Land Surface:

Introduction to warm processes

Vegetation, carbon & ECLand

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Outlines

- **Vegetation**
  - **Role of vegetation in NWP**
  - **Tiled approach and current data**
  - **Evolution of vegetation parametrization and practical cases**
- **Carbon**
  - **Why are we interested in carbon?**
  - **Parametrization and feedback from the atmosphere**
  - **Comparison with Jarvis approach and interaction with the atmosphere**
- **Improved Soil discretization and perspectives**

Vegetation state affects

- **Energy/water budgets**
  - **Evapotranspiration**
  - **Interception evaporation**
  - **Surface albedo (net radiation at the surface)**
  - **Aerodynamic exchange through surface roughness**

- **Carbon budget**
  - **Plant Respiration**
  - **Photosynthesis**

Vegetation role: some recaps

- **Energy balance equation**

$$(1-a)R_S^{\downarrow} + \varepsilon_g R_T^{\downarrow} - \varepsilon_g \sigma T_{sk}^4 + H + \lambda E = G$$

=>Albedo (a) and emissivity (ε) depend on the surface/vegetation condition

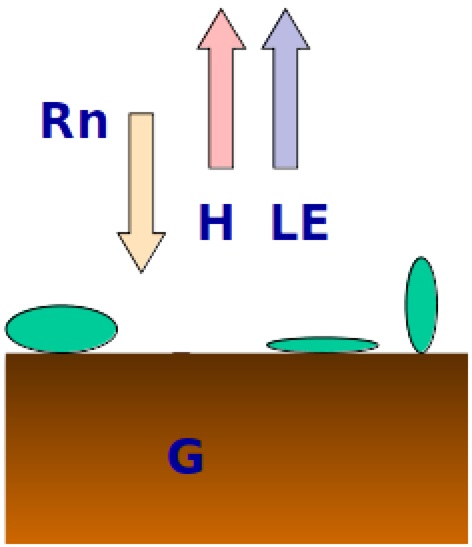


Table 3.1
Radiative Properties of Natural Surfaces[a]

Surface type	Other specifications	Albedo (a)	Emissivity (ε)
Water	Small zenith angle	0.03–0.10	0.92–0.97
	Large zenith angle	0.10–0.50	0.92–0.97
Snow	Old	0.40–0.70	0.82–0.89
	Fresh	0.45–0.95	0.90–0.99
Ice	Sea	0.30–0.40	0.92–0.97
	Glacier	0.20–0.40	
Bare sand	Dry	0.35–0.45	0.84–0.90
	Wet	0.20–0.30	0.91–0.95
Bare soil	Dry clay	0.20–0.35	0.95
	Moist clay	0.10–0.20	0.97
	Wet fallow field	0.05–0.07	
Paved	Concrete	0.17–0.27	0.71–0.88
	Black gravel road	0.05–0.10	0.88–0.95
Grass	Long (1 m) Short (0.02 m)	0.16–0.26	0.90–0.95
Agricultural	Wheat, rice, etc.	0.10–0.25	0.90–0.99
	Orchards	0.15–0.20	0.90–0.95
Forests	Deciduous	0.10–0.20	0.97–0.98
	Coniferous	0.05–0.15	0.97–0.99

[a] Compiled from Sellers (1965), Kondratyev (1969), and Oke (1978).

Arya, 1988

Vegetation role: some recaps

- **Energy balance equation**

$$(1-a)R_S^{\downarrow} + \varepsilon_g R_T^{\downarrow} - \varepsilon_g \sigma T_{sk}^4 + H + \lambda E = G$$

=> Latent heat (LE) is related to vegetation through:

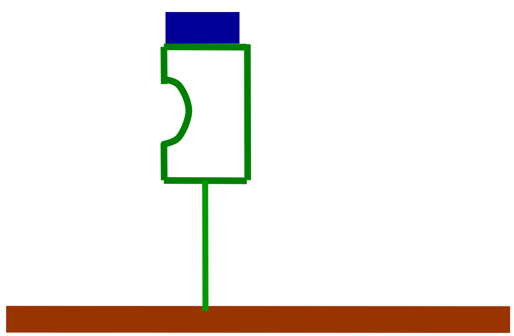
Evapotranspiration and momentum exchange

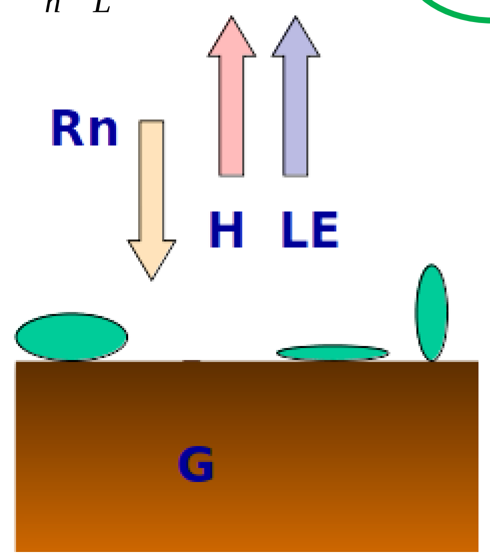
$$\begin{cases} E = \dfrac{\rho_a}{r_c + r_a}\left[q_a - q_{sat}(T_{sk})\right] \\ r_c = \dfrac{r_{s,\min}}{LAI} f_1\left(R_S^{\downarrow}\right) f_3\left(\overline{\theta}\right) f_4\left(D_a\right) \\ r_a = \dfrac{1}{C_h u_L}, C_h = f(Ri_B, z_{oh}, z_{om}) \end{cases}$$

Interception evaporation=

f(Interception reservoir)=>f(LAI))

Wet vegetation



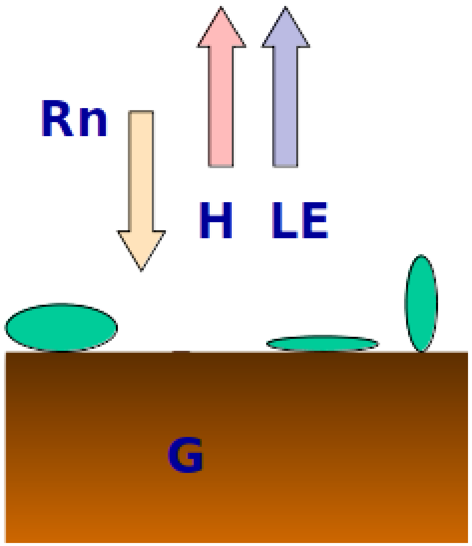


Vegetation role: some recaps

- **Energy balance equation**

$$(1-a)R_S^{\downarrow} + \varepsilon_g R_T^{\downarrow} - \varepsilon_g \sigma T_{sk}^4 + H + \lambda E = G$$

=> Sensible heat (H) is also related to vegetation through its relative partition with LE and the aerodynamic exchange specific to surface/vegetation type



Sensible heat flux

$$H = \rho C_h u_L (C_p T_L + gz - C_p T_{sk})$$

$$C_h = f(Ri_B, z_{oh}, z_{om})$$

z_{oh}, z_{om}

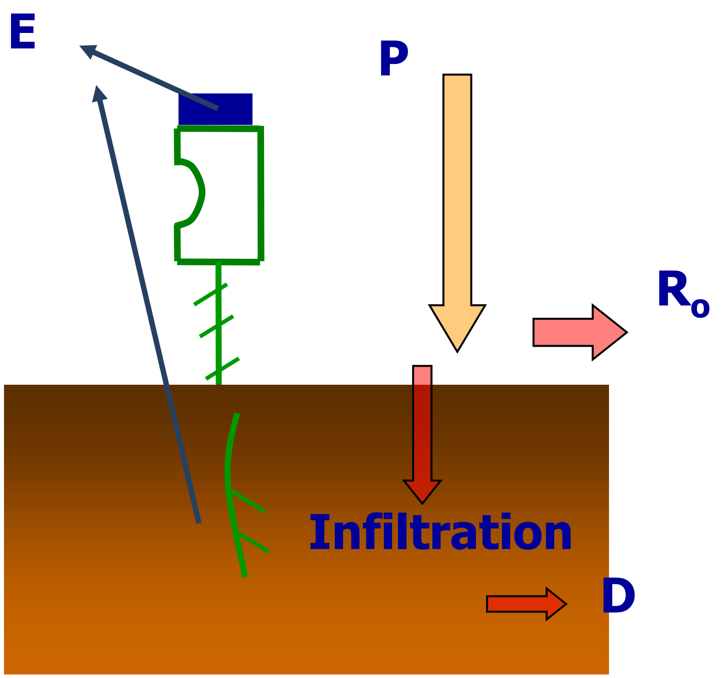
Roughness length for heat and momentum

Dependent on surface/vegetation type

Vegetation role: some recaps

- **Water balance equation**

$$\partial W/\partial t = P - E - Ro - I - D$$



$\partial W/\partial t$ = change in water storage
P= precipitation
E= evapotranspiration
Ro= runoff
I =Infiltration
D=lateral diffusion

Evaporation from:

Bare soil

Interception layer

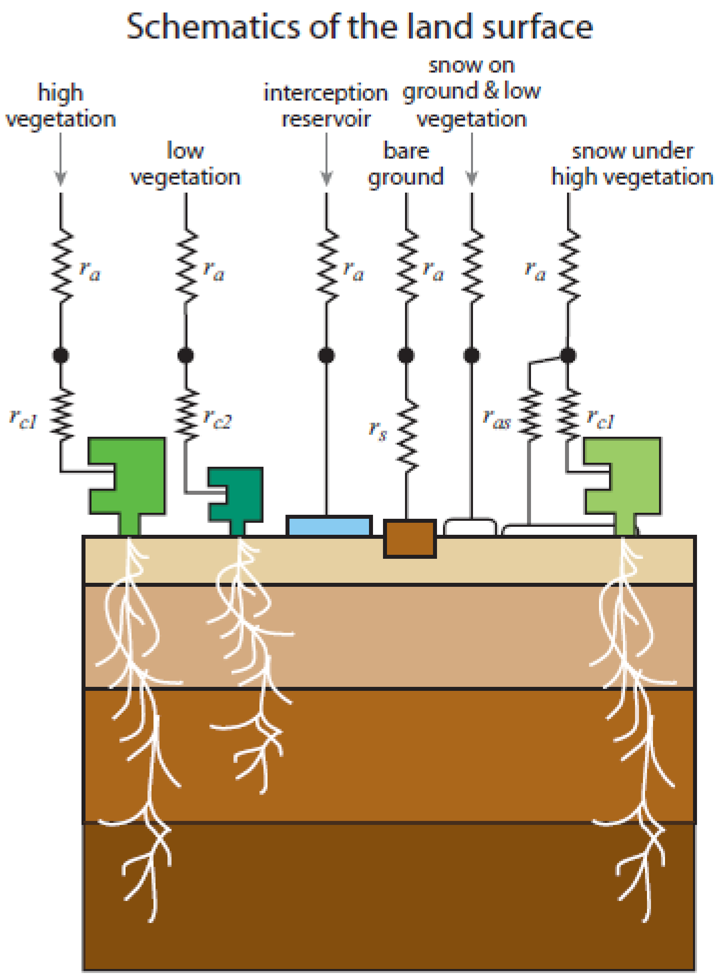
Root transpiration

Infiltration also depend on through fall amount

Vegetation heterogeneity

- Land surface is heterogeneous blend of vegetation at many scales
  - forest/cropland/urban area
  - within forest: different trees/moss/understories
- Most LSMs use set of parallel “plant functional types” (PFTs) with specific properties
  - gridbox mean or tiled
  - Some ecological models treat species competition and dynamics within PFTs
- Properties of PFTs
  - LAI
  - rooting depth
  - roughness
  - albedo
  - emission/absorption of organic compounds

ECLand :a tiles approach



Land/vegetation	Sea and ice
High vegetation	Open sea / unfrozen lakes
Low vegetation	Sea ice / frozen lakes
High vegetation with snow	
Snow on low vegetation	+ new urban tile
Bare ground	
Interception layer	

ECLand geographic characteristics

Fields	ERA15	TESSEL	CHTESSEL
Vegetation	Fraction	Fraction of low Fraction of high	Fraction of low Fraction of high
Vegetation type	Global constant (grass)	Dominant low type Dominant high type	Dominant low type Dominant high type
Albedo	Annual	**Monthly**	Monthly
LAI r_{smin}	Global constants	**Annual, Dependent on vegetation type**	**Monthly**
Root depth	1 m	**Annual, Dependent on vegetation type**	Annual, Dependent on vegetation type
Root profile	Global constant		

More realistic vegetation dynamic: Seasonal varying Leaf Area Index

Vegetation types dependent parameters

Index	Vegetation type	H/L	$r_{s,min}$ (sm^{-1})	LAI (m^2m^{-2})	c_{veg}	g_D (hPa^{-1})	a_r	b_r
1	Crops, mixed farming	L	180	3	0.90	0	5.558	2.614
2	Short grass	L	110	2	0.85	0	10.739	2.608
3	Evergreen needleleaf trees	H	500	5	0.90	0.03	6.706	2.175
4	Deciduous needleleaf trees	H	500	5	0.90	0.03	7.066	1.953
5	Deciduous broadleaf trees	H	175	5	0.90	0.03	5.990	1.955
6	Evergreen broadleaf trees	H	240	6	0.99	0.03	7.344	1.303
7	Tall grass	L	100	2	0.70	0	8.235	1.627
8	Desert	–	250					
9	Tundra	L	80					
10	Irrigated crops	L	180					
11	Semidesert	L	150					
12	Ice caps and glaciers	–	–					
13	Bogs and marshes	L	240					
14	Inland water	–	–					
15	Ocean	–	–					
16	Evergreen shrubs	L	225					
17	Deciduous shrubs	L	225					
18	Mixed forest/woodland	H	250					
19	Interrupted forest	H	175					
20	Water and land mixtures	L	150					

Era Interim

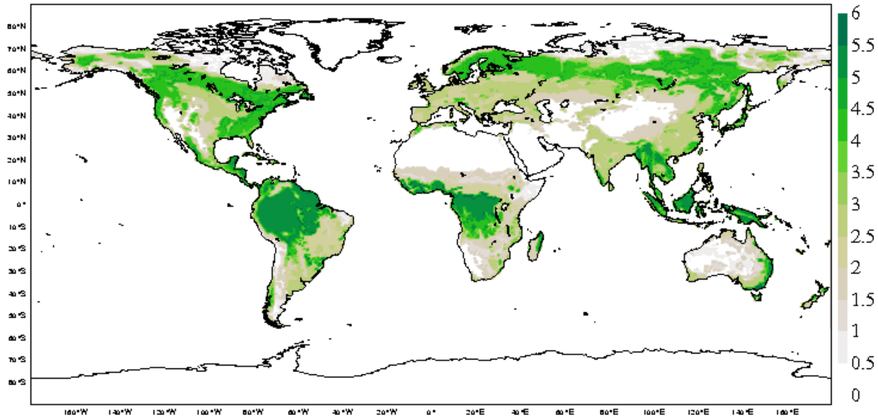
Cycle 48r1

Index	Vegetation type	H/L	$r_{s,min}$ (sm^{-1})	c_{veg}	g_D (hPa^{-1})	a_r	b_r
1	Crops, mixed farming	L	100	0.90	0	5.558	2.614
2	Short grass	L	100	0.85	0	10.739	2.608
3	Evergreen needleleaf trees	H	250	0.90	0.03	6.706	2.175
4	Deciduous needleleaf trees	H	250	0.90	0.03	7.066	1.953
5	Deciduous broadleaf trees	H	175	0.90	0.03	5.990	1.955
6	Evergreen broadleaf trees	H	240	0.99	0.03	7.344	1.303
7	Tall grass	L	100	0.70	0	8.235	1.627
8	Desert	–	250	0	0	4.372	0.978
9	Tundra	L	80	0.50	0	8.992	8.992
10	Irrigated crops	L	180	0.90	0	5.558	2.614
11	Semidesert	L	150	0.10	0	4.372	0.978
12	Ice caps and glaciers	–	–	–	–	–	–
13	Bogs and marshes	L	240	0.60	0	7.344	1.303
14	Inland water	–	–	–	–	–	–
15	Ocean	–	–	–	–	–	–
16	Evergreen shrubs	L	225	0.50	0	6.326	1.567
17	Deciduous shrubs	L	225	0.50	0	6.326	1.567
18	Mixed forest/woodland	H	250	0.90	0.03	4.453	1.631
19	Interrupted forest	H	175	0.90	0.03	4.453	1.631
20	Water and land mixtures	L	150	0.60	0	–	–

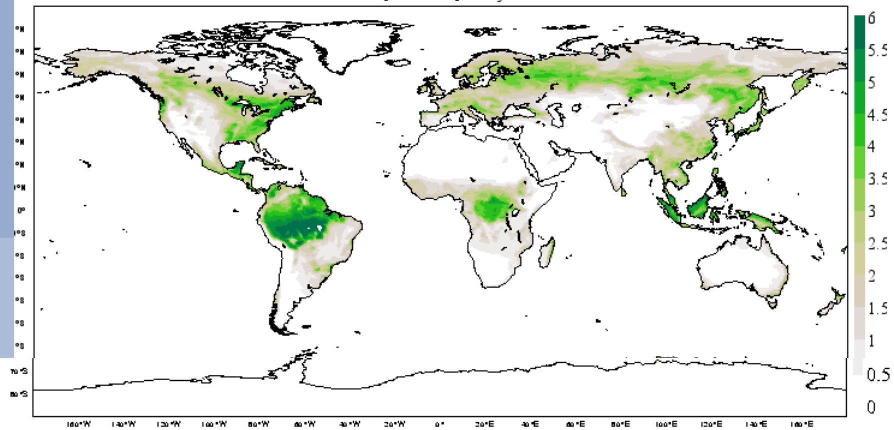
Era 5

Seasonal Varying Leaf Area Index

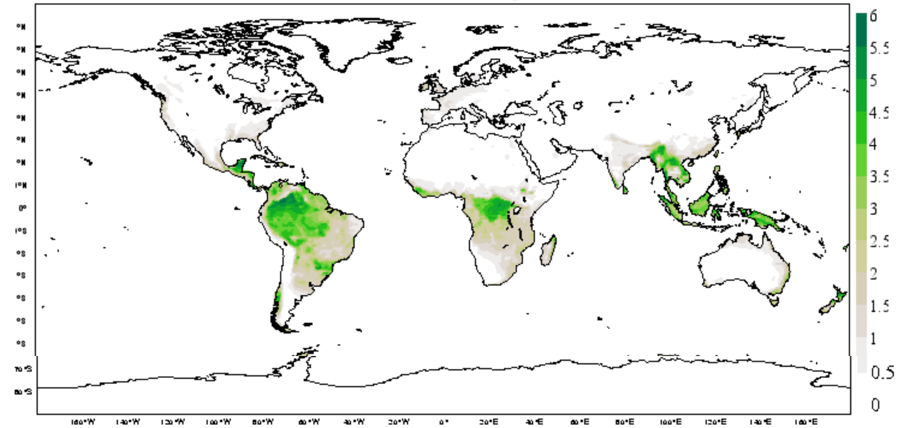
Total LAI [m2 m-2] -Operational



Total LAI [m2 m-2] -July MODIS



Total LAI [m2 m-2] -January MODIS



Obtained by the inversion of a 3D radiative transfer model which compute the LAI and FPAR based on the biome type and an atmospherically corrected surface reflectance thanks to a look-up-table

=>derived 8years (2000-2008) climatological time serie

Expected LAI impact on screen level Temperature

For vegetated area the evapotranspiration is parameterized as:

$$E_i = \frac{\rho_\mathrm{a}}{r_\mathrm{a} + r_\mathrm{c}} [q_\mathrm{L} - q_\mathrm{sat}(T_{\mathrm{sk},i})]$$

Where the canopy resistance r_c is defined following Jarvis(1976) as:

$$r_\mathrm{c} = \frac{r_{\mathrm{S,min}}}{LAI} f_1(R_\mathrm{s}) f_2(\bar{\theta}) f_3(D_\mathrm{a})$$

Where $\boldsymbol{r_{s,min}}$ is the minimum stomatal resistance, ***LAI*** is the leaf area index and $f1$, $f2$, $f3$ are respectively stress functions for the downward shortwave radiation $\boldsymbol{R_s}$, soil moisture $\boldsymbol{\theta}$ and vapour deficit $\boldsymbol{D_a}$

If *LAI* ↘ then r_c ↗ and *E* ↘ so *T2m* ↗

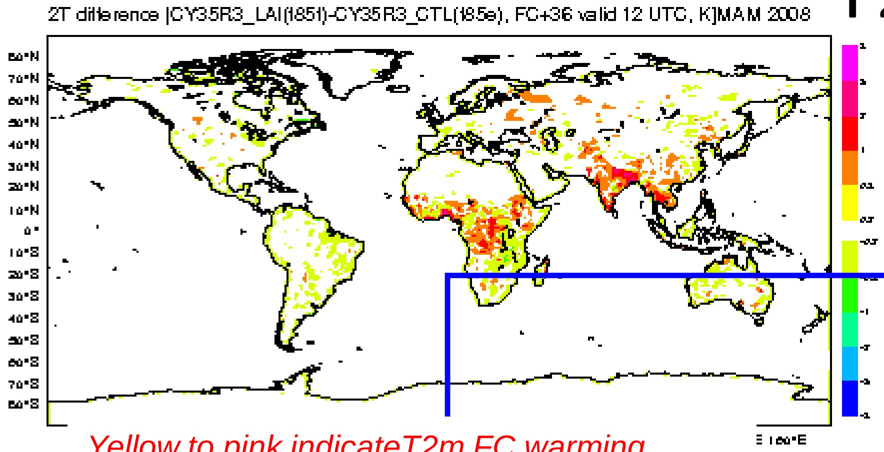
If *LAI* ↗ then r_c ↘ and *E* ↗ so *T2m* ↘

Analysis experiment: fc experiment validation

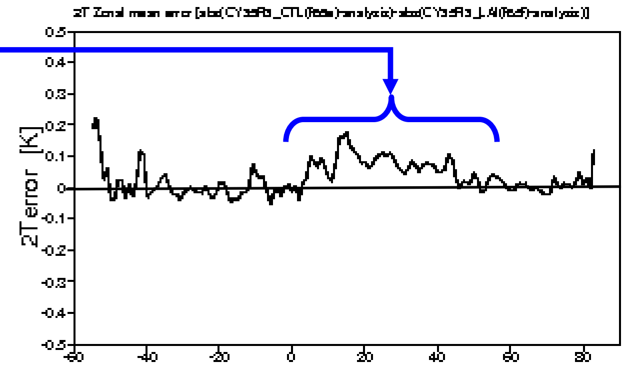
Setup:	T255
	14/02/2008 -1/09/2008
	Seasonal LAI vs fixed LAI

Sensitivity

T 2m

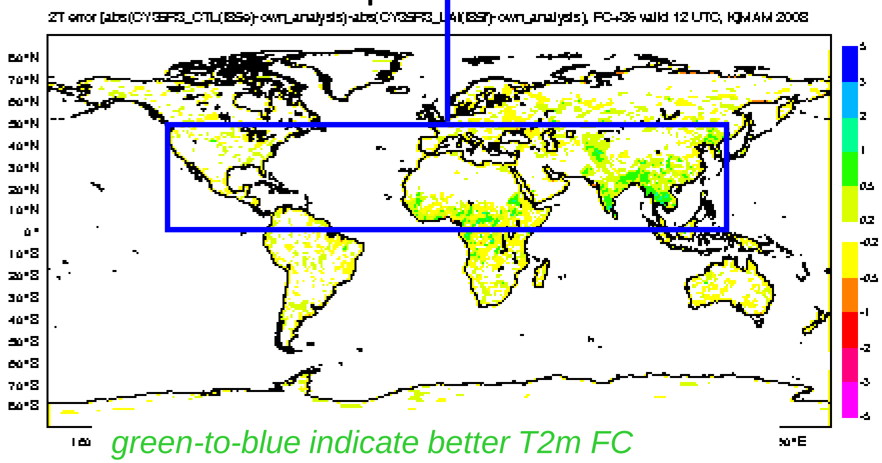


Yellow to pink indicateT2m FC warming



Zonal mean impact

Impact



green-to-blue indicate better T2m FC

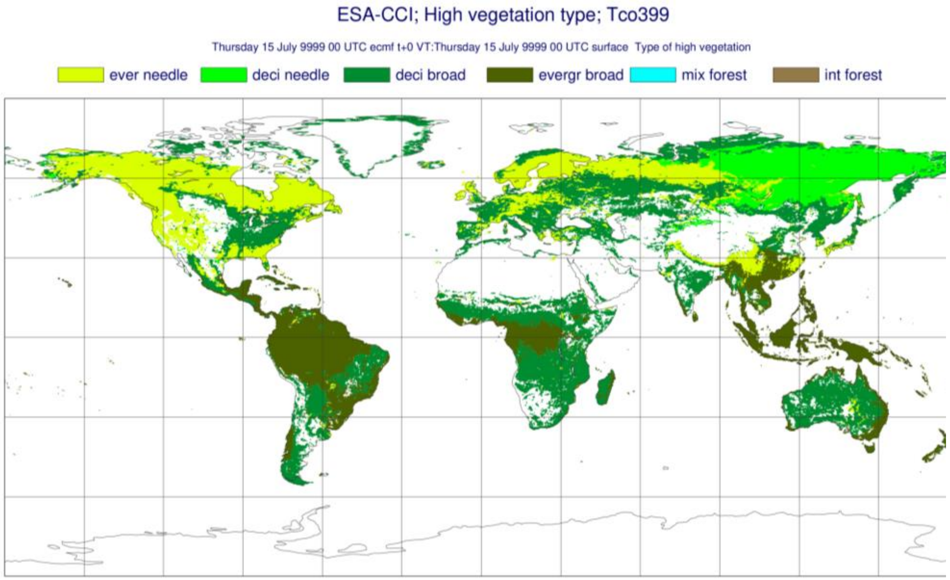
The MODIS LAI introduces a consistent warming seen in FC36h (12UTC) due to reduction of LAI in spring, (increasing vegetation resistance to ET).

This has beneficial impact on near surface temperature forecast (green being positive impact in reducing t2m bias by ~0.5degree)

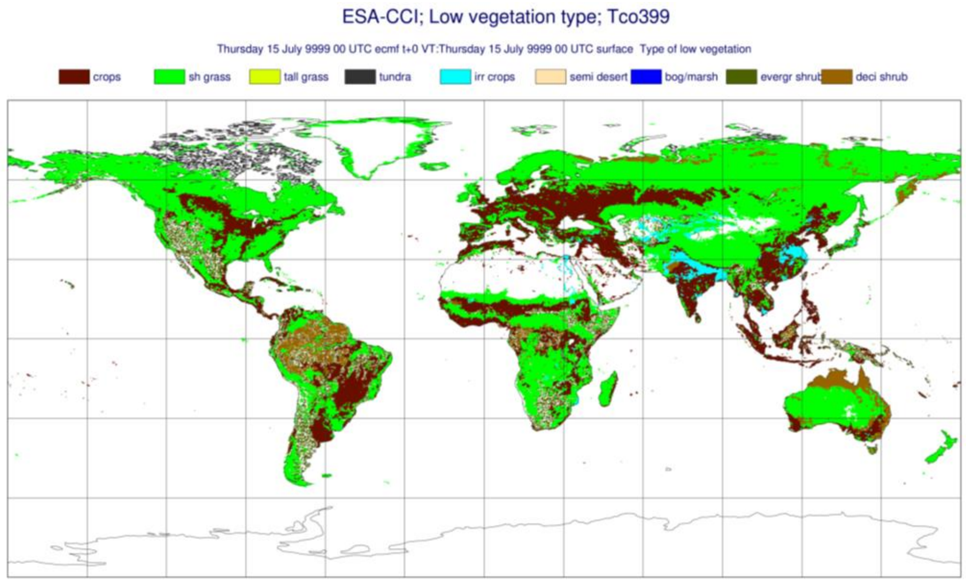
Updated and more realistic vegetation fields

