

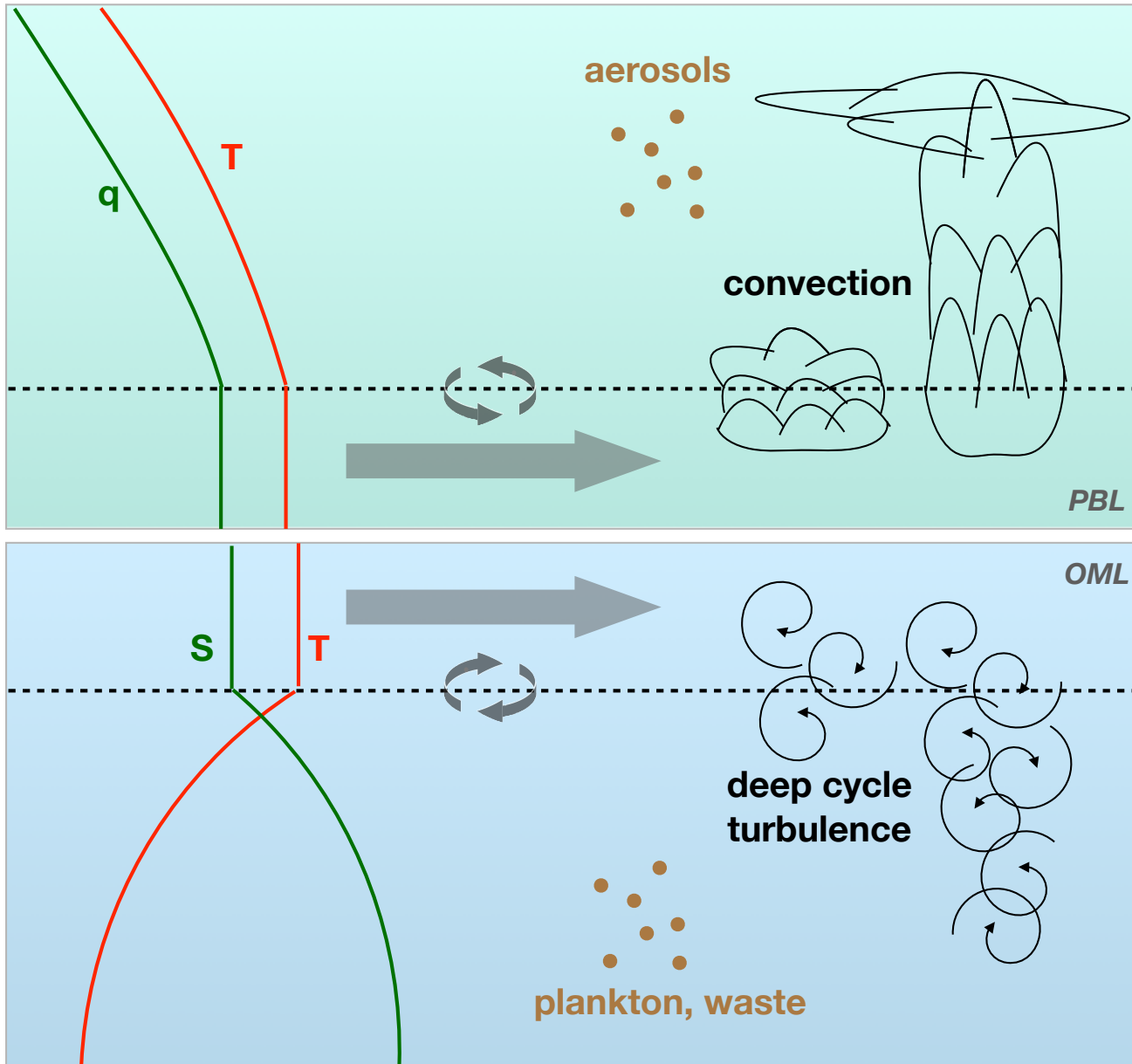


# Challenging physics in seamless forecasts: Air-sea interactions in the tropics

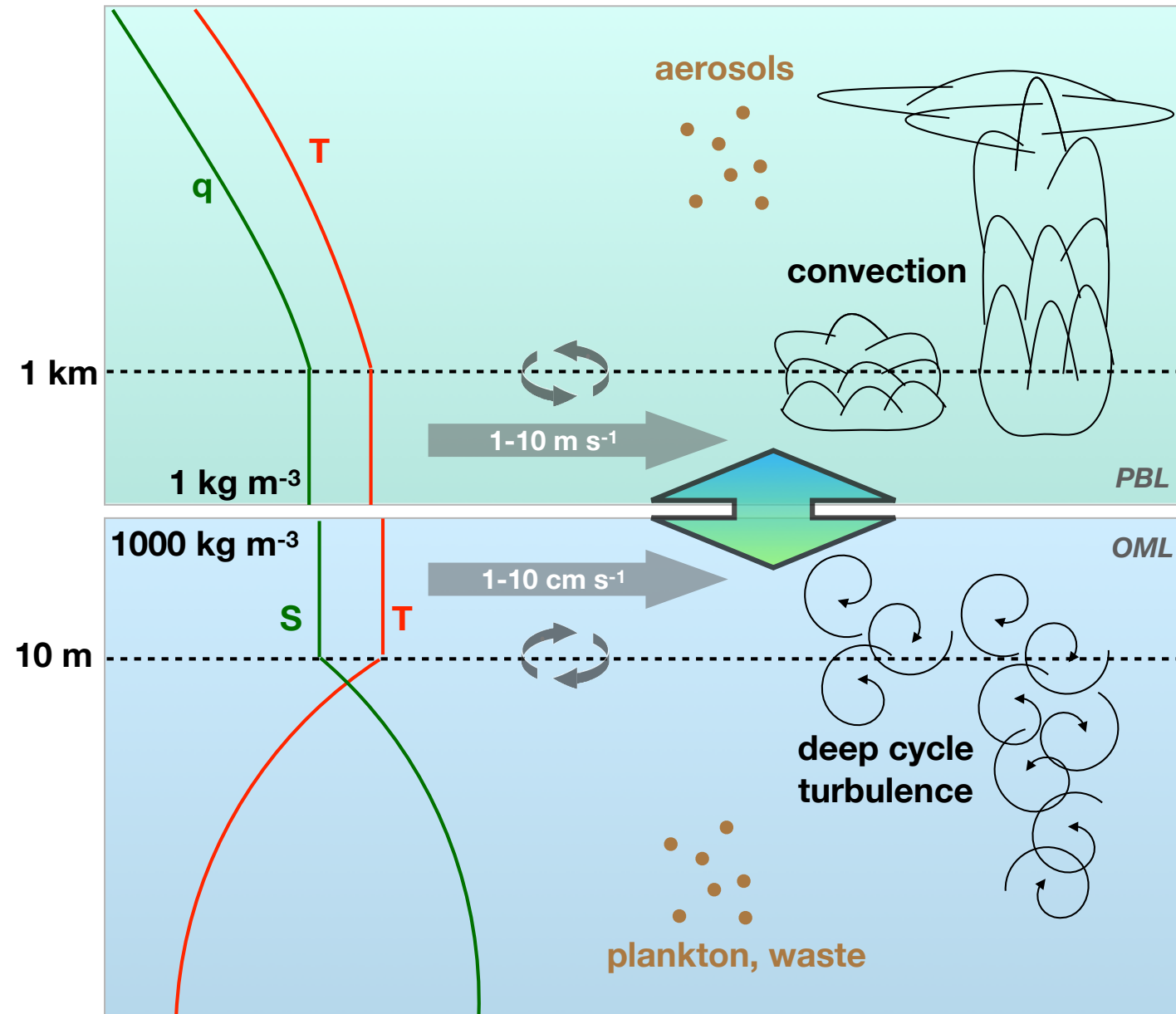
Charlotte A. DeMott  
Colorado State University

**overview**

# ocean-atmosphere similarities



# ocean-atmosphere differences



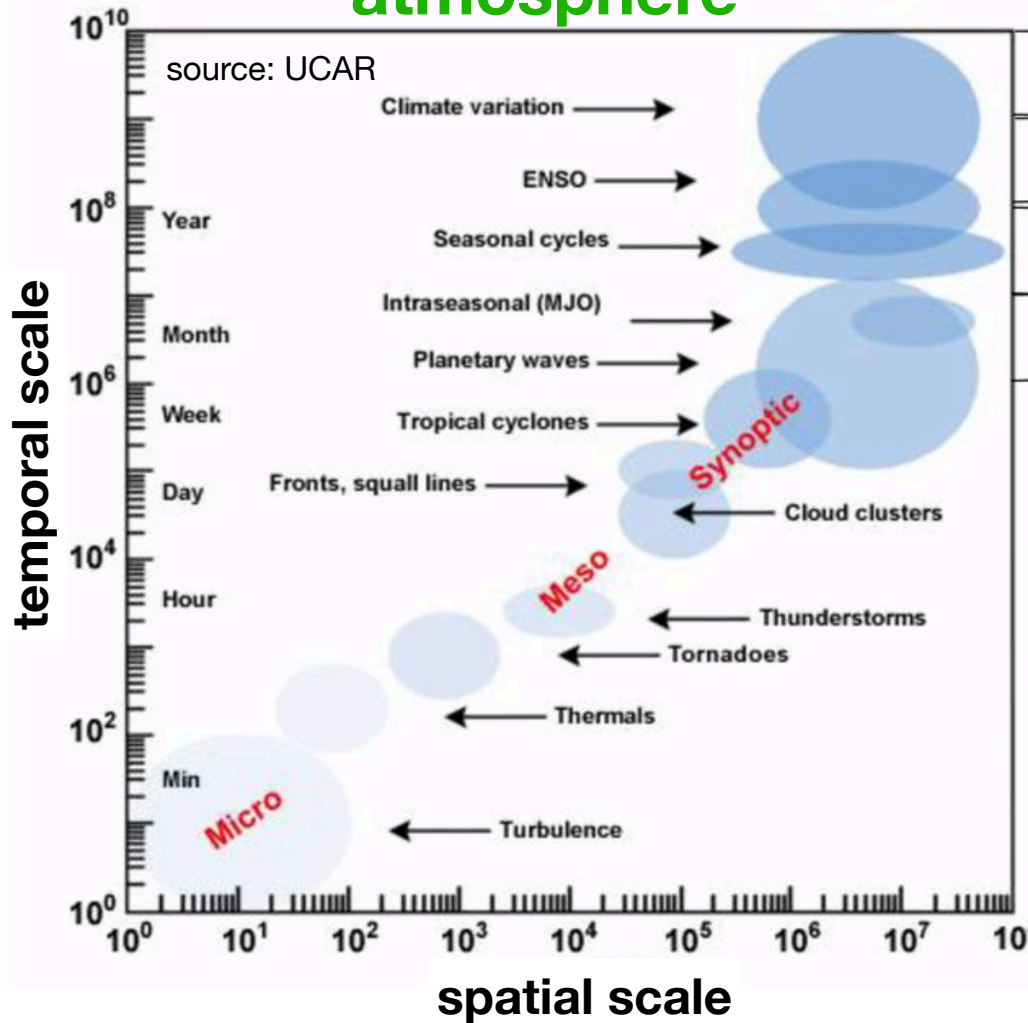
- density:  $\delta \sim 10^3$
- ML depth:  $\delta \sim 10^2$
- $u, v$ :  $\delta \sim 10^2$
- clouds are global; DCT mostly on Equator
- clouds regulate radiation; DCT does not

*O-A density differences are responsible for different adjustment timescales*

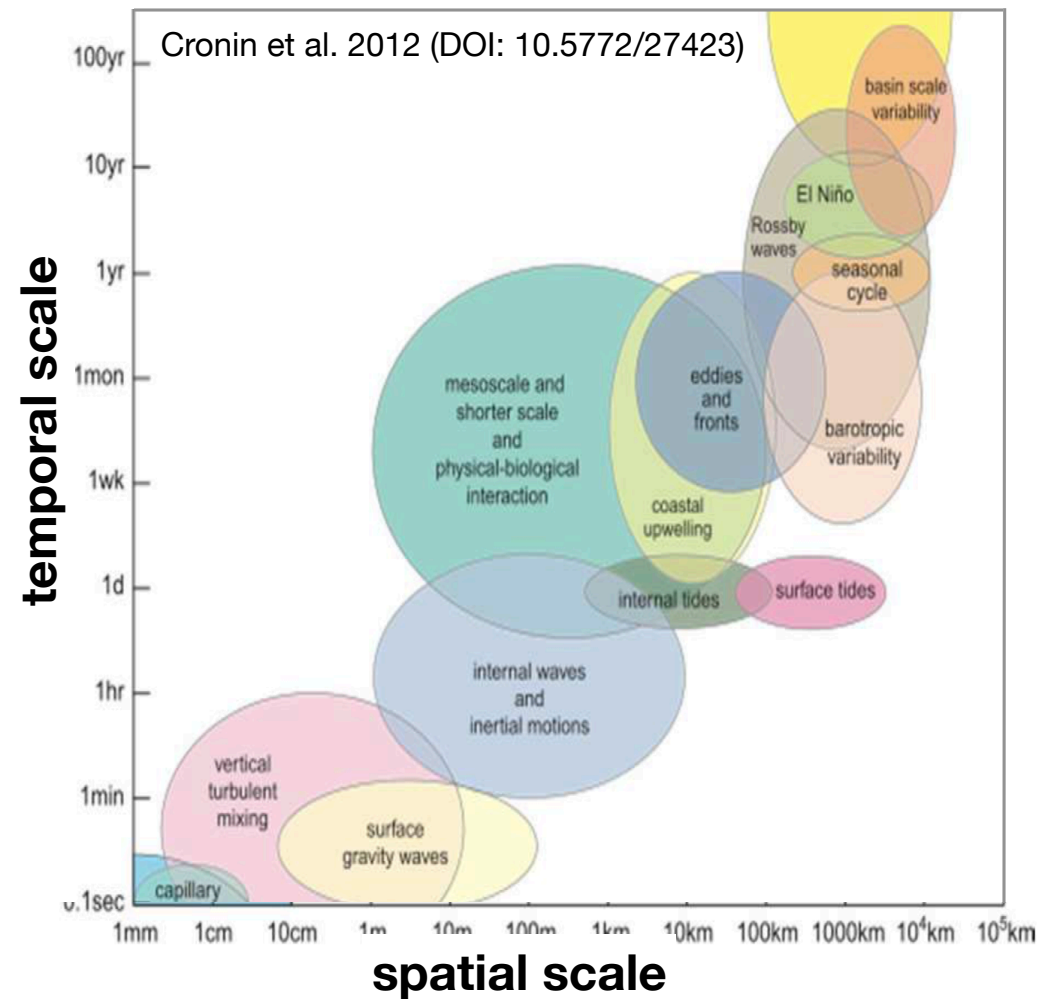
*surface fluxes are the result of O-A constantly trying to equilibrate*

# Scales of variability: atmosphere & ocean

## atmosphere



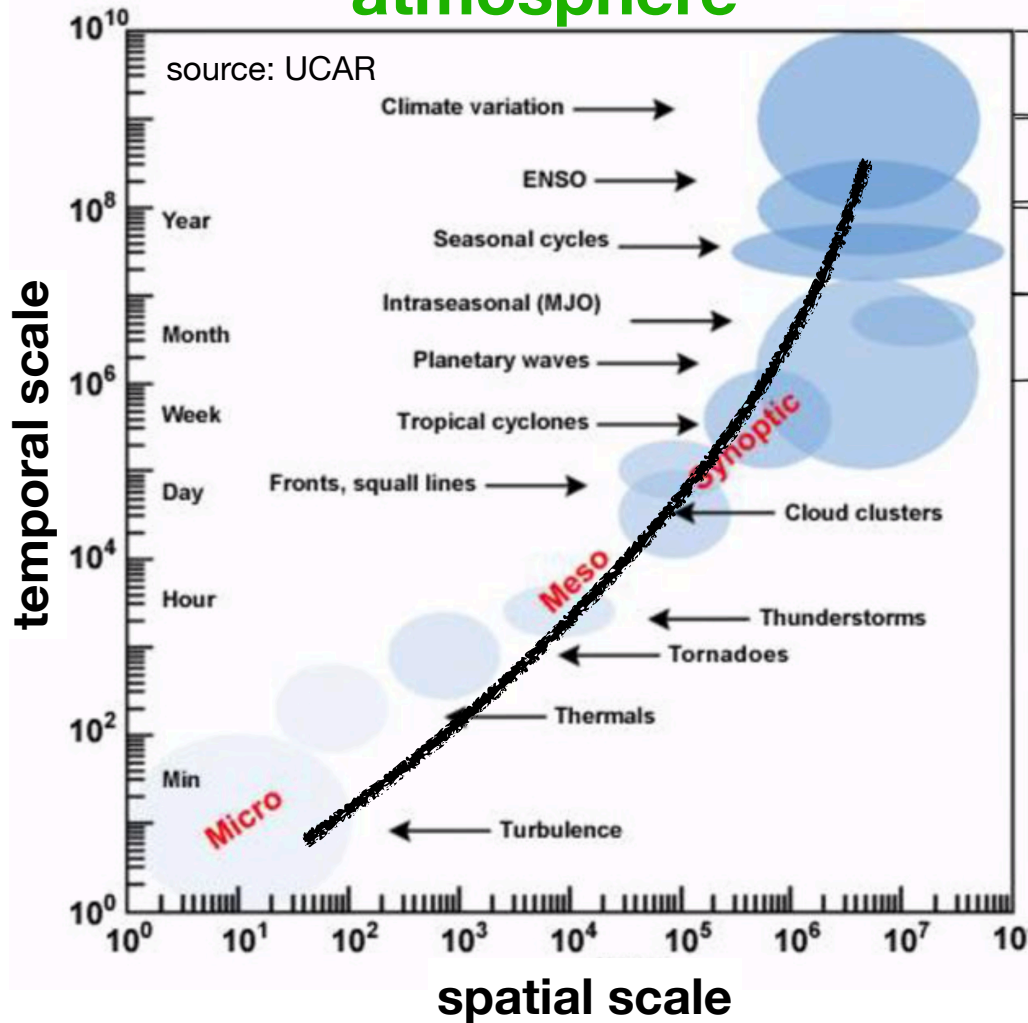
## ocean



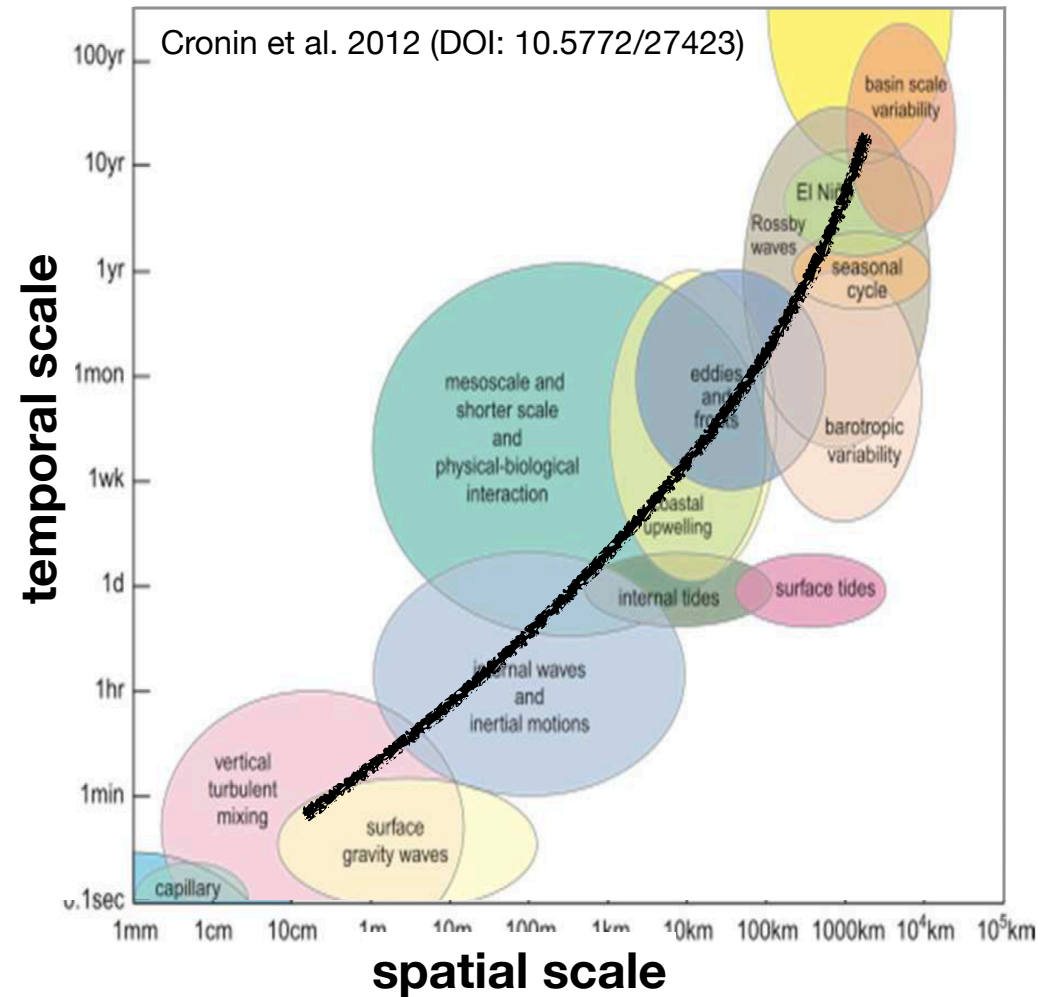
Each process is a link in a long chain

# Scales of variability: atmosphere & ocean

## atmosphere

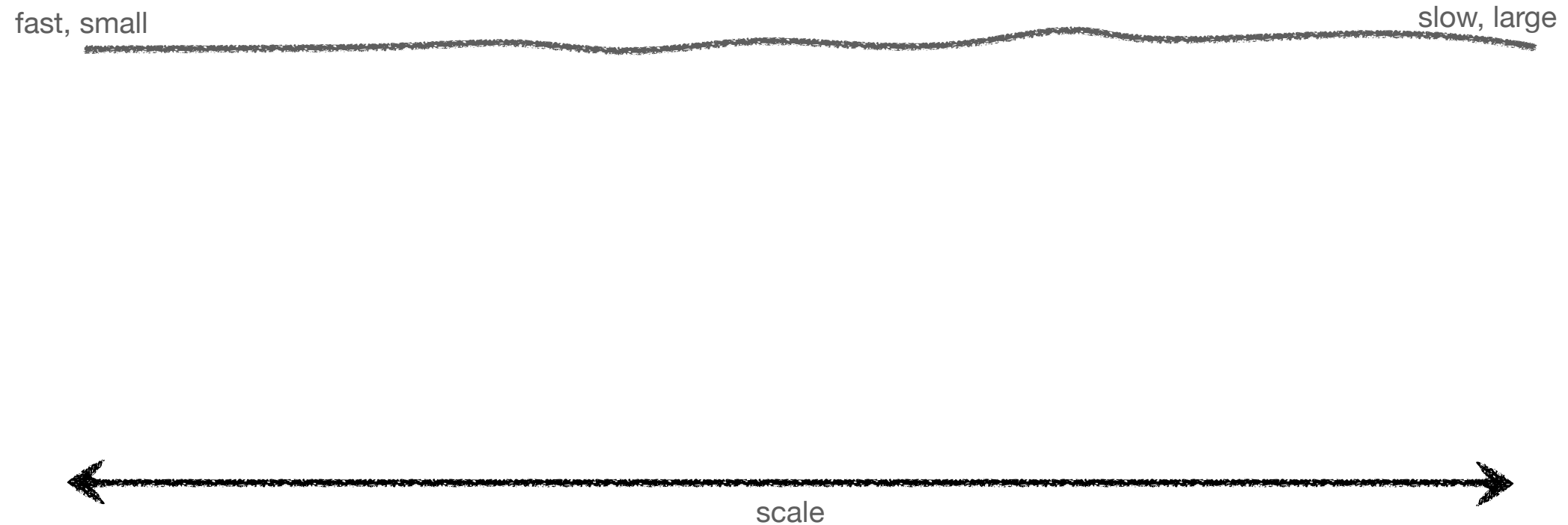


## ocean



Each process is a link in a long chain

# coupled cross-scale interactions



# coupled cross-scale interactions

atmosphere

fast, small

slow, large

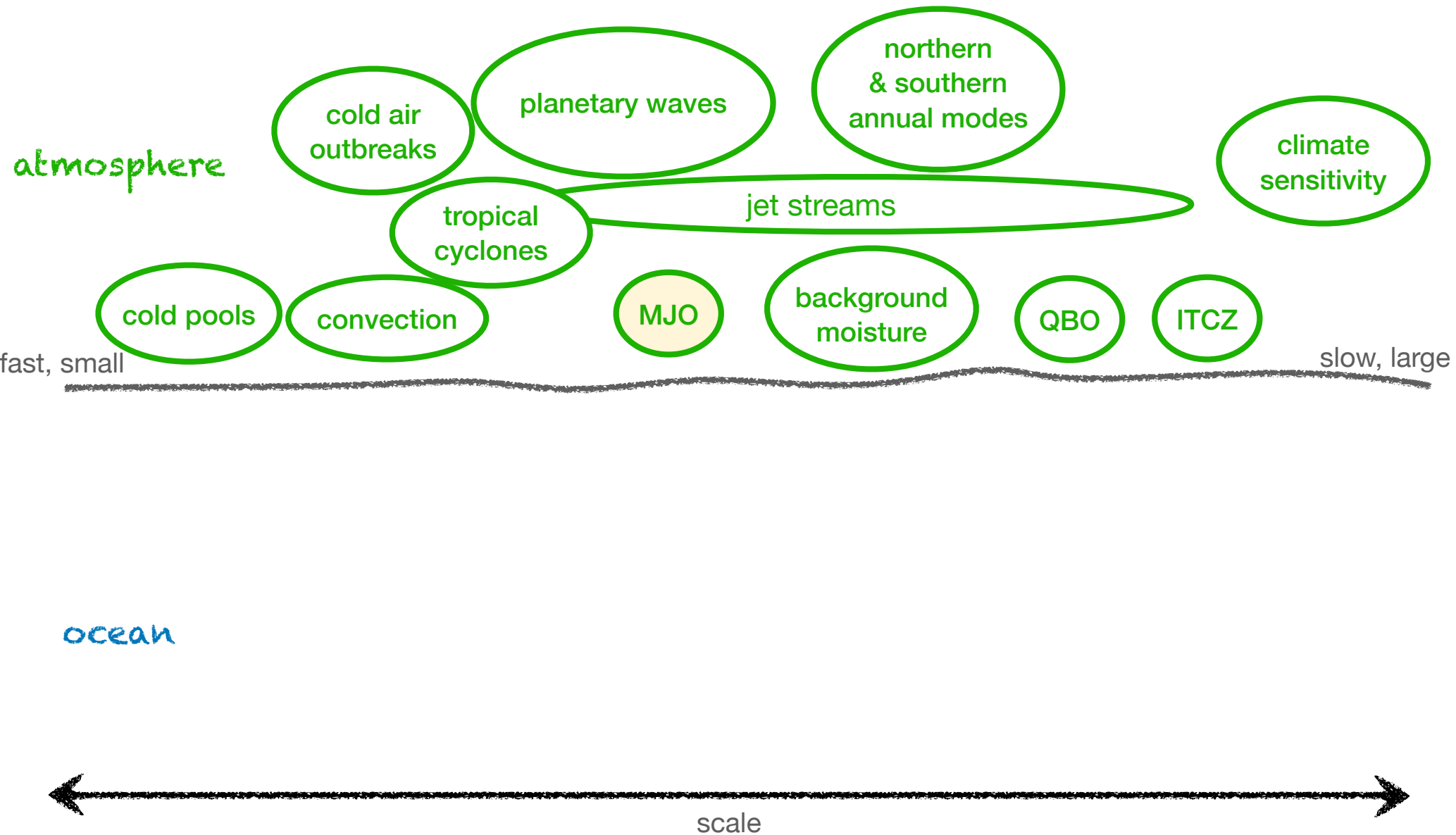
ocean

scale

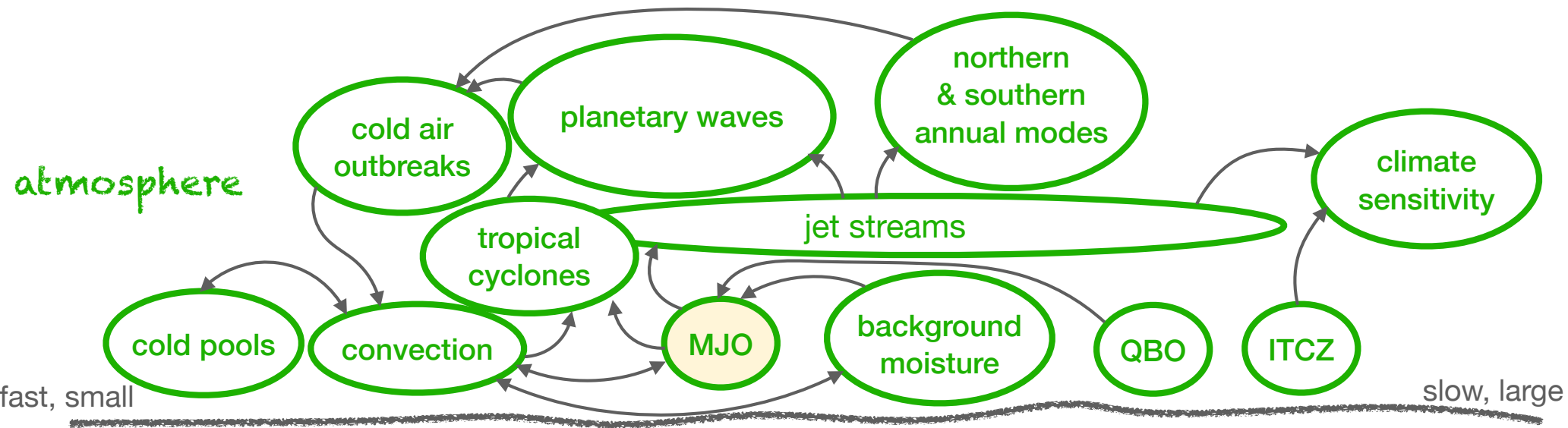




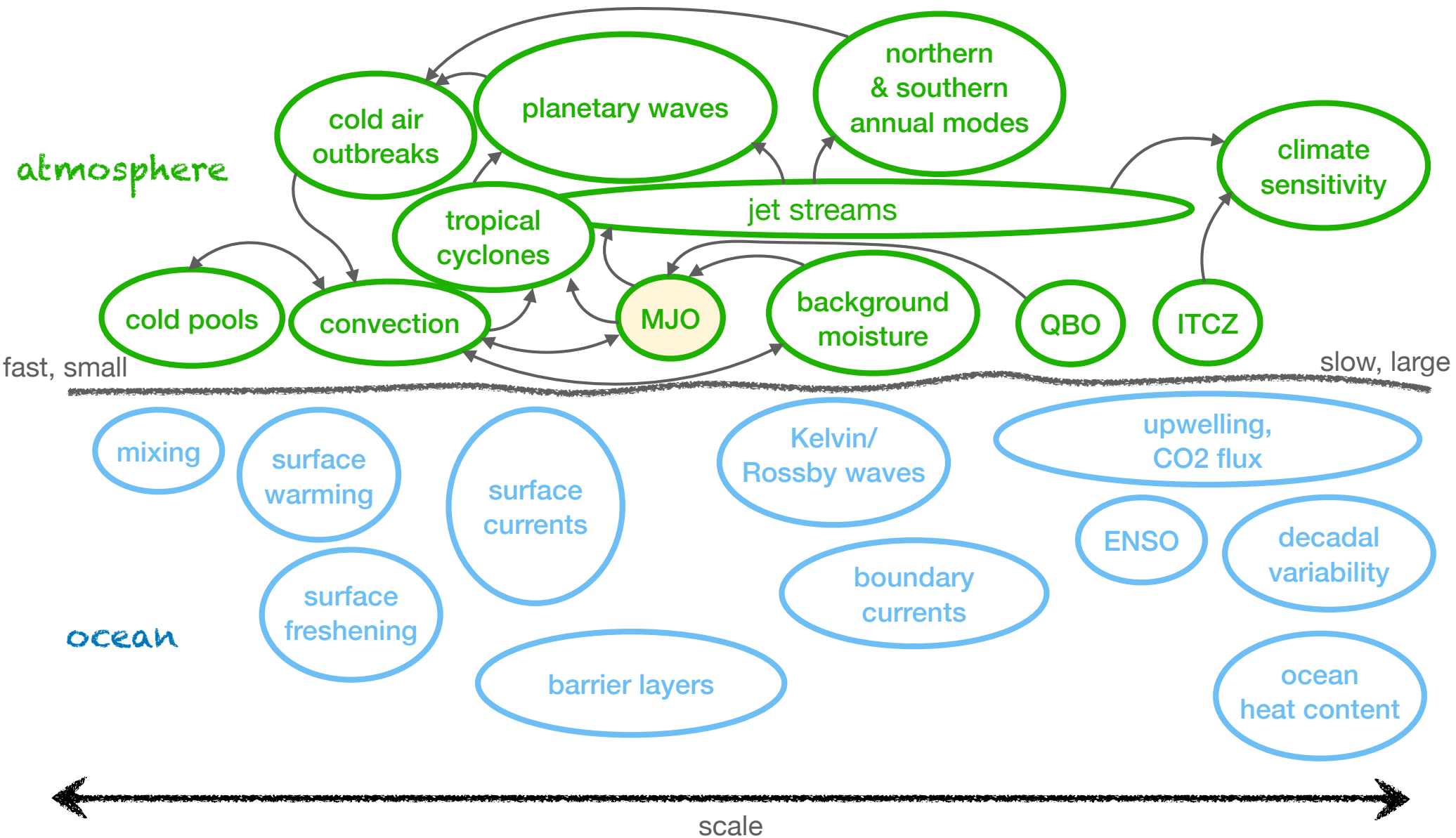
# coupled cross-scale interactions



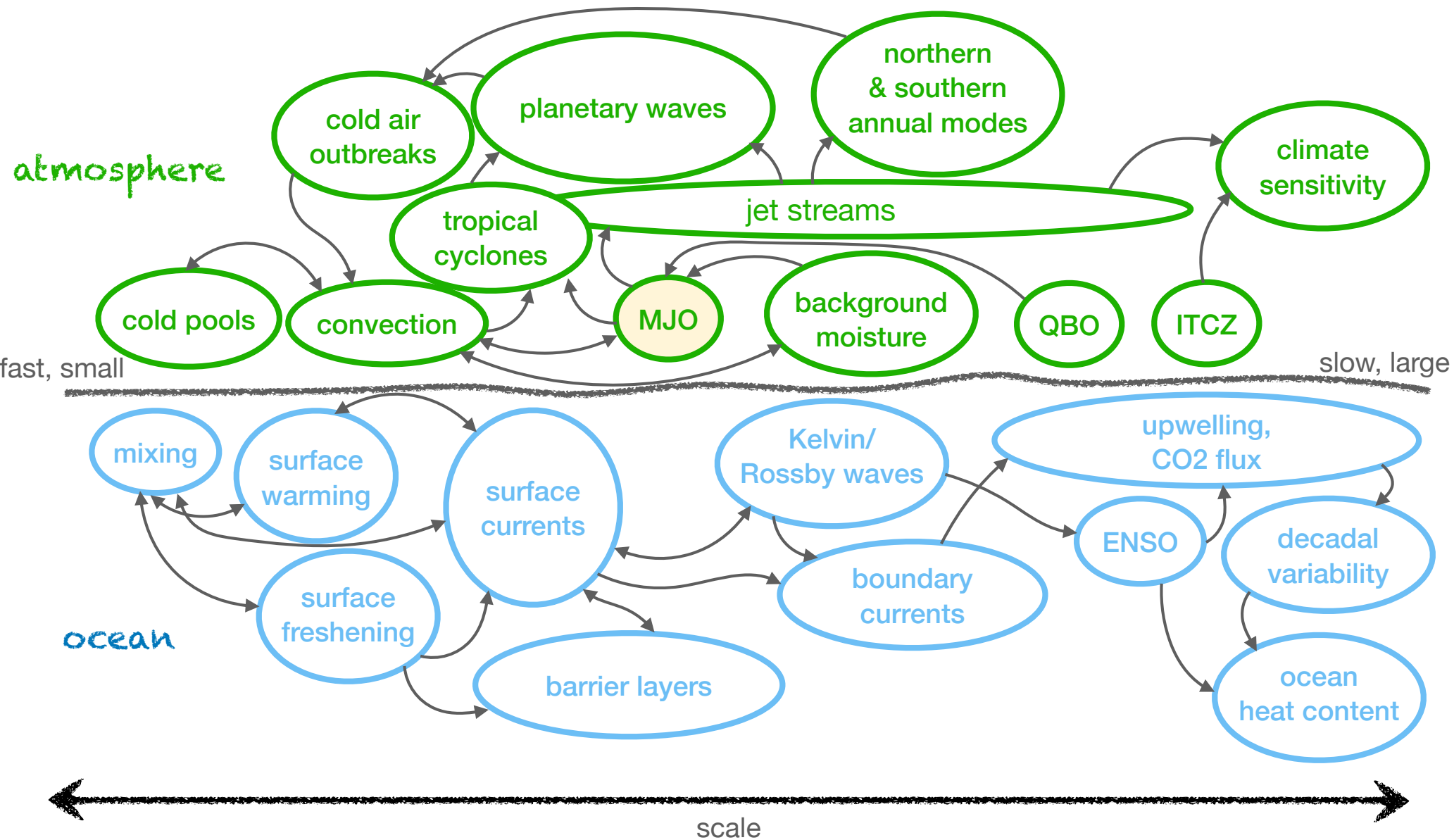
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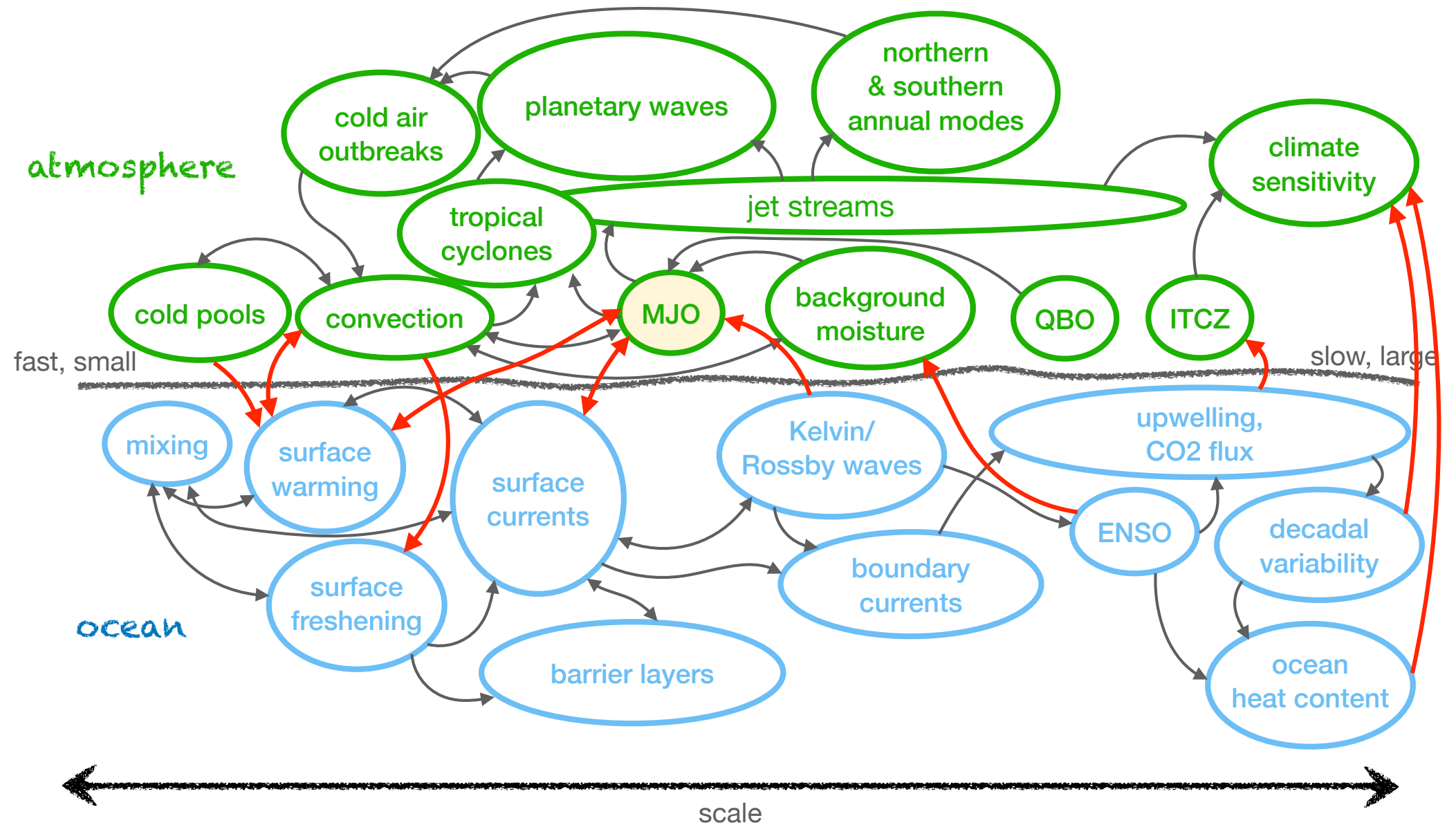
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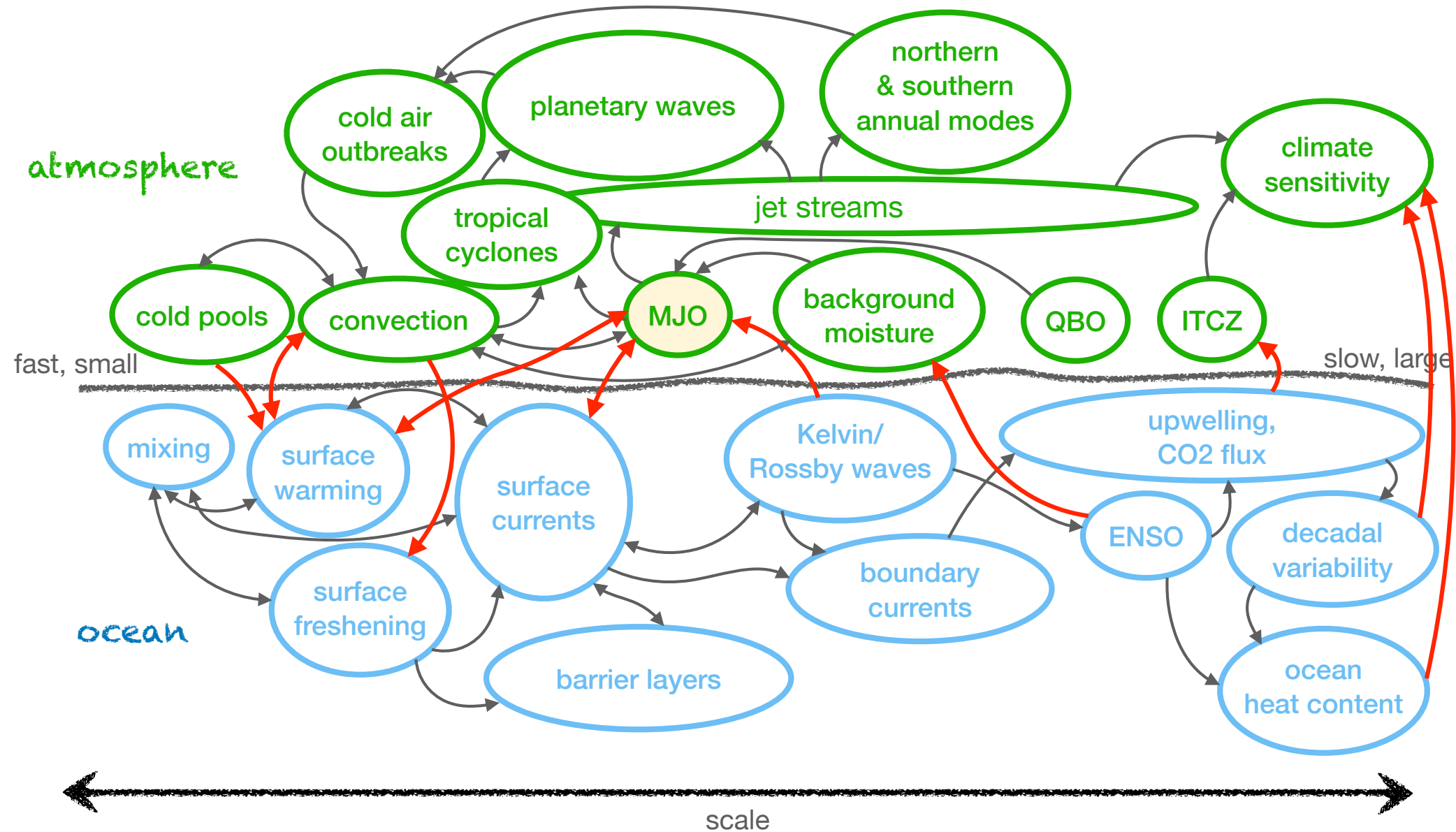


# coupled cross-scale interactions



# coupled cross-scale interactions

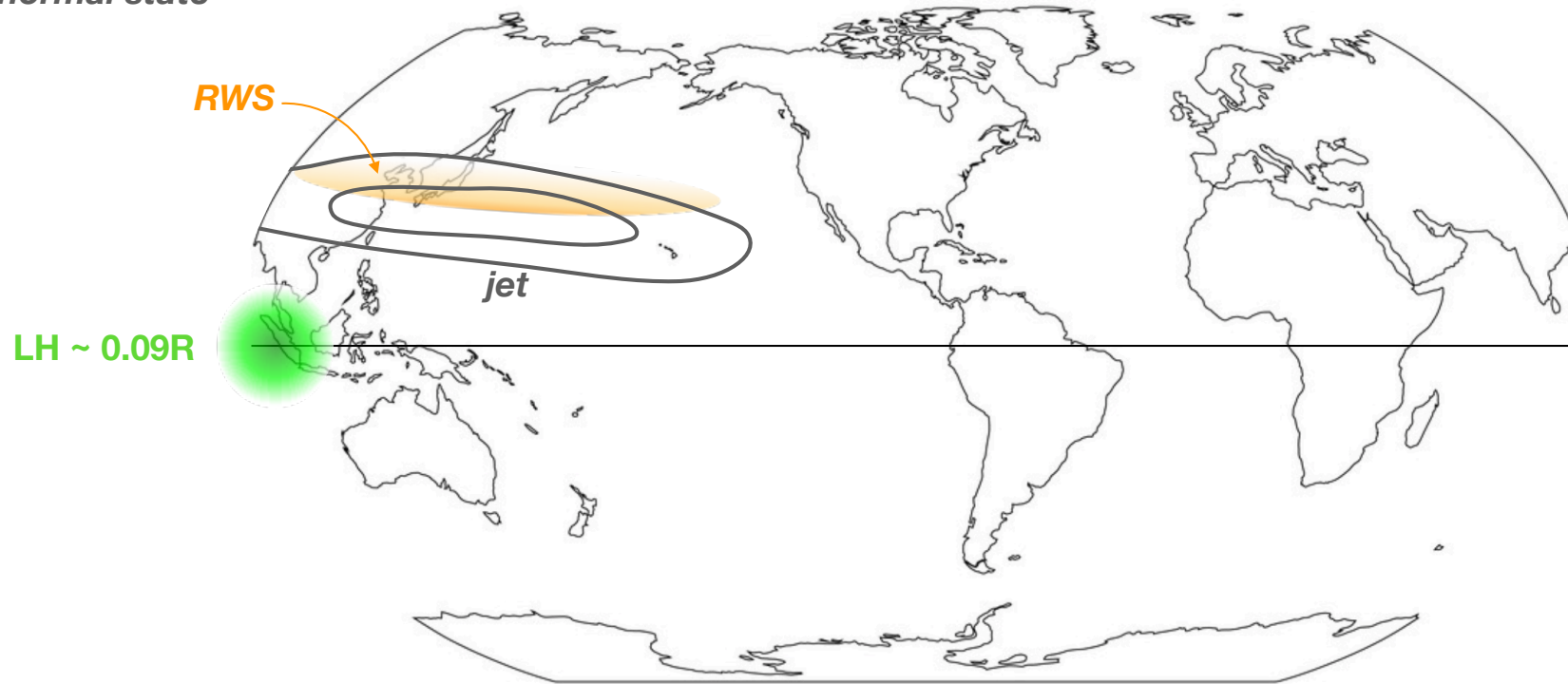
coupling is regulated by fluxes of heat, water, and momentum



# coupled feedbacks in weather & climate simulation

## Example: MJO teleconnections

*normal state*

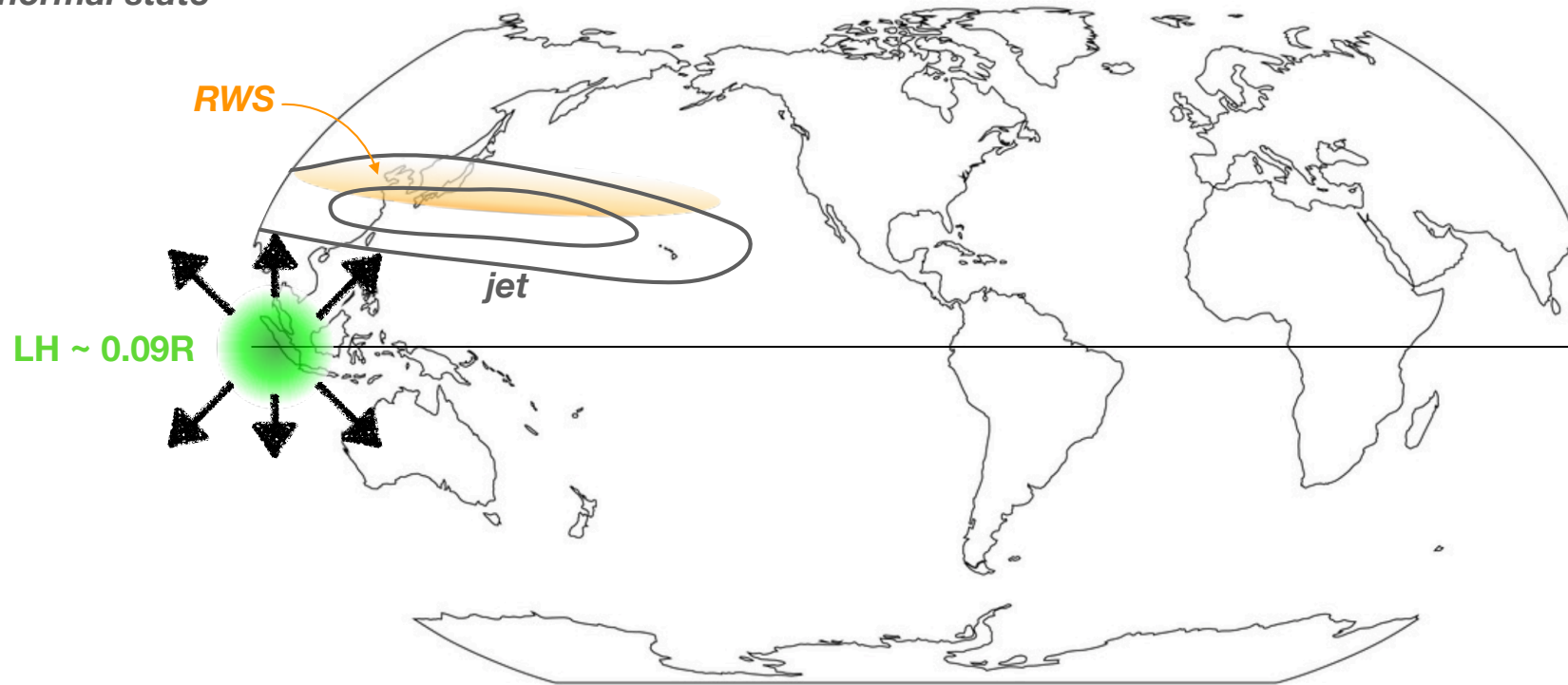


*animations adapted from Henderson et al. 2017*

# coupled feedbacks in weather & climate simulation

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*normal state*



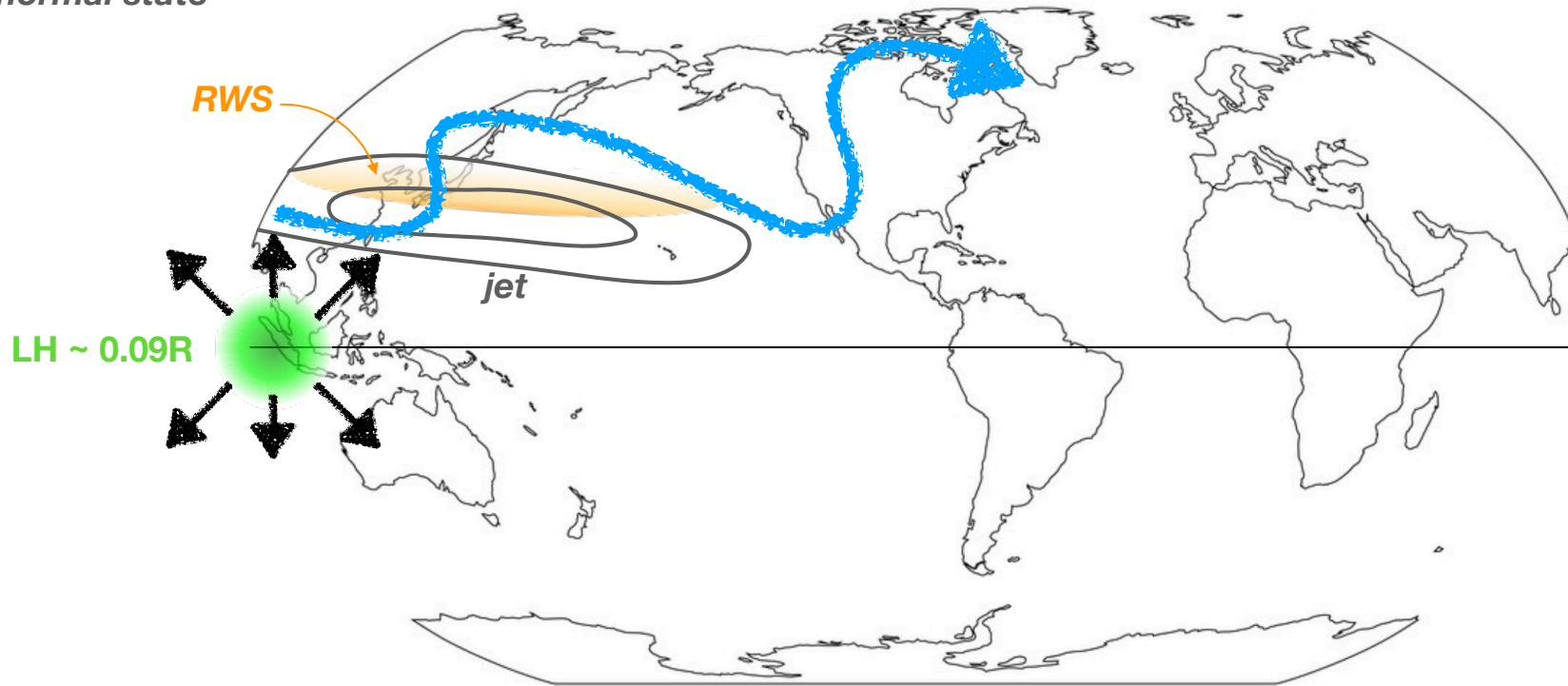
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# coupled feedbacks in weather & climate simulation

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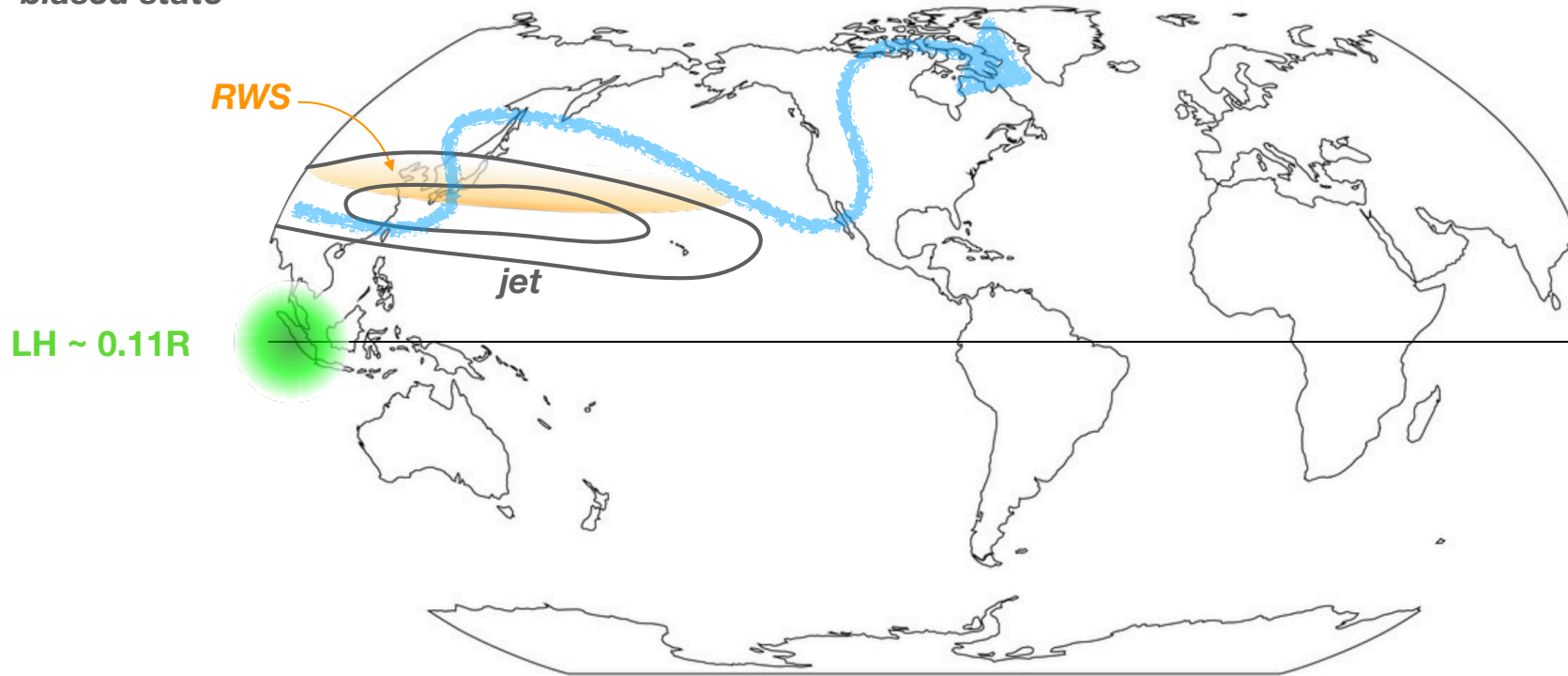


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# coupled feedbacks in weather & climate simulation

## Example: MJO teleconnections

*biased state*

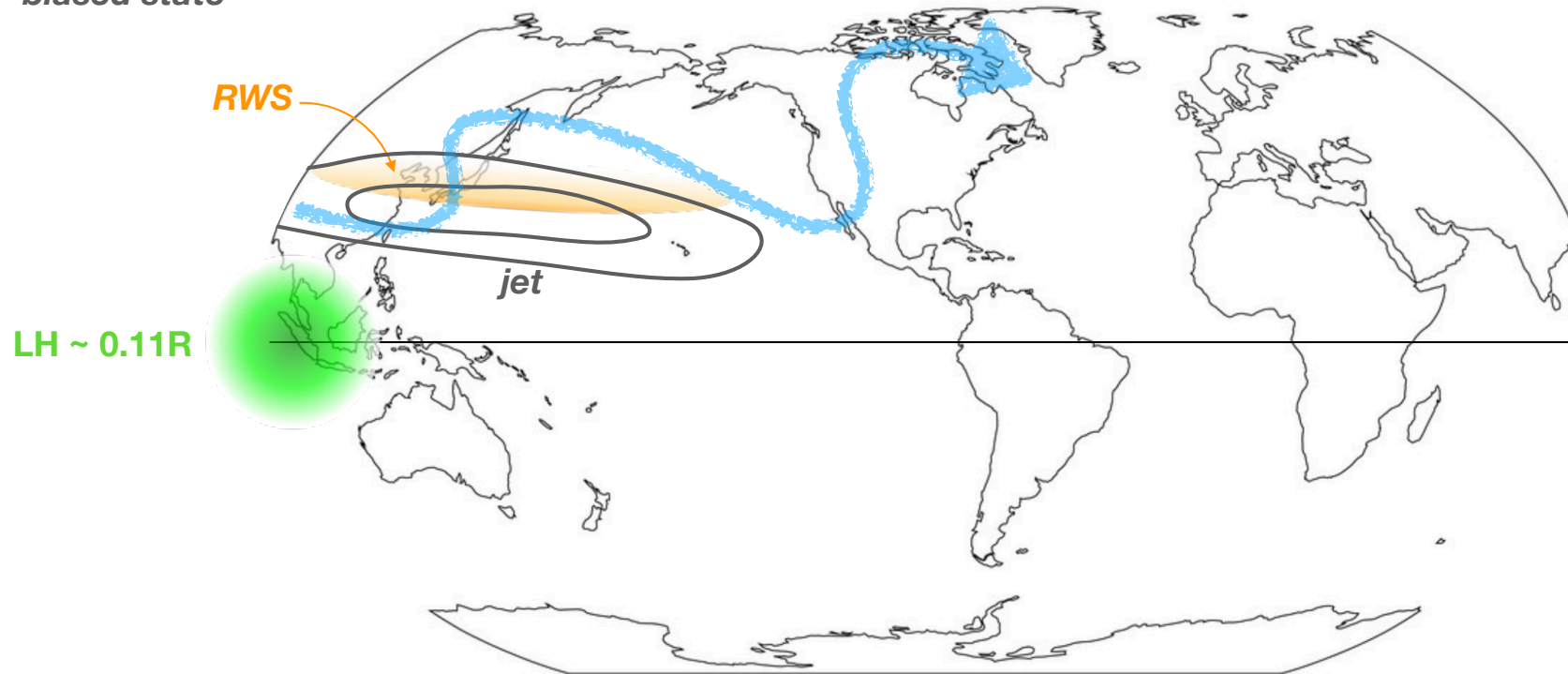


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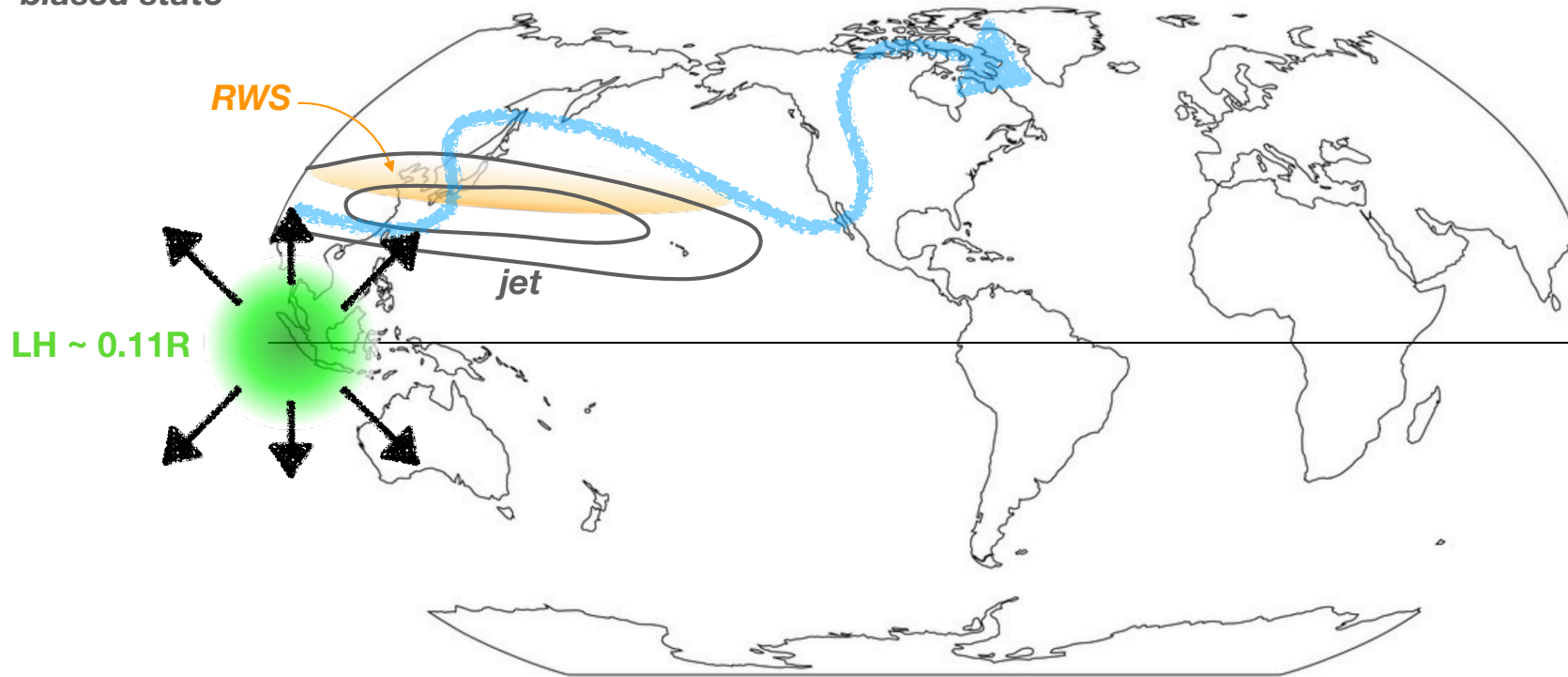


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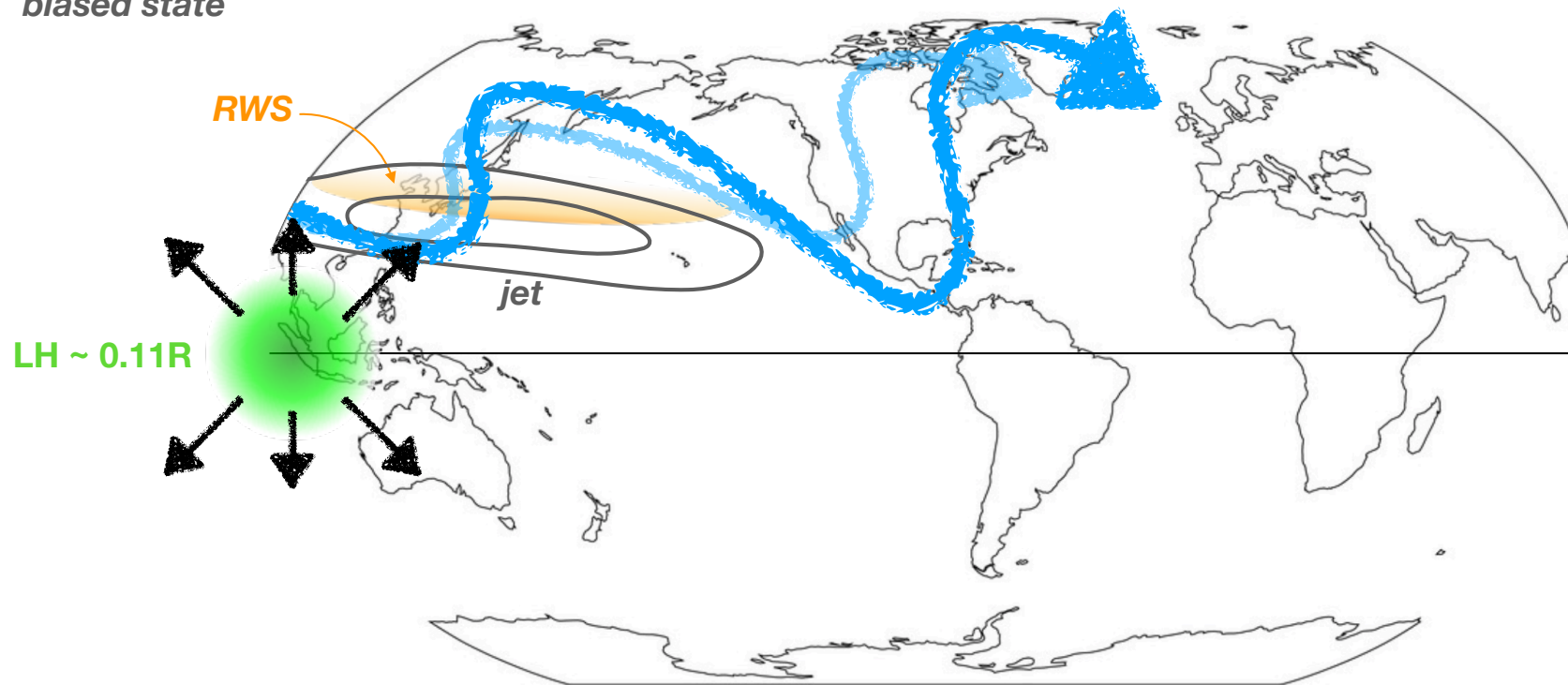


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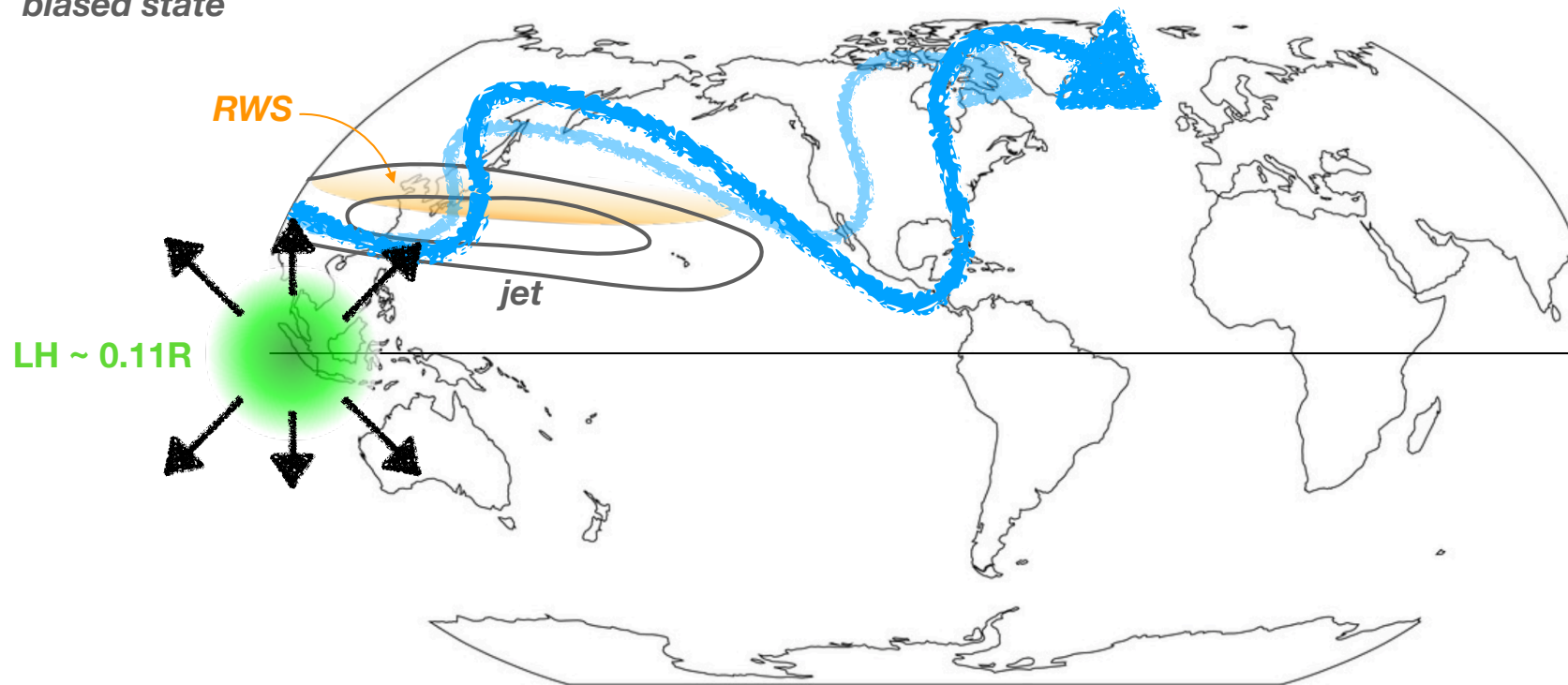


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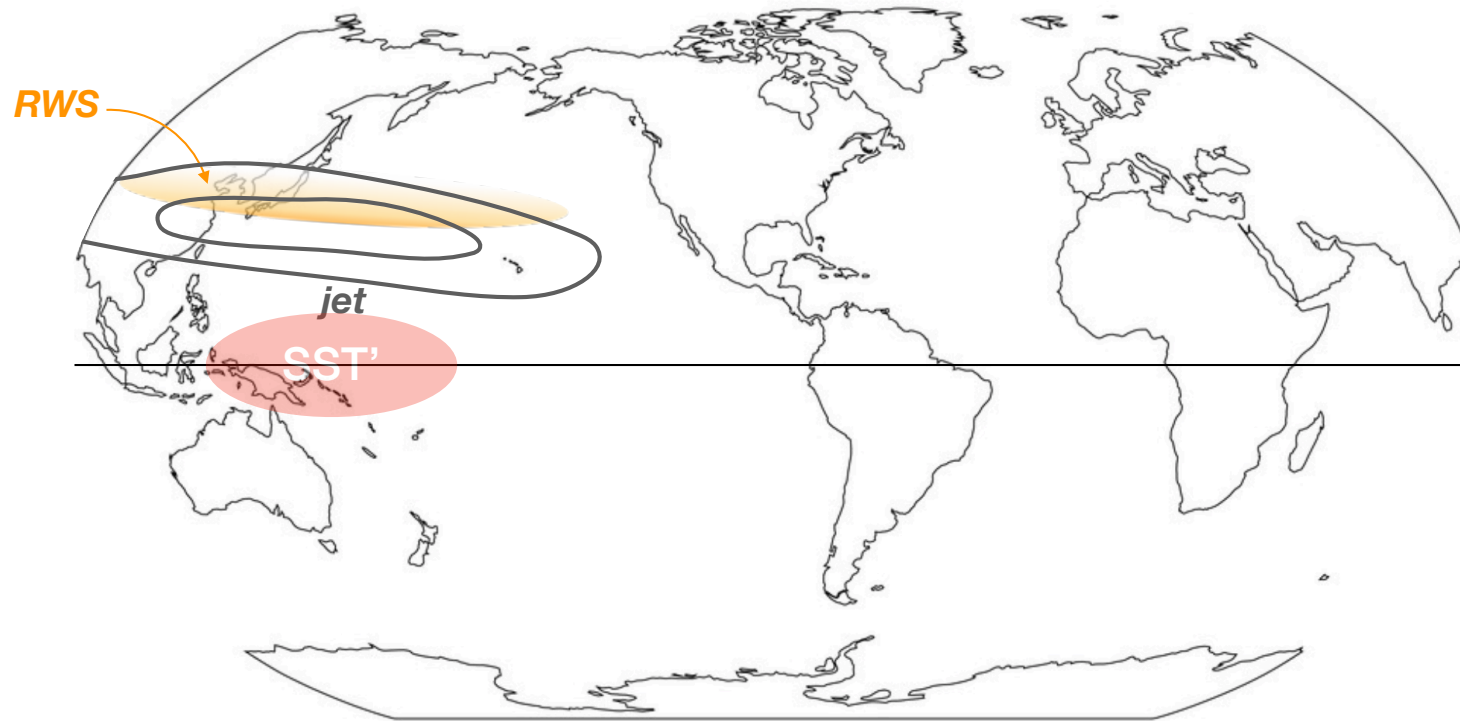


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- MJO amplitude biases -> jet stream biases

# coupled feedbacks in weather & climate simulation

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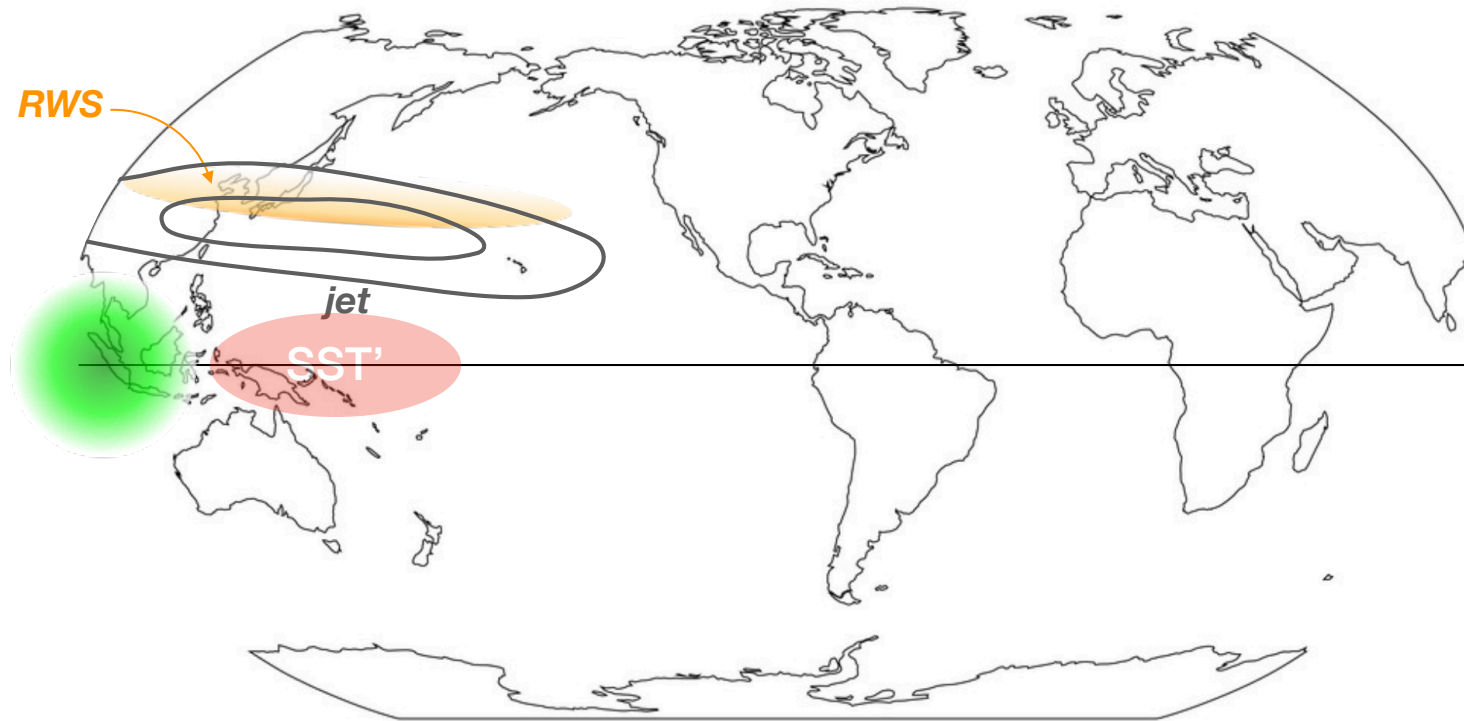


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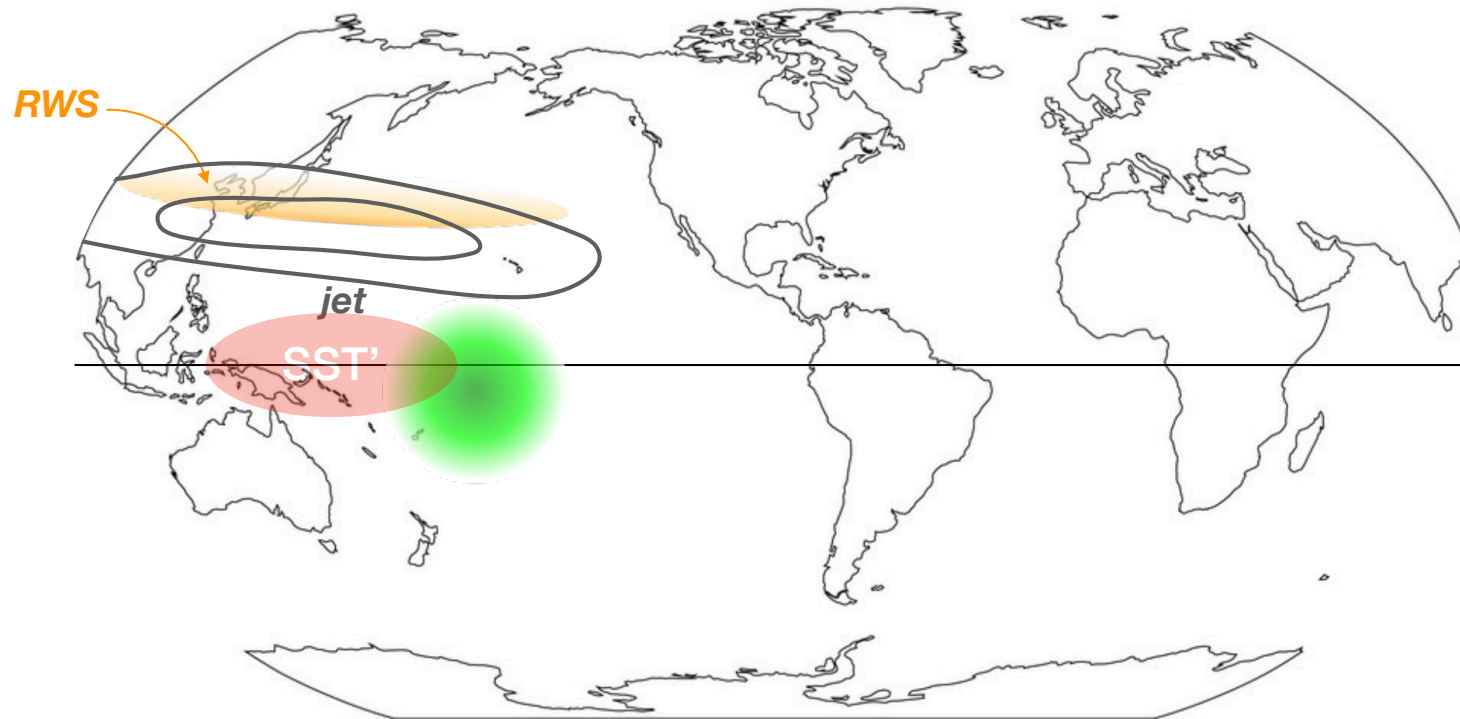
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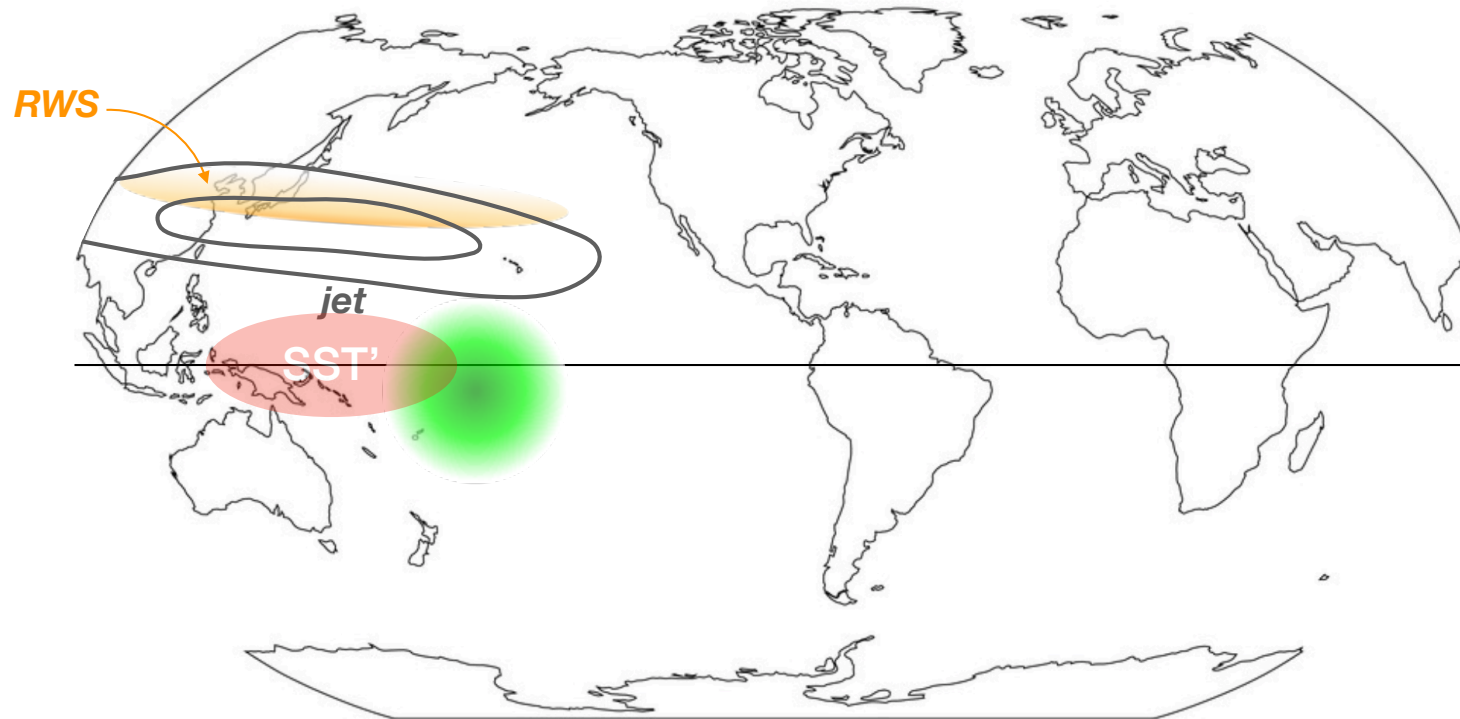


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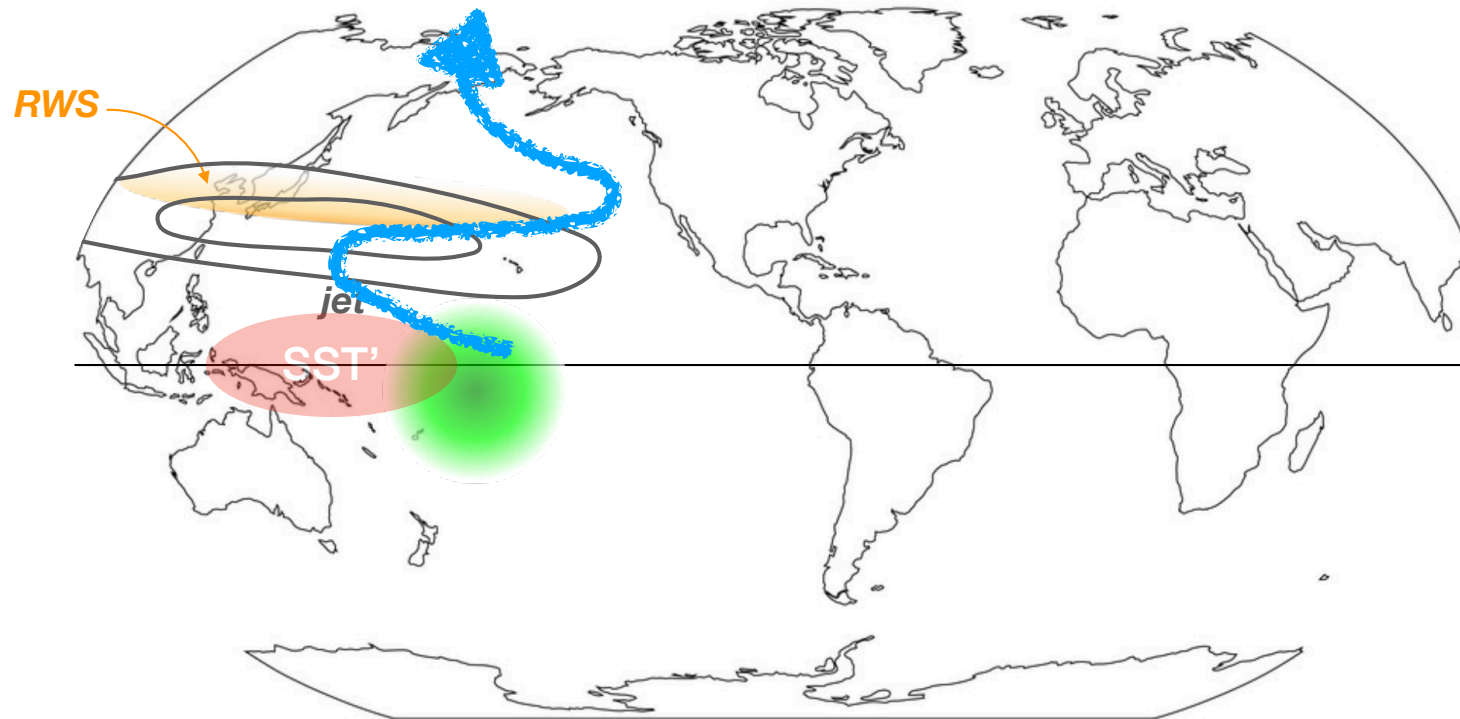


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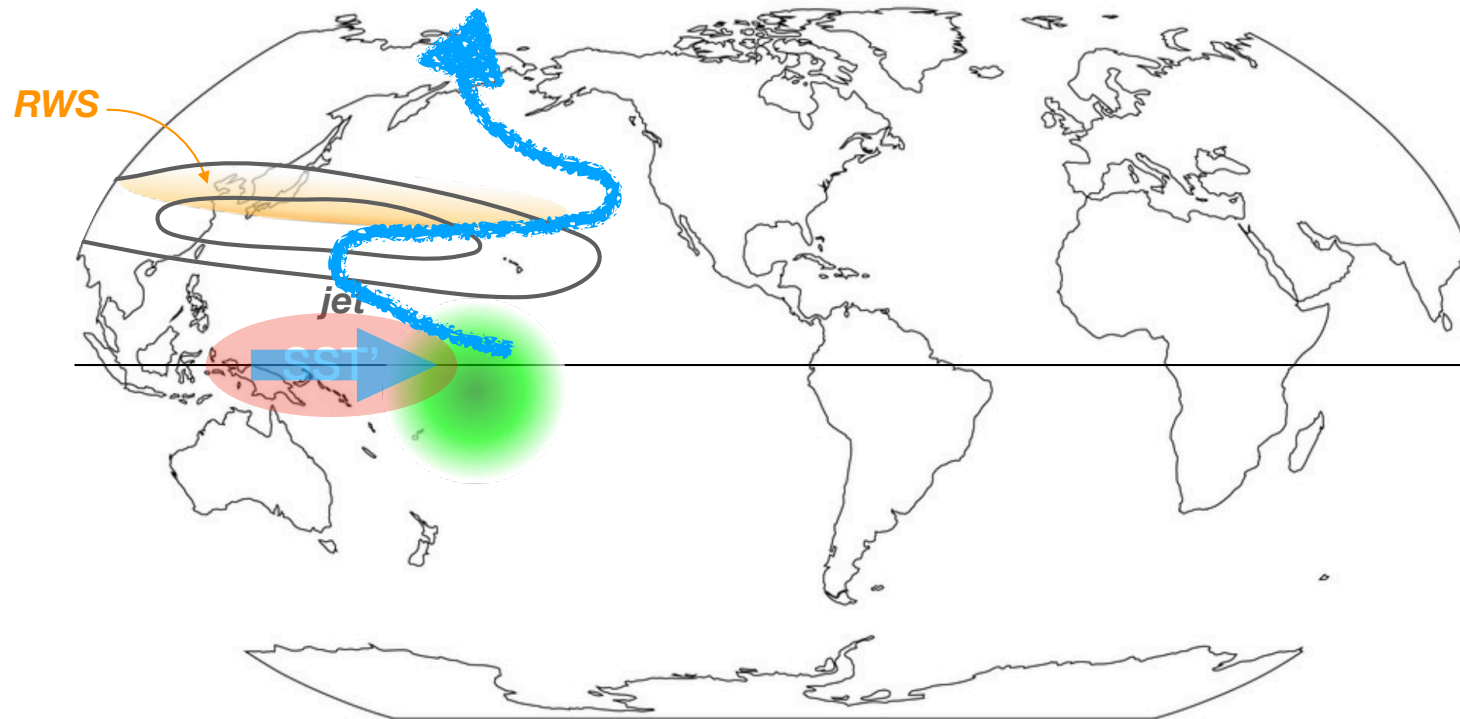


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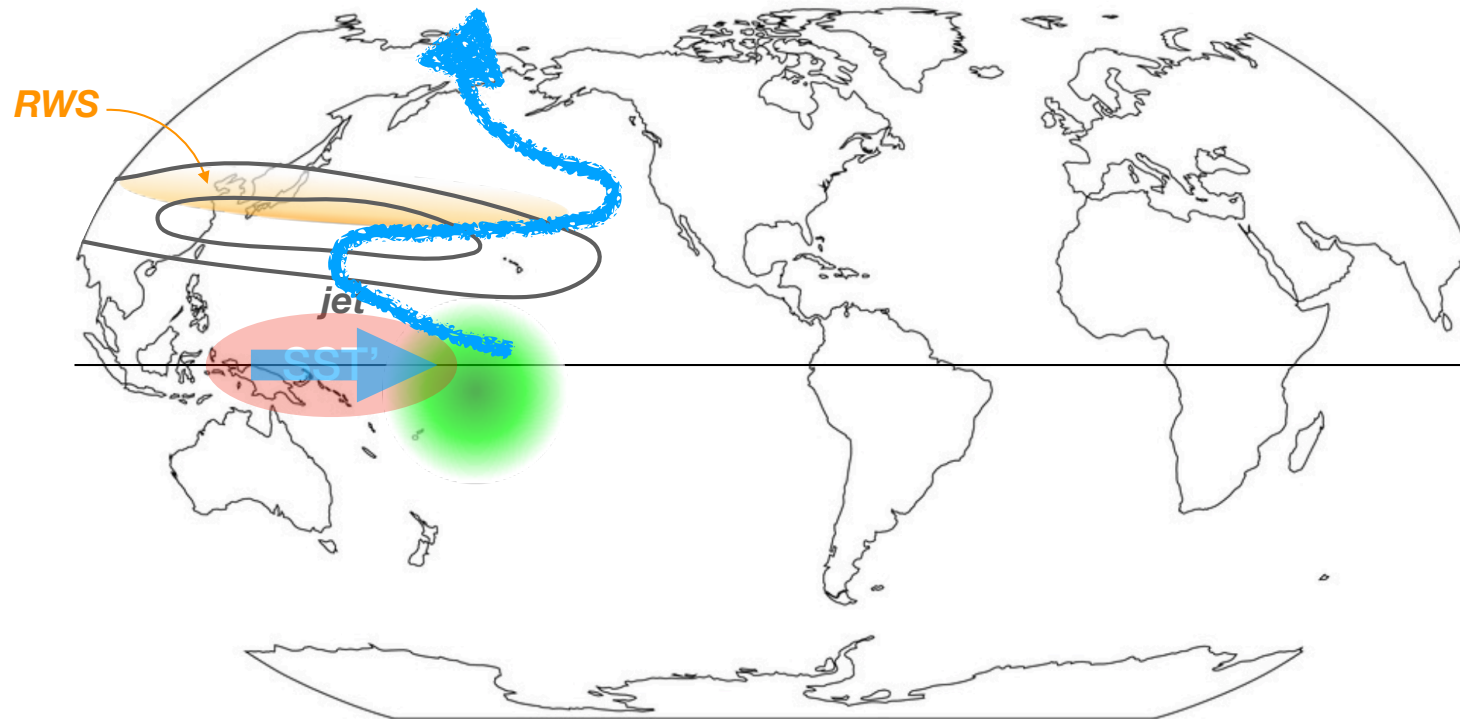


