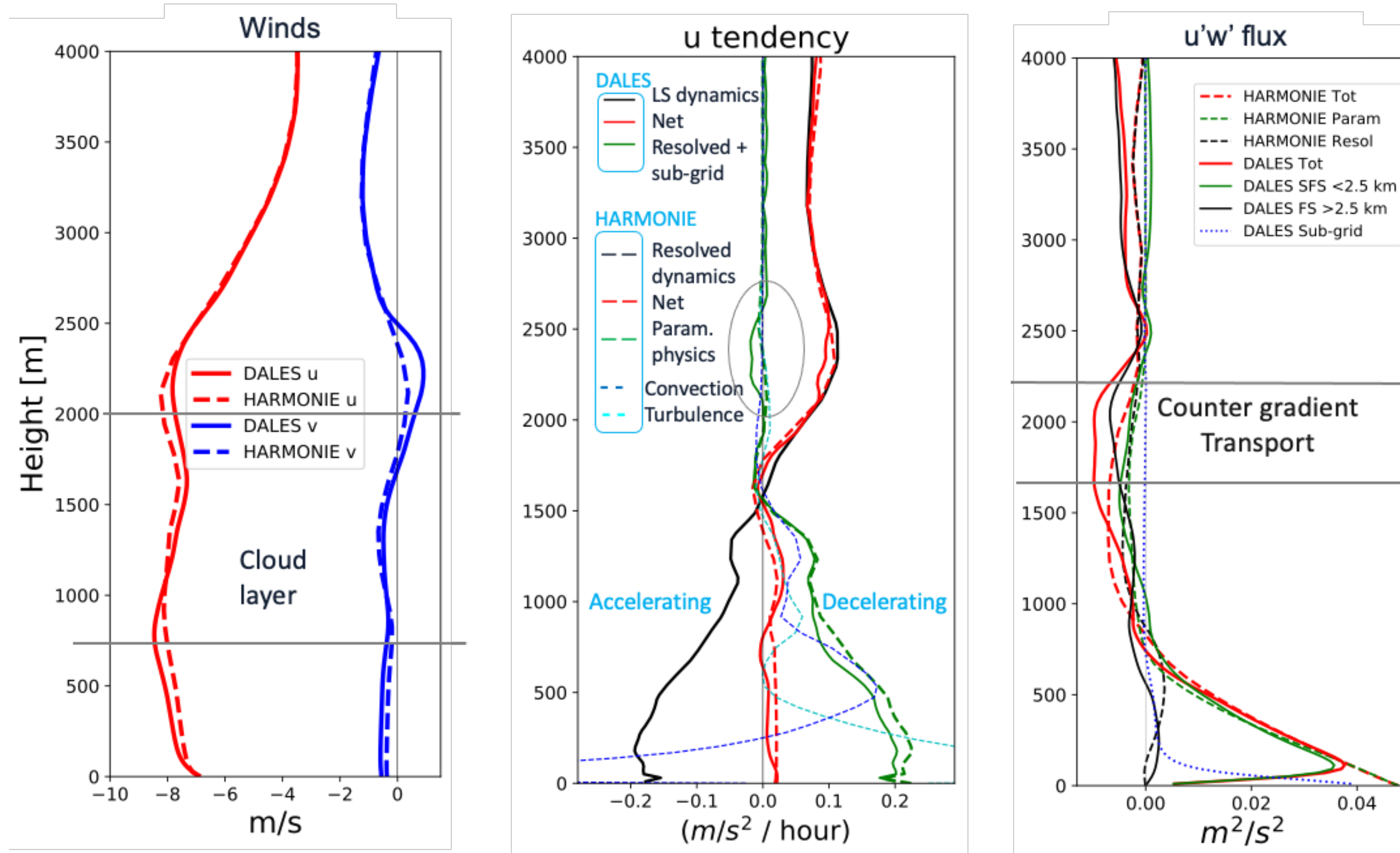
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Mesoscale momentum transport important for upgradient transport and acceleration near cloud tops, but also in the mixed-layer