Updated vegetation types

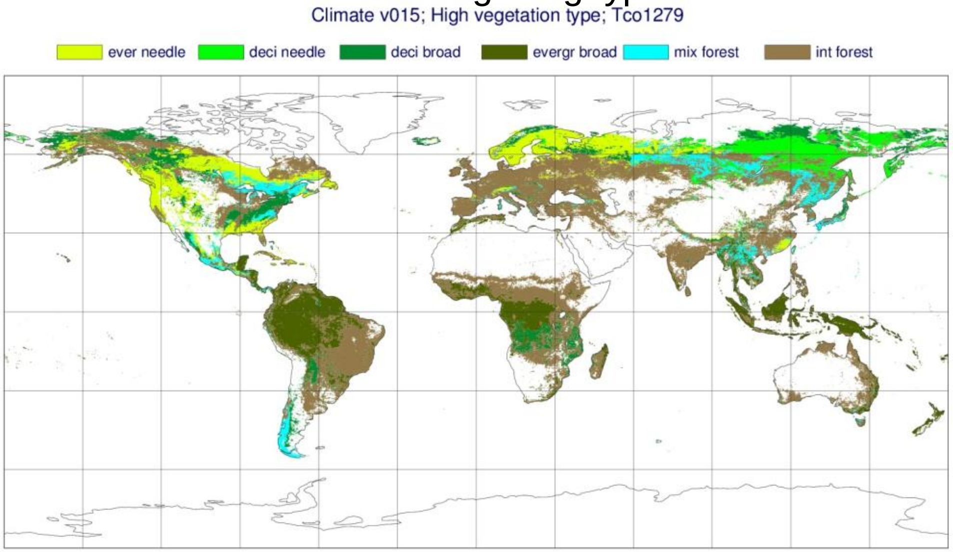
ESA-CCI high veg type



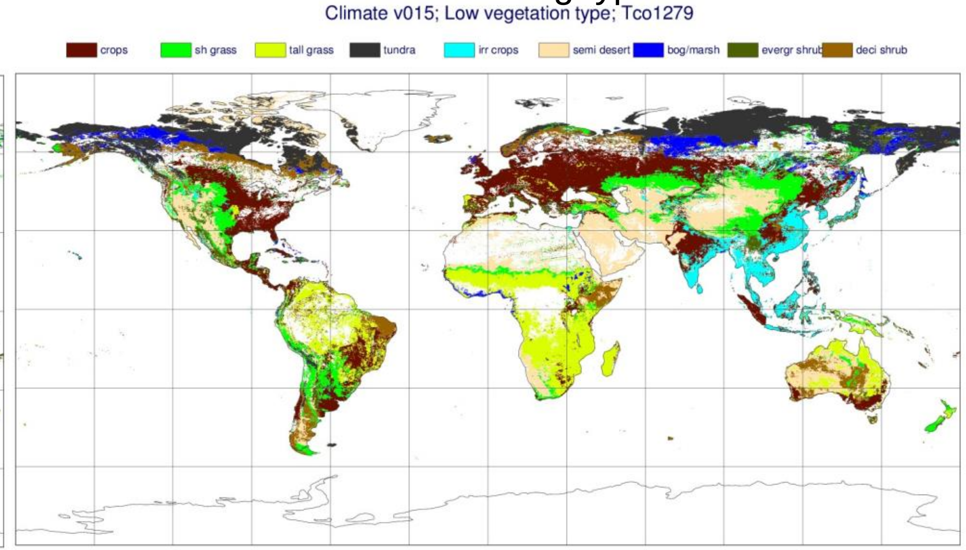
ESA-CCI low veg type



GLCC1.2 high veg type

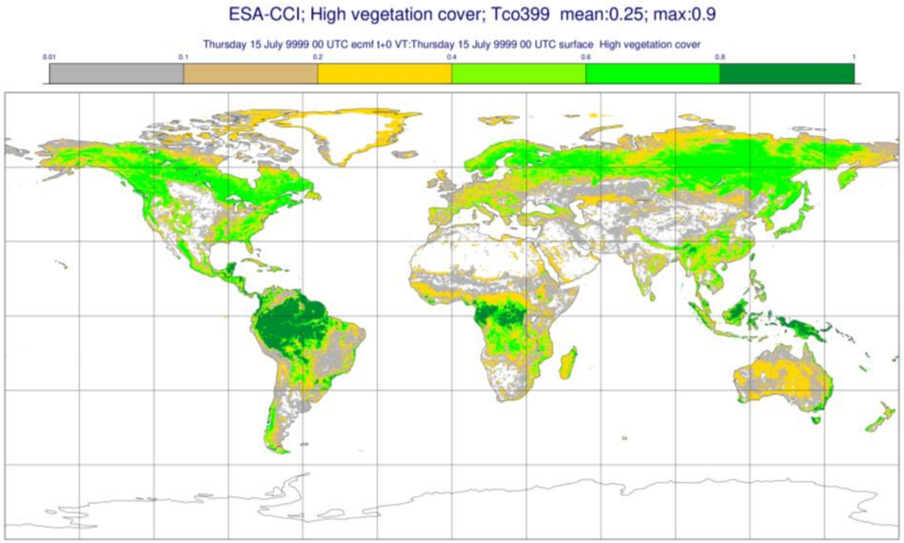


GLCC1.2 low veg type

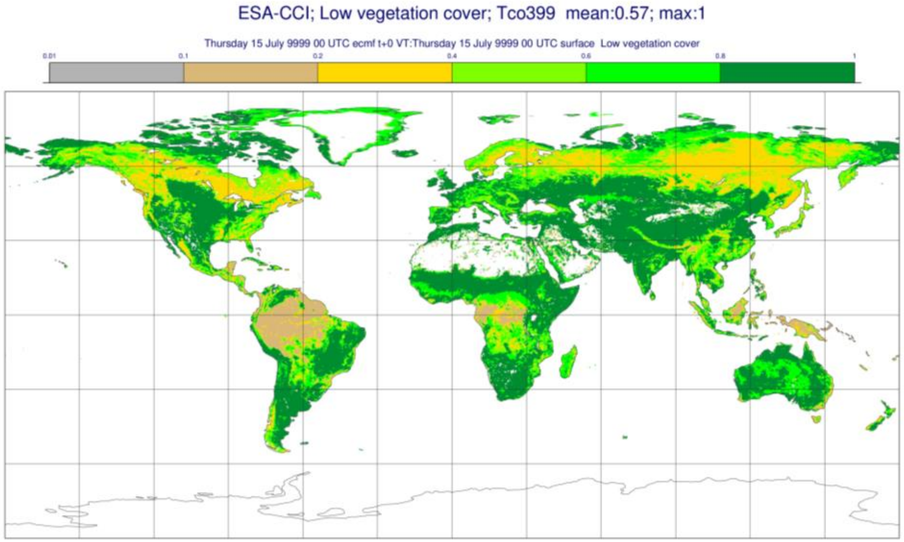


Updated vegetation cover

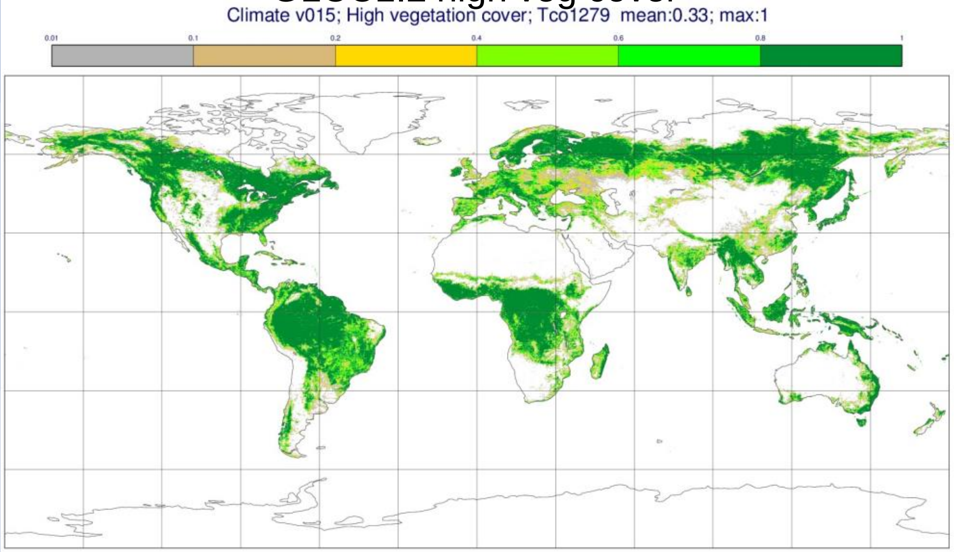
ESA-CCI high veg cover



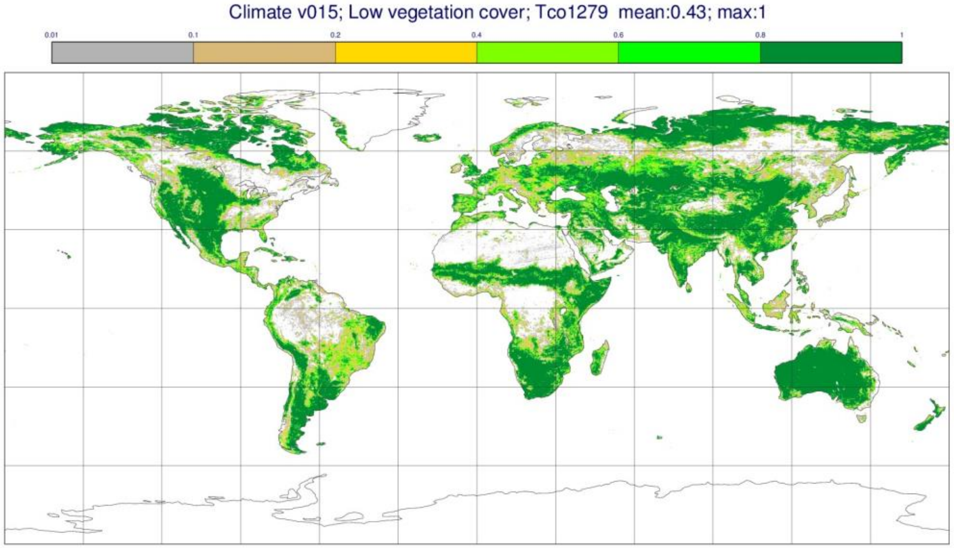
ESA-CCI low veg cover



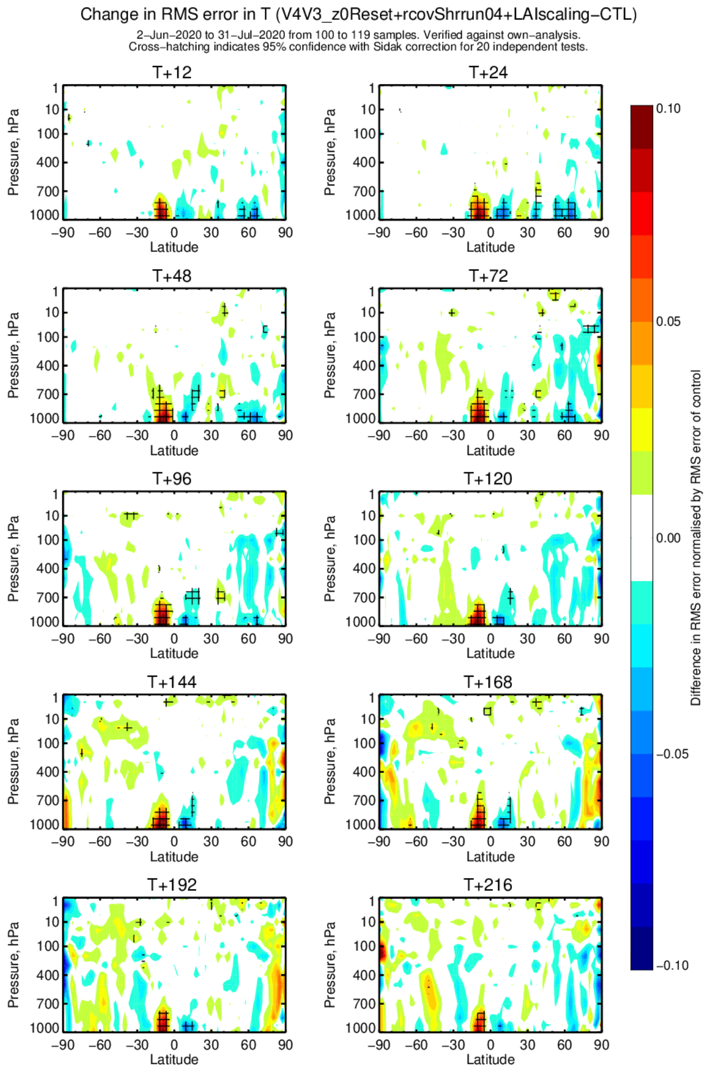
GLCC1.2 high veg cover



GLCC1.2 low veg cover



Atmospheric impact I: vegetation parameters revision



Stratified optimization with Vegetation types

Updated Vegetation:
C3S LULC + CGLS LAI (+ **new disaggregation operator**)

1^{st} revision of Canopy resistance and momentum parameters (rsmin, cveg)

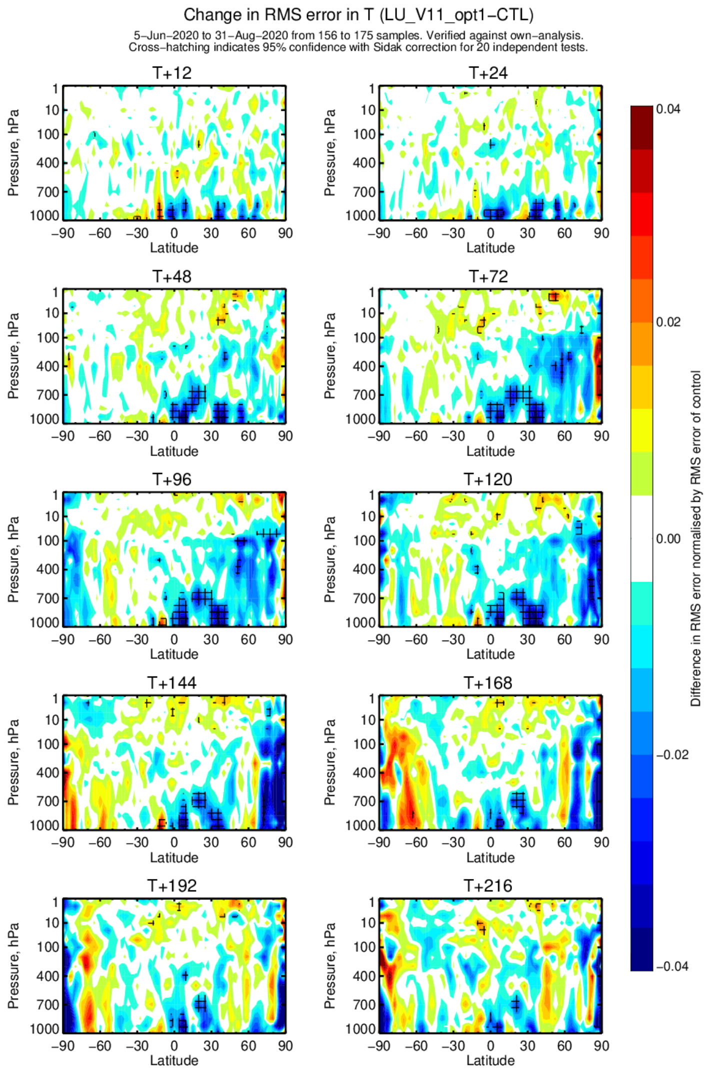
Updated Soil moisture stress

2^{nd} revision of Canopy resistance parameters (rsmin, cveg)

Time-varying LULC &/or LAI
(Offline surface model)

Prognostic LAI
(Offline surface model)

Atmospheric impact II: parametrisation update



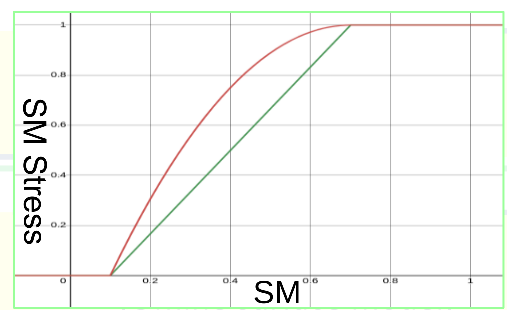
New Soil moisture Stress function

Updated Vegetation:
C3S LULC + CGLS LAI (+ **new disaggregation operator**)

1st revision of Canopy resistance and momentum parameters (rsmin, cveg)

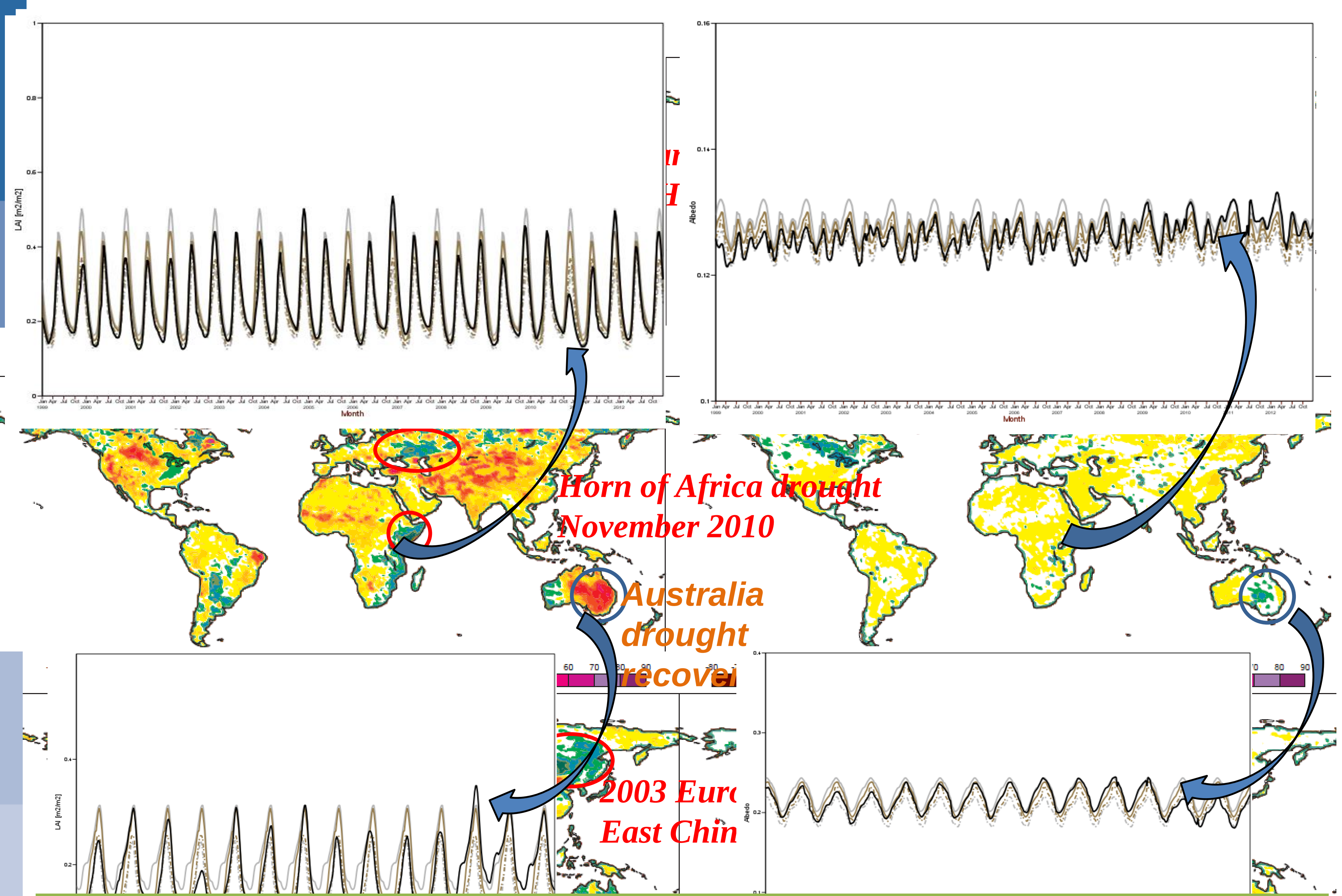
Updated Soil moisture stress

2nd revision of Canopy resistance parameters and momentum (rsmin, cveg)



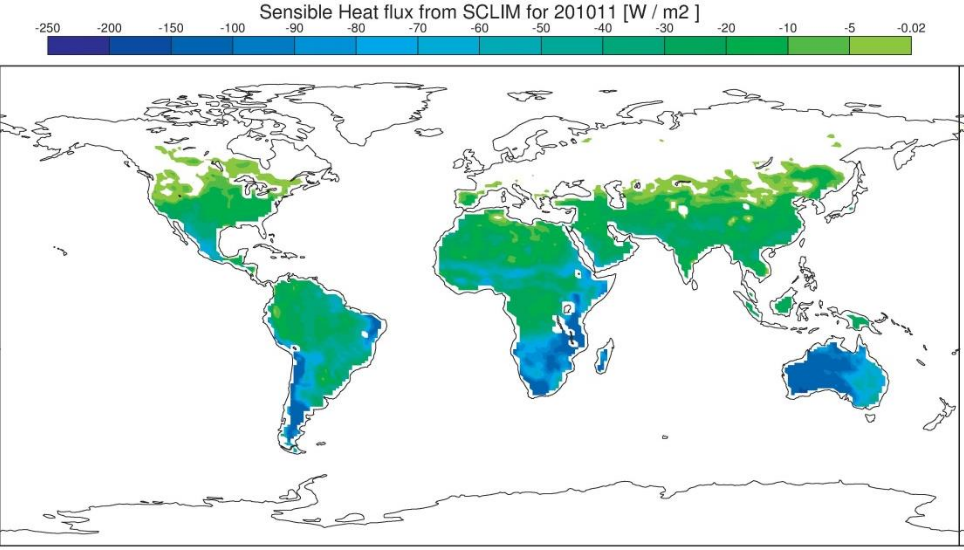
* Physically consistent with the soil matric potential
* Allow higher evapotranspiration under drier conditions

More and more realistic vegetation dynamic: Assimilation of Near Real Time LAI/Albedo

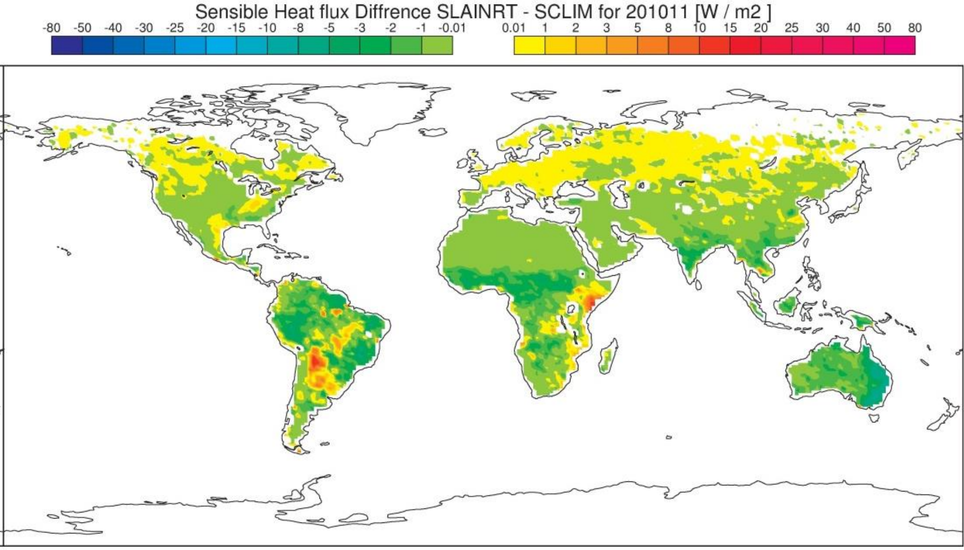


=> NRT analysed LAI is able to fairly detect/monitor anomalous year
=> The analysed LAI and albedo signal can be covariant mainly during wet year.

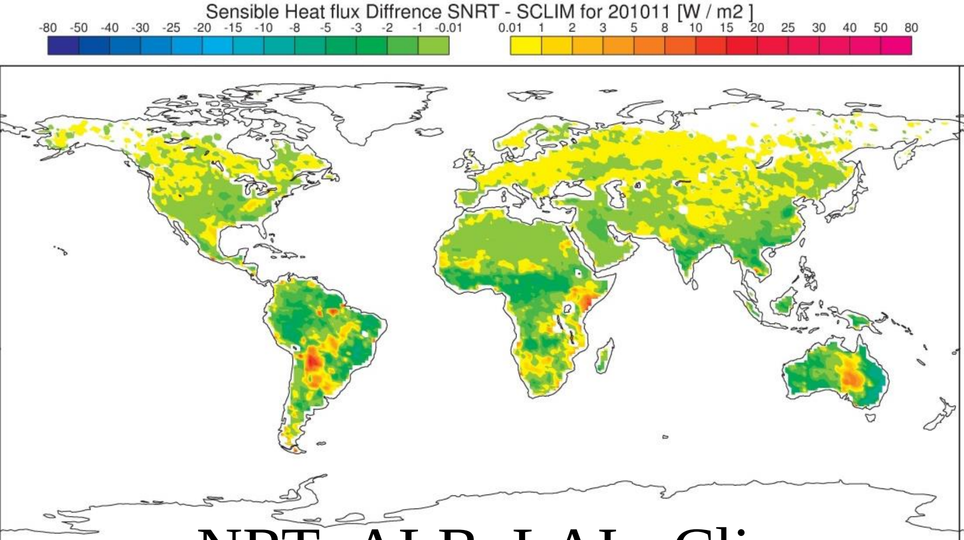
Sensible Heat flux



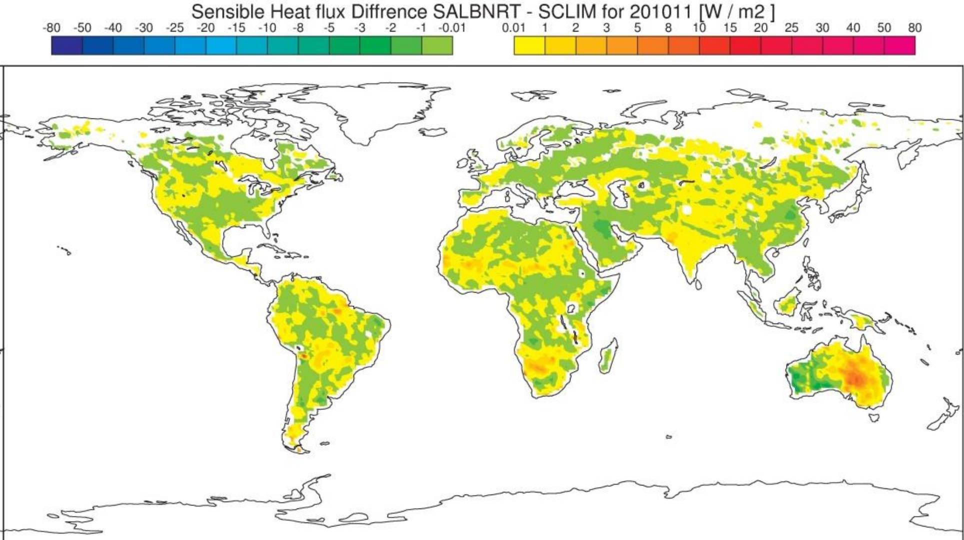
Clim



NRT_LAI - Clim

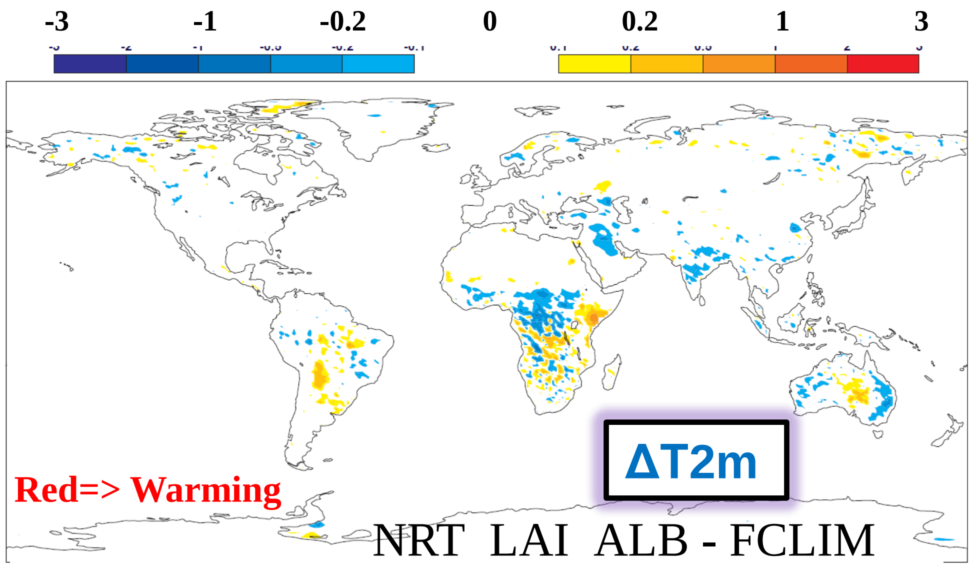


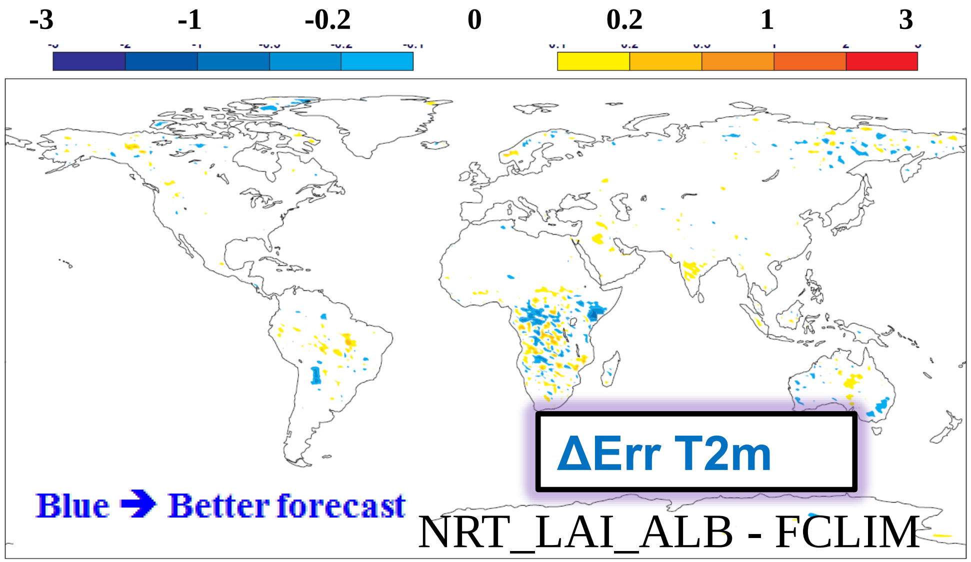
NRT_ALB_LAI - Clim



NRT_ALB - Clim

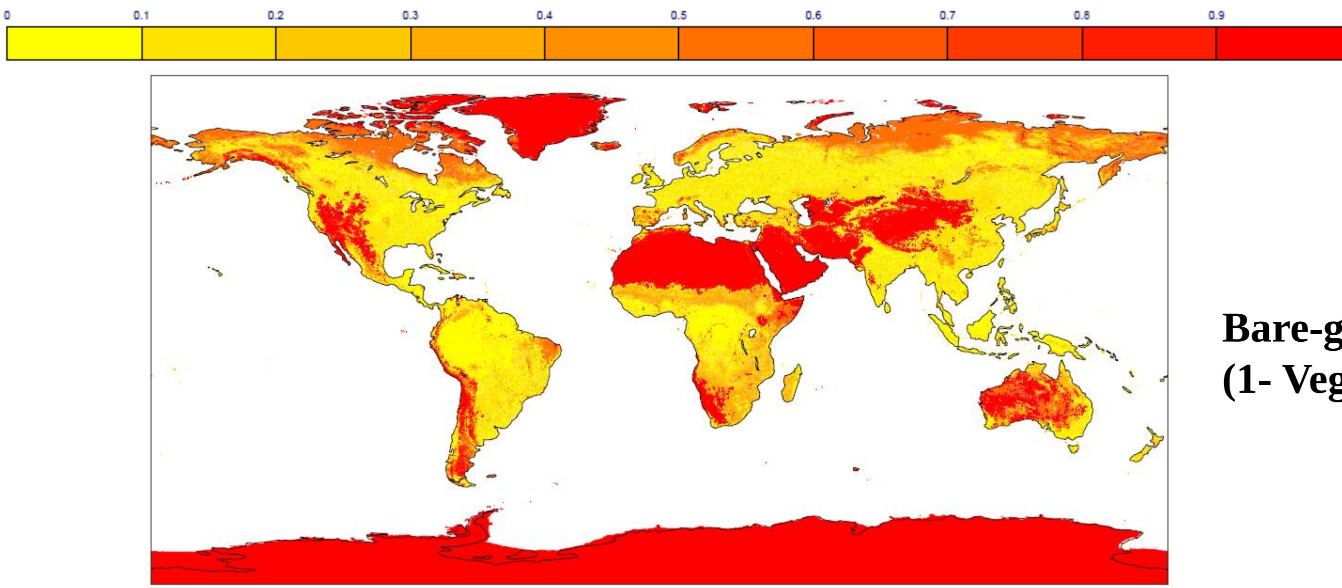
2m temperature sensitivity in coupled run