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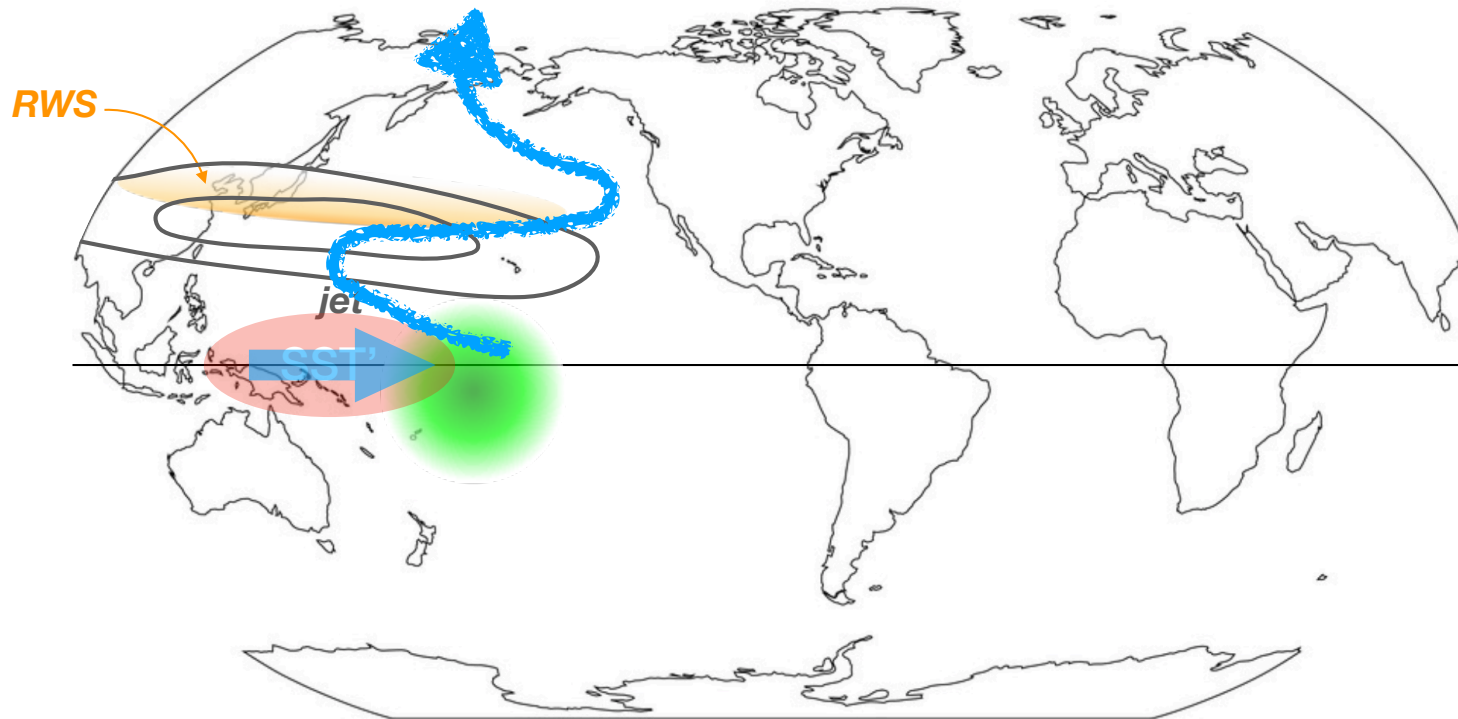


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- MJO propagation biases -> ocean Kelvin wave, ENSO biases

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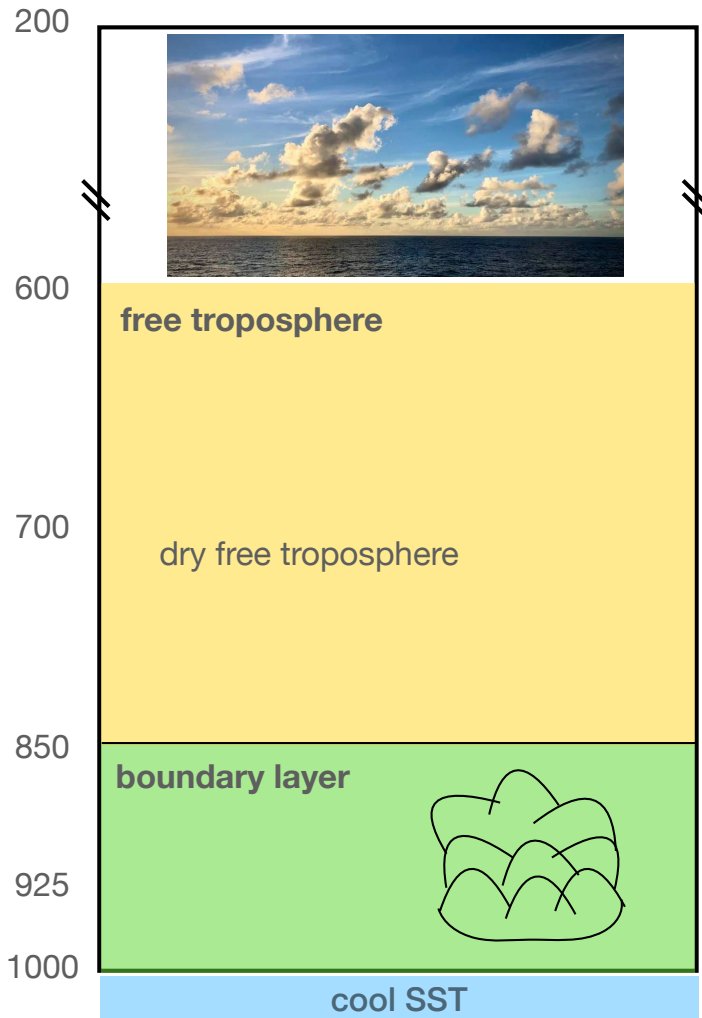
## Example: MJO teleconnections



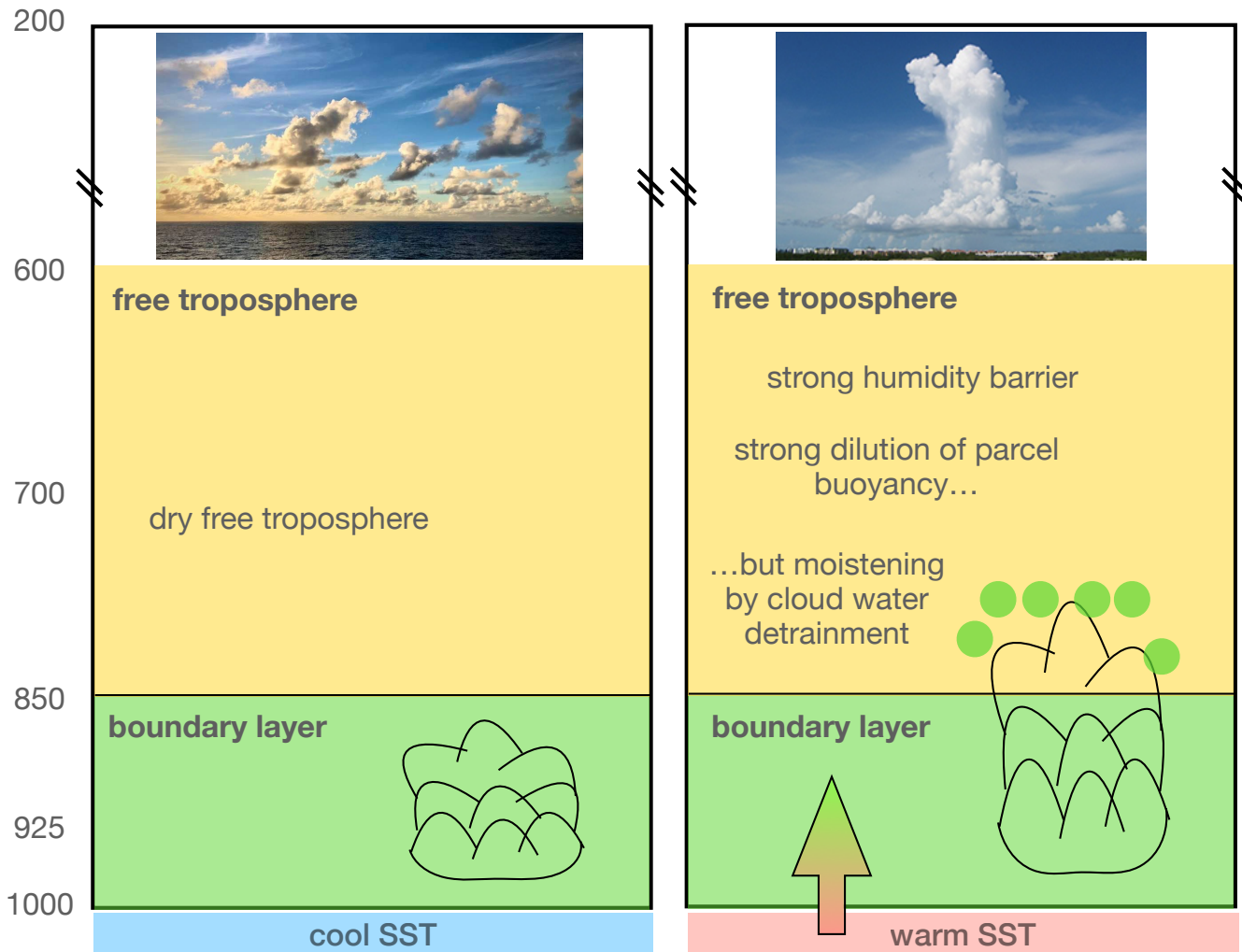
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# buoyancy perspective of tropical convection

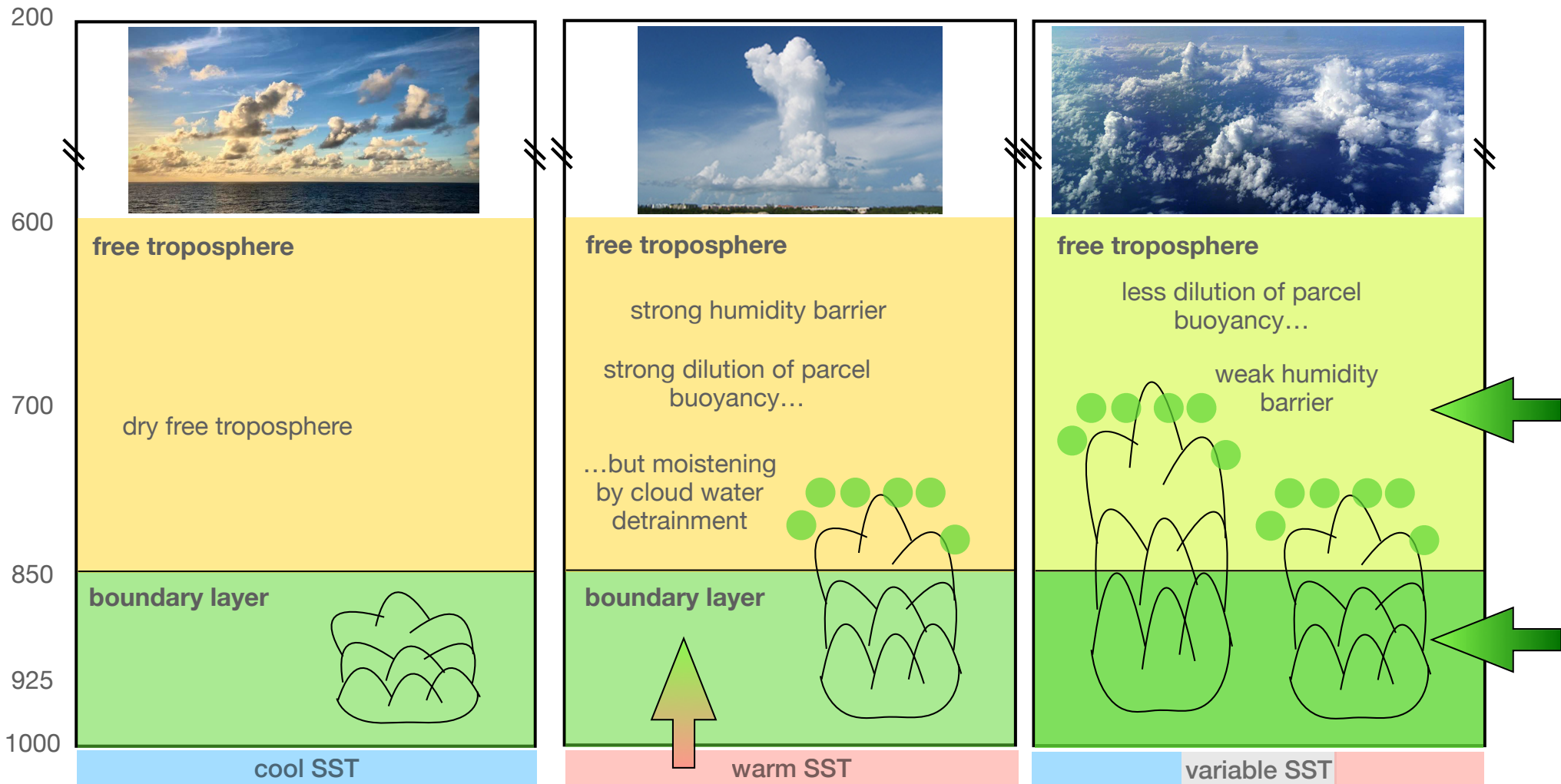


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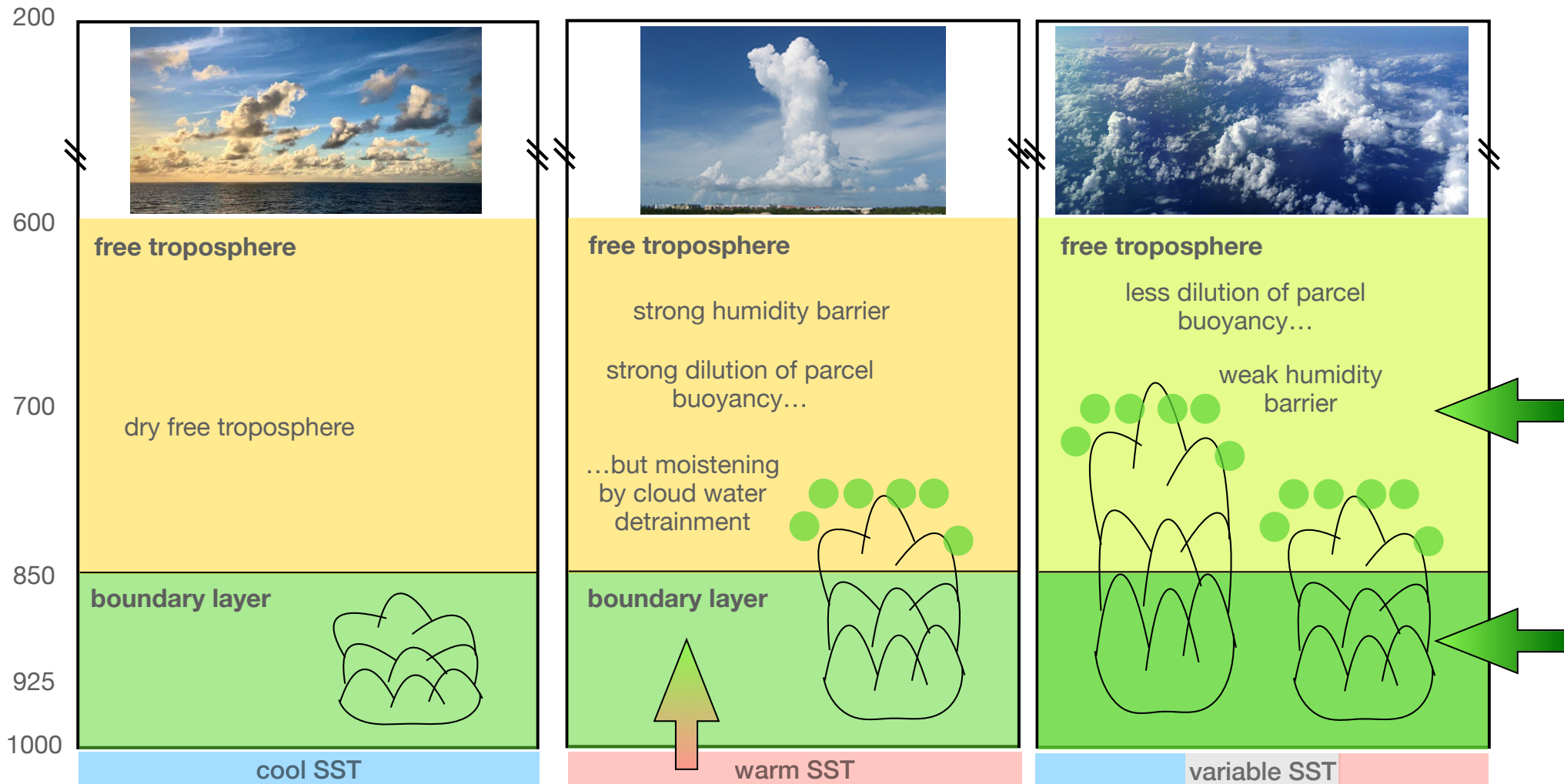




