Satellite wind information on the ocean surface: Scatterometer & Altimeter
(& SMOS)

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with contributions from Saleh Abdalla, Hao Zuo, Magdalena Balmaseda, Tony McNally, Linus Magnusson, Patrick Laloyaux, Cristina Lupu, Nicolas Reul (Ifremer)

ECMWF

(andrej67/iStock/Getty Images)

ECMWF

Outline

Scatterometer Winds

- The importance of scatterometer wind observations
- Scatterometer principles
- Data usage at ECMWF and their impact
- How we can improve usage and impact

Altimeter Wind, Waves, Sea Surface Height

- Altimeter principles
- Use of altimeter data in the wave and ocean models
- Altimeter data impacts

SMOS winds

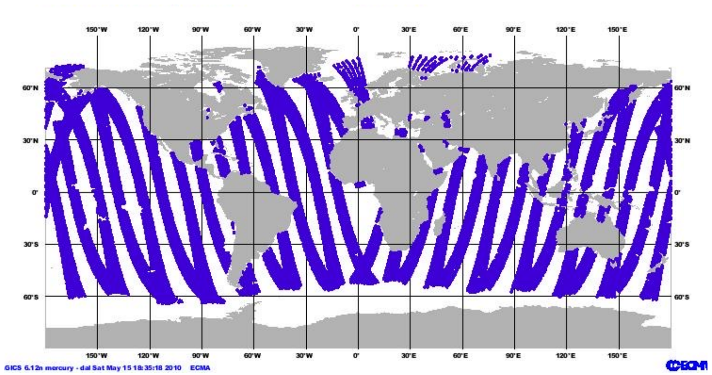
- Principles
- Database available
- TC examples

Why is Scatterometer important?

The scatterometer measures the ocean surface winds (ocean wind vector).

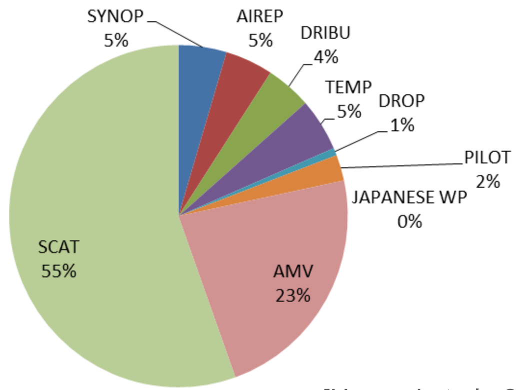
Ocean surface winds:

- affect the full range of ocean movement
- modulate air-sea exchanges of heat, momentum, gases, and particulates
- direct impact on human activities



Wide daily coverage of ocean surface winds
Ex: 1 day of ASCAT-A data

Wind observations below 850 hPa
FSO values relative quantities (in %)

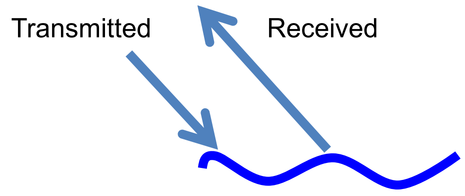


[Horanyi et al., 2013]

Scatterometer

A Scatterometer is an active microwave instrument (side-looking radar)

- Day and night acquisition
- Not affected by clouds



$$\sigma \sim \frac{\text{received power}}{\text{transmitted power}}$$

The return signal, *backscatter* (σ_0 *sigma-nought),* is sensitive to:

- Surface wind (ocean)
- Soil moisture (land)
- Ice age (ice)

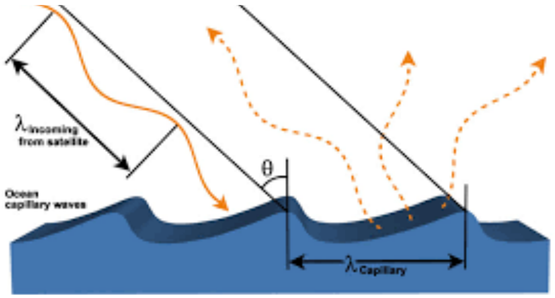
Scatterometer was originally designed to measure ocean wind vectors:

- Measurements sensitive to the ocean-surface roughness due to capillary gravity waves generated by local wind conditions (surface stress)

- Observations from different look angles: wind direction

Scatterometer

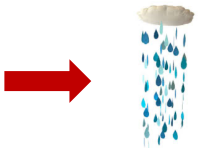
Bragg scattering occurs from the ocean capillary-gravity waves (cm-range) that are in resonance with the microwaves



The amount of backscatter depends on:

The frequency and polarization of the emitted wave

- C-band (5.3 GHz): λ ~ 5.7 cm
- Ku-band (13.5 GHz): λ ~ 2.1 cm



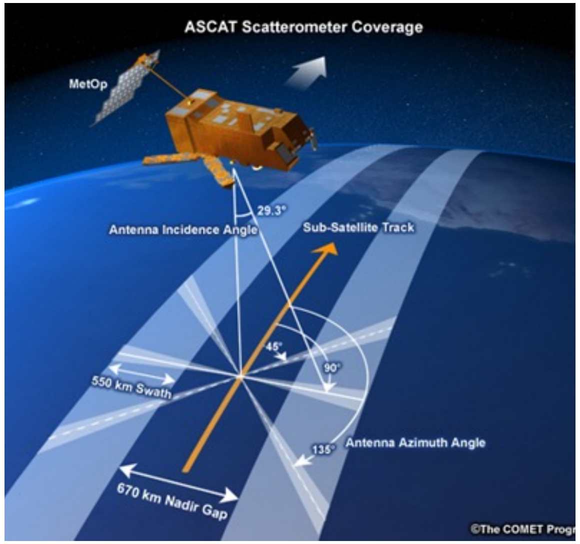
Backscatter *highly* depends on:

- Incidence angle (largest sensitivity to changes in winds between 30 and 60 deg)
- Wind speed
- Relative direction between the surface wind and look angle

C- band scatterometers (Fan beam)

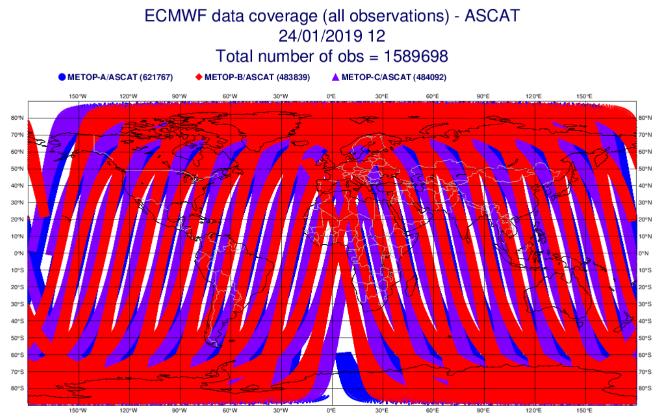
Used on European platforms (1991 onwards):

- ✓ SCAT on ERS-1, ERS-2 by ESA
- ✓ ASCAT on Metop-A, Metop-B, Metop-C by EUMETSAT
  - f~5.3 GHz (λ~5.7 cm)
  - Two sets of three antennas
  - σ_0 on a 12.5 km or 25 km grid



Pros and cons:

- ✓ Hardly affected by rain
- ✓ High quality wind direction (especially ASCAT)
- ✓ Two nearly opposite wind solutions
- ✓ Rather narrow swath:
  - ERