

# Atmospheric energy spectra in kilometer-scale global simulations

Claudia C. Stephan

Max Planck Institute for Meteorology

Julia Duras, Lucas Harris, Daniel Klocke, William M. Putman, Mark Taylor,  
Nils P. Wedi, Nedjeljka Žagar, Florian Ziemer

Tellus A: Dynamic Meteorology and Oceanography, **74**, 280–299, doi: [10.16993/tellusa.26/](https://doi.org/10.16993/tellusa.26/)

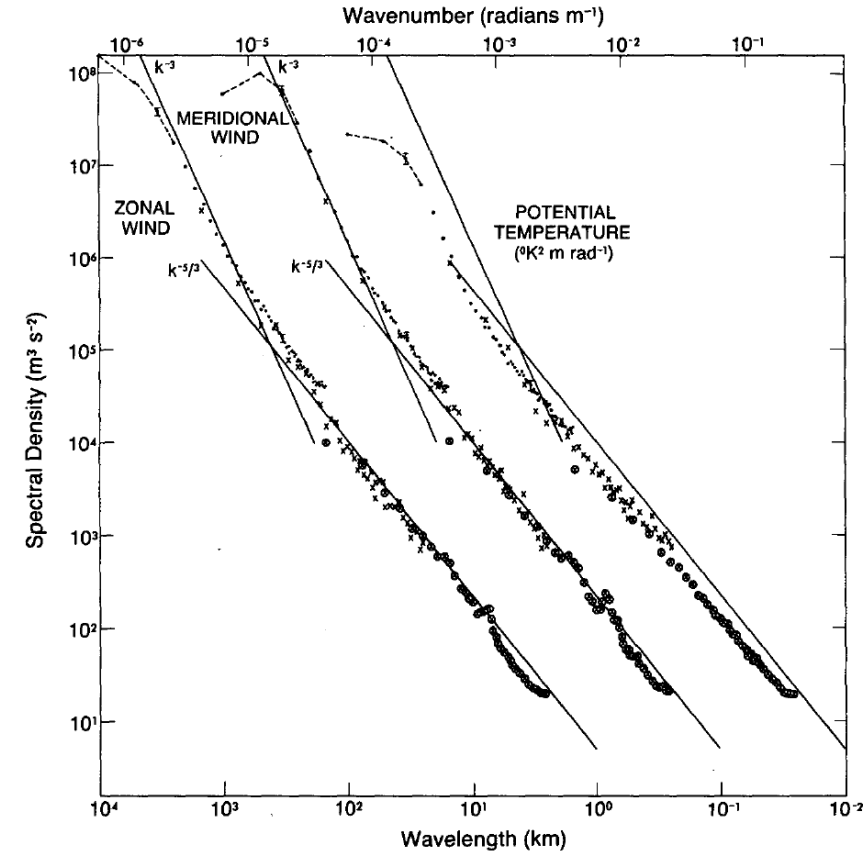
# Horizontal kinetic energy spectra

Synoptic scales:  $k^{-3}$

- non-divergent motions (Fjørtoft 1953)

Mesoscale:  $k^{-5/3}$

- non-linearly interacting inertia-gravity waves (Dewan 1979, VanZandt 1982)
- forward energy cascade (Lindborg 2006)
- non-linear advection (Lindborg and Mohanan 2017)



# Horizontal kinetic energy spectra

Synoptic scales:  $k^{-3}$

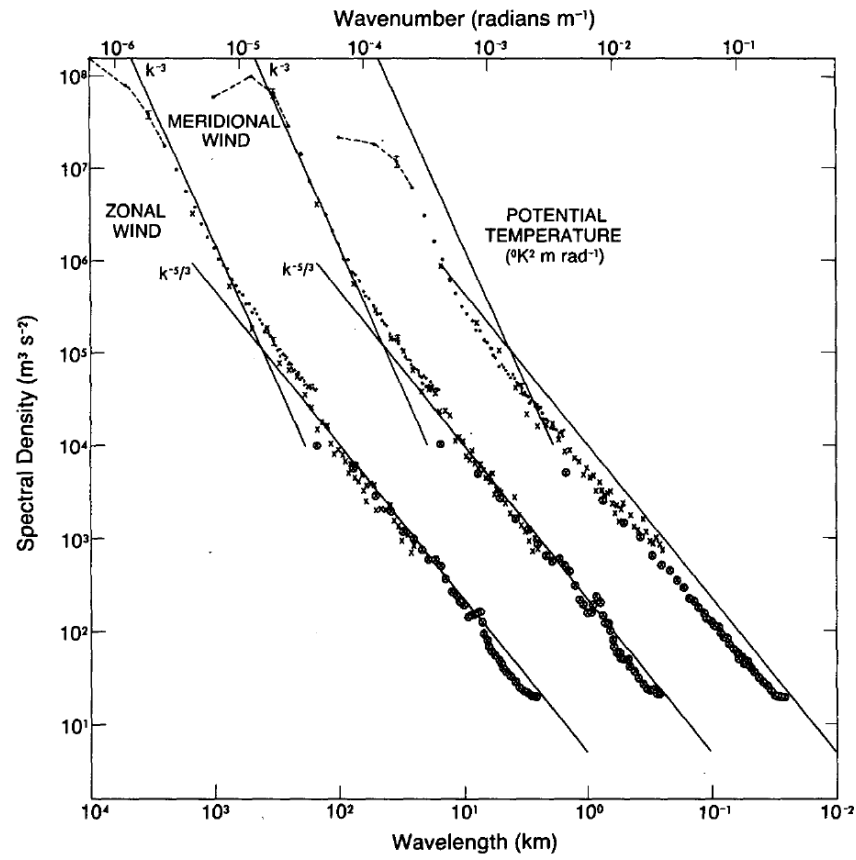
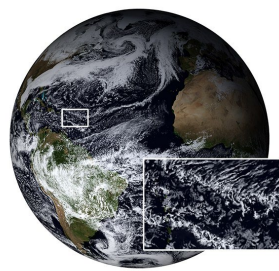
- non-divergent motions (Fjørtoft 1953)

Mesoscale:  $k^{-5/3}$

- non-linearly interacting inertia-gravity waves (Dewan 1979, VanZandt 1982)
- forward energy cascade (Lindborg 2006)
- non-linear advection (Lindborg and Mohanan 2017)

DYAMOND simulations:  $\Delta x < 9 \text{ km}$

- Dynamics of convective storms
- Influence of topography on circulation



# Horizontal kinetic energy spectra

Synoptic scales:  $k^{-3}$

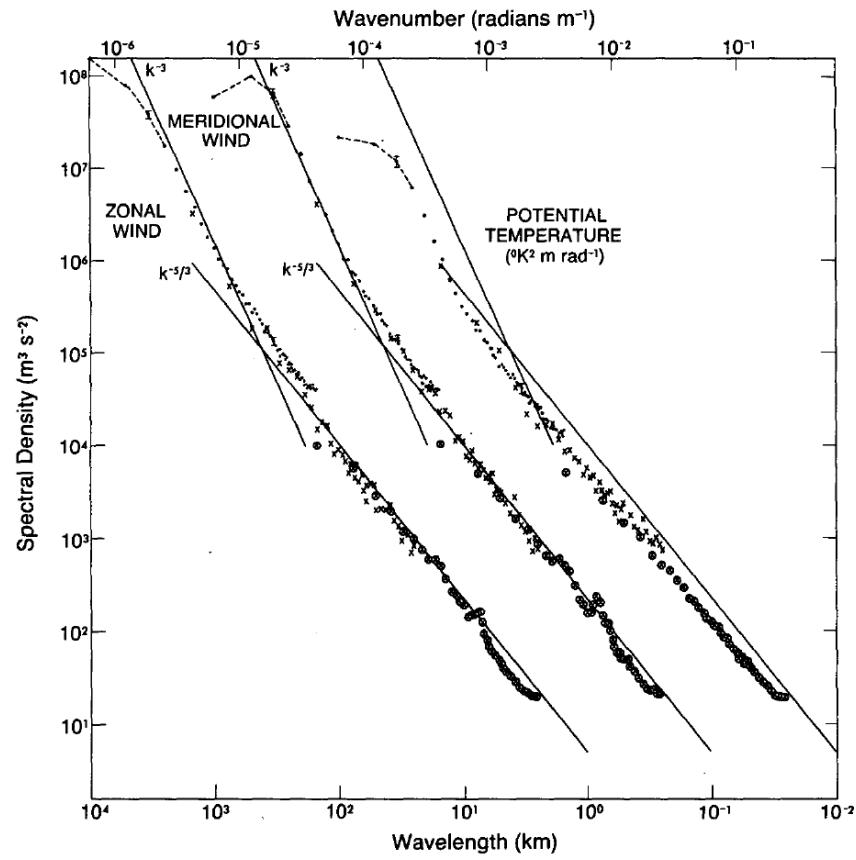
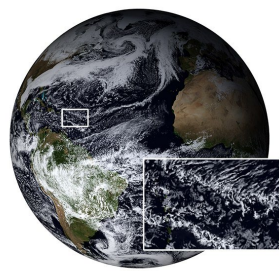
- non-divergent motions (Fjørtoft 1953)

Mesoscale:  $k^{-5/3}$

- non-linearly interacting inertia-gravity waves (Dewan 1979, VanZandt 1982)
- forward energy cascade (Lindborg 2006)
- non-linear advection (Lindborg and Mohanan 2017)

DYAMOND simulations:  $\Delta x < 9$  km

- Dynamics of convective storms
- Influence of topography on circulation
- Microphysics
- Vertical diffusion
- (Convection)



# Horizontal kinetic energy spectra

Synoptic scales:  $k^{-3}$

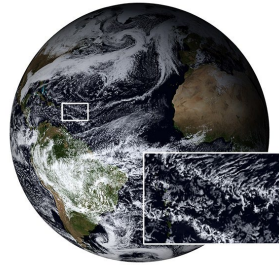
- non-divergent motions (Fjørtoft 1953)

Mesoscale:  $k^{-5/3}$

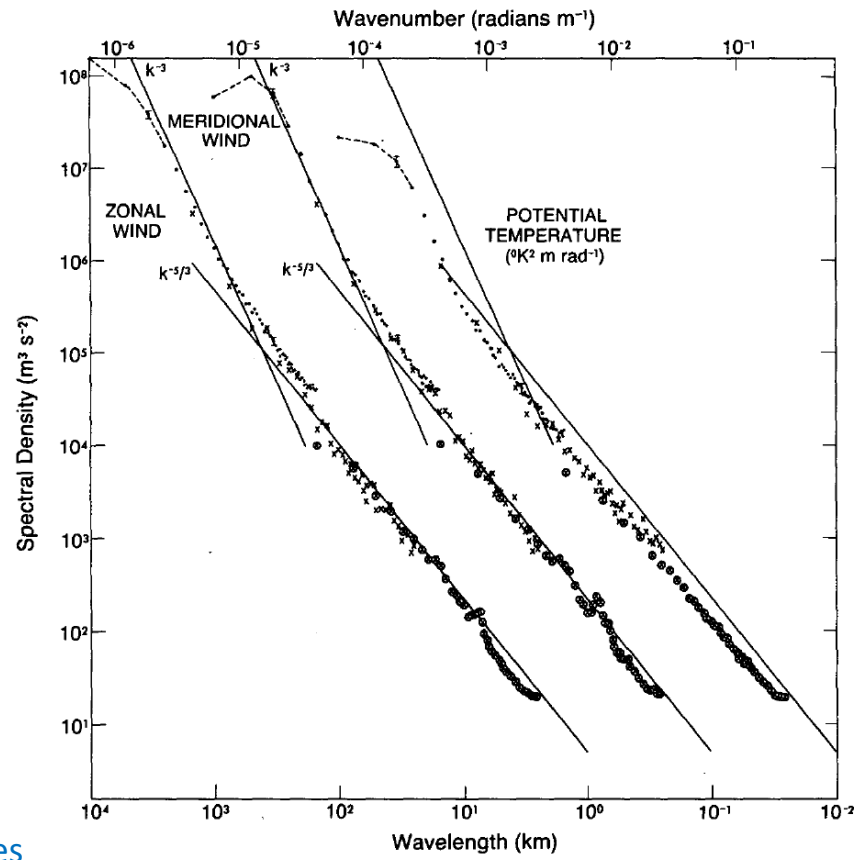
- non-linearly interacting inertia-gravity waves (Dewan 1979, VanZandt 1982)
- forward energy cascade (Lindborg 2006)
- non-linear advection (Lindborg and Mohanan 2017)

DYAMOND simulations:  $\Delta x < 9$  km

- Dynamics of convective storms
- Influence of topography on circulation
- Microphysics
- Vertical diffusion
- (Convection)



Vertical momentum and energy exchanges are to some degree modelled explicitly.