Even more realistic vegetation dynamic: Variable vegetation cover

February **May** **July** **October**

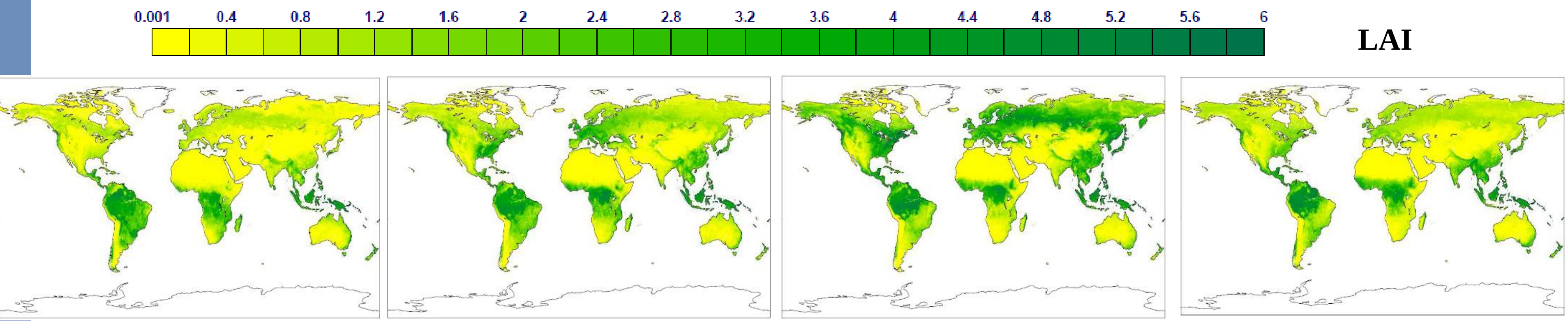


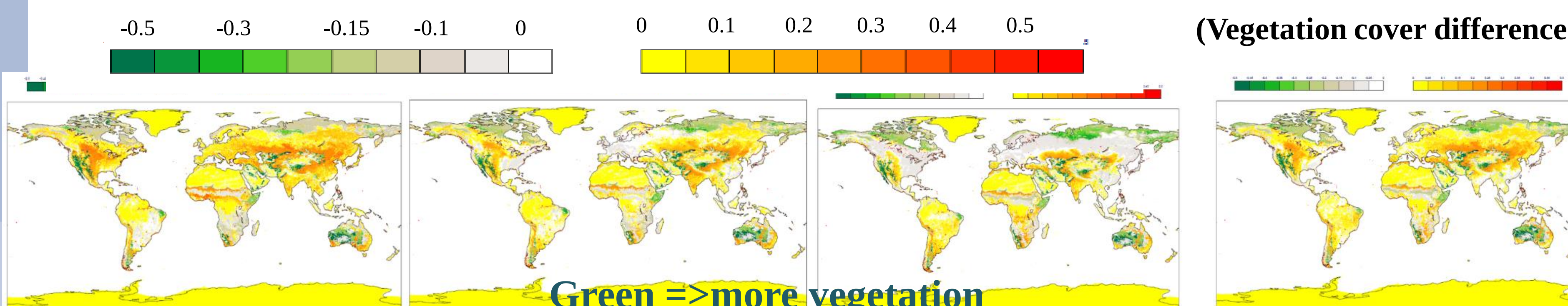
(1- Vegetation fraction)

Bare-ground/snow cover (1- Vegetation fraction)

➔ vegetation cover variation based on satellite observation of Leaf Area Index according to a modified Beer-Lambert law with clumping

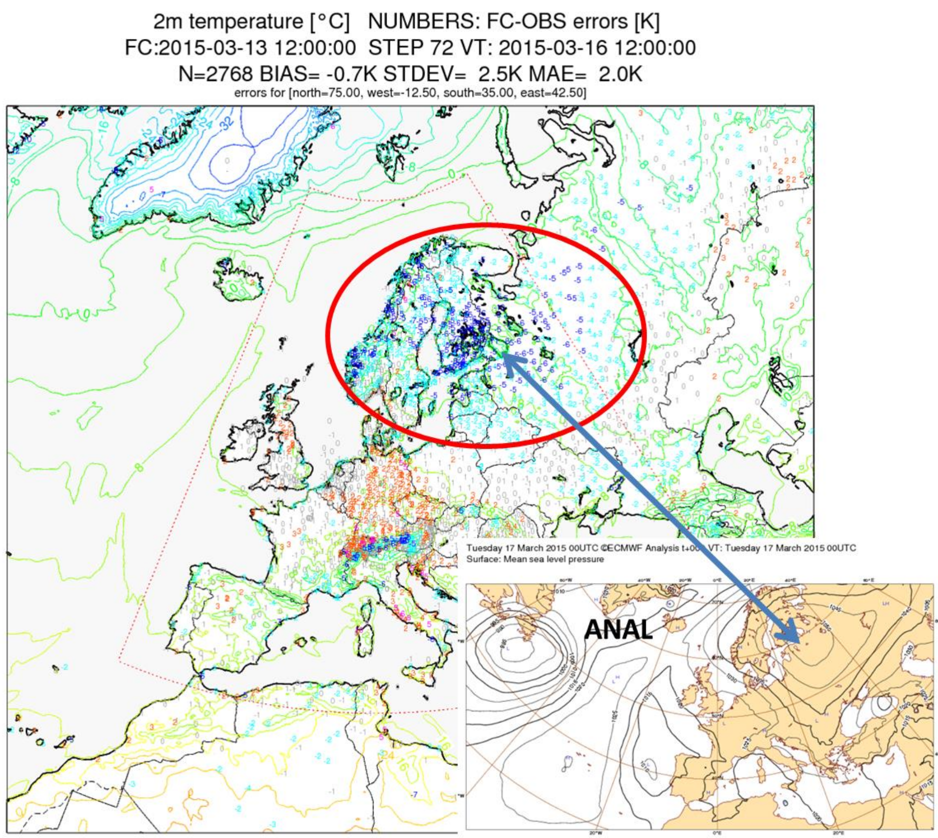
$$C_{veg} = 1 - e^{-0.5\omega LAI}$$

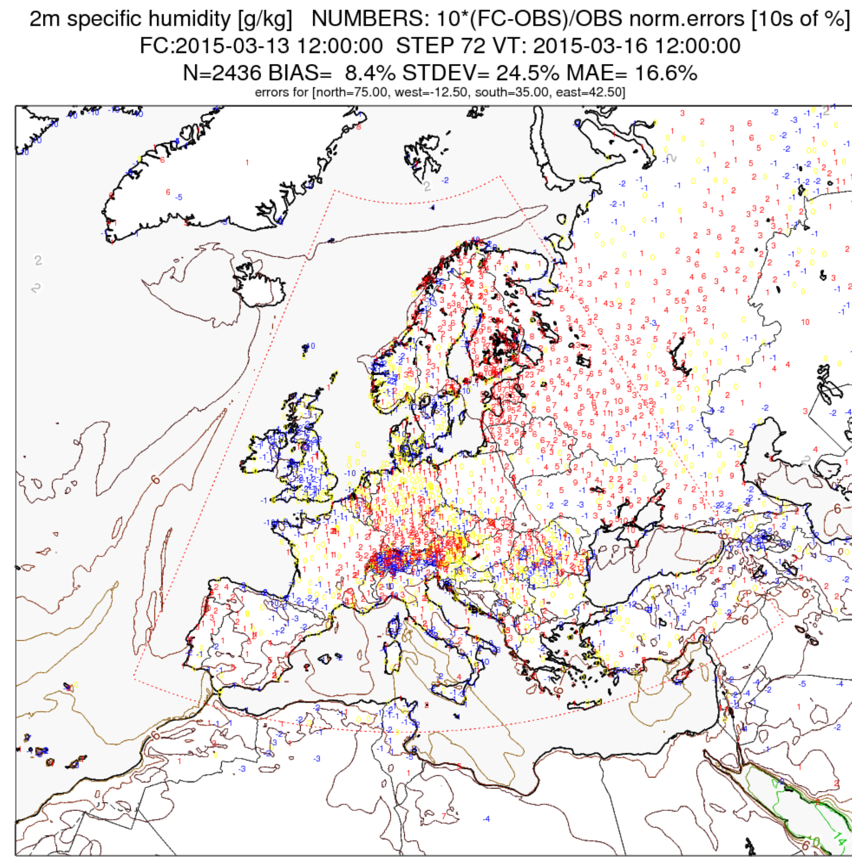




=> Physically-based seasonal variability of the vegetation cover

Impact in weather forecast mode



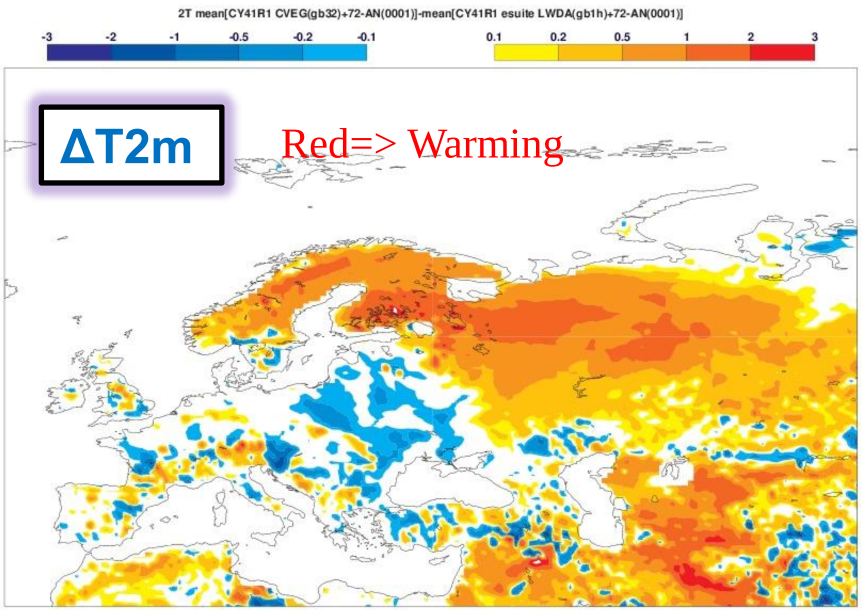


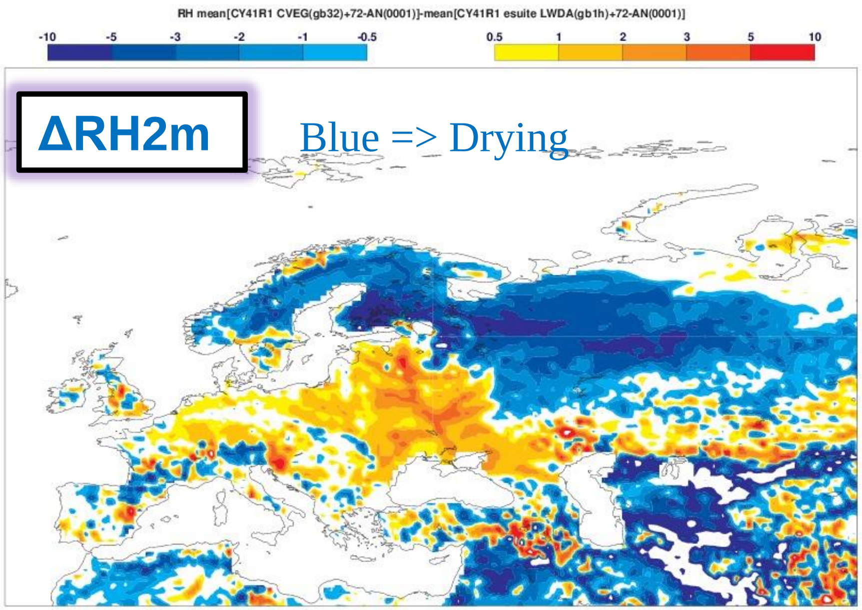
Cold bias on 2m Temperature 4K on average