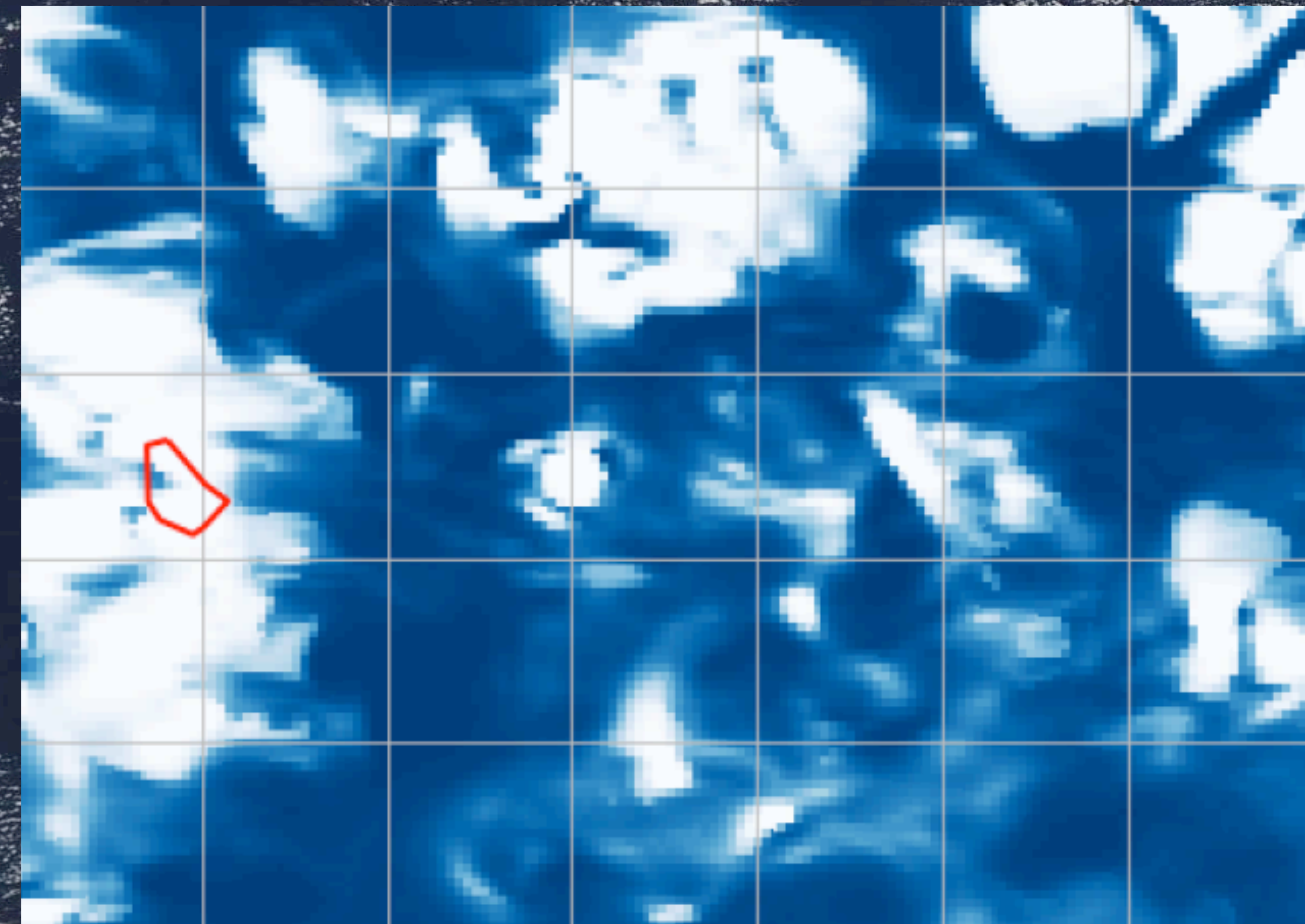
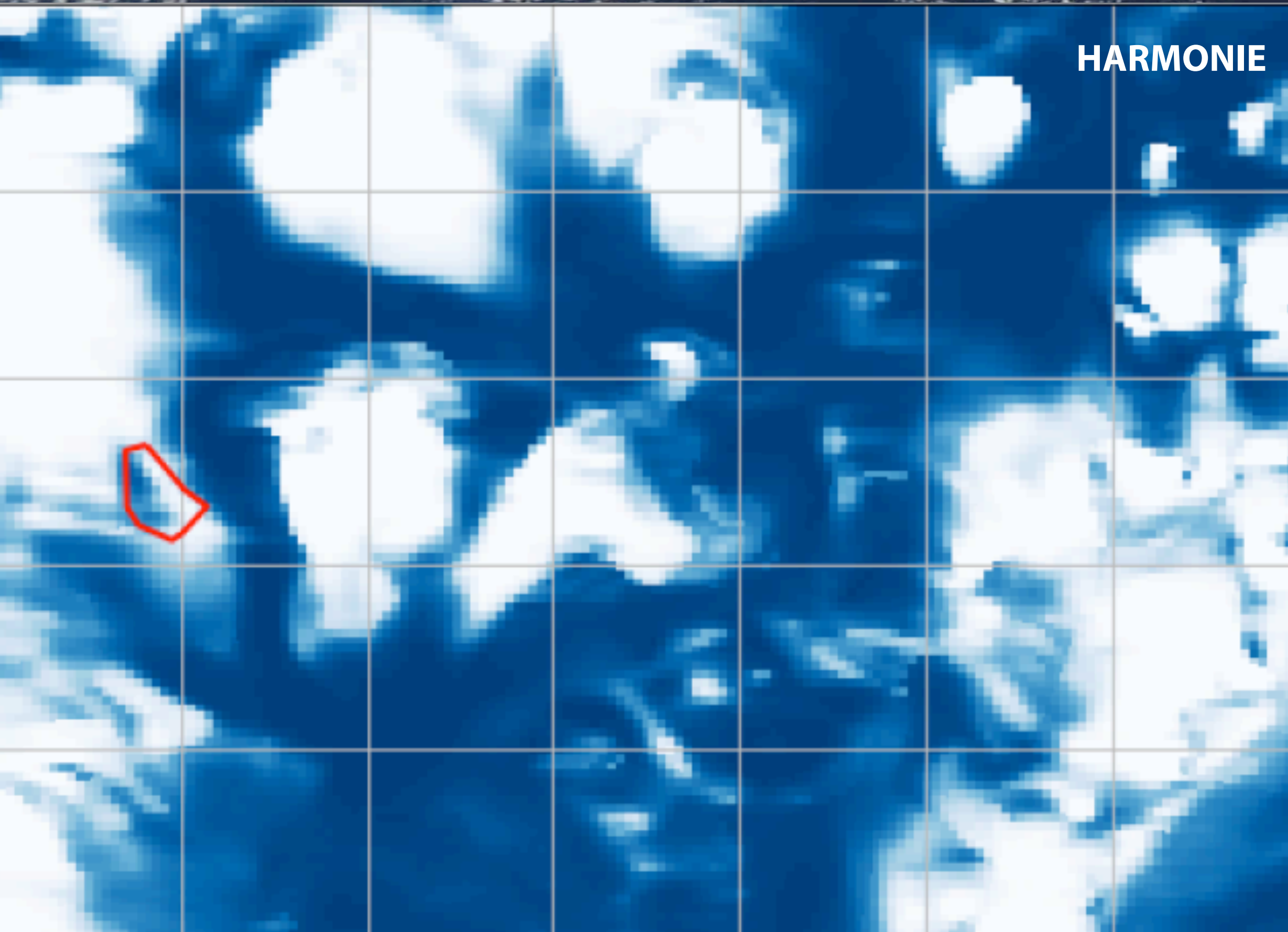


50% ≤ 2.5 km, 50% > 2.5 km

Mesoscale momentum transport important for upgradient transport and acceleration near cloud tops, but also in the mixed-layer



no shallow convective momentum transport



no shallow convective momentum transport