# Expectations based on DYAMOND summer

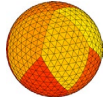
ICON

$\Delta x$ : 5 & 2.5 km



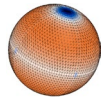
NICAM

$\Delta x$ : 7 & 3.5 km



IFS

$\Delta x$ : 9 & 4 km

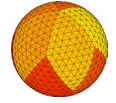


# Expectations based on DYAMOND summer

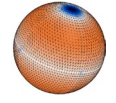
ICON  
 $\Delta x$ : 5 & 2.5 km



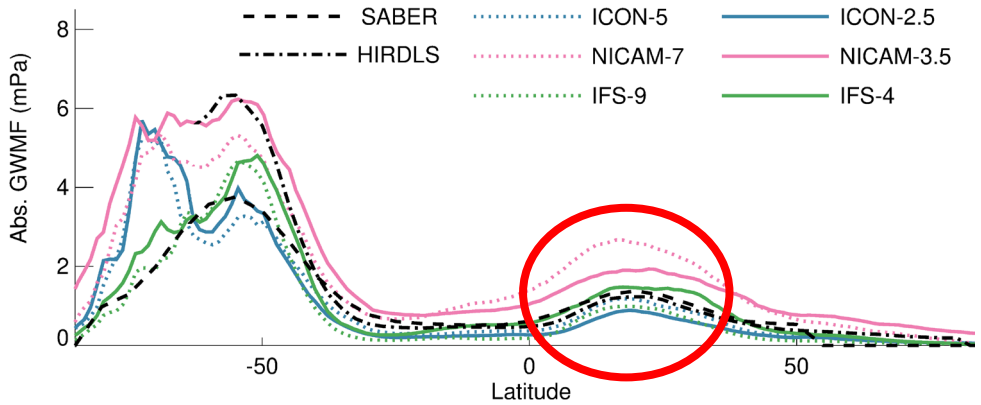
NICAM  
 $\Delta x$ : 7 & 3.5 km



IFS  
 $\Delta x$ : 9 & 4 km



Zonal mean  
absolute GWMF:  
Simulations differ

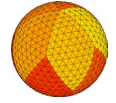


# Expectations based on DYAMOND summer

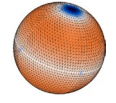
ICON  
 $\Delta x$ : 5 & 2.5 km



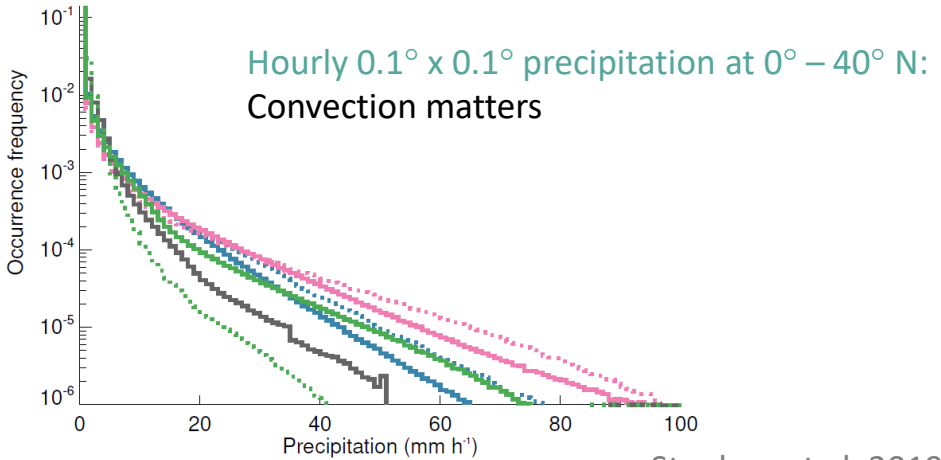
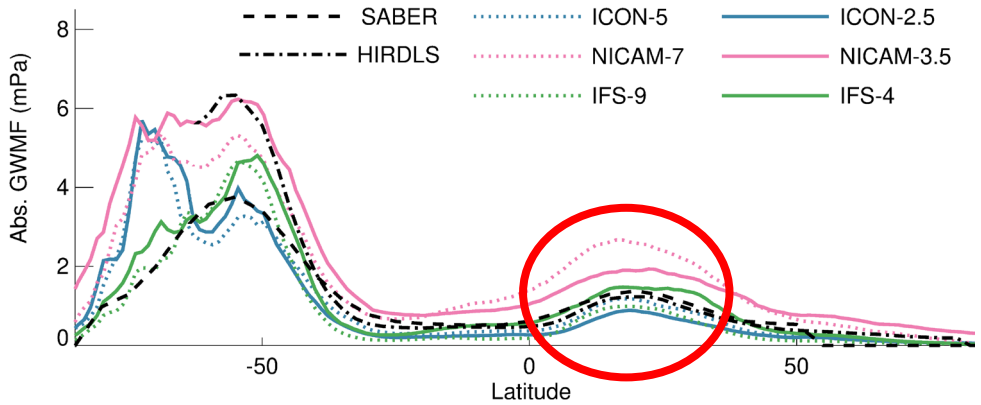
NICAM  
 $\Delta x$ : 7 & 3.5 km



IFS  
 $\Delta x$ : 9 & 4 km


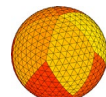
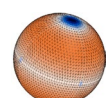


Zonal mean  
 absolute GWMF:  
 Simulations differ

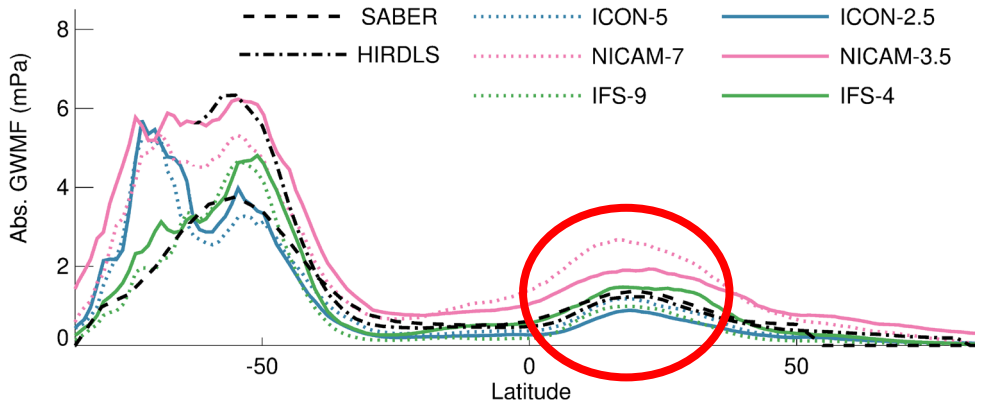




# Expectations based on DYAMOND summer

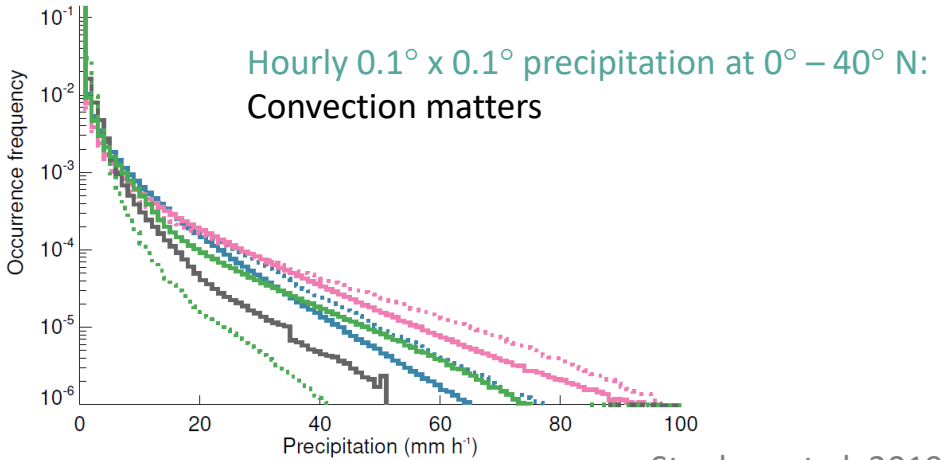
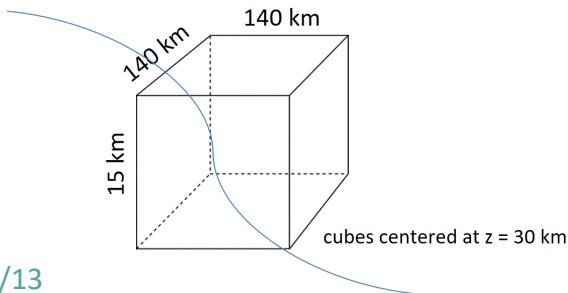
- ICON**  
 $\Delta x$ : 5 & 2.5 km
 
- NICAM**  
 $\Delta x$ : 7 & 3.5 km
 
- IFS**  
 $\Delta x$ : 9 & 4 km
 

Zonal mean  
absolute GWMF:  
Simulations differ



**Limitations of S3D** (Lehmann et al. 2012):

- Focus on predefined wavelengths
- Imperfect filtering of gravity waves
- Focus on a single height level



# Today: a different perspective

## **Normal mode function (NMF) decomposition:**

Using the software package MODES (Žagar et al. 2015)

It projects the **3D** fields of geopotential height and horizontal winds onto an orthogonal set of NMFs.

## **Properties:**

- the NMFs are eigensolutions to the linearized primitive equations
- clean separation into balanced and unbalanced circulation

# Today: a different perspective

## Normal mode function (NMF) decomposition:

Using the software package MODES (Žagar et al. 2015)

It projects the **3D** fields of geopotential height and horizontal winds onto an orthogonal set of NMFs.

## Properties:

- the NMFs are eigensolutions to the linearized primitive equations
- clean separation into **balanced** and **unbalanced circulation**

Rossby waves (RWs)

Inertia-gravity waves (IGWs)

# Today: a different perspective

## Normal mode function (NMF) decomposition:

Using the software package MODES (Žagar et al. 2015)

It projects the **3D** fields of geopotential height and horizontal winds onto an orthogonal set of NMFs.

## Properties:

- the NMFs are eigensolutions to the linearized primitive equations
- clean separation into **balanced** and **unbalanced circulation**

Rossby waves (RWs)

Inertia-gravity waves (IGWs)

**Important:** We derive the **3D** kinetic plus potential available energy spectra of horizontal motions.

# Today: a different perspective

## Normal mode function (NMF) decomposition:

Using the software package MODES (Žagar et al. 2015)

It projects the **3D** fields of geopotential height and horizontal winds onto an orthogonal set of NMFs.

## Properties:

- the NMFs are eigensolutions to the linearized primitive equations
- clean separation into **balanced** and **unbalanced circulation**

Rossby waves (RWs)

Inertia-gravity waves (IGWs)

**Important:** We derive the **3D** kinetic plus potential available energy spectra of horizontal motions.

This is the first time NMF decomposition is used for inter-comparing high-resolution models.

- How closely do the models produce the canonical spectra?
- What is different in terms of total energy levels / synoptic and sub-synoptic slopes / the crossing scale ?
- Can we understand the origin of these differences?

# The simulations

Initialized 20th January 2020 with the global 9 km meteorological analysis from the ECMWF  
Freely evolving until 1st March 2020.

Simulation	$\sqrt{A_{mean}}$	$H_t$	Grid	Coupled	Conv.	BL	SSO	comments
IFS-9	9	80	Octo	yes	F	K	yes	hydrostatic
IFS-4	4.5	80	Octo	yes	S	"	"	
ICON-nwp	2.5	75	Icoso	no	-	TKE	no	
ICON-sap	5	"	"	yes	"	S	"	
ICON-sap+	"	"	"	yes	"	"	"	continuation of ICON-sap
ICON-vdu	"	"	"	no	"	TTE	"	
ICON-vdc	"	"	"	yes	"	"	"	
ICON-vda	"	"	"	yes	"	"	"	increased albedo
GEOS	3	80	Cube	no	F	K	yes	deep plumes disabled
SHIELD	3	40	"	mixed-layer ocean	S	TKE	yes	
SCREAM	3	40	"	no	-	SHOC	no	

# The simulations

Initialized 20th January 2020 with the global 9 km meteorological analysis from the ECMWF  
Freely evolving until 1st March 2020.

Simulation	$\sqrt{A_{mean}}$	$H_t$	Grid	Coupled	Conv.	BL	SSO	comments
IFS-9	9	80	Octo	yes	F	K	yes	hydrostatic
IFS-4	4.5	80	Octo	yes	S	"	"	
ICON-nwp	2.5	75	Icoso	no	-	TKE	no	
ICON-sap	5	"	"	yes	"	S	"	
ICON-sap+	"	"	"	yes	"	"	"	continuation of ICON-sap
ICON-vdu	"	"	"	no	"	TTE	"	
ICON-vdc	"	"	"	yes	"	"	"	
ICON-vda	"	"	"	yes	"	"	"	increased albedo
GEOS	3	80	Cube	no	F	K	yes	deep plumes disabled
SHIELD	3	40	"	mixed-layer ocean	S	TKE	yes	
SCREAM	3	40	"	no	-	SHOC	no	

# The simulations

Initialized 20th January 2020 with the global 9 km meteorological analysis from the ECMWF  
 Freely evolving until 1st March 2020.

Simulation	$\sqrt{A_{mean}}$	$H_t$	Grid	Coupled	Conv.	BL	SSO	comments
IFS-9	9	80	Octo	yes	F	K	yes	hydrostatic
IFS-4	4.5	80	Octo	yes	S	"	"	
ICON-nwp	2.5	75	Icoso	no	-	TKE	no	
ICON-sap	5	"	"	yes	"	S	"	
ICON-sap+	"	"	"	yes	"	"	"	continuation of ICON-sap
ICON-vdu	"	"	"	no	"	TTE	"	
ICON-vdc	"	"	"	yes	"	"	"	
ICON-vda	"	"	"	yes	"	"	"	increased albedo
GEOS	3	80	Cube	no	F	K	yes	deep plumes disabled
SHIELD	3	40	"	mixed-layer ocean	S	TKE	yes	
SCREAM	3	40	"	no	-	SHOC	no	

K: diagnostic eddy diffusivity  
 TKE: prognostic turbulent kinetic energy  
 TTE: prognostic turbulent total energy  
 S: Smagorinsky scheme  
 SHOC: Simplified Higher Order Closure



# The simulations

Initialized 20th January 2020 with the global 9 km meteorological analysis from the ECMWF  
 Freely evolving until 1st March 2020.

Simulation	$\sqrt{A_{mean}}$	$H_t$	Grid	Coupled	Conv.	BL	SSO	comments
IFS-9	9	80	Octo	yes	F	K	yes	hydrostatic
IFS-4	4.5	80	Octo	yes	S	"	"	
ICON-nwp	2.5	75	Icoso	no	-	TKE	no	
ICON-sap	5	"	"	yes	"	S	"	
ICON-sap+	"	"	"	yes	"	"	"	continuation of ICON-sap
ICON-vdu	"	"	"	no	"	TTE	"	} Overly diffusive
ICON-vdc	"	"	"	yes	"	"	"	
ICON-vda	"	"	"	yes	"	"	"	
GEOS	3	80	Cube	no	F	K	yes	deep plumes disabled
SHIELD	3	40	"	mixed-layer ocean	S	TKE	yes	
SCREAM	3	40	"	no	-	SHOC	no	

K: diagnostic eddy diffusivity  
 TKE: prognostic turbulent kinetic energy  
 TTE: prognostic turbulent total energy  
 S: Smagorinsky scheme  
 SHOC: Simplified Higher Order Closure

Overly diffusive

# Convection

Expectations based on DYAMOND summer

Stronger and deeper convection is usually associated with stronger vertical velocities and larger gravity wave momentum flux.