Moist bias on 2m specific humidity 1g/kg on average

Weather forecasts sensitivity

=>Check the T 2m and RH on short term forecast fc+72 valid 12 UTC, March 2015



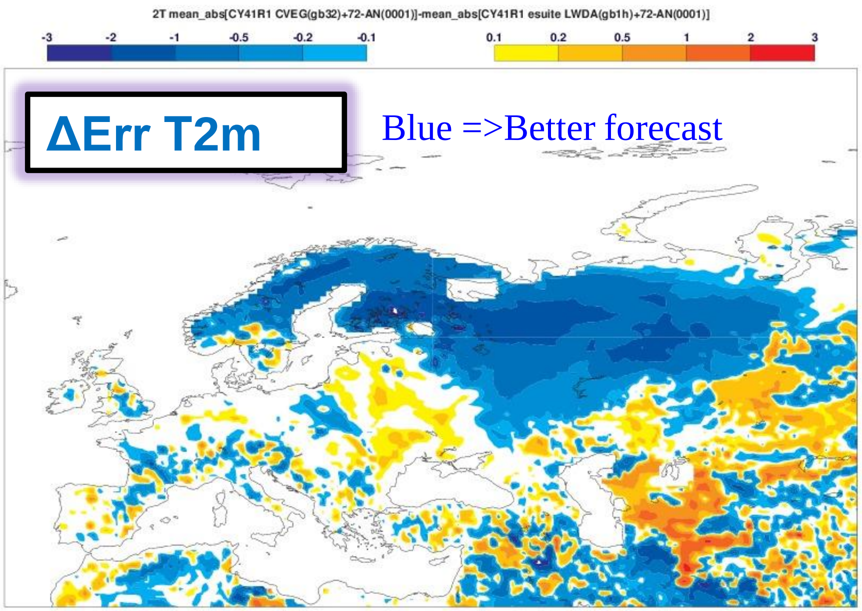


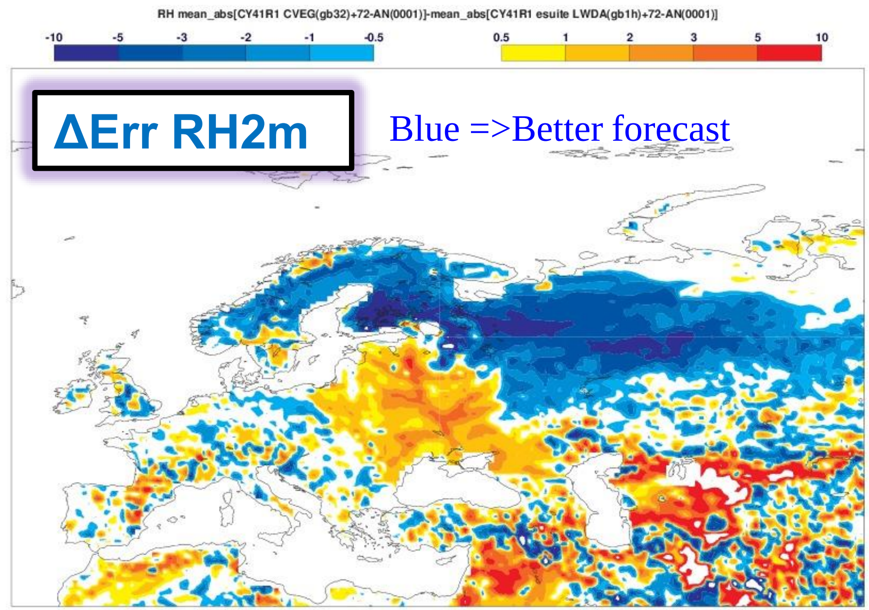
Sensitivity = *CVEG* - *CTL* ,

if >0 => Warming / adding moisture

if <0 => Cooling / removing moisture

Weather forecasts impact



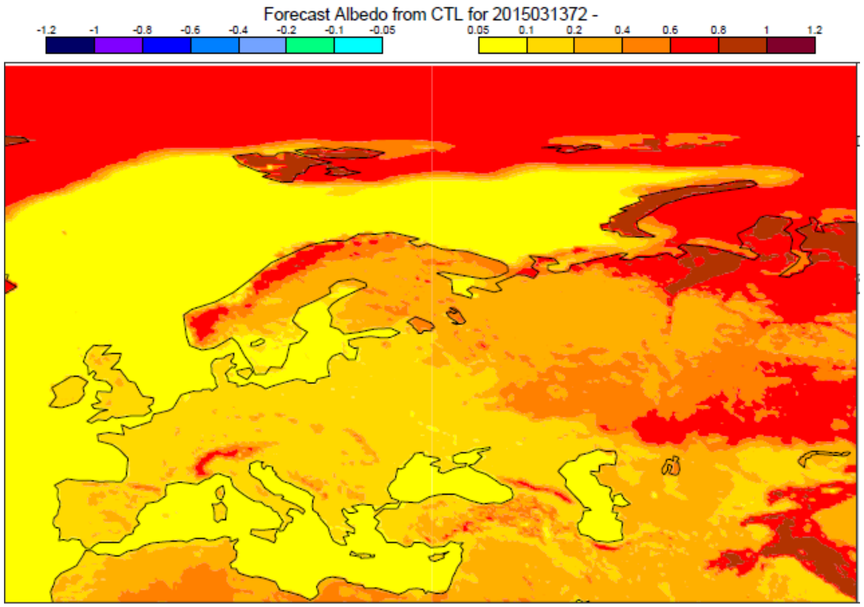


Impact = |*CTL* – *analysis*| - |*CVEG* – *analysis*| ,

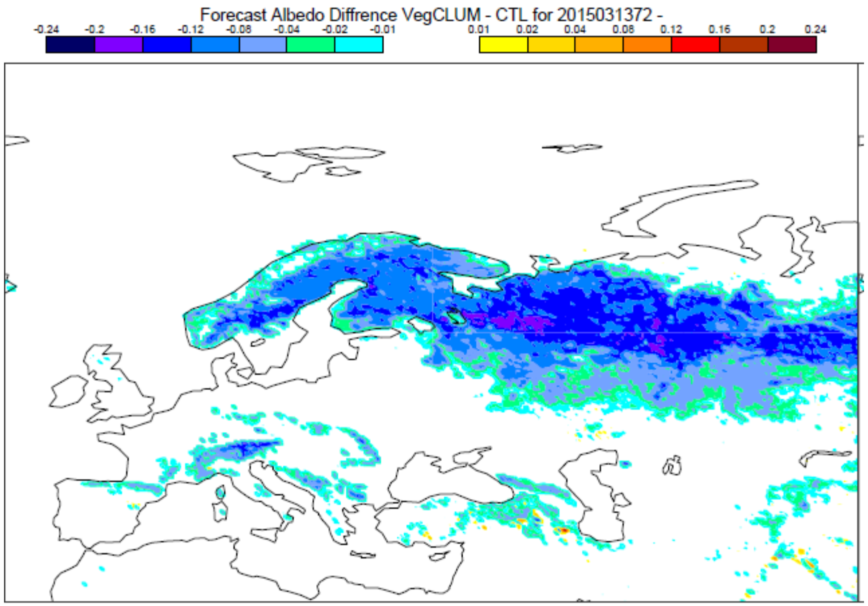
if >0 => relative error reduction from the analysis (positive impact)

if <0 => relative error increase from the analysis (negative impact)

Behind the scene: change in the forest albedo



Forecast Albedo for CTL

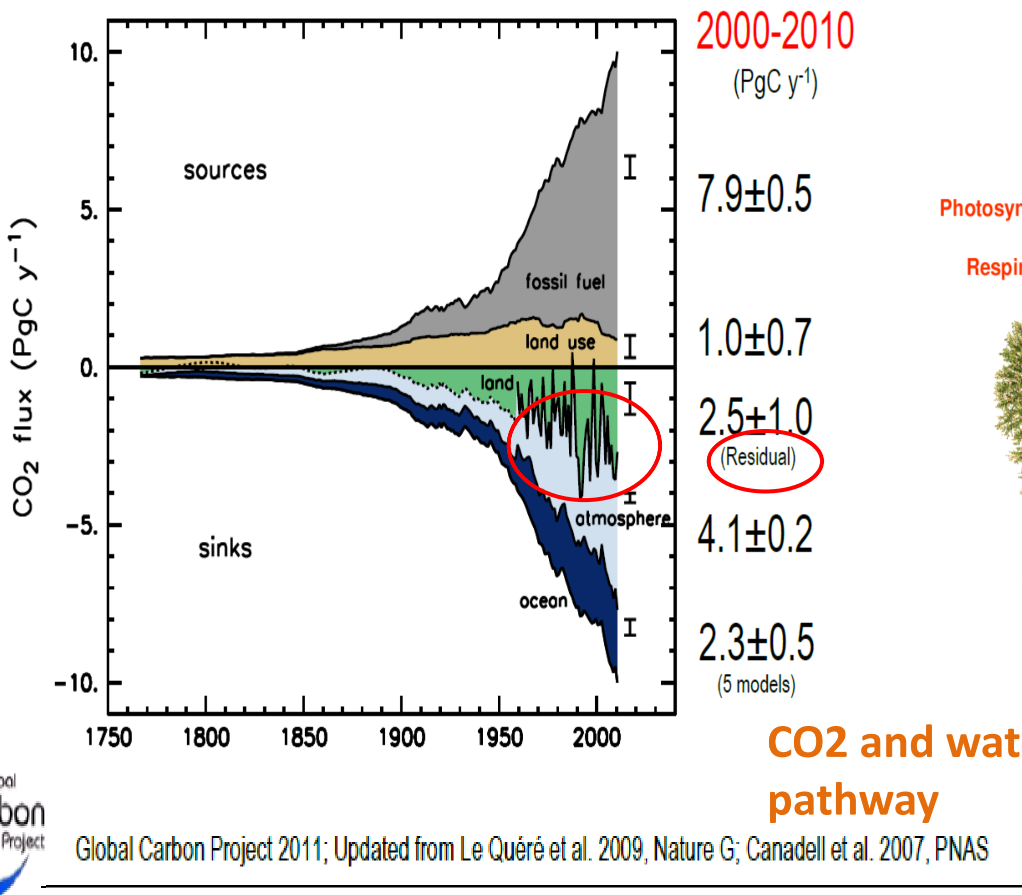


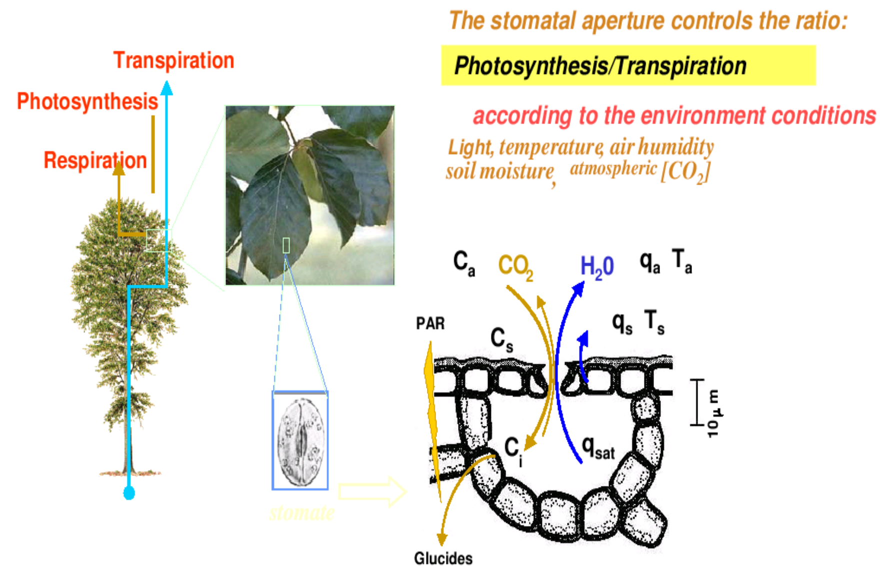
CVEG albedo – CTL albedo

=> Change in the vegetation cover is linked with a change in the forest albedo in presence of snow (in this case)

Introducing Land Carbon parametrisation

Why increasing complexity?





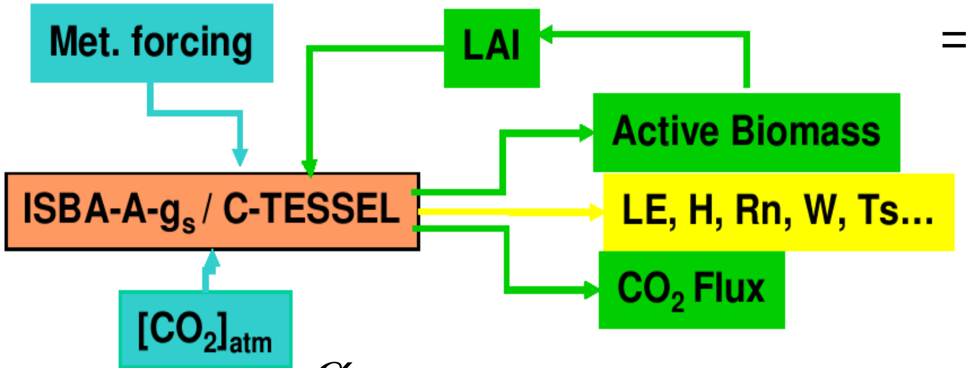
CO2 and water vapour share the same pathway

The land surface natural contribution to the global carbon budget is highly uncertain

=> A better representation of the vegetation processes

=> And also attempt to reduce uncertainties from the global carbon budget

Land carbon/photosynthesis-based canopy resistance parametrisation



$$A_n = \frac{\alpha}{r_{cc}}(C_s - C_i)$$

$$E = \frac{\beta}{r_c + r_a}(q_a - q_{sat}), r_c = f(r_{cc})$$

- $A_n = \rho\, f(\text{soil m})\, \Delta CO_2 / r_c$

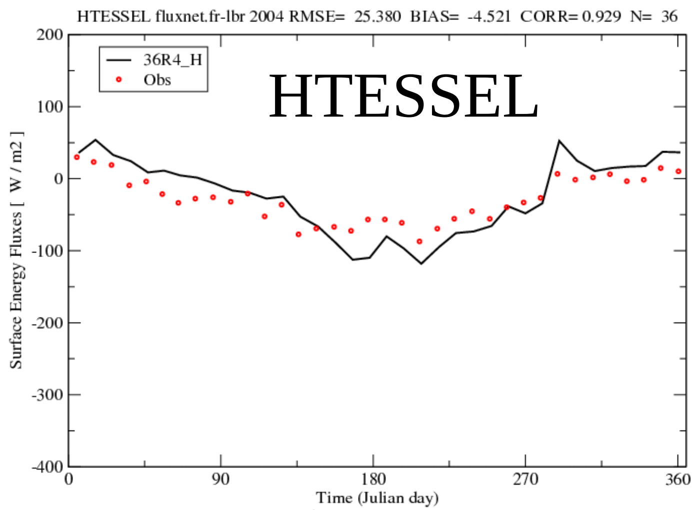
=> r_c back-calculated from

- Empirical soil moisture dependence
- CO_2-gradient ΔCO_2 is also $f(q_{sat} - q)$
- Net photosynthetic rate A_n
  - $A_{n,max}$
  - Photosynthetic active Radiation (PAR)
  - temperature
  - $[CO_2]$

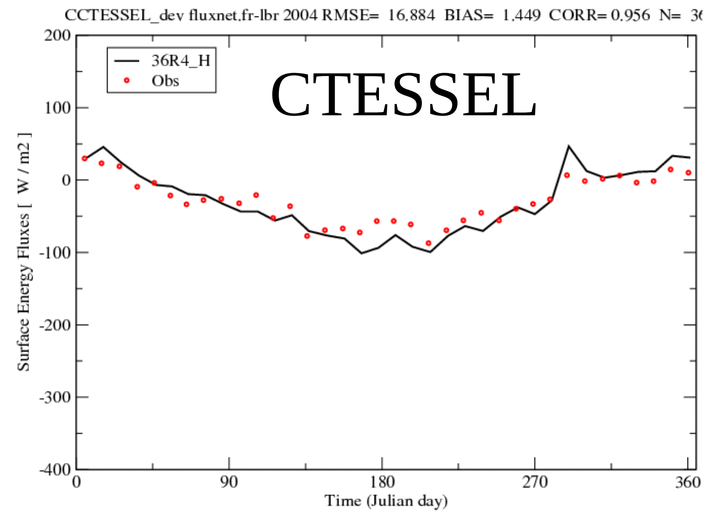
CTESSEL combines HTESSEL (Balsamo et al. 2009) with the A-gs model used within the ISBA-Ags (Calvet et al.1998) and developed by Jacobs et al. (1996);

- ➔ **Account for the effect of CO2 concentration and the interactions between all environment factors on the stomatal aperture.**
- ➔ **Replaces the Jarvis-type stomata conductance by a photosynthesis dependant-type stomata conductance (Jacobs et al.1996)**
- ➔ **The model can account for the vegetation response to the radiation at the surface, temperature, soil moisture stress**
- ➔ **Vegetation Assimilation of CO2 can be used to drive a vegetation growth module to simulate LAI**
- ➔ **The Ecosystem Respiration is parametrized as a function of soil temperature, soil moisture and biome type via a reference respiration parameter**

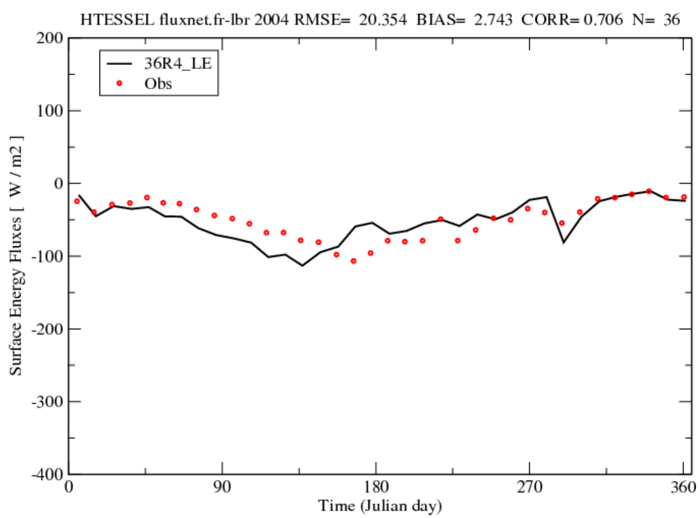
Jarvis Vs photosynthesis-based evapo-transpiration (offline run)



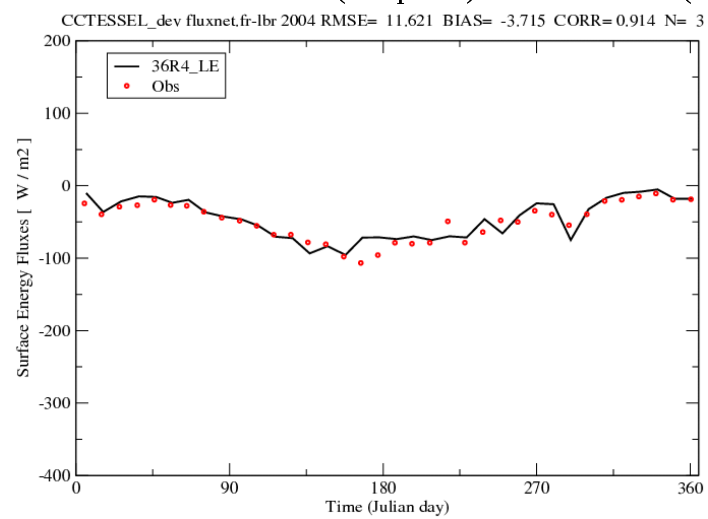
H



Surface sensible heat flux (W/m^2) compared with flux-tower observations over Fr-LBr for HTESSEL (left panel) and CTESSEL (right panel)



LE



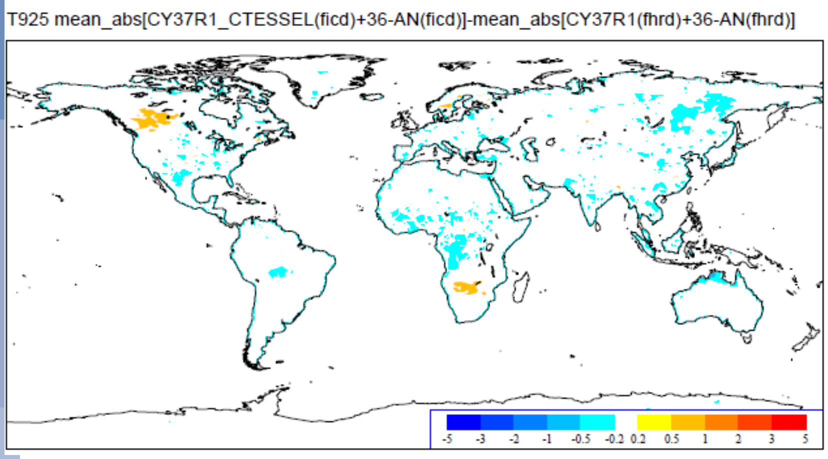
Surface laten heat flux (W/m^2) compared with flux-tower observations over Fr-LBr for HTESSEL (left panel) and CTESSEL (right panel).