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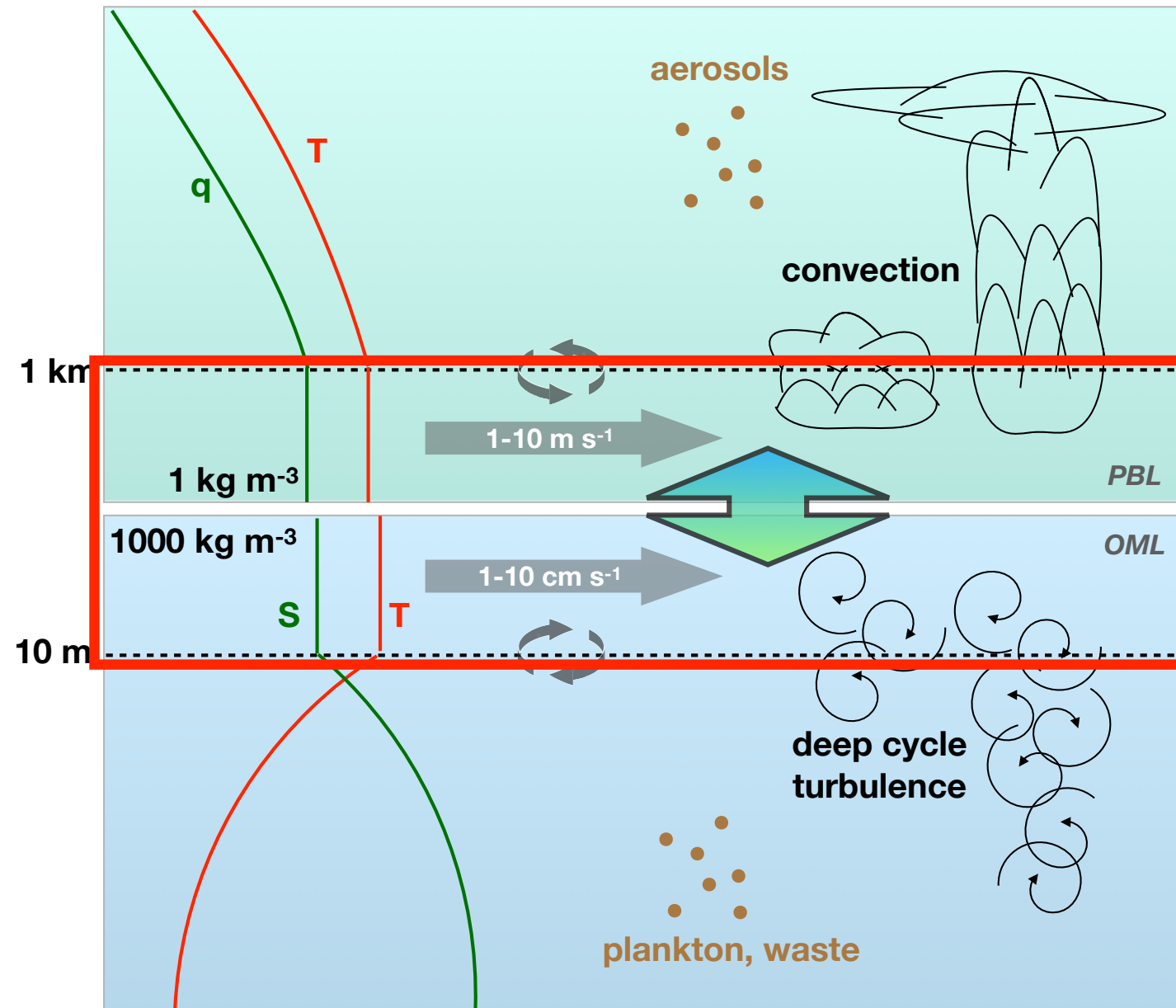
# buoyancy perspective of tropical convection



buoyancy profiles affected by:

- SST (ocean processes)
- T, q in MABL
- T, q above MABL

# ocean-atmosphere differences



- density:  $\delta \sim 10^3$
- ML depth:  $\delta \sim 10^2$
- $u, v$ :  $\delta \sim 10^2$
- clouds are global; DCT mostly on Equator
- clouds regulate radiation; DCT does not

*O-A density differences are responsible for different adjustment timescales*

*surface fluxes are the result of O-A constantly trying to equilibrate*

# summary of introduction

- lower atmosphere, upper ocean vary on different timescales
- O-A surface fluxes destabilize atmospheric boundary layer
- convection “connects” boundary layer and lower free troposphere
- tropical convection initiates teleconnections

# **SST drift in forecast models**

# How might ocean forecasts affect the atmosphere?

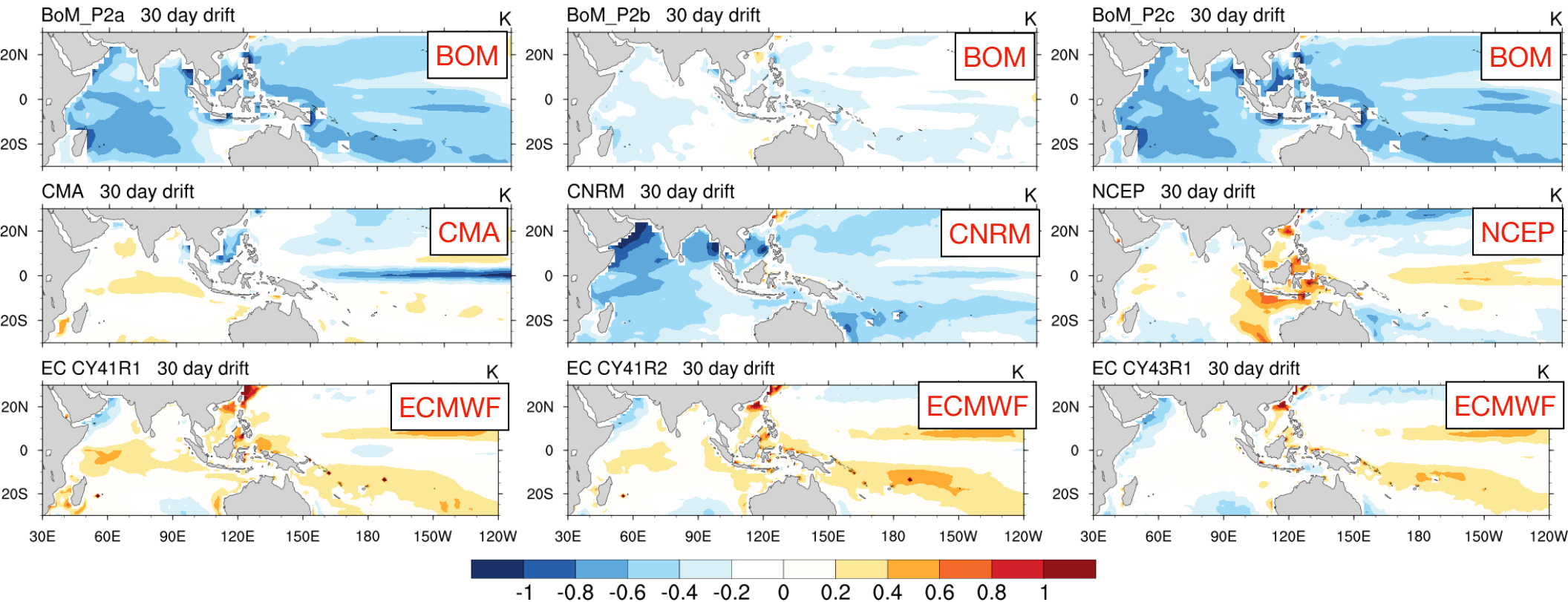
- **model:** ocean initialization
- **model:** SST drift
- **nature:** ocean initial state
- **nature:** ocean evolution

# How might ocean forecasts affect the atmosphere?

- **model**: ocean initialization
  - **model**: SST drift
  - **nature**: ocean initial state
  - **nature**: ocean evolution
- } my focus today

# average 30-day SST drift in S2S database models

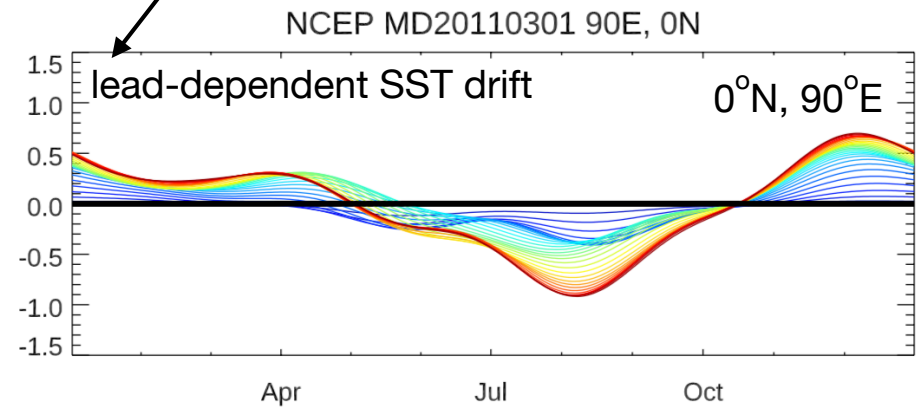
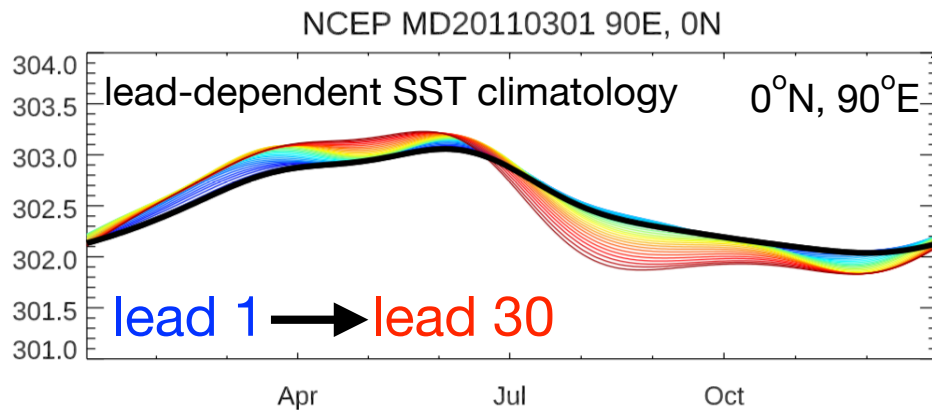
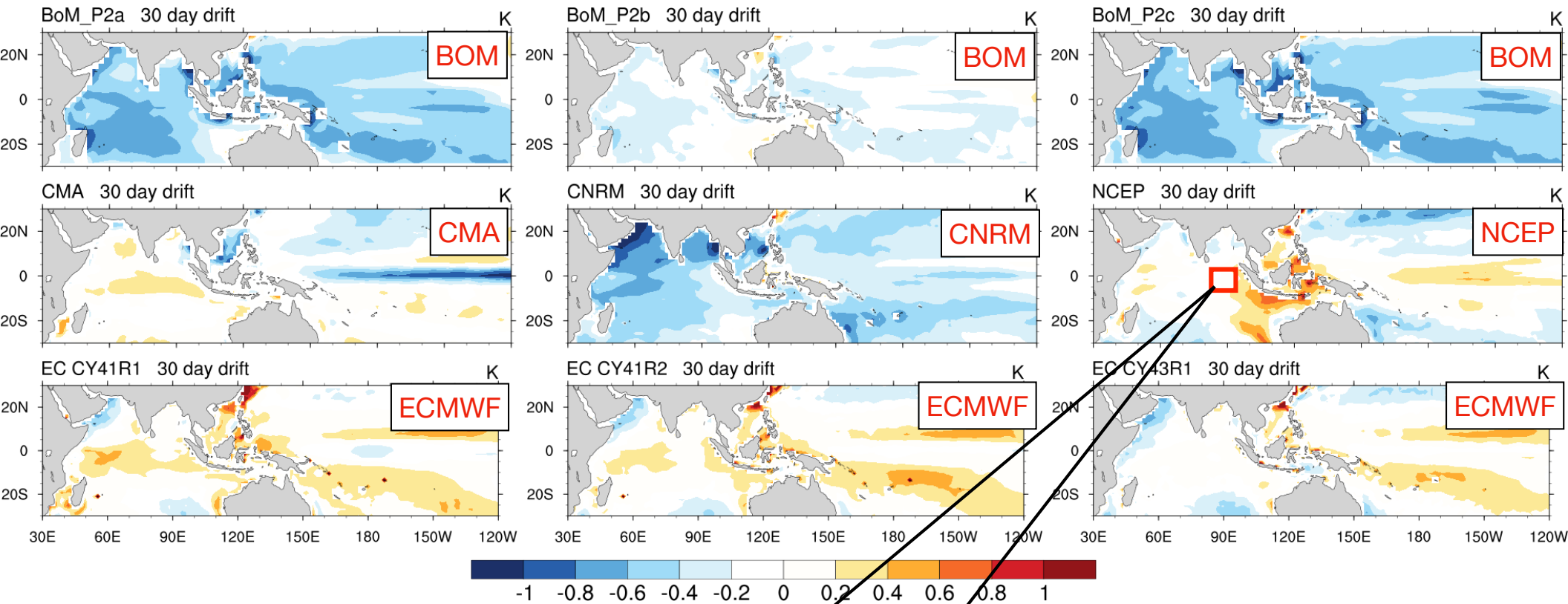
(Nov-Apr climatology)



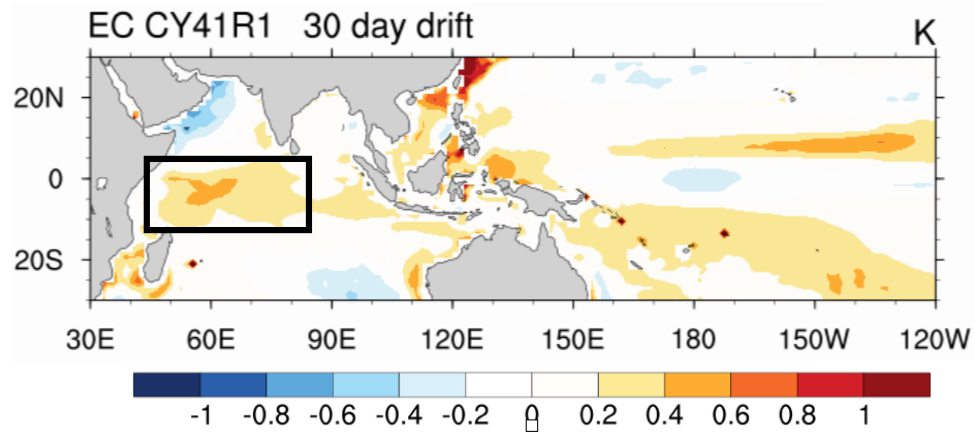


# average 30-day SST drift in S2S database models

(Nov-Apr climatology)



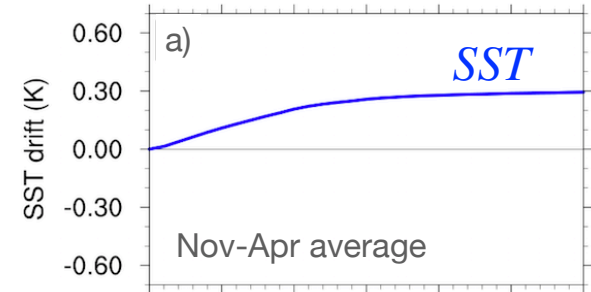
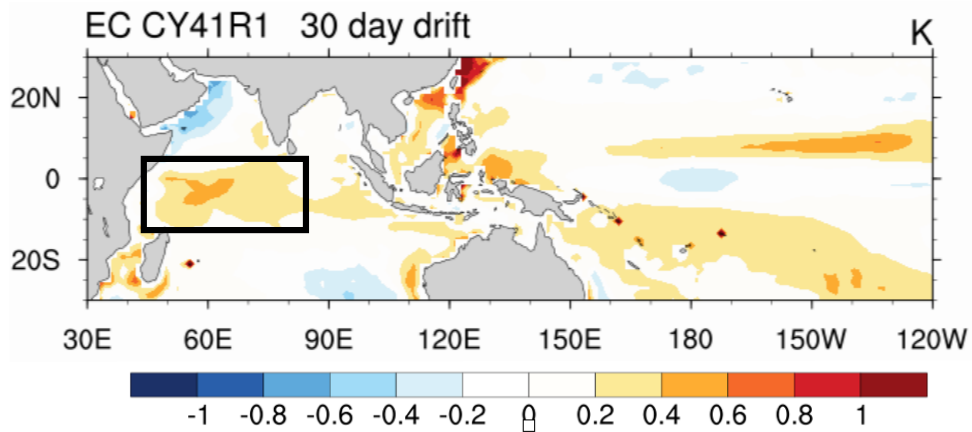
# diagnosing mean state SST drift



SST drift can arise from drift in net surface heating ( $Q_{net}$ ) or ocean processes:

$$\underline{\partial SST / \partial t} \sim \underline{Q_{net}} + \underline{ocnproc}$$

# diagnosing mean state SST drift EC CY41R1



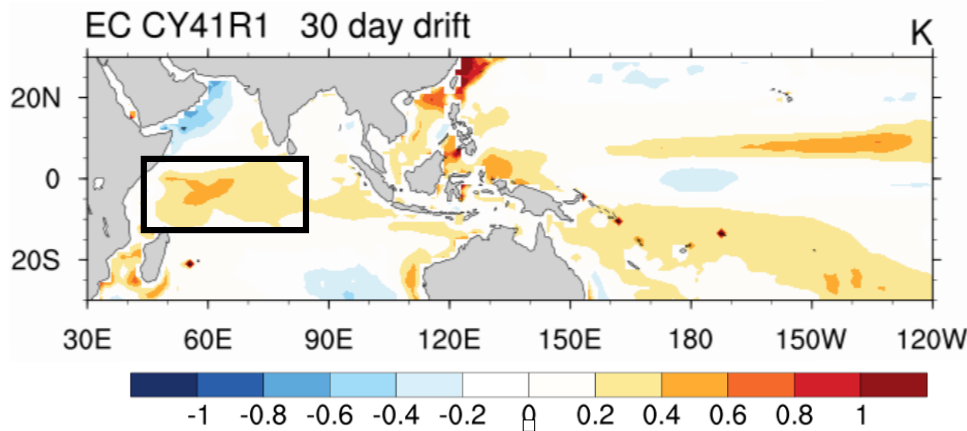
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a) compute SST lead-dependent climatology

0 5 10 15 20 25 30  
lead (days)