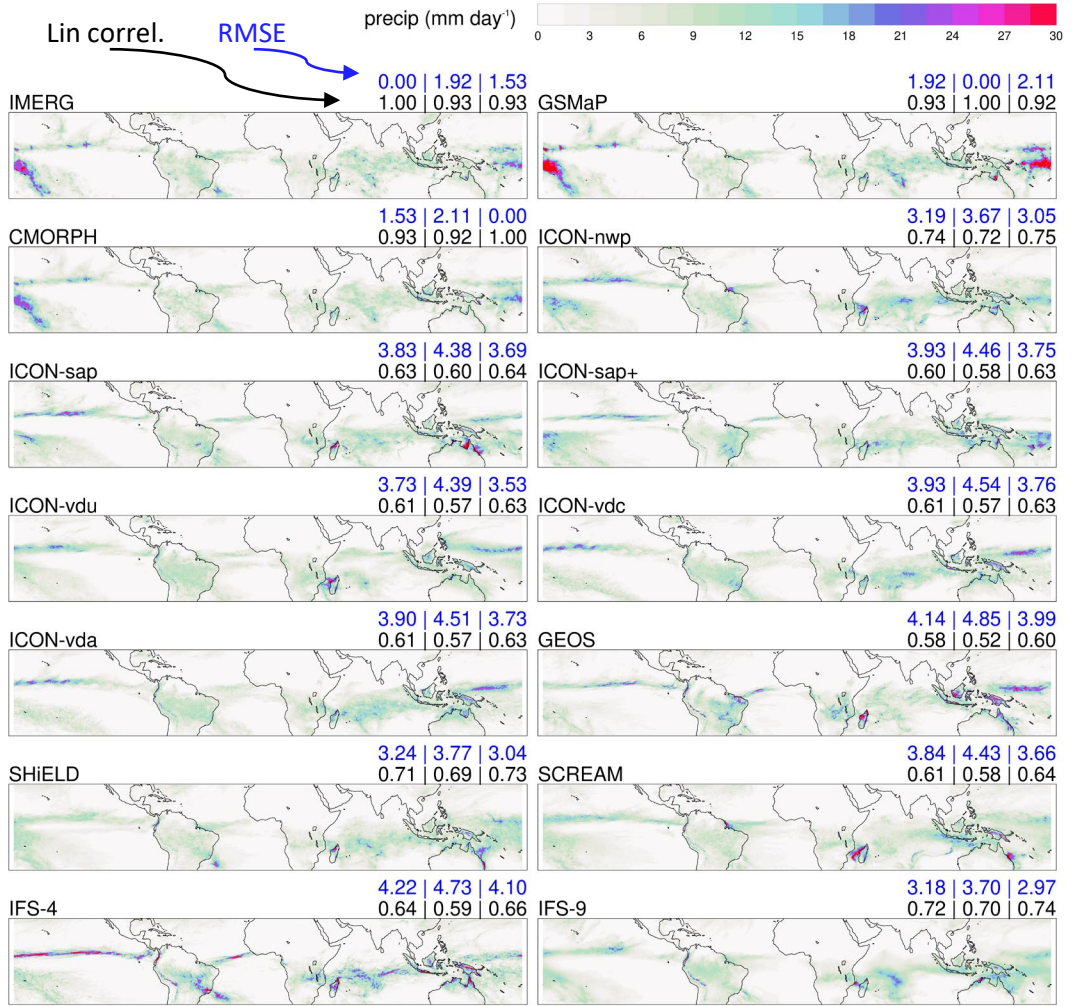
(Müller et al. 2018; Stephan et al. 2019a,b)

⇒ (How) does this reflect in spectra?

# Convection

Expectations based on DYAMOND summer  
 Stronger and deeper convection is usually associated with stronger vertical velocities and larger gravity wave momentum flux.  
 (Müller et al. 2018; Stephan et al. 2019a,b)

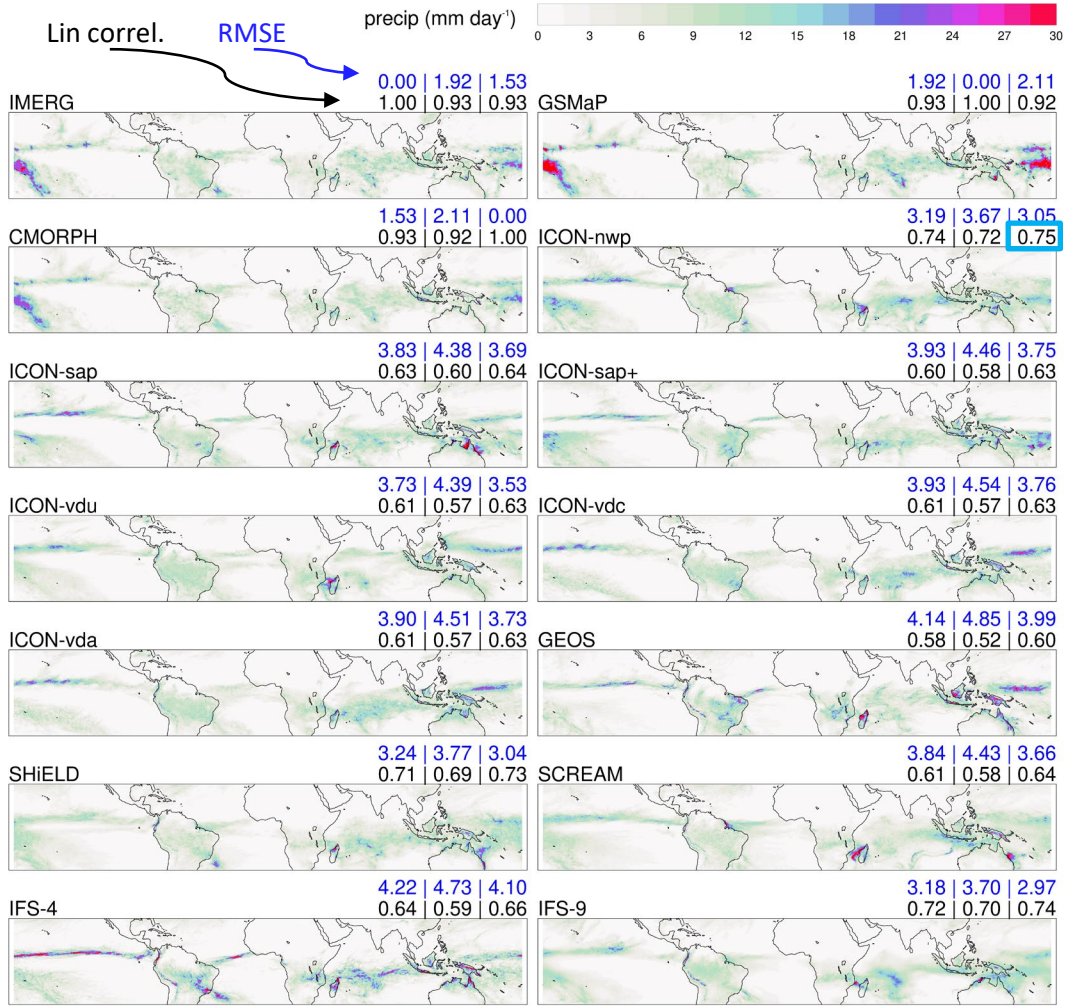
⇒ (How) does this reflect in spectra?



# Convection

Expectations based on DYAMOND summer  
 Stronger and deeper convection is usually associated with stronger vertical velocities and larger gravity wave momentum flux.  
 (Müller et al. 2018; Stephan et al. 2019a,b)

⇒ (How) does this reflect in spectra?

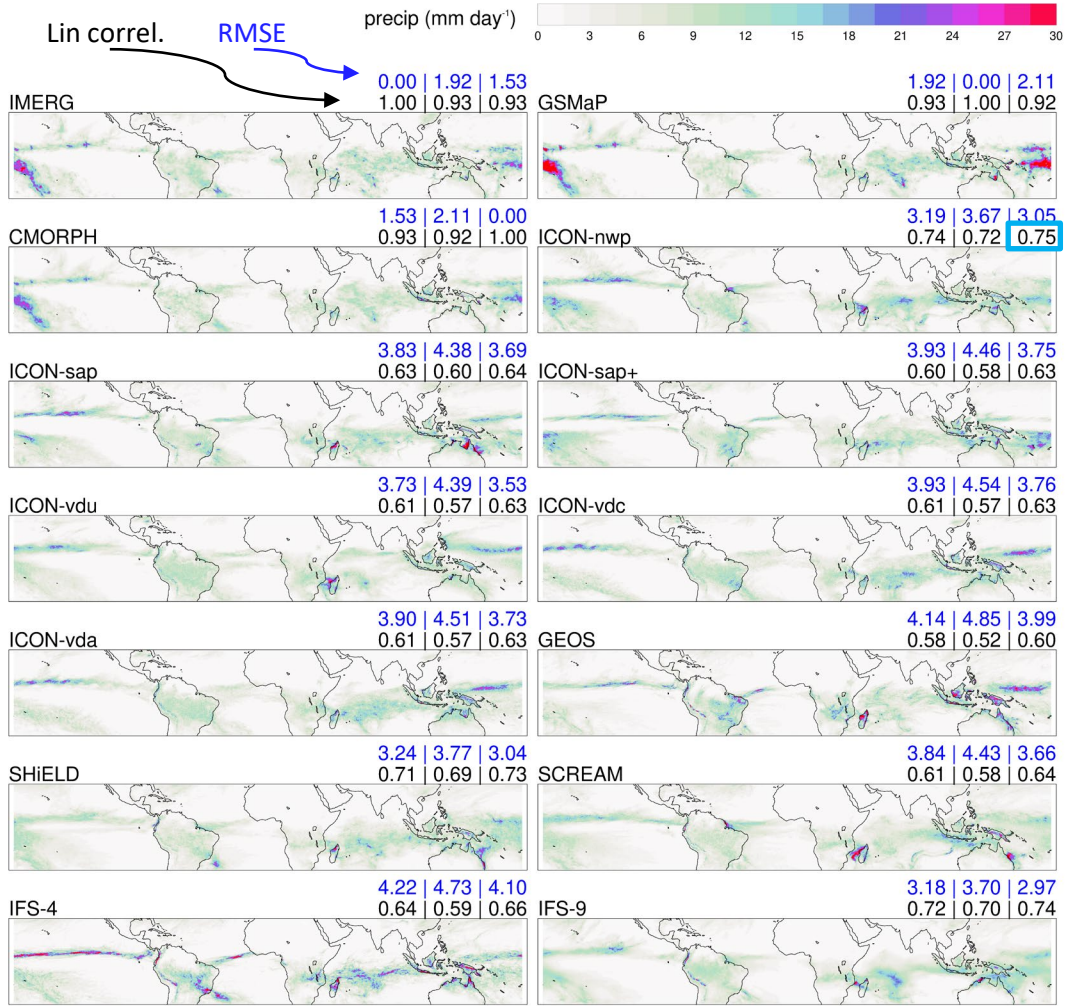


# Convection

Expectations based on DYAMOND summer  
 Stronger and deeper convection is usually associated with stronger vertical velocities and larger gravity wave momentum flux.  
 (Müller et al. 2018; Stephan et al. 2019a,b)

⇒ (How) does this reflect in spectra?

All simulations agree best with CMORPH and least with GSMaP.





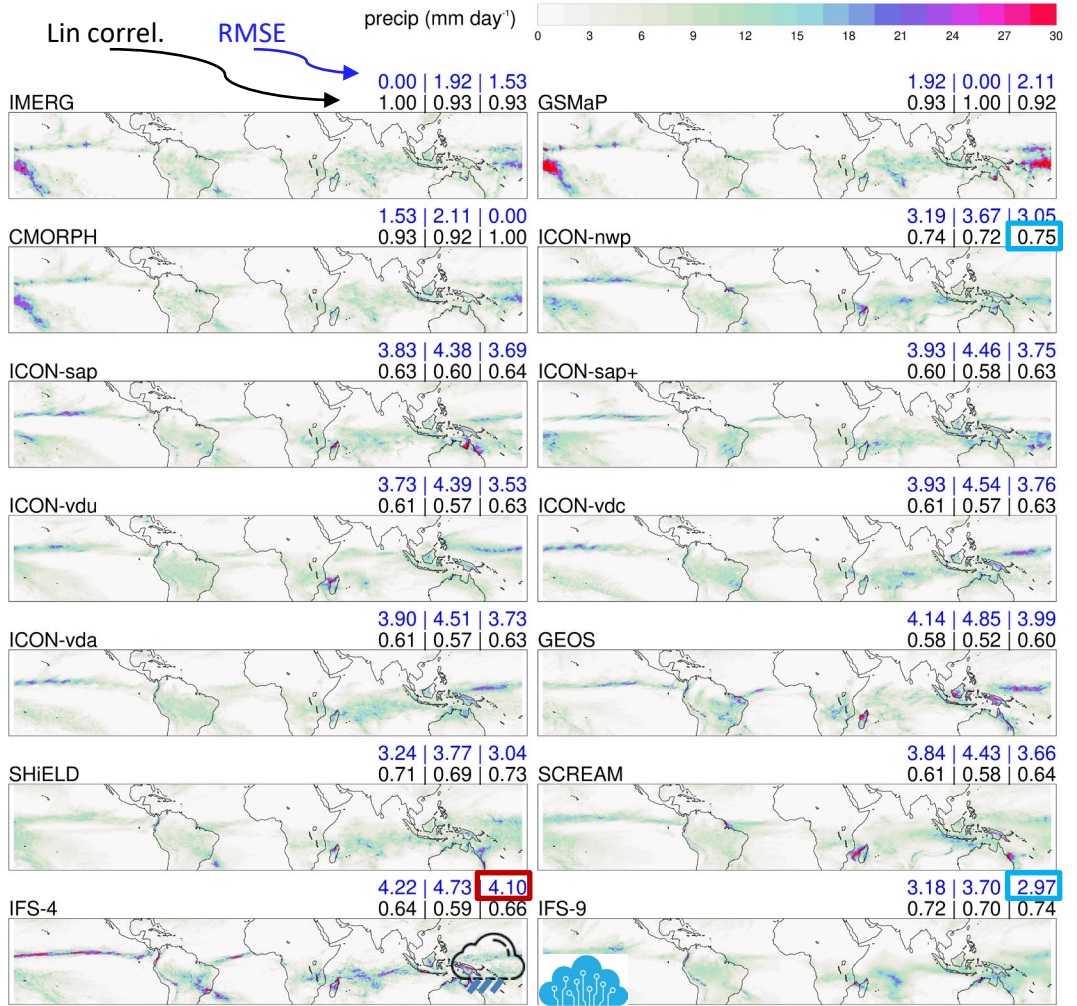
# Convection

Expectations based on DYAMOND summer  
 Stronger and deeper convection is usually associated with stronger vertical velocities and larger gravity wave momentum flux.  
 (Müller et al. 2018; Stephan et al. 2019a,b)

⇒ (How) does this reflect in spectra?

All simulations agree best with CMORPH and least with GSMaP.

Convection is under-resolved in some cases.  
 (Stephan et al. 2019a,b, Wedi et al. 2020)



# Convection

Expectations based on DYAMOND summer

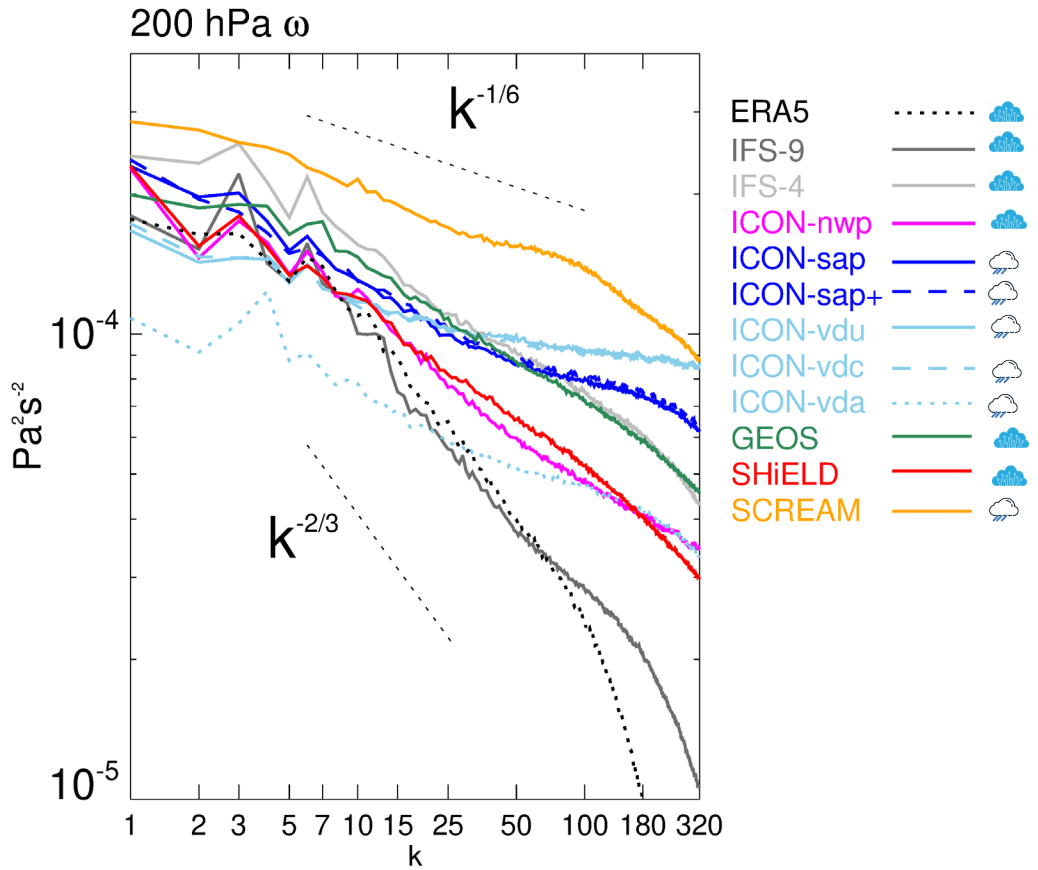
Stronger and deeper convection is usually associated with stronger vertical velocities and larger gravity wave momentum flux. (Müller et al. 2018; Stephan et al. 2019a,b)

⇒ (How) does this reflect in spectra?

All simulations agree best with CMORPH and least with GSMaP.

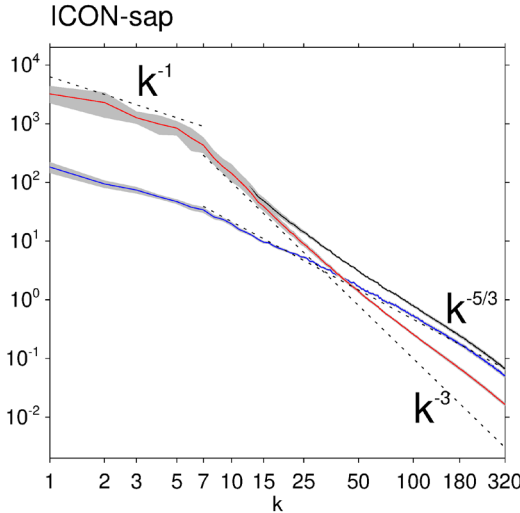
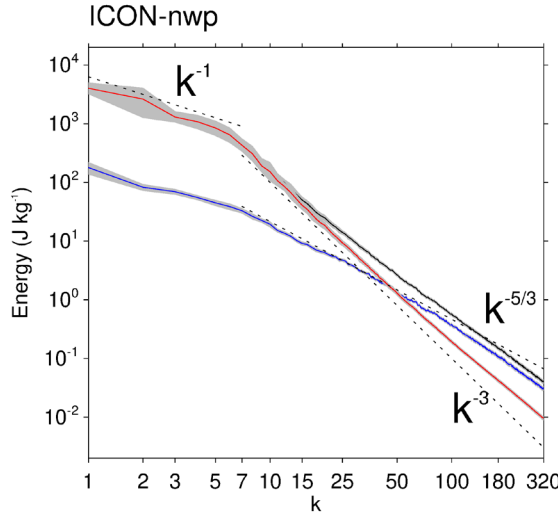
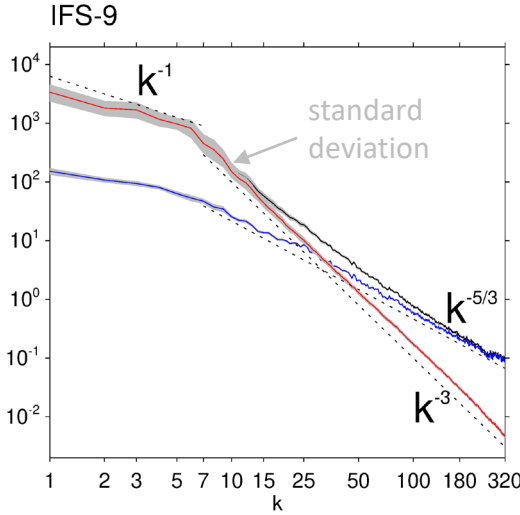
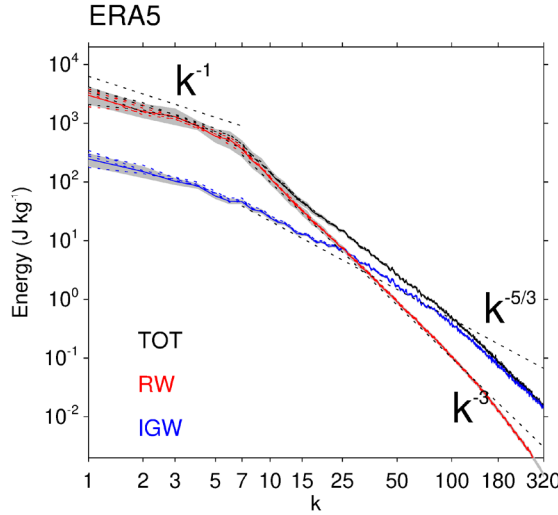
Convection is under-resolved in some cases. (Stephan et al. 2019a,b, Wedi et al. 2020)

Model formulation (☁️ 🌧️) affects the small scales.



# Energy spectra

ERA5: Little interannual variability



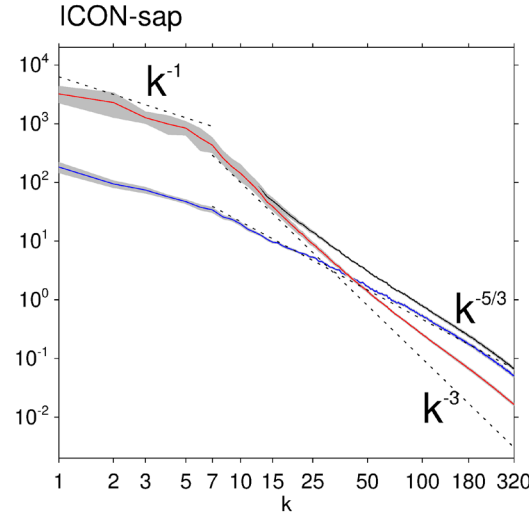
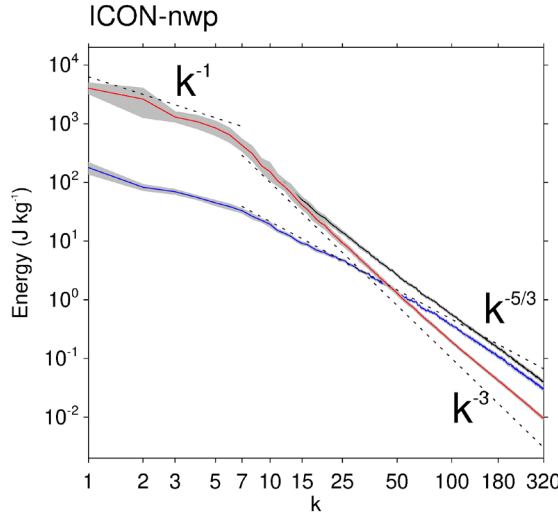
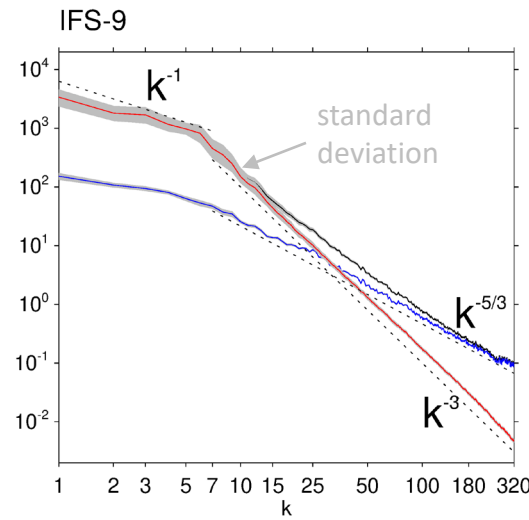
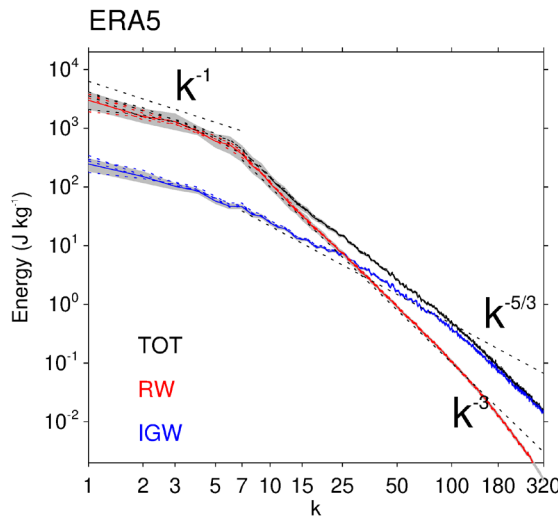


# Energy spectra

**ERA5:** Little interannual variability

**Planetary scales:** still depend on initialization.  
We focus on  $k > 7$ .

Spectra are robust footprints!  
(Boer and Shepherd 1983; Malardel and Wedi 2016)



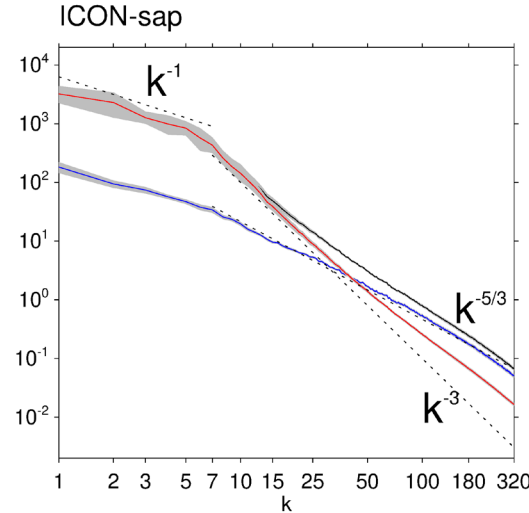
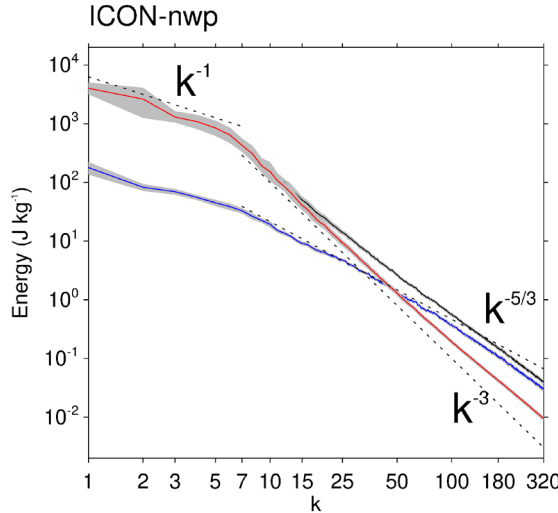
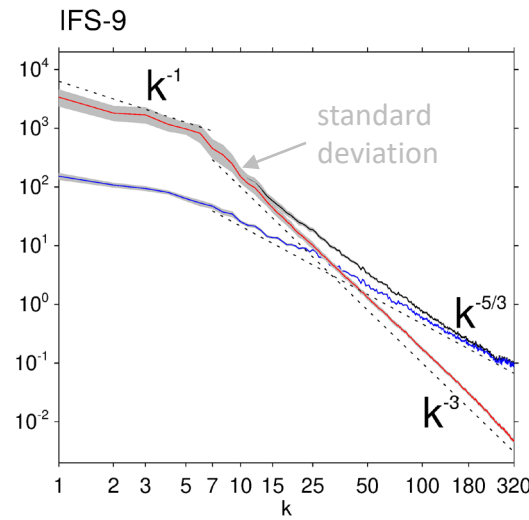
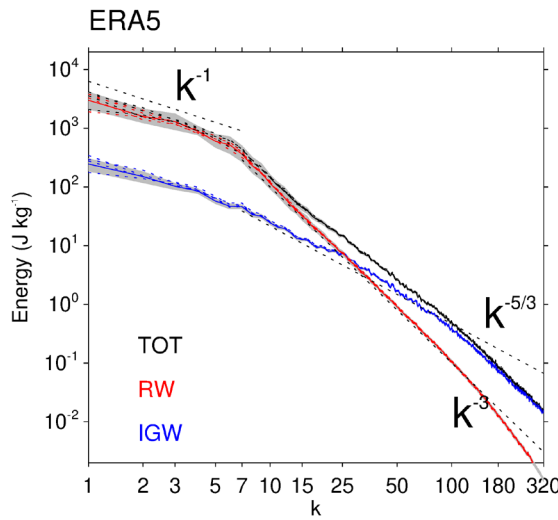
# Energy spectra

**ERA5:** Little interannual variability

**Planetary scales:** still depend on initialization.  
We focus on  $k > 7$ .

Spectra are robust footprints!  
(Boer and Shepherd 1983; Malardel and Wedi 2016)

- ⇒ Deviations from the canonical spectrum must be due to model formulation.
- ⇒ What factors shape the energy spectra in kilometer-scale models?



# Energy spectra

**ERA5:** Little interannual variability

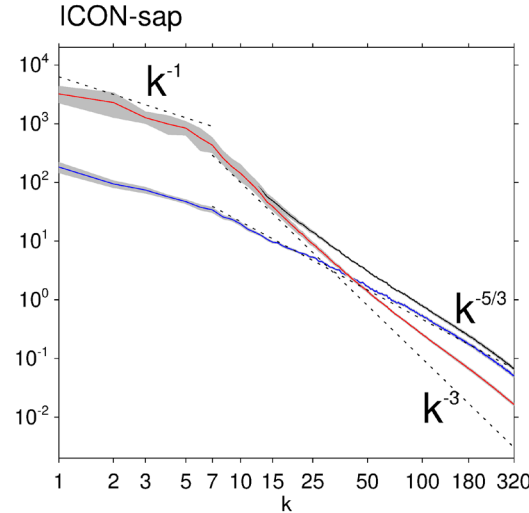
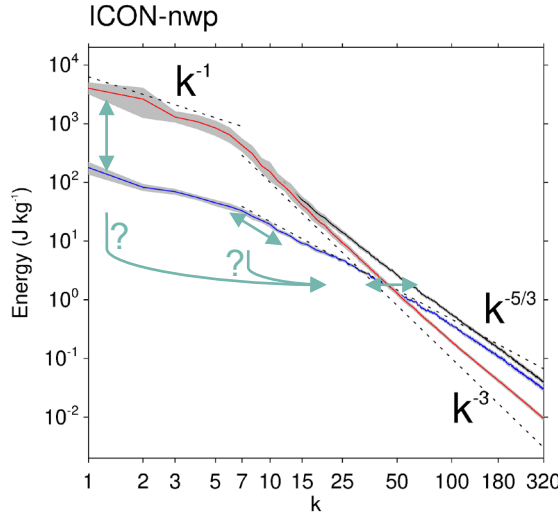
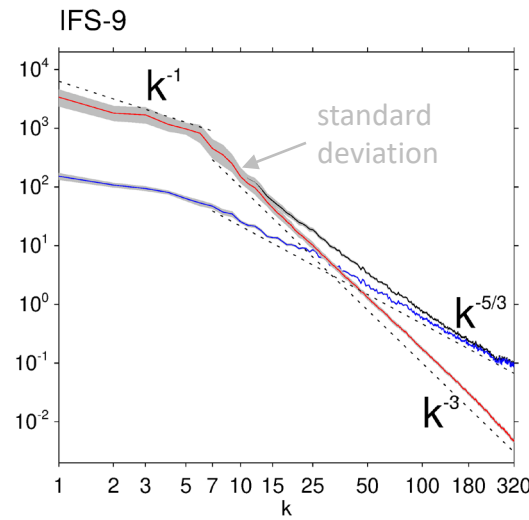
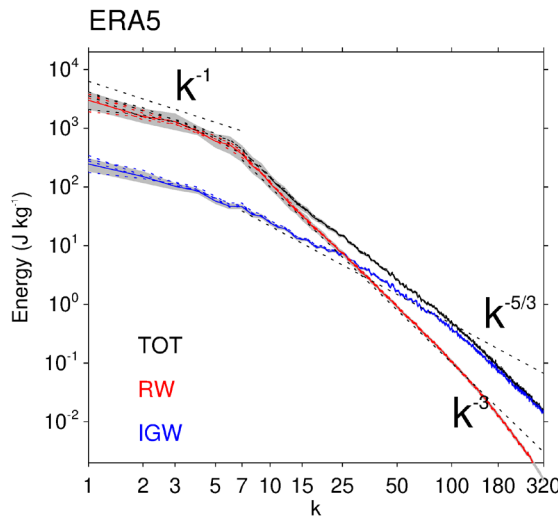
**Planetary scales:** still depend on initialization.  
We focus on  $k > 7$ .

Spectra are robust footprints!  
(Boer and Shepherd 1983; Malardel and Wedi 2016)

- ⇒ Deviations from the canonical spectrum must be due to model formulation.
- ⇒ What factors shape the energy spectra in kilometer-scale models?

The details are different:

- Offset
- Slopes
- Crossing



# Total wave energies

## Energy in $k = 1-320$

Operational analyses struggled for a long time to have an appropriate energy partitioning between RWs and IGWs.

(Tanaka et al. 1986; Tanaka and Kung 1988; Tanaka and Ji 1995; Žagar et al. 2009, 2012)

Modern analyses: 9-15% IGW

(Žagar et al. 2009)

Note: depends on top and level density

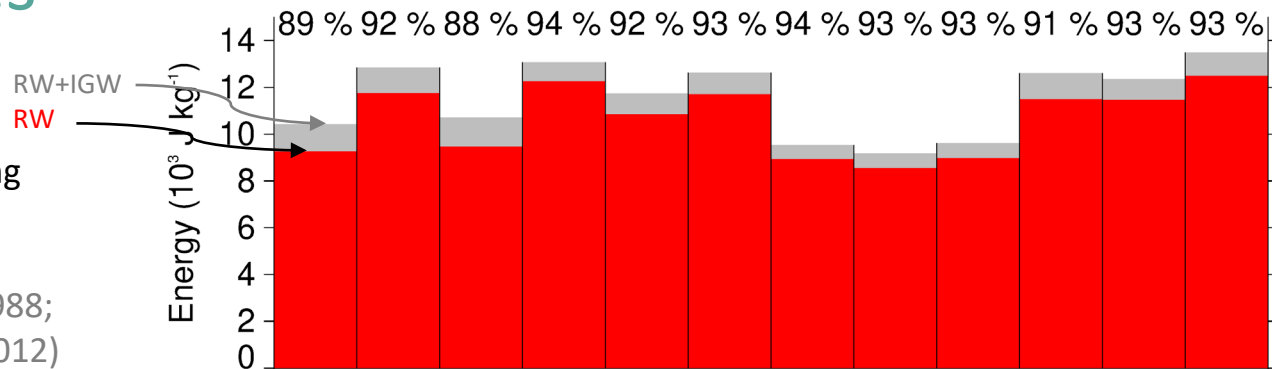
(Žagar et al. 2012)

# Total wave energies

## Energy in $k = 1-320$

Operational analyses struggled for a long time to have an appropriate energy partitioning between RWs and IGWs.

(Tanaka et al. 1986; Tanaka and Kung 1988; Tanaka and Ji 1995; Žagar et al. 2009, 2012)

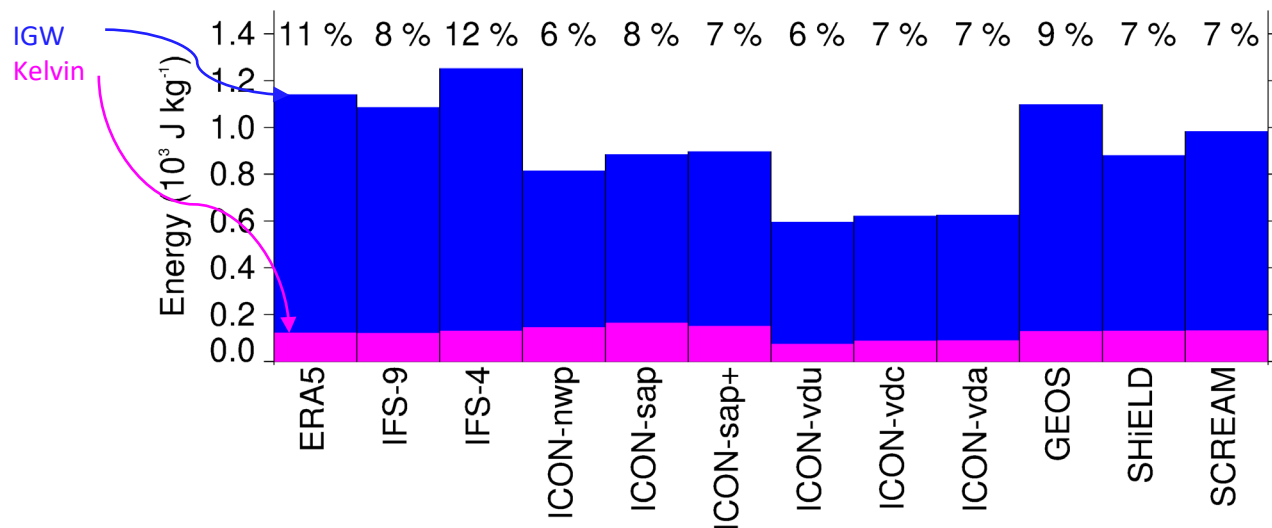


## Modern analyses: 9-15% IGW

(Žagar et al. 2009)

Note: depends on top and level density

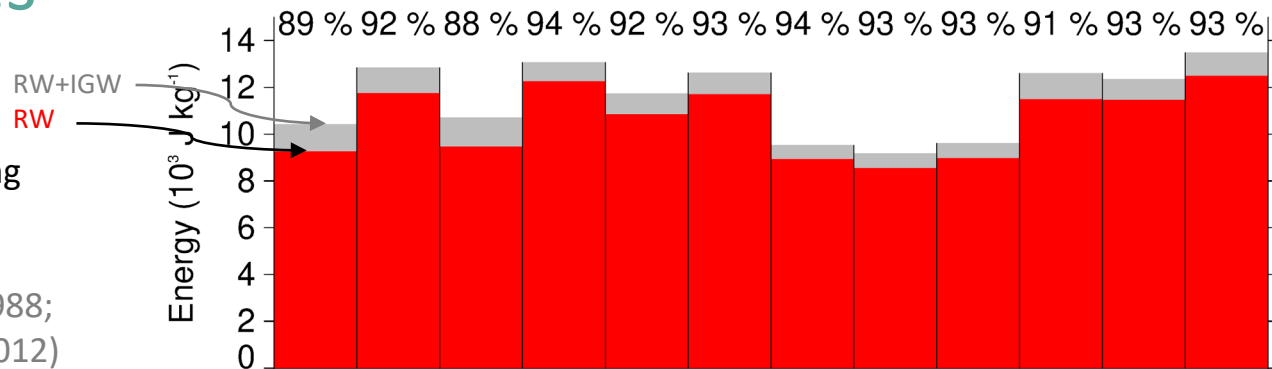
(Žagar et al. 2012)



# Total wave energies

## Energy in $k = 1-320$

Operational analyses struggled for a long time to have an appropriate energy partitioning between RWs and IGWs.  
(Tanaka et al. 1986; Tanaka and Kung 1988; Tanaka and Ji 1995; Žagar et al. 2009, 2012)

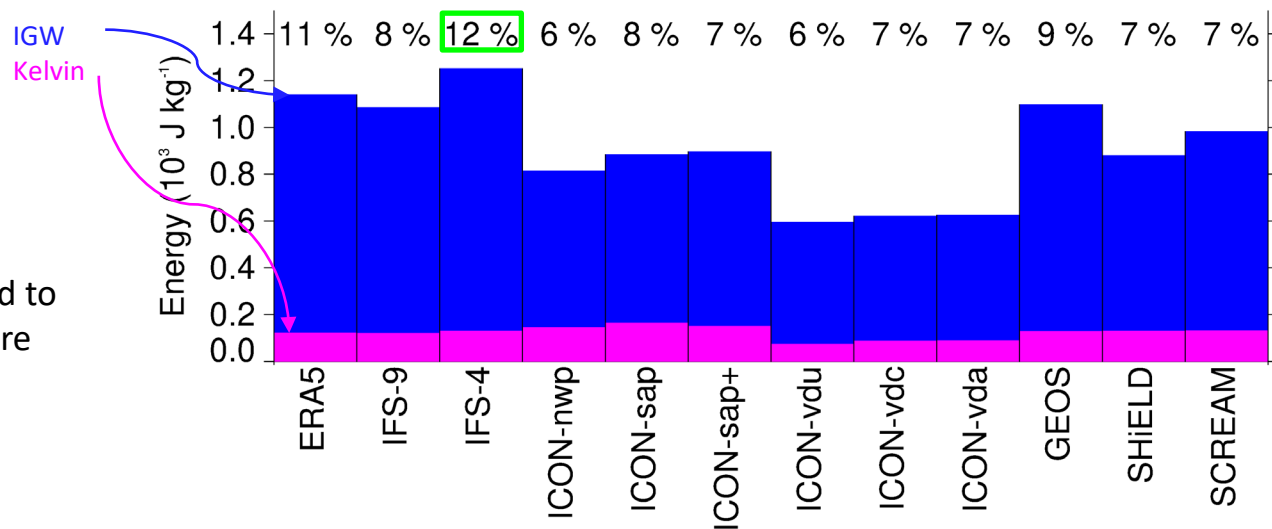


## Modern analyses: 9-15% IGW

(Žagar et al. 2009)

Note: depends on top and level density

(Žagar et al. 2012)

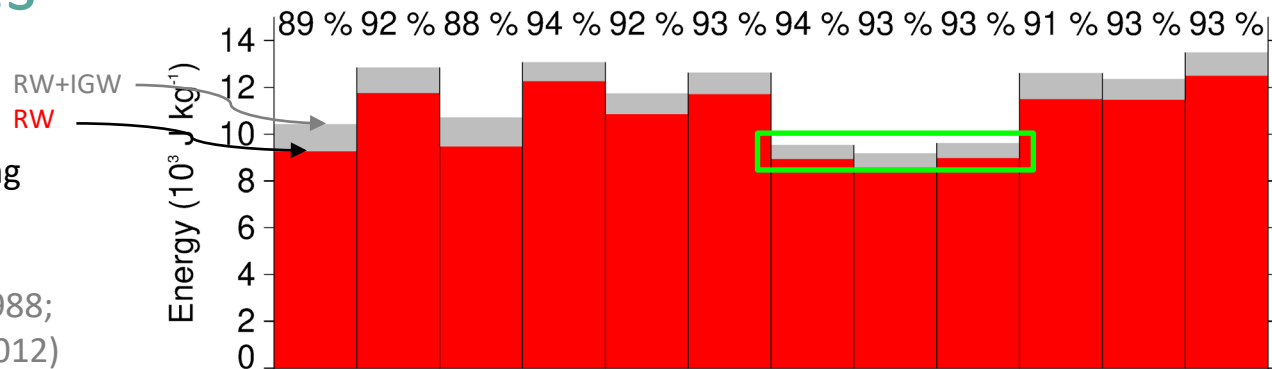


Compared to ERA5, the simulations tend to have less energy in IGW modes, but more energy in RW modes.

# Total wave energies

## Energy in $k = 1-320$

Operational analyses struggled for a long time to have an appropriate energy partitioning between RWs and IGWs.  
(Tanaka et al. 1986; Tanaka and Kung 1988; Tanaka and Ji 1995; Žagar et al. 2009, 2012)

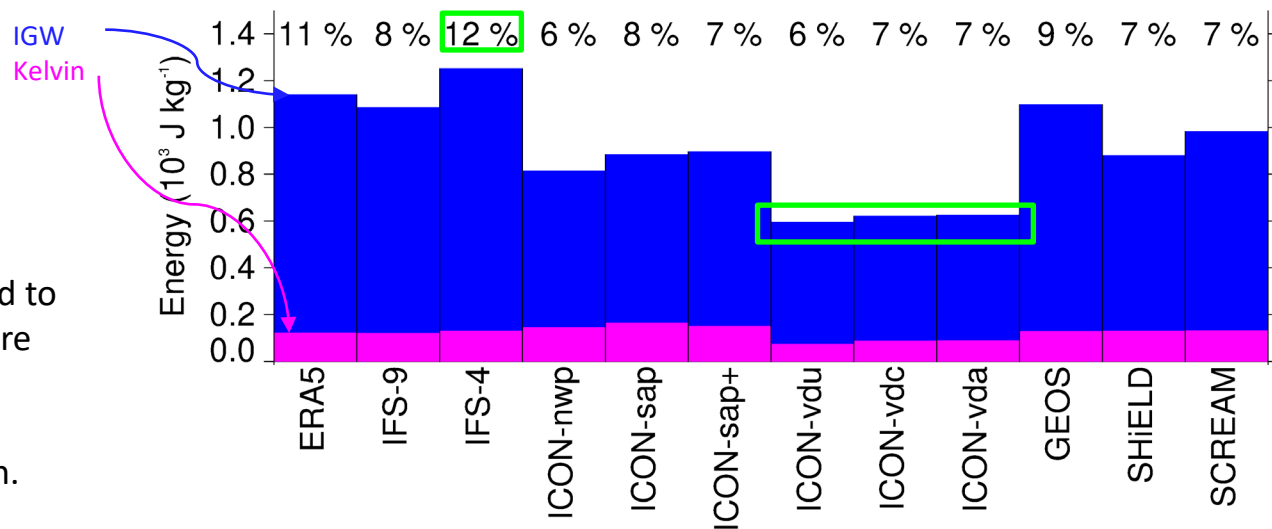


## Modern analyses: 9-15% IGW

(Žagar et al. 2009)

Note: depends on top and level density

(Žagar et al. 2012)



Compared to ERA5, the simulations tend to have less energy in IGW modes, but more energy in RW modes.

There is a sensitivity to vertical diffusion.

# Spectral slopes

wavenumbers:	TOT			RW			IGW		
	1-7	8-50	51-320	1-7	8-50	51-320	1-7	8-50	51-320
ERA5	-1.1	-2.5	-2.8	-1.1	-3.0	-3.7	-0.9	-1.6	-2.6
IFS-9	-1.0	-2.6	-1.9	-1.0	-3.1	-3.0	-0.6	-1.6	-1.7
IFS-4	-1.1	-2.5	-2.2	-1.2	-3.0	-2.6	-0.9	-1.6	-2.1
ICON-nwp	-1.1	-2.6	-2.3	-1.1	-3.0	-2.6	-0.9	-1.6	-2.0
ICON-sap	-1.0	-2.5	-2.1	-1.0	-2.9	-2.4	-0.9	-1.5	-1.9
ICON-sap+	-1.1	-2.5	-2.1	-1.1	-2.9	-2.4	-0.8	-1.5	-1.9
ICON-vdu	-1.2	-2.4	-2.1	-1.2	-2.8	-2.4	-0.9	-1.5	-2.0
ICON-vdc	-1.2	-2.5	-2.1	-1.2	-2.9	-2.4	-0.9	-1.5	-2.0
ICON-vda	-1.2	-2.5	-2.1	-1.2	-2.9	-2.4	-0.9	-1.5	-2.0
GEOS	-1.0	-2.6	-2.4	-1.0	-3.0	-2.8	-0.6	-1.6	-2.3
SHIELD	-1.1	-2.6	-2.4	-1.2	-3.0	-2.8	-0.7	-1.7	-2.3
SCREAM	-1.2	-2.5	-2.3	-1.3	-3.0	-2.6	-0.7	-1.5	-2.1



# Spectral slopes

In contrast to integrated energies, changes in vertical diffusion do not appear to affect spectral slopes.

Wave energy at small scales is underestimated in ERA5 due to limitations in data assimilation procedures.

	TOT			RW			IGW		
	1-7	8-50	51-320	1-7	8-50	51-320	1-7	8-50	51-320
wavenumbers:									
ERA5	-1.1	-2.5	-2.8	-1.1	-3.0	-3.7	-0.9	-1.6	-2.6
IFS-9	-1.0	-2.6	-1.9	-1.0	-3.1	-3.0	-0.6	-1.6	-1.7
IFS-4	-1.1	-2.5	-2.2	-1.2	-3.0	-2.6	-0.9	-1.6	-2.1
ICON-nwp	-1.1	-2.6	-2.3	-1.1	-3.0	-2.6	-0.9	-1.6	-2.0
ICON-sap	-1.0	-2.5	-2.1	-1.0	-2.9	-2.4	-0.9	-1.5	-1.9
ICON-sap+	-1.1	-2.5	-2.1	-1.1	-2.9	-2.4	-0.8	-1.5	-1.9
ICON-vdu	-1.2	-2.4	-2.1	-1.2	-2.8	-2.4	-0.9	-1.5	-2.0
ICON-vdc	-1.2	-2.5	-2.1	-1.2	-2.9	-2.4	-0.9	-1.5	-2.0
ICON-vda	-1.2	-2.5	-2.1	-1.2	-2.9	-2.4	-0.9	-1.5	-2.0
GEOS	-1.0	-2.6	-2.4	-1.0	-3.0	-2.8	-0.6	-1.6	-2.3
SHIELD	-1.1	-2.6	-2.4	-1.2	-3.0	-2.8	-0.7	-1.7	-2.3
SCREAM	-1.2	-2.5	-2.3	-1.3	-3.0	-2.6	-0.7	-1.5	-2.1

# Spectral slopes

In contrast to integrated energies, changes in vertical diffusion do not appear to affect spectral slopes.

Wave energy at small scales is underestimated in ERA5 due to limitations in data assimilation procedures.

RW spectra flatten towards small scales and IGW spectra steepen.

	TOT			RW			IGW		
	1-7	8-50	51-320	1-7	8-50	51-320	1-7	8-50	51-320
wavenumbers:									
ERA5	-1.1	-2.5	-2.8	-1.1	-3.0	-3.7	-0.9	-1.6	-2.6
IFS-9	-1.0	-2.6	-1.9	-1.0	-3.1	-3.0	-0.6	-1.6	-1.7
IFS-4	-1.1	-2.5	-2.2	-1.2	-3.0	-2.6	-0.9	-1.6	-2.1
ICON-nwp	-1.1	-2.6	-2.3	-1.1	-3.0	-2.6	-0.9	-1.6	-2.0
ICON-sap	-1.0	-2.5	-2.1	-1.0	-2.9	-2.4	-0.9	-1.5	-1.9
ICON-sap+	-1.1	-2.5	-2.1	-1.1	-2.9	-2.4	-0.8	-1.5	-1.9
ICON-vdu	-1.2	-2.4	-2.1	-1.2	-2.8	-2.4	-0.9	-1.5	-2.0
ICON-vdc	-1.2	-2.5	-2.1	-1.2	-2.9	-2.4	-0.9	-1.5	-2.0
ICON-vda	-1.2	-2.5	-2.1	-1.2	-2.9	-2.4	-0.9	-1.5	-2.0
GEOS	-1.0	-2.6	-2.4	-1.0	-3.0	-2.8	-0.6	-1.6	-2.3
SHIELD	-1.1	-2.6	-2.4	-1.2	-3.0	-2.8	-0.7	-1.7	-2.3
SCREAM	-1.2	-2.5	-2.3	-1.3	-3.0	-2.6	-0.7	-1.5	-2.1

# Spectral slopes

In contrast to integrated energies, changes in vertical diffusion do not appear to affect spectral slopes.

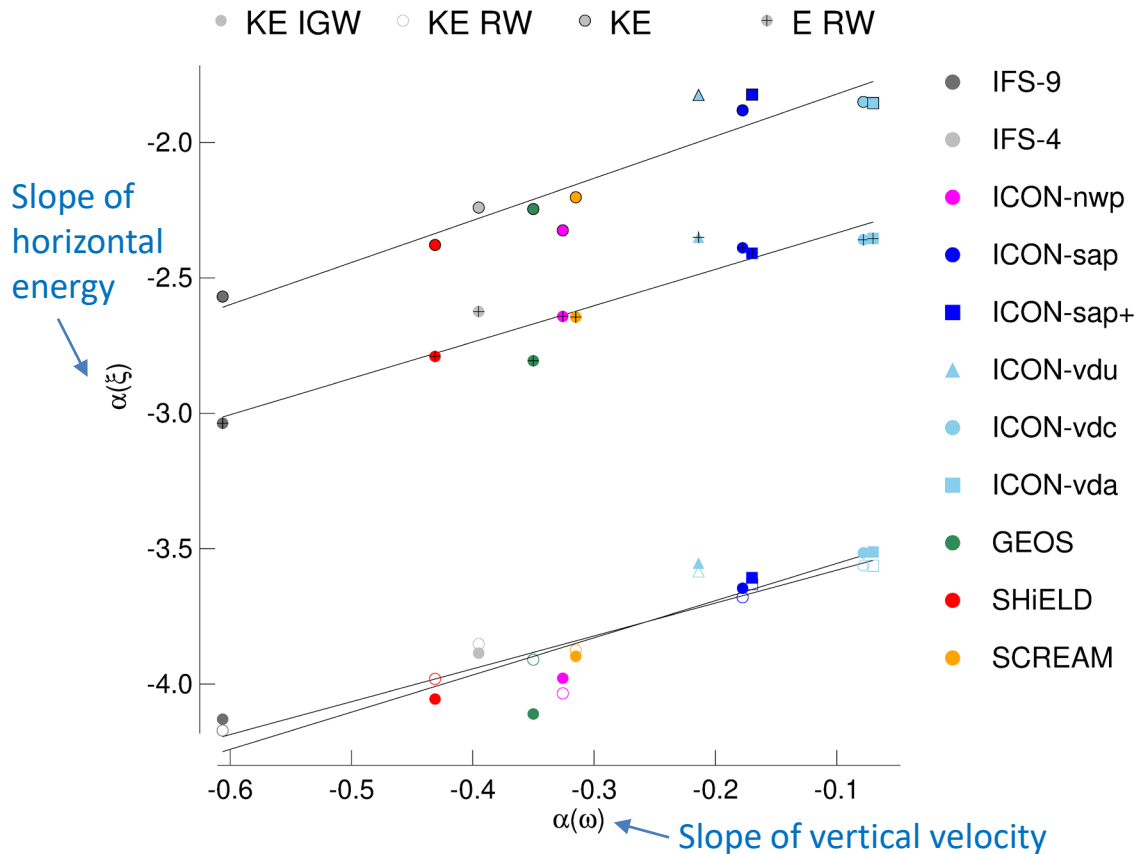
Wave energy at small scales is underestimated in ERA5 due to limitations in data assimilation procedures.

RW spectra flatten towards small scales and IGW spectra steepen.

What is the relationship to vertical velocity / relationship to convection?

	TOT			RW			IGW		
	1-7	8-50	51-320	1-7	8-50	51-320	1-7	8-50	51-320
wavenumbers:									
ERA5	-1.1	-2.5	-2.8	-1.1	-3.0	-3.7	-0.9	-1.6	-2.6
IFS-9	-1.0	-2.6	-1.9	-1.0	-3.1	-3.0	-0.6	-1.6	-1.7
IFS-4	-1.1	-2.5	-2.2	-1.2	-3.0	-2.6	-0.9	-1.6	-2.1
ICON-nwp	-1.1	-2.6	-2.3	-1.1	-3.0	-2.6	-0.9	-1.6	-2.0
ICON-sap	-1.0	-2.5	-2.1	-1.0	-2.9	-2.4	-0.9	-1.5	-1.9
ICON-sap+	-1.1	-2.5	-2.1	-1.1	-2.9	-2.4	-0.8	-1.5	-1.9
ICON-vdu	-1.2	-2.4	-2.1	-1.2	-2.8	-2.4	-0.9	-1.5	-2.0
ICON-vdc	-1.2	-2.5	-2.1	-1.2	-2.9	-2.4	-0.9	-1.5	-2.0
ICON-vda	-1.2	-2.5	-2.1	-1.2	-2.9	-2.4	-0.9	-1.5	-2.0
GEOS	-1.0	-2.6	-2.4	-1.0	-3.0	-2.8	-0.6	-1.6	-2.3
SHIELD	-1.1	-2.6	-2.4	-1.2	-3.0	-2.8	-0.7	-1.7	-2.3
SCREAM	-1.2	-2.5	-2.3	-1.3	-3.0	-2.6	-0.7	-1.5	-2.1

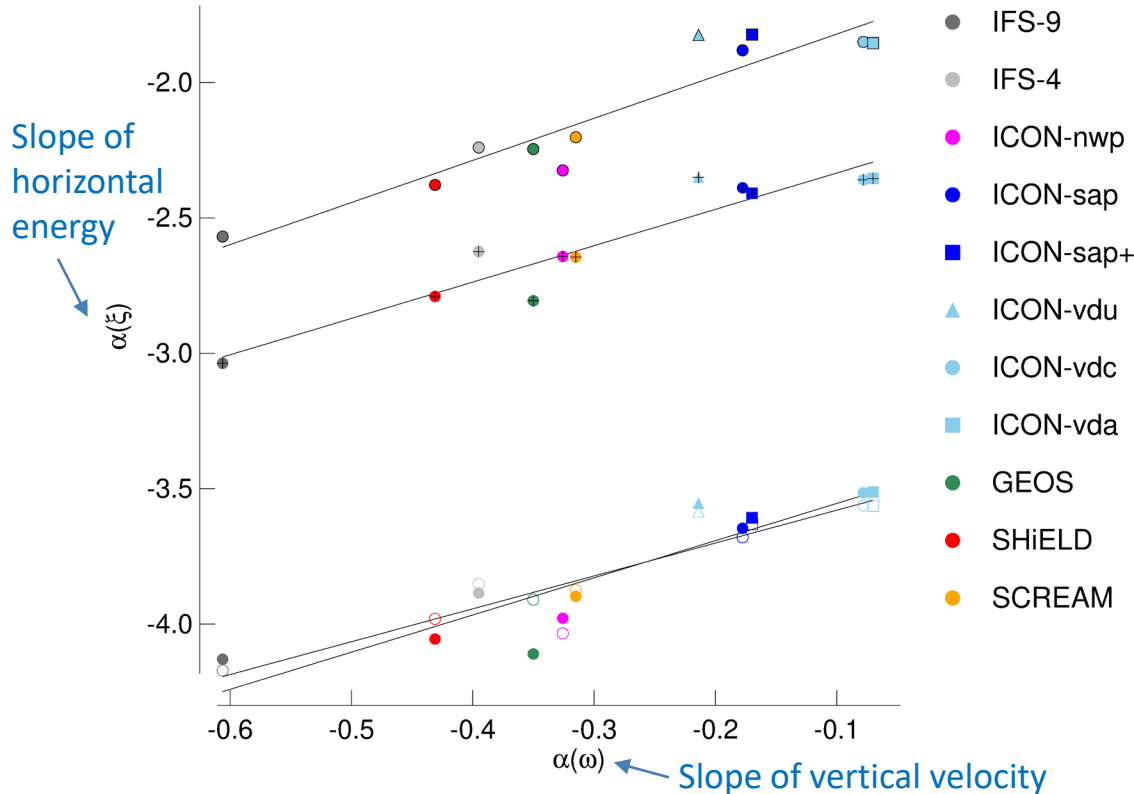
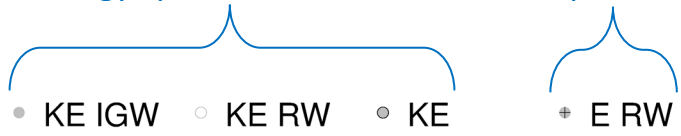
# Small scales and convection