- **CTESSEL improves the LE/H simulations (Photosynthesis-based vs Jarvis approach).**

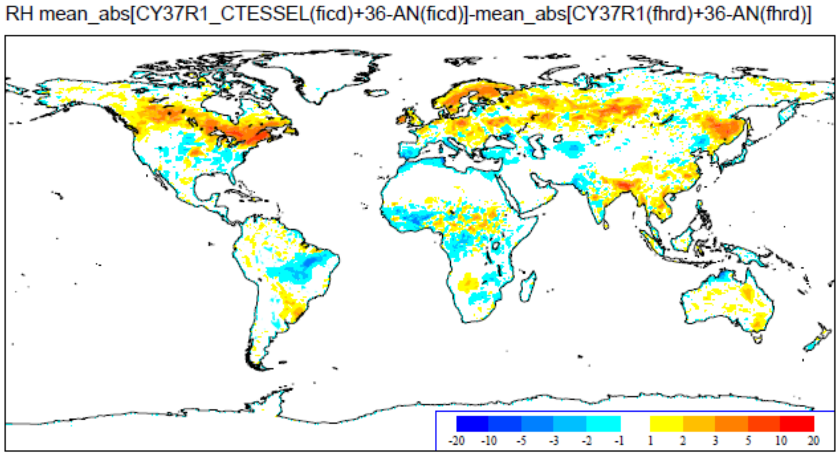
LE/H: When "good" is not enough?

(Interaction with the atmosphere)

2m T Error differences from the CTL



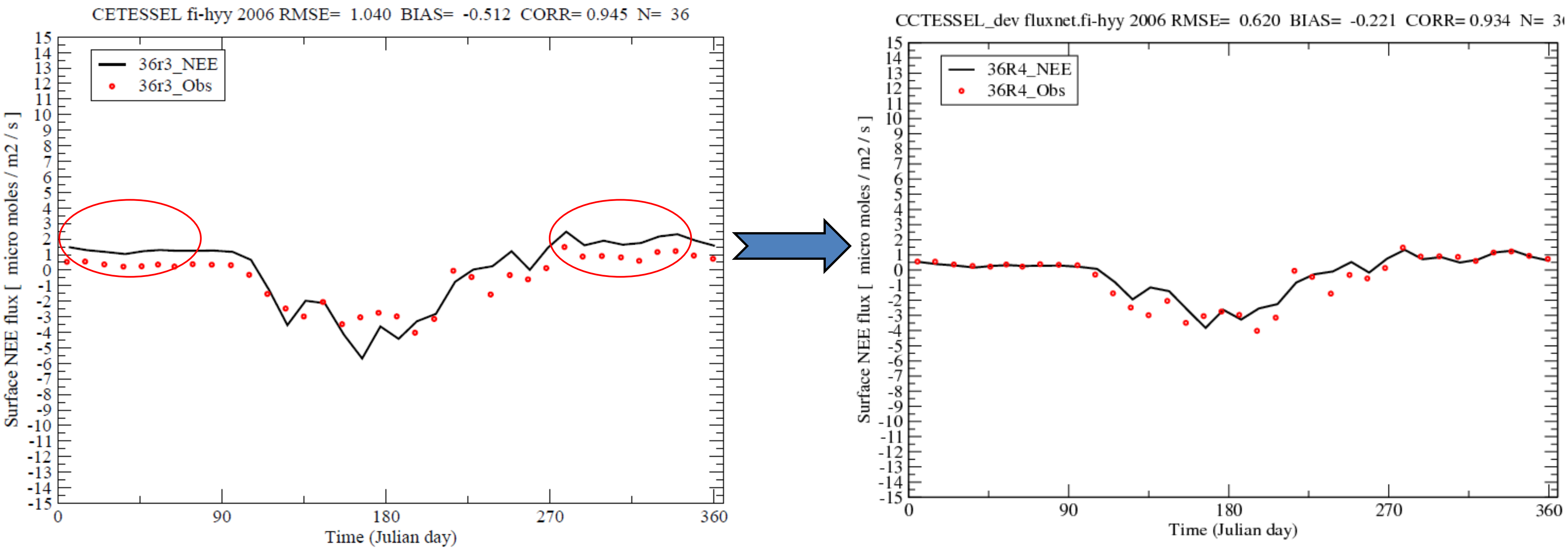
2m Rh Error differences from the CTL



Having better LE/H heat flux from the surface does not always lead to a better atmospheric prediction =>interaction with other processes and compensating errors?

Soil Respiration improvement for winter season

$$NEE = A_n - R_{soil}$$



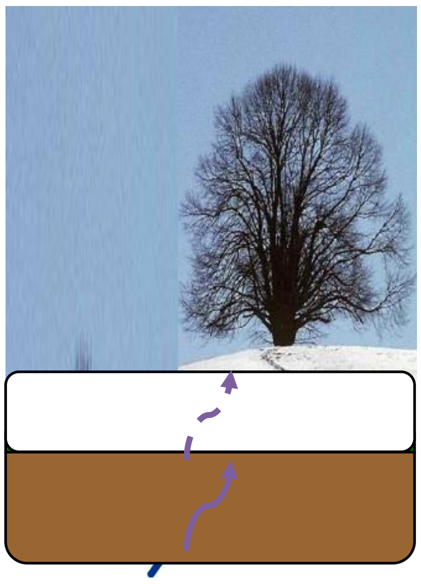
Example of NEE (micro moles /m^2/s) predicted over the site Fi-Hyy taking the cold process into account (right) and previous simulation (left) by CTESSEL (black line) and observed (red dots)

Feedback from the atmosphere can contribute to improve the physical understanding and adjust the contribution from the surface

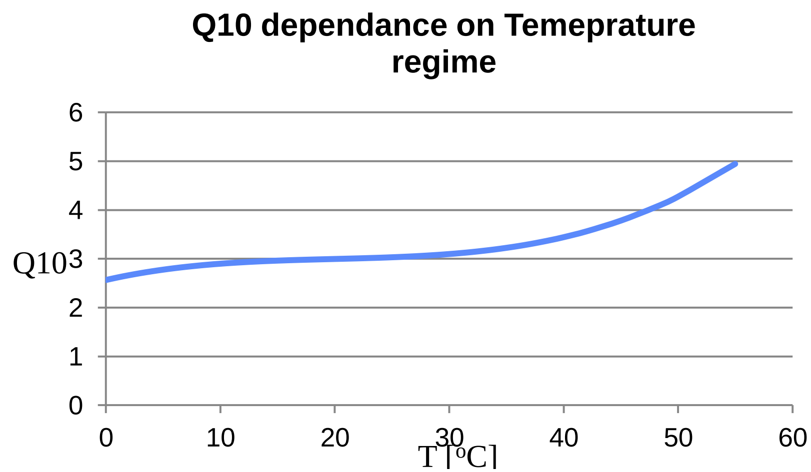
Soil Respiration and winter improvement

$$R_{soil} = R_0 Q_{10}^{(0.1(T_{soil}-25))} f_{sm}$$

$$R_{soil} = R_0 e^{-\alpha . Z_{snow}} Q_{10}^{(0.1(T_{soil}-25))} f_{sm}$$

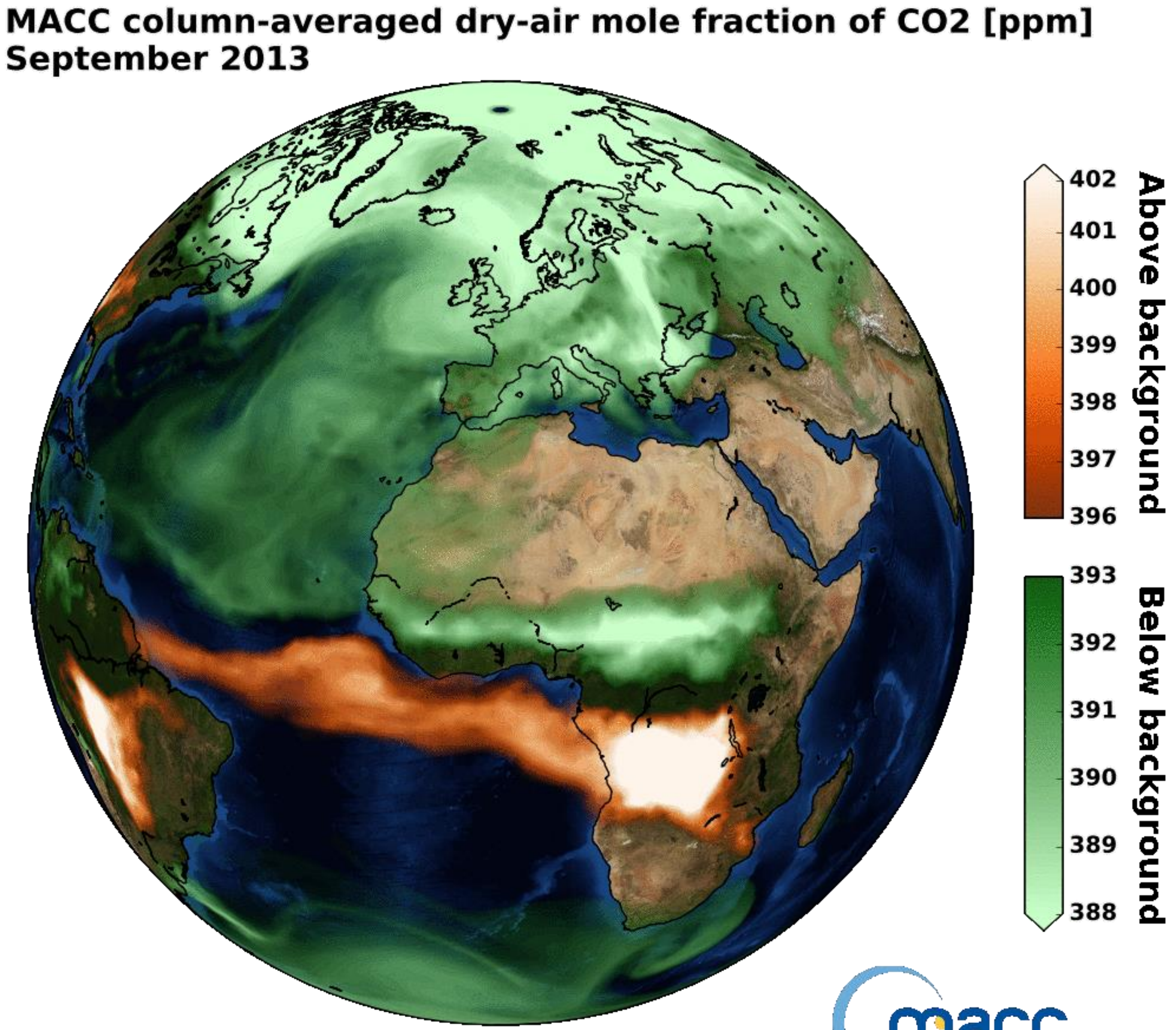


Including a snow attenuation effect on the soil CO2 emission



Including a temeperature dependancy on the Q10 parameter (McGuire et al., 1992)

Near Real Time CO_2 concentration modelled in CAMS

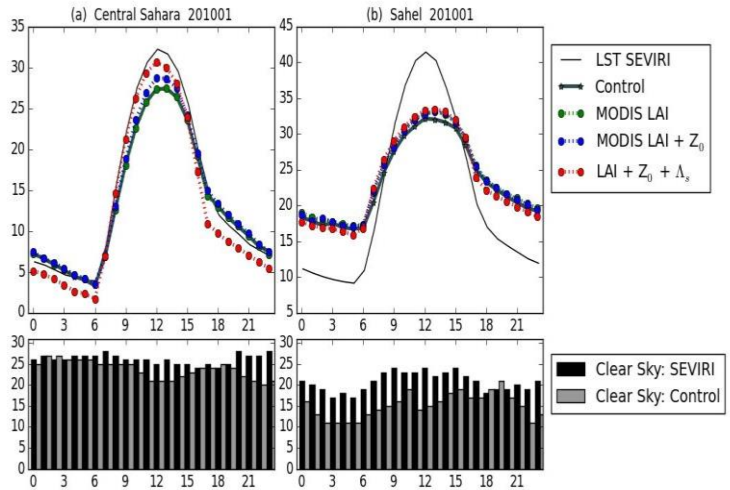


- CHTESSEL fluxes used in MACC-II (CAMS) to forecast CO2 atmospheric concentrations (16 km global simulation)
- Green colours highlight effects of photosynthetic uptake by vegetation
- Diurnal cycle (fluxes driven) and Synoptic variability (Weather driven) are crucial elements for simulating the CO2 of the Earth system.

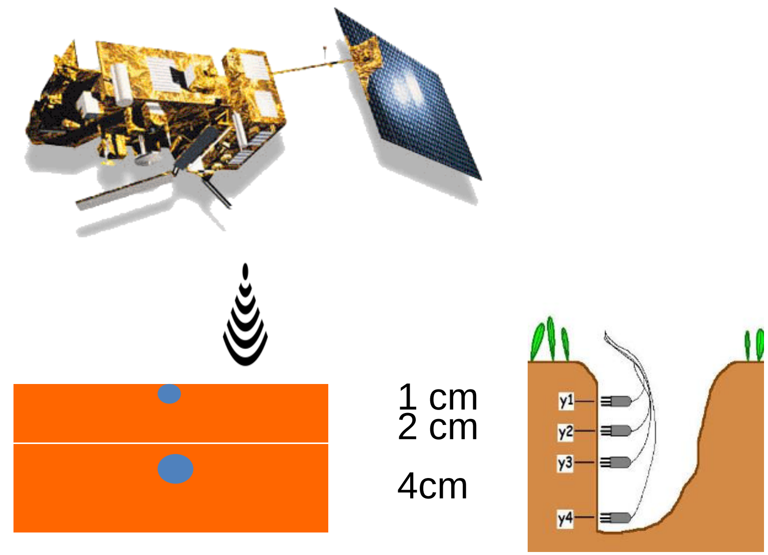
Agusti-Panareda et al. (2014, ACP), Boussetta et al. (2013 JGR)

Improved soil vertical discretization

Why increasing complexity?