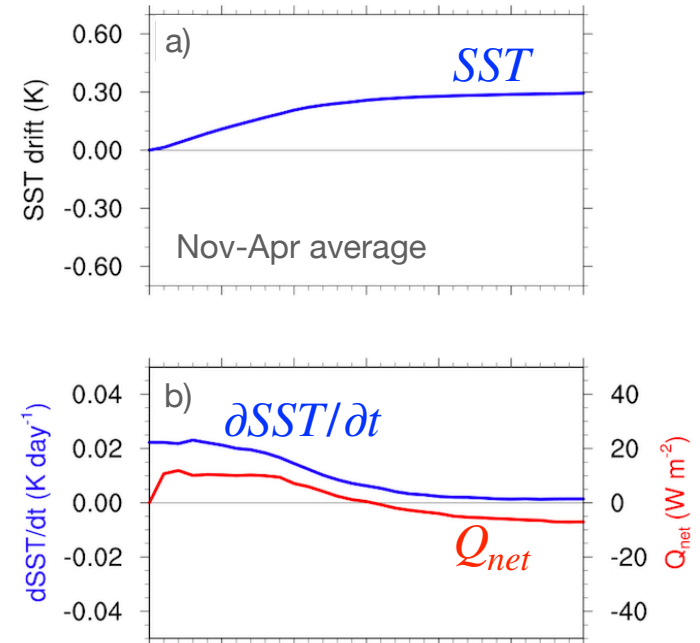
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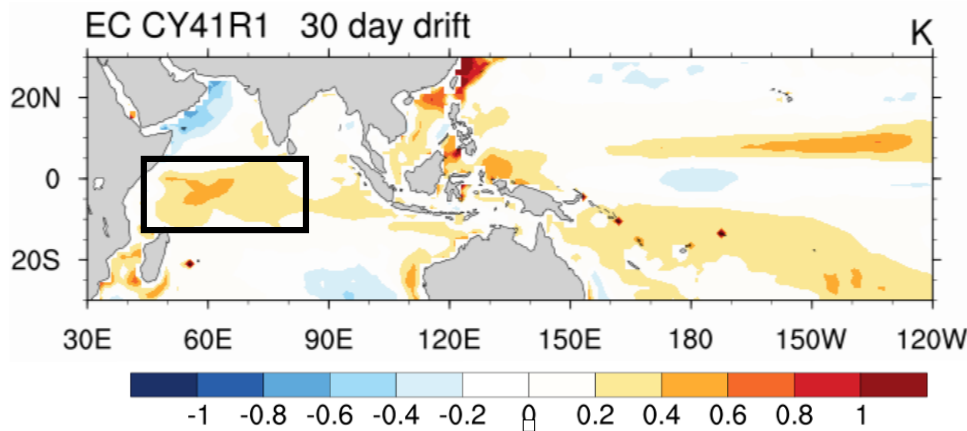
$$\underline{\partial SST / \partial t} \sim \underline{Q_{net}} + \underline{ocnproc}$$

- a) compute SST lead-dependent climatology  
 b) regress  $\partial SST / \partial t$  onto  $Q_{net}$



0 5 10 15 20 25 30  
lead (days)

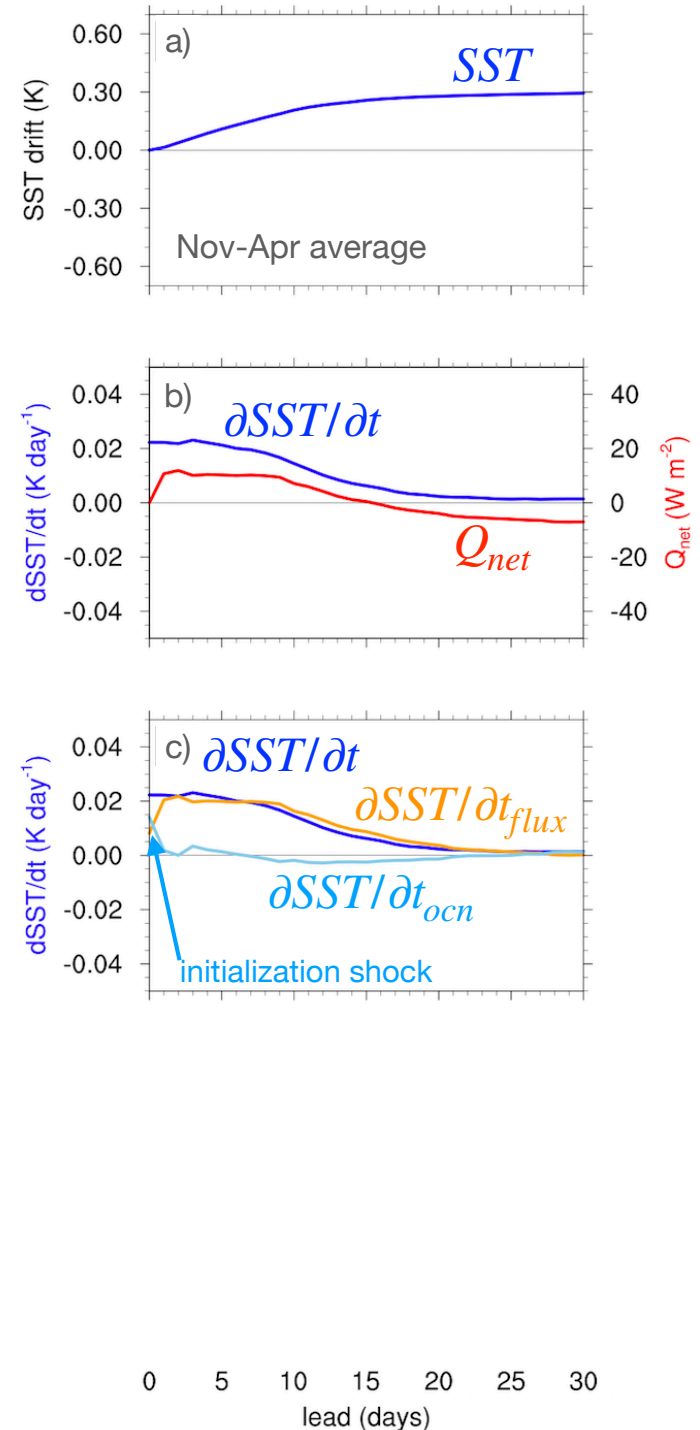
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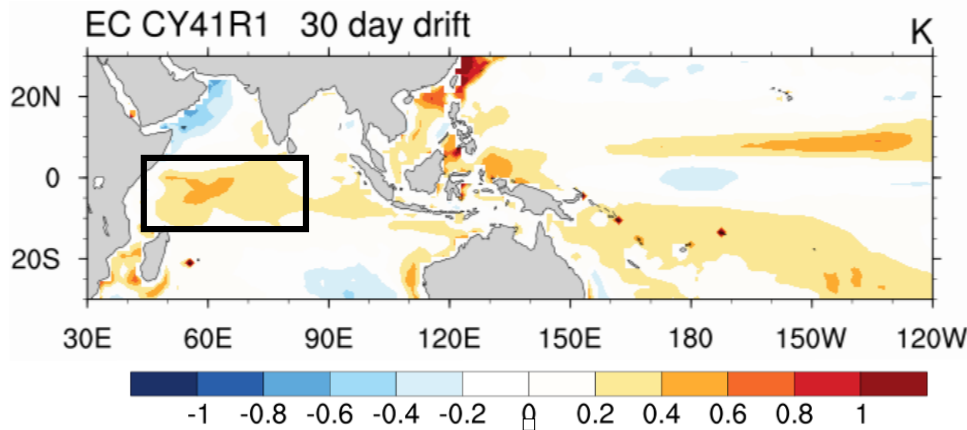
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- c) residual is ocean-driven  $\partial SST / \partial t$



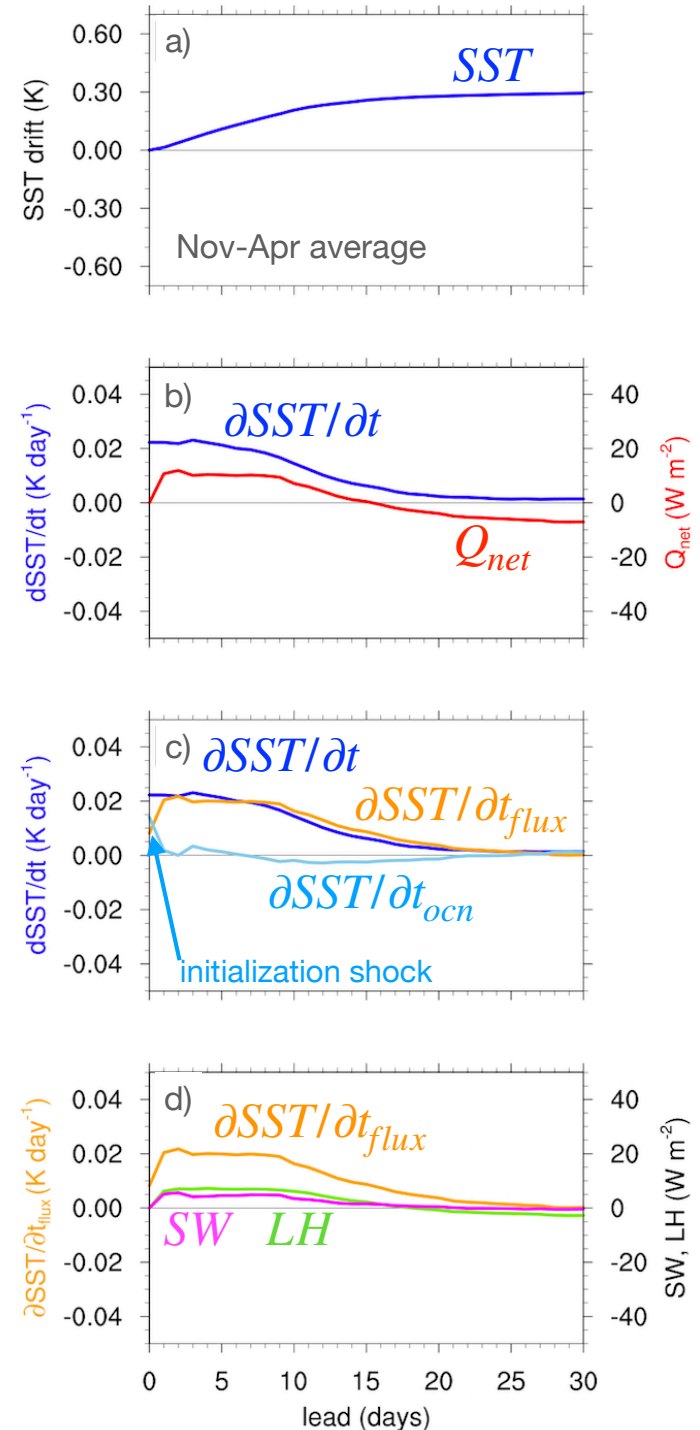
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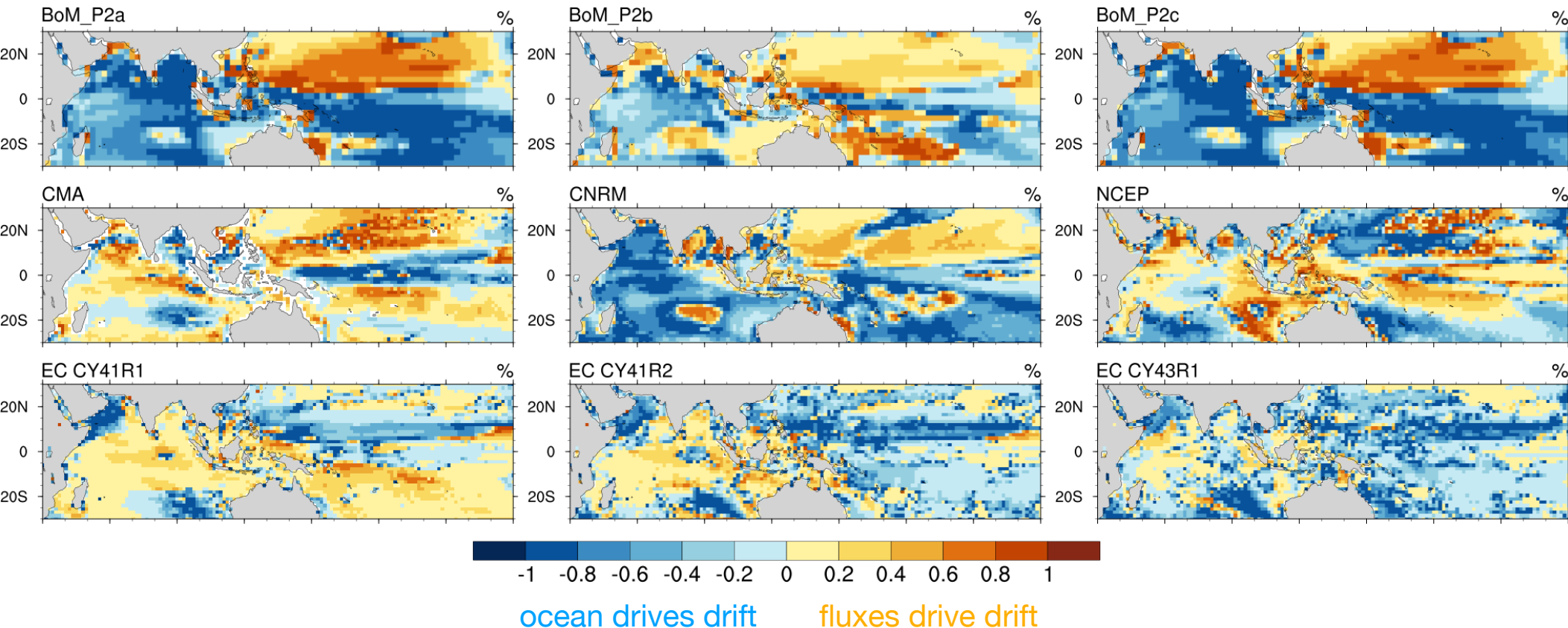
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- c) residual is ocean-driven  $\partial SST / \partial t$
- d) for ECMWF,  $Q_{net}$  contributions equally driven by SW and LH drift (LW, SH contributions are small)



# surface flux vs ocean dynamics for SST drift



as determined by a “balance factor” that compares RMSE of each time series (Halkides et al. 2015)

- large variety of SST drift sources across models
- ocean sources of drift: initialization shock, model biases, insufficient observations for initialization?

# model biases that contribute to SST drift

- erroneous **(de)stabilization** of water column
  - surface heating
  - surface freshening
- erroneous **momentum flux** across air-sea interface
  - surface winds
  - representation of surface momentum flux: bulk parameterization, wave effects
  - representation of ocean mixing: affects surface currents, MLD



**prospects for improvement**

# S2S ocean output variables to identify sources of bias

- **SST**: sea surface temperature
- **SSH**: sea surface height
- **SSS**: sea surface salinity
- *u, v* **surface currents**
- **MLD**: mixed layer depth
- **H300**: 0-300m ocean heat content
- **T20D**: 20C isotherm depth
- **sea ice concentration**

*Ben Webber, Eliza Karłowska; U. East Anglia  
Chris Roberts; ECMWF*

*Michael Jacox, NOAA PSL  
Ángel Muñoz, Columbia/IRI  
Dillon Amaya, NOAA PSL  
Juliana Dias, NOAA PSL*

# future directions: a coupled observing strategy?



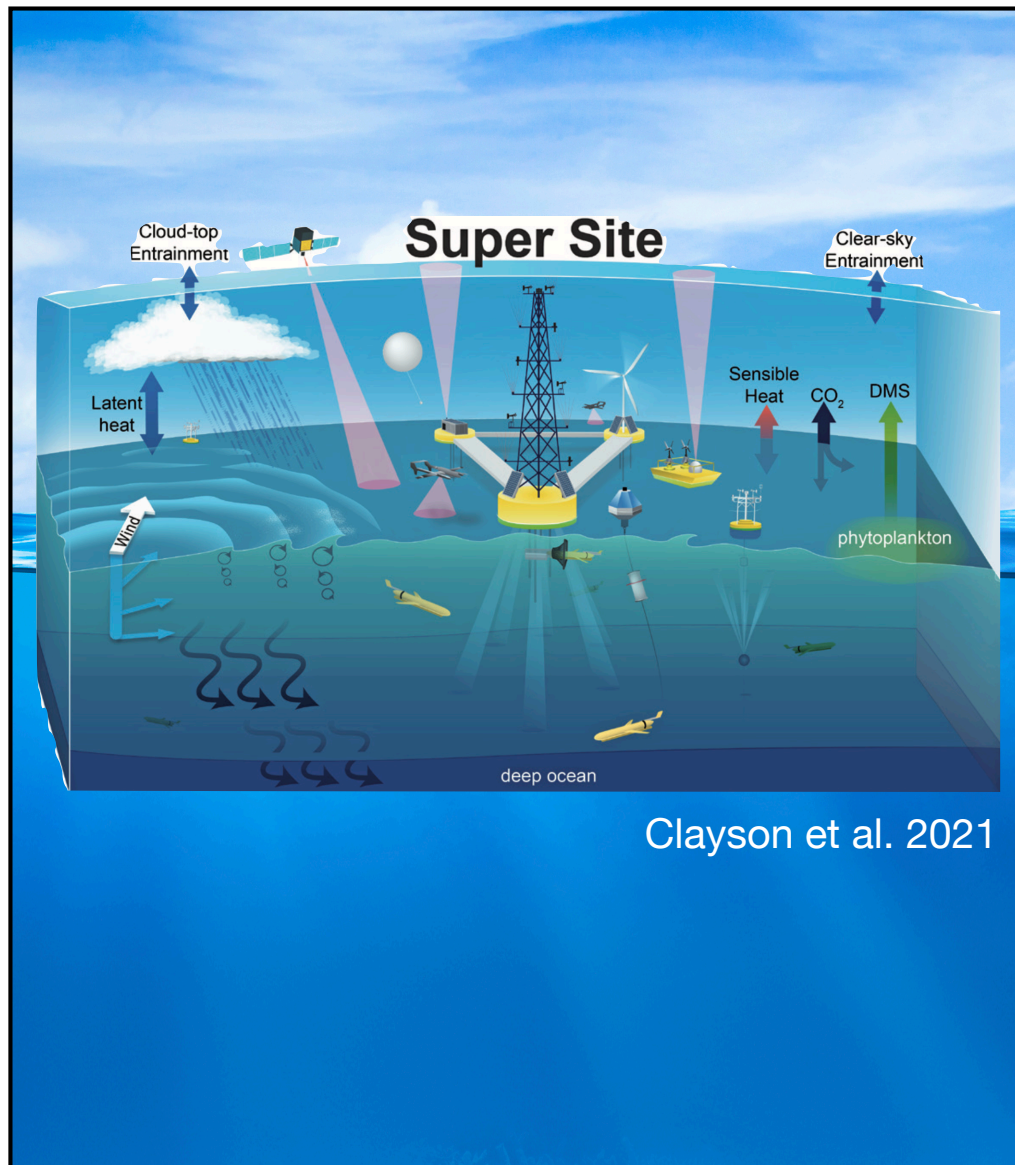
“Whole column” approach to measuring the upper ocean and lower atmosphere in process studies.

Sustained colocated, synchronous measurements of the tropical upper ocean and lower atmosphere are needed to:

- advance process understanding
- improve process representation in climate and forecast models
- accelerate progress in coupled data assimilation for improved forecast reliability

*leveraging ideas put forth by TPOS2020, OASIS, US CLIVAR Tropical Pacific Observing Needs Workshop, and the US CLIVAR Air-Sea Transition Zone Study Group*

# future directions: a coupled observing strategy?



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# Summary

- ocean forecast skill is important for forecasting atmospheric phenomena, and for commercial activity planning.
- ocean simulation (**model**) and ocean state (**nature**) can both affect MJO forecast skill.
- identifying sources of ocean forecast errors is difficult when limited to only SST.
- newly available ocean output variables have the potential to improve ocean process understanding and representation in forecast models