# Small scales and convection

Tropical horizontal kinetic energy spectra at 200 hPa

Global RW spectrum



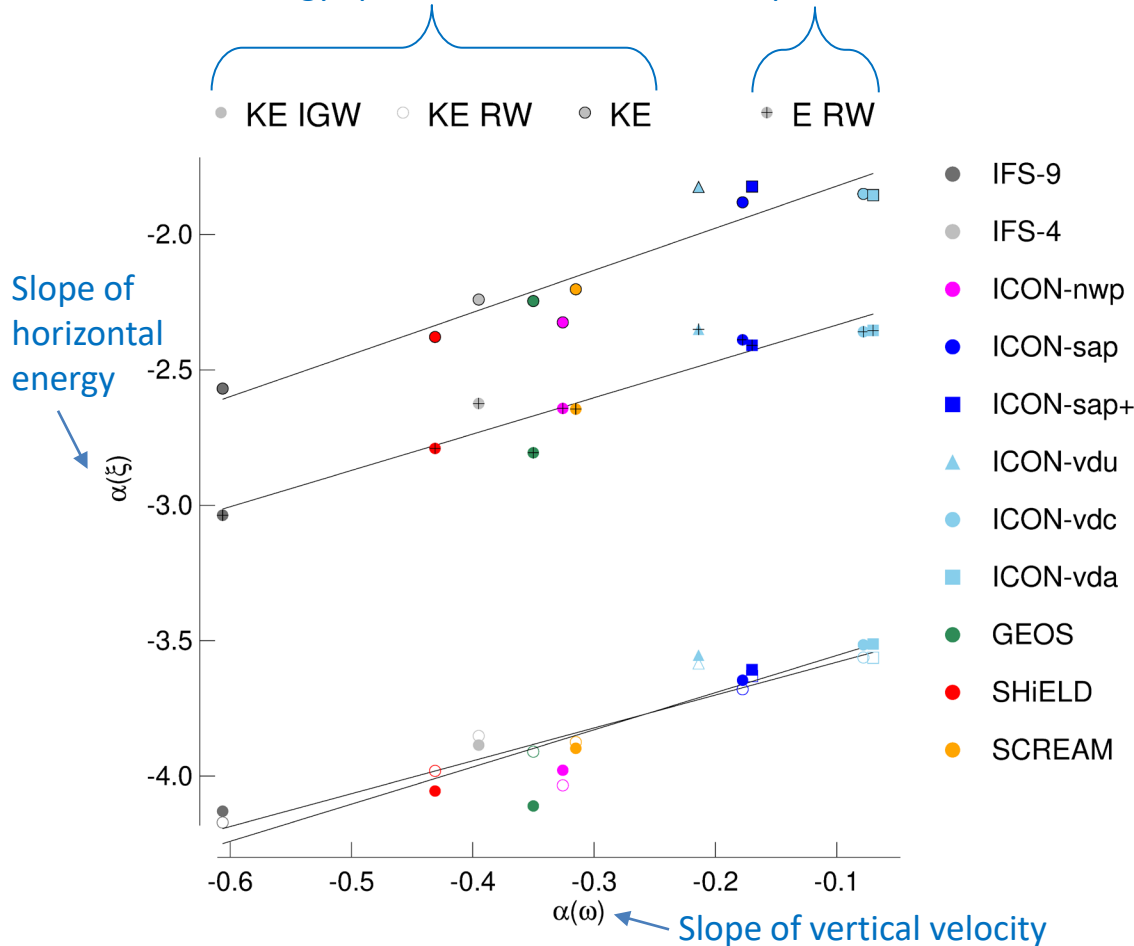
# Small scales and convection

Slopes of horizontal energy and vertical velocity are strongly correlated at  $k = 50-180$ .

Plausible: transient tropical heating is a source of RWs that propagate within as well as out of the tropics, and of equatorially trapped IGWs.

Tropical horizontal kinetic energy spectra at 200 hPa

Global RW spectrum



# Small scales and convection

Slopes of horizontal energy and vertical velocity are strongly correlated at  $k = 50-180$ .

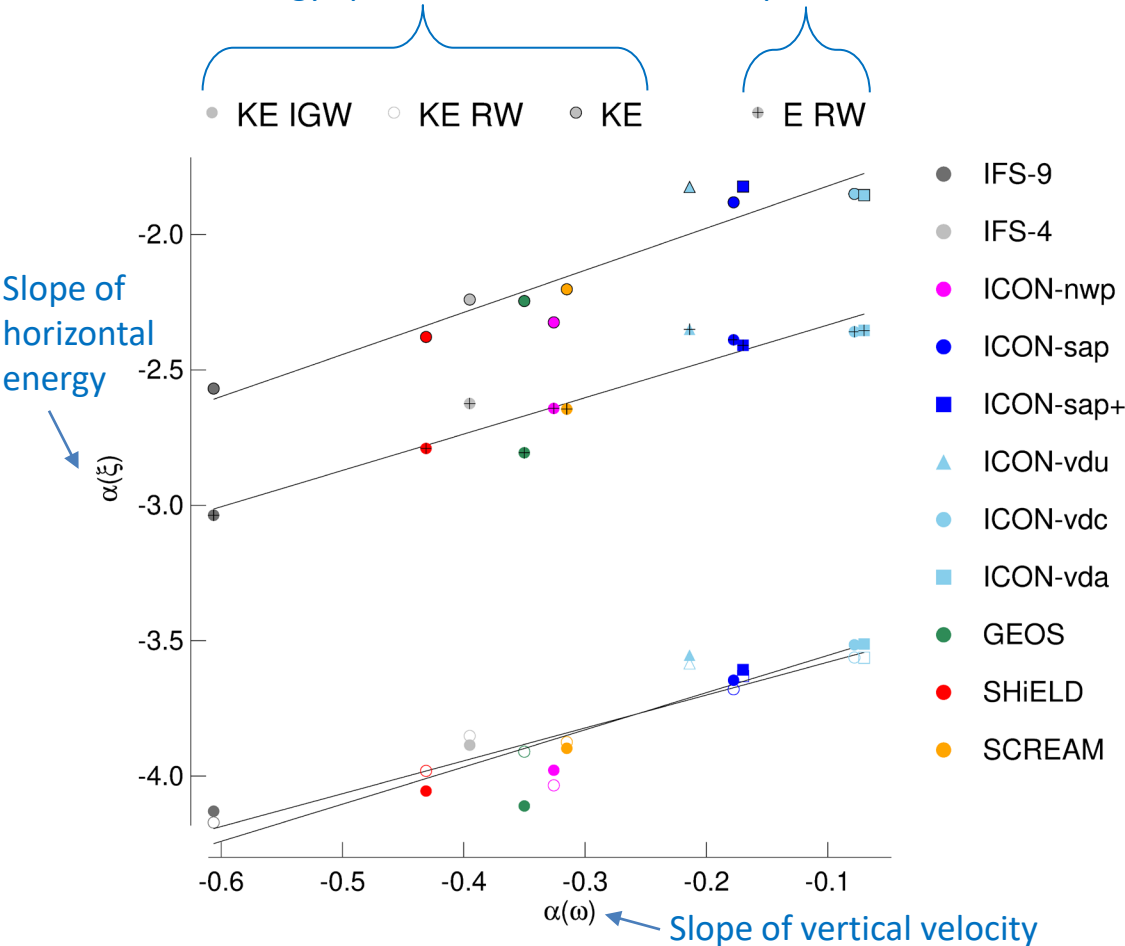
Plausible: transient tropical heating is a source of RWs that propagate within as well as out of the tropics, and of equatorially trapped IGWs.

The slope of the global IGW spectrum is **not** correlated with  $\alpha(\omega)$ .

Plausible: equatorial trapping, importance of extratropical IGW sources, relatively greater contribution of stratospheric levels.

Tropical horizontal kinetic energy spectra at 200 hPa

Global RW spectrum



# Crossing scale

Crossing scales have important implications for the applicability of spatial averaging, commonly used for decomposing motions into background and waves.

	$k_c$	$L_c$ (km)	
		$\phi = 0$	$\phi = \pm 45^\circ$
ERA5	25	1601	1132
IFS-9	32	1251	885
IFS-4	24	1668	1179
ICON-nwp	49	817	578
ICON-sap	41	976	690
ICON-sap+	41	976	690
ICON-vdu	42	953	674
ICON-vdc	42	953	674
ICON-vda	41	976	690
GEOS	29	1380	976
SHIELD	37	1082	765
SCREAM	32	1251	885



# Crossing scale

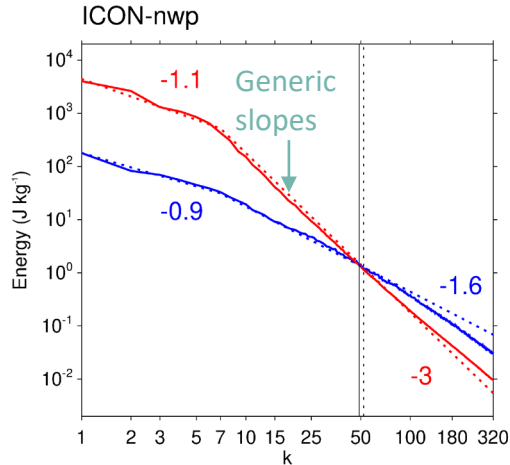
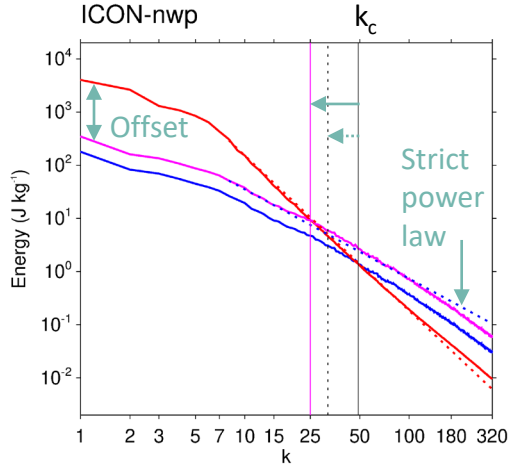
Crossing scales have important implications for the applicability of spatial averaging, commonly used for decomposing motions into background and waves.

	$k_c$	$L_c$ (km)	
		$\phi = 0$	$\phi = \pm 45^\circ$
ERA5	25	1601	1132
IFS-9	32	1251	885
IFS-4	24	1668	1179
ICON-nwp	49	817	578
ICON-sap	41	976	690
ICON-sap+	41	976	690
ICON-vdu	42	953	674
ICON-vdc	42	953	674
ICON-vda	41	976	690
GEOS	29	1380	976
SHIELD	37	1082	765
SCREAM	32	1251	885

# Crossing scale

Crossing scales have important implications for the applicability of spatial averaging, commonly used for decomposing motions into background and waves.

What determines the crossing scale?



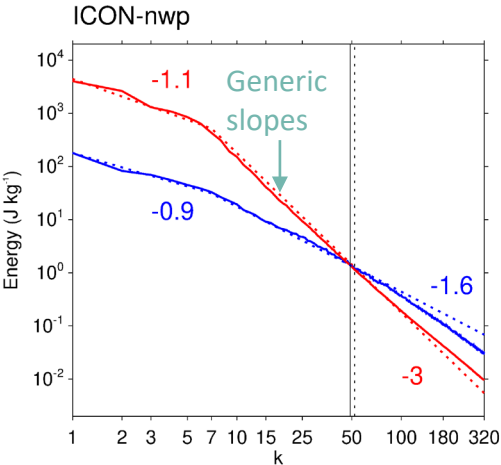
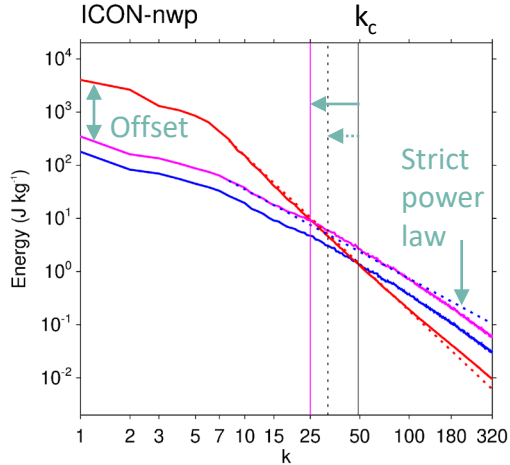
	$k_c$	$L_c$ (km)	
		$\phi = 0$	$\phi = \pm 45^\circ$

ERA5	25	1601	1132
IFS-9	32	1251	885
IFS-4	24	1668	1179
ICON-nwp	49	817	578
ICON-sap	41	976	690
ICON-sap+	41	976	690
ICON-vdu	42	953	674
ICON-vdc	42	953	674
ICON-vda	41	976	690
GEOS	29	1380	976
SHIELD	37	1082	765
SCREAM	32	1251	885

# Crossing scale

Crossing scales have important implications for the applicability of spatial averaging, commonly used for decomposing motions into background and waves.

What determines the crossing scale?



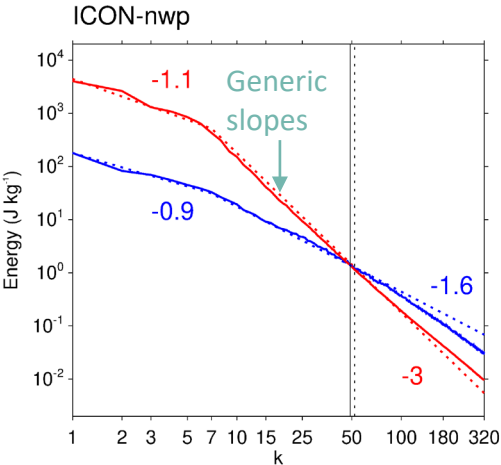
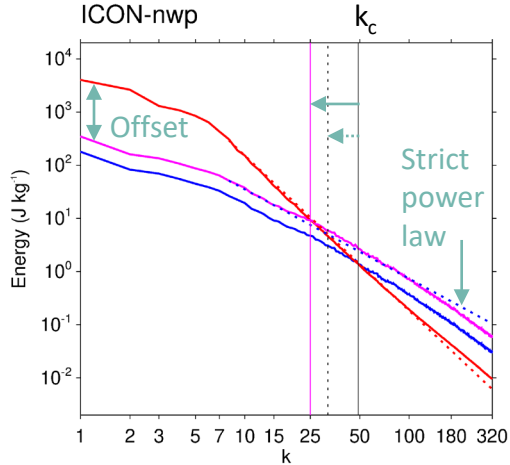
$k_c$  is to first order determined by the fraction of large-scale RW to IGW energy and much less by spectral slopes and shape.

	$k_c$	$L_c$ (km)	
		$\phi = 0$	$\phi = \pm 45^\circ$
ERA5	25	1601	1132
IFS-9	32	1251	885
IFS-4	24	1668	1179
ICON-nwp	49	817	578
ICON-sap	41	976	690
ICON-sap+	41	976	690
ICON-vdu	42	953	674
ICON-vdc	42	953	674
ICON-vda	41	976	690
GEOS	29	1380	976
SHIELD	37	1082	765
SCREAM	32	1251	885

# Crossing scale

Crossing scales have important implications for the applicability of spatial averaging, commonly used for decomposing motions into background and waves.

What determines the crossing scale?



$k_c$  is to first order determined by the fraction of large-scale RW to IGW energy and much less by spectral slopes and shape.

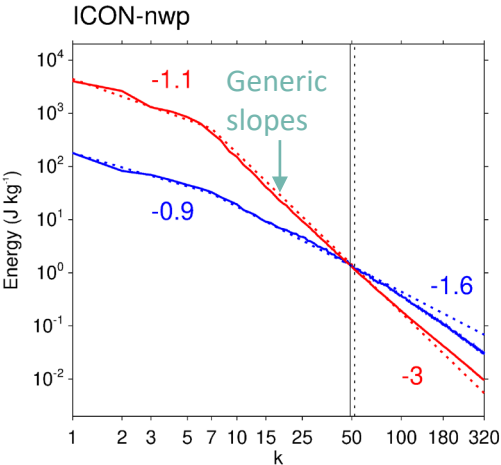
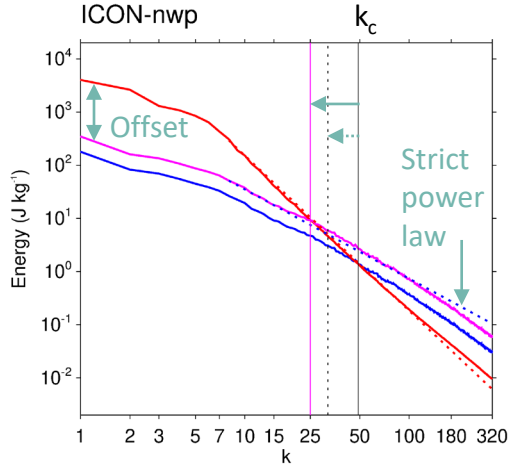
$k_c$  appears to be sensitive to the boundary layer parameterization.

	$k_c$	$L_c$ (km)	
		$\phi = 0$	$\phi = \pm 45^\circ$
ERA5	25	1601	1132
IFS-9	32	1251	885
IFS-4	24	1668	1179
ICON-nwp	49	817	578
ICON-sap	41	976	690
ICON-sap+	41	976	690
ICON-vdu	42	953	674
ICON-vdc	42	953	674
ICON-vda	41	976	690
GEOS	29	1380	976
SHIELD	37	1082	765
SCREAM	32	1251	885

# Crossing scale

Crossing scales have important implications for the applicability of spatial averaging, commonly used for decomposing motions into background and waves.

What determines the crossing scale?



$k_c$  is to first order determined by the fraction of large-scale RW to IGW energy and much less by spectral slopes and shape.

$k_c$  appears to be sensitive to the boundary layer parameterization.

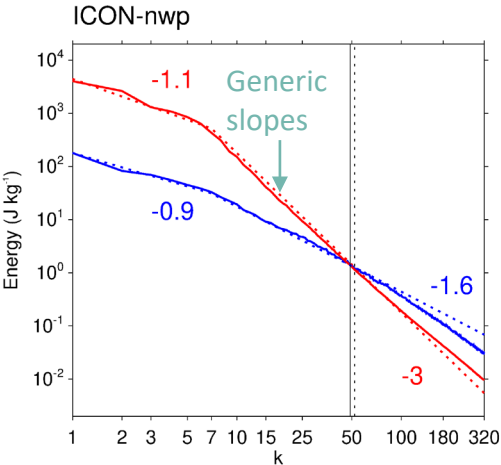
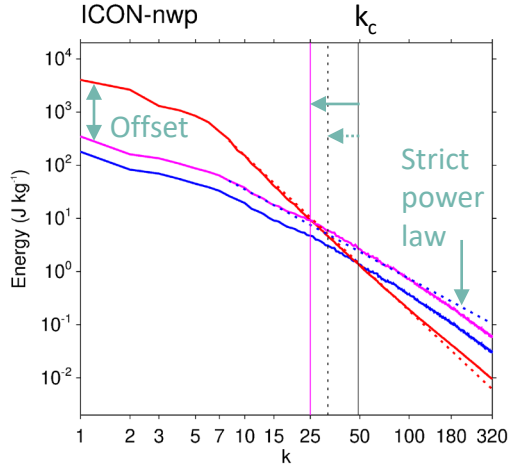
3D Smag.

	$k_c$	$L_c$ (km)	
		$\phi = 0$	$\phi = \pm 45^\circ$
ERA5	25	1601	1132
IFS-9	32	1251	885
IFS-4	24	1668	1179
ICON-nwp	49	817	578
ICON-sap	41	976	690
ICON-sap+	41	976	690
ICON-vdu	42	953	674
ICON-vdc	42	953	674
ICON-vda	41	976	690
GEOS	29	1380	976
SHIELD	37	1082	765
SCREAM	32	1251	885

# Crossing scale

Crossing scales have important implications for the applicability of spatial averaging, commonly used for decomposing motions into background and waves.

What determines the crossing scale?



$k_c$  is to first order determined by the fraction of large-scale RW to IGW energy and much less by spectral slopes and shape.

$k_c$  appears to be sensitive to the boundary layer parameterization.

3D Smag.

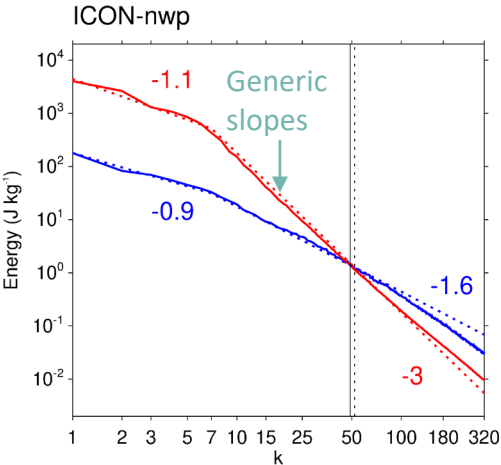
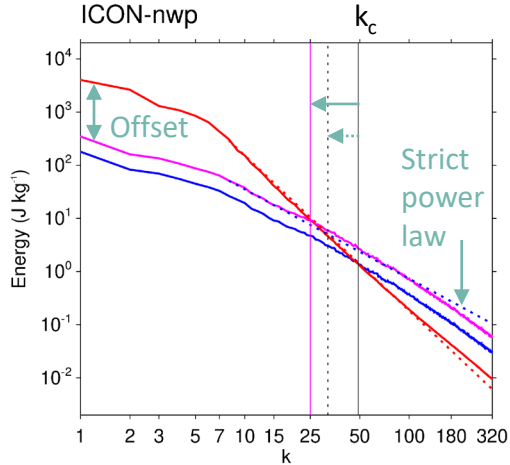
TTE/TKE

	$k_c$	$L_c$ (km)	
		$\phi = 0$	$\phi = \pm 45^\circ$
ERA5	25	1601	1132
IFS-9	32	1251	885
IFS-4	24	1668	1179
ICON-nwp	49	817	578
ICON-sap	41	976	690
ICON-sap+	41	976	690
ICON-vdu	42	953	674
ICON-vdc	42	953	674
ICON-vda	41	976	690
GEOS	29	1380	976
SHIELD	37	1082	765
SCREAM	32	1251	885

# Crossing scale

Crossing scales have important implications for the applicability of spatial averaging, commonly used for decomposing motions into background and waves.

What determines the crossing scale?



$k_c$  is to first order determined by the fraction of large-scale RW to IGW energy and much less by spectral slopes and shape.

$k_c$  appears to be sensitive to the boundary layer parameterization.

$k_c$	$L_c$ (km)	
	$\phi = 0$	$\phi = \pm 45^\circ$

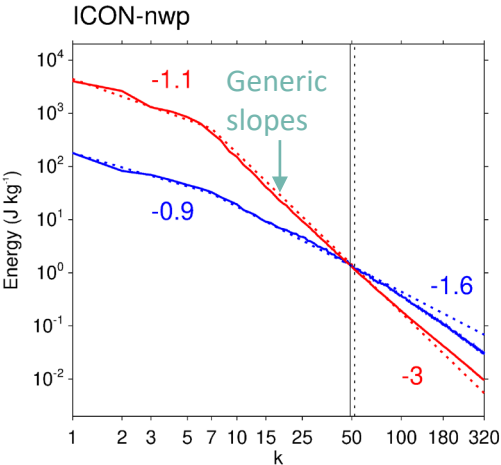
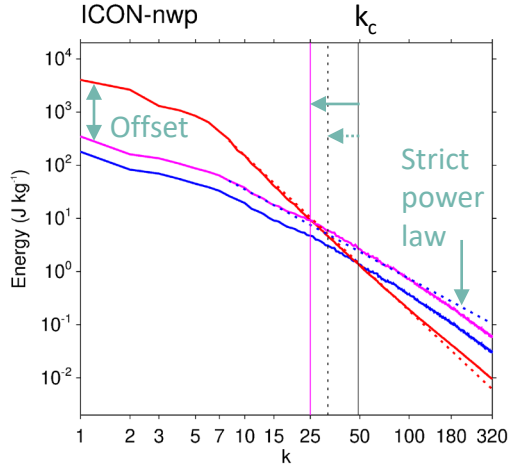
ERA5	25	1601	1132
IFS-9	32	1251	885
IFS-4	24	1668	1179
ICON-nwp	49	817	578
ICON-sap	41	976	690
ICON-sap+	41	976	690
ICON-vdu	42	953	674
ICON-vdc	42	953	674
ICON-vda	41	976	690
GEOS	29	1380	976
SHIELD	37	1082	765
SCREAM	32	1251	885

- 3D Smag.
- TTE/TKE
- PDF-based

# Crossing scale

Crossing scales have important implications for the applicability of spatial averaging, commonly used for decomposing motions into background and waves.

What determines the crossing scale?



$k_c$  is to first order determined by the fraction of large-scale RW to IGW energy and much less by spectral slopes and shape.

$k_c$  appears to be sensitive to the boundary layer parameterization.

$k_c$	$L_c$ (km)	
	$\phi = 0$	$\phi = \pm 45^\circ$

- 3D Smag.
- TTE/TKE
- PDF-based
- K-closure

ERA5	25	1601	1132
IFS-9	32	1251	885
IFS-4	24	1668	1179
ICON-nwp	49	817	578
ICON-sap	41	976	690
ICON-sap+	41	976	690
ICON-vdu	42	953	674
ICON-vdc	42	953	674
ICON-vda	41	976	690
GEOS	29	1380	976
SHIELD	37	1082	765
SCREAM	32	1251	885



# Summary – Wave spectra of 11 DYAMOND simulations

---

The spectra are robust characteristics of the simulations.

All simulations produce the expected canonical shape of the spectra:

- $k = 8-50$  the simulations agree well on both RW and IGW slopes: encouraging

# Summary – Wave spectra of 11 DYAMOND simulations

---

The spectra are robust characteristics of the simulations.

All simulations produce the expected canonical shape of the spectra:

- $k = 8-50$  the simulations agree well on both RW and IGW slopes: encouraging

**Significant differences:**

# Summary – Wave spectra of 11 DYAMOND simulations

The spectra are robust characteristics of the simulations.

All simulations produce the expected canonical shape of the spectra:

- $k = 8-50$  the simulations agree well on both RW and IGW slopes: encouraging

**Significant differences:**

## Total energy levels

- Total wave energies differ by 21% among the simulations (excluding the sensitivity experiments with stronger vertical diffusion).
- Differences in IGW energy levels reach 35%.

# Summary – Wave spectra of 11 DYAMOND simulations

The spectra are robust characteristics of the simulations.

All simulations produce the expected canonical shape of the spectra:

- $k = 8-50$  the simulations agree well on both RW and IGW slopes: encouraging

## Significant differences:

### Total energy levels

- Total wave energies differ by 21% among the simulations (excluding the sensitivity experiments with stronger vertical diffusion).
- Differences in IGW energy levels reach 35%.

### Spectral slopes

- At small scales: RW spectra flatten, IGW spectra steepen.
- Substantial differences in the spectra of upper tropospheric vertical velocity
- Strong correlation between slopes of tropical horizontal KE and vertical velocity; also global RW

# Summary – Wave spectra of 11 DYAMOND simulations

The spectra are robust characteristics of the simulations.

All simulations produce the expected canonical shape of the spectra:

- $k = 8-50$  the simulations agree well on both RW and IGW slopes: encouraging

## Significant differences:

### Total energy levels

- Total wave energies differ by 21% among the simulations (excluding the sensitivity experiments with stronger vertical diffusion).
- Differences in IGW energy levels reach 35%.

### Spectral slopes

- At small scales: RW spectra flatten, IGW spectra steepen.
- Substantial differences in the spectra of upper tropospheric vertical velocity
- Strong correlation between slopes of tropical horizontal KE and vertical velocity; also global RW

### Spectral crossing scales

- Partitioning of total energy: most important factor for determining the spectral crossing scale
- Crossing scales differ by a factor of about 2; models group by type of turbulence closure scheme (no indication for sensitivity to different horizontal or vertical resolutions, or hydrostatic versus non-hydrostatic dynamics)

# Outlook

---

## Relevance of small scales

This study looked at global horizontal plus potential available energy – **small scales** contribute very little

But:

- small scales are important for driving the global mean circulation
- horizontally short gravity waves with **large vertical velocities** | locally strong momentum flux

# Outlook

---

## Relevance of small scales

This study looked at global horizontal plus potential available energy – **small scales** contribute very little

But:

- small scales are important for driving the global mean circulation
- horizontally short gravity waves with **large vertical velocities** | locally strong momentum flux

## High-resolution models

Small scales differ greatly in DYAMOND! (**convection, vertical velocity, horizontal spectra**)

- How fit are DYAMOND models for multi-decade predictions? (Dependence on turbulence parameterization)
- What are the limitations for 'learning' from DYAMOND models? (e.g. AI-parameterizations)

# Outlook

---

## Relevance of small scales

This study looked at global horizontal plus potential available energy – **small scales** contribute very little

But:

- small scales are important for driving the global mean circulation
- horizontally short gravity waves with **large vertical velocities** | locally strong momentum flux

## High-resolution models

Small scales differ greatly in DYAMOND! (**convection, vertical velocity, horizontal spectra**)

- How fit are DYAMOND models for multi-decade predictions? (Dependence on turbulence parameterization)
- What are the limitations for ‘learning’ from DYAMOND models? (e.g. AI-parameterizations)

## Observations

This study supports our 2019 results, which were based on a completely different approach: Both report the same factor of 3 differences in DYAMOND (2019 for GWMF, here for vertical velocity)

- Currently comparing to observations (balloon measurements) – work by Laura Köhler



# Outlook

## Relevance of small scales

This study looked at global horizontal plus potential available energy – **small scales** contribute very little

But:

- small scales are important for driving the global mean circulation
- horizontally short gravity waves with **large vertical velocities** | locally strong momentum flux

## High-resolution models

Small scales differ greatly in DYAMOND! (**convection, vertical velocity, horizontal spectra**)

- How fit are DYAMOND models for multi-decade predictions? (Dependence on turbulence parameterization)
- What are the limitations for ‘learning’ from DYAMOND models? (e.g. AI-parameterizations)

## Observations

This study supports our 2019 results, which were based on a completely different approach: Both report the same factor of 3 differences in DYAMOND (2019 for GWMF, here for vertical velocity)

- Currently comparing to observations (balloon measurements) – work by Laura Köhler

## Follow-up study is happening

We ran the models again 😊, writing out i) physics tendencies ii) everything required to compute non-linear spectral transfer (following Augier and Lindborg 2013) – work by Yanmichel Morfa Avalos