The model bias in Tskin amplitude (*Trigo et al. 2015)*

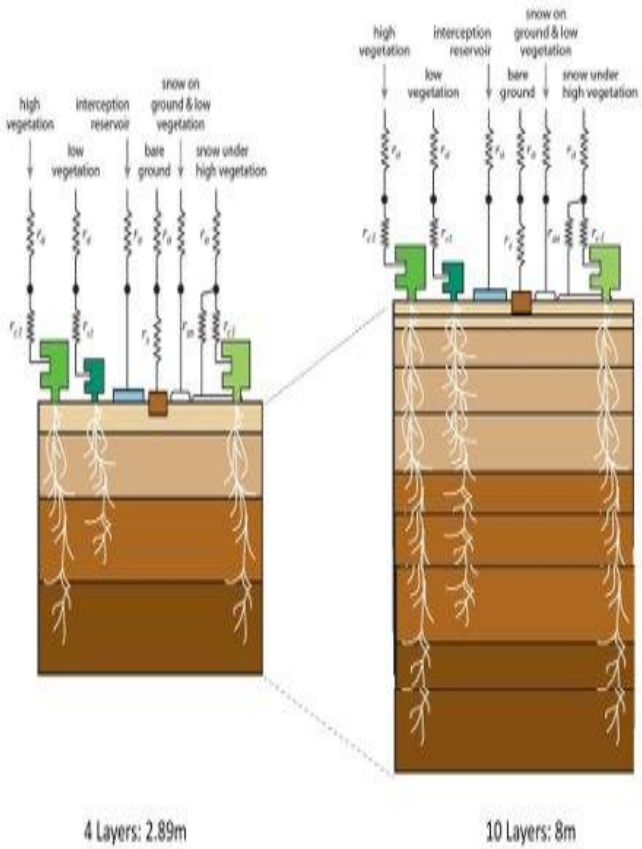


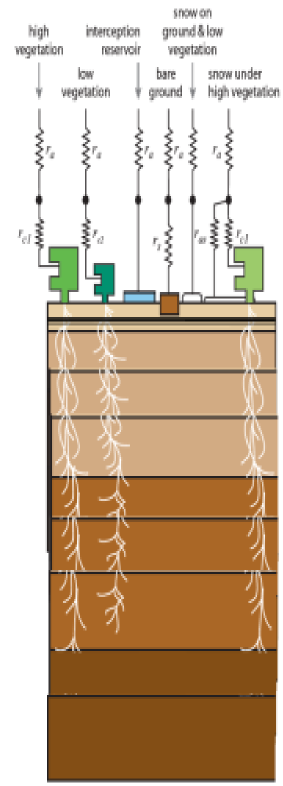
Thinner surface layers

vertical discretization would improve the match with Observation

Dirmeyer et al. 2021, also showed the importance of an accurate soil representation for a proper L-A feedback that could simulate drought such us the European 2018 one

Which configuration is good?





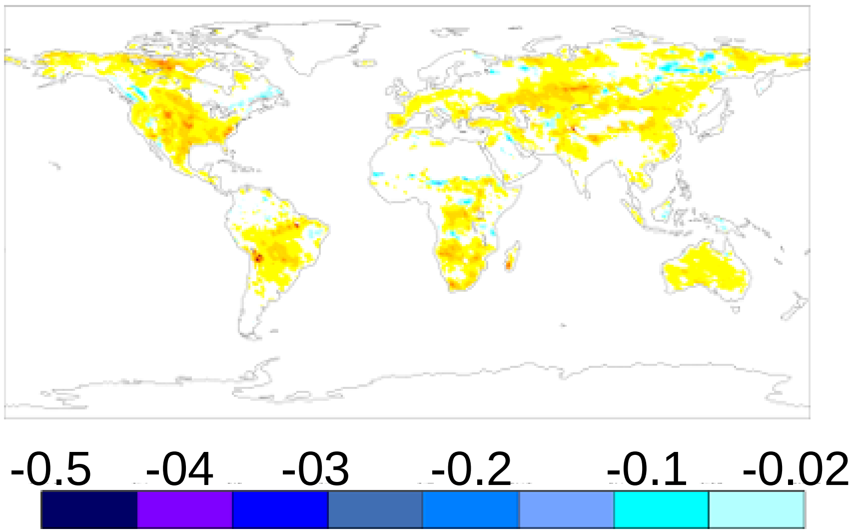
10 Layers; 8m depth for water +temperature

10 Layers; 2.89m depth for water + 8m temperature

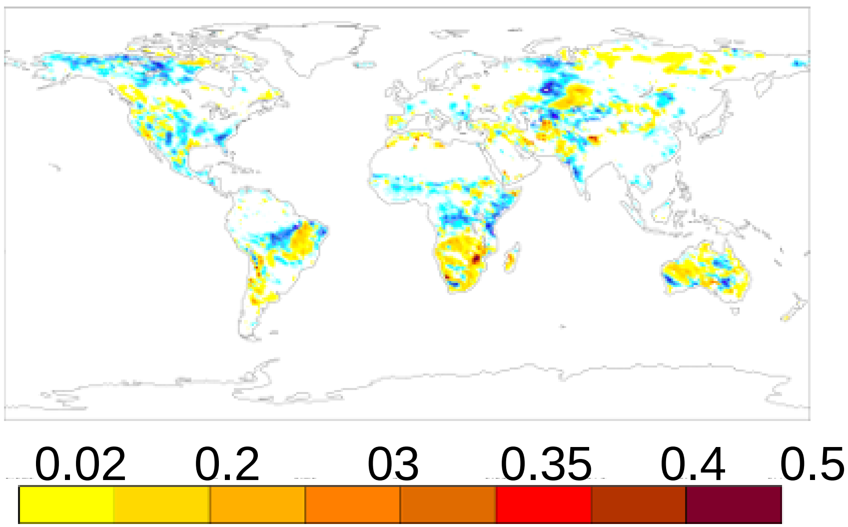
Impact on surface fluxes and L-A interaction

Difference in correlation with FluxCom between the 10-layers soil experiment and the control 4-layers experiment

a) Sensible Heat flux

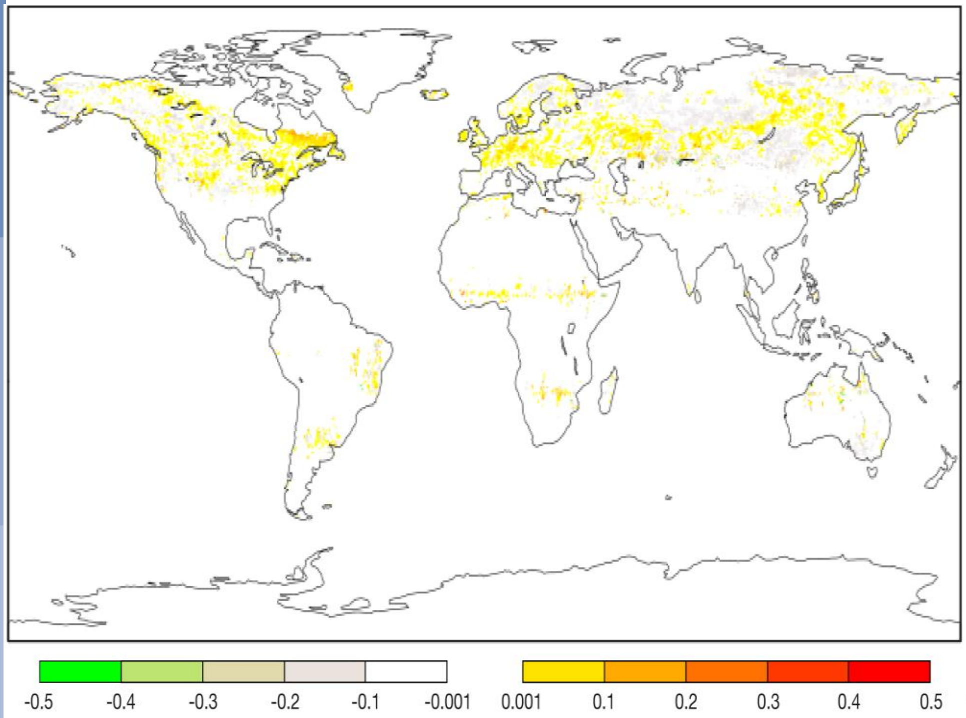


b) Latent Heat flux

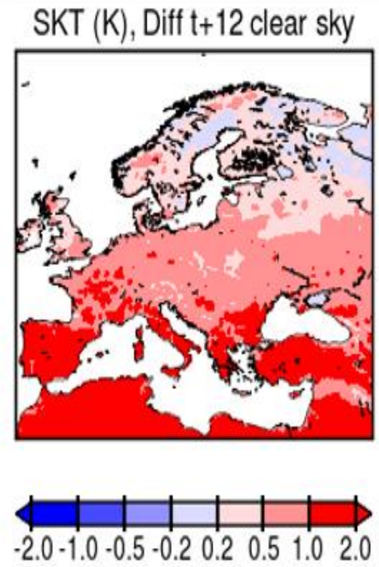


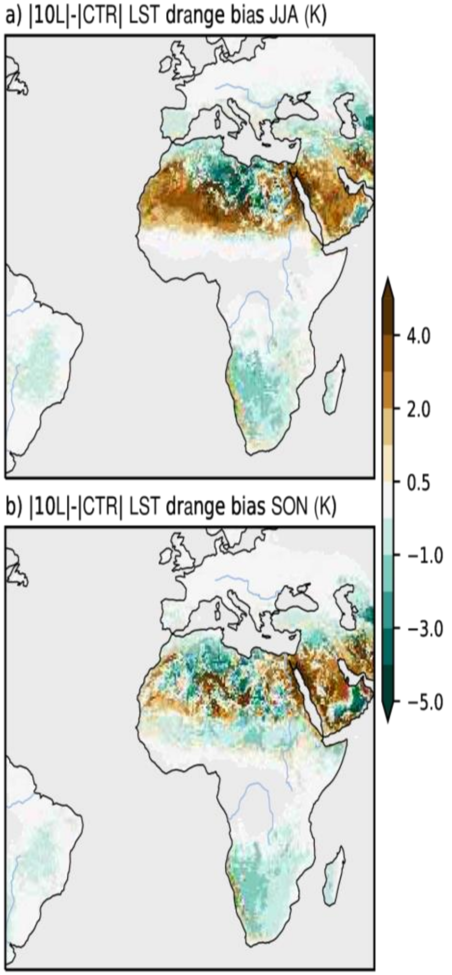
Better correlation with FluxCom sensible heat flux ==> could infer a better L-A interaction
Latent heat flux shows an overall decrease of correlation and a slight increase over some arid areas

Impact on Soil moisture and match with Satellite obs



Mean JJA difference of amplitude of diurnal cycle SKT 10L-control, JJA, clear sky





Layer1 SM correlate better with ESA-CCI SM (plot for JJA 2018: Corr_10L –Corr_CTL)

Increased SKT amplitude and better match with Land-SAF LST

Additional thoughts

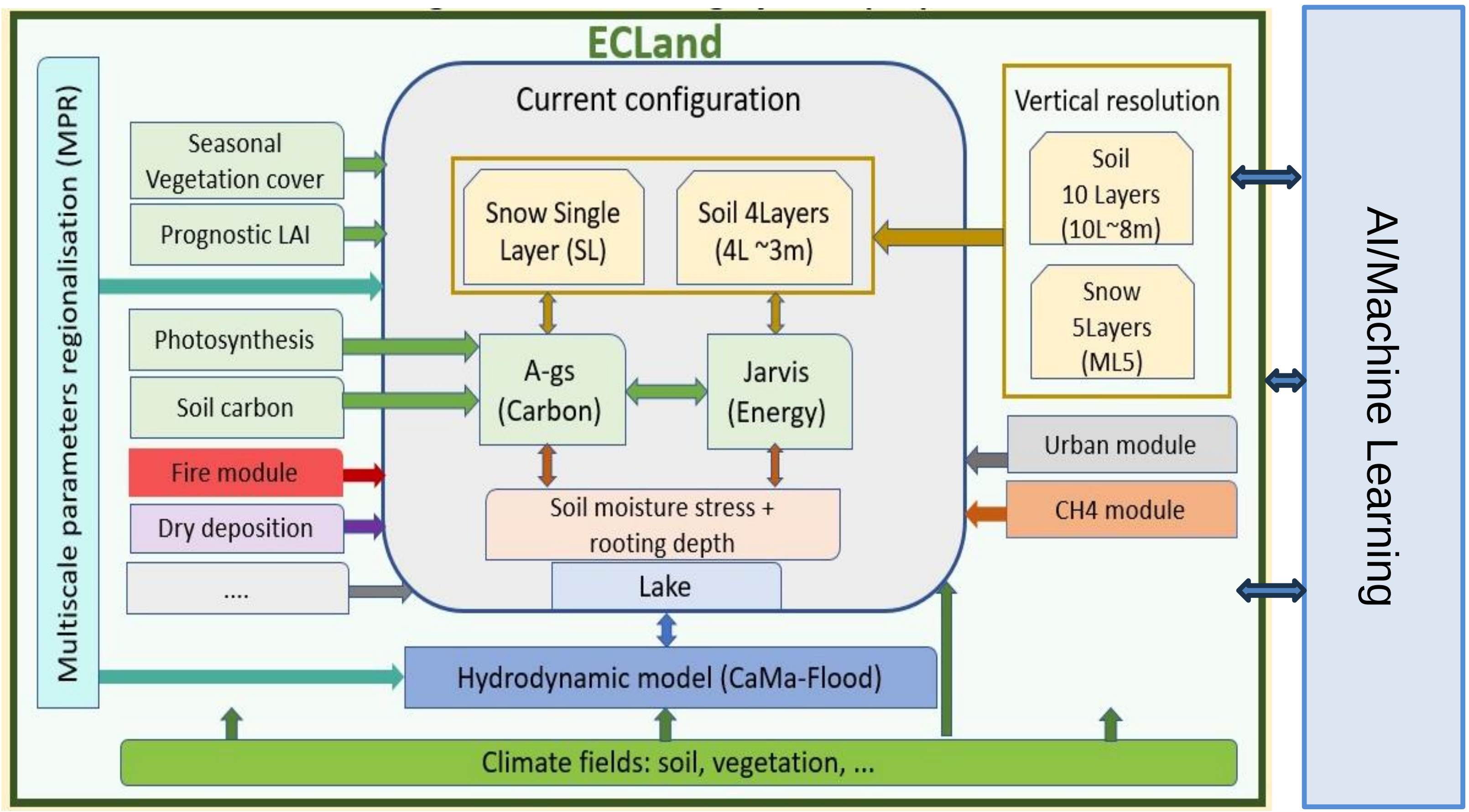
- Taking into account realistic vegetation dynamics is important for accurate representation of surface fluxes and eventually better atmospheric predictability.

- Carbon, Hydrology and Energy cycles are tightly coupled and an integrated treatment of these processes is a challenge to achieve the necessary accuracy in simulating Net Ecosystem Exchange (CO_2 flux) in global models (and as a component of the global carbon budget).

- Enhanced connections between albedo, LAI (and roughness) in Earth System Models (ESMs) will most likely increase the sensitivity to vegetation dynamics, and with increased surface related satellite observation products there is potential for further improvements of NWP systems linked with land surface. (better initialisation/ better process description/ possibility to better tune non-observable model parameters)

- With increased resolution ESMs will have to take into account additional layer of physical complexity such as
  - vegetation interaction with snow/frozen soil,
  - better vegetation dynamics,
  - better representation of the soil profile,
  - surface- atmosphere coupling and the link with satellite LST,
  - CO_2/evapo-transpiration coupled processes and satellite fluorescence observation.

ECLand platform within the IFS and perspectives



Evaluation & Benchmarking

Thank you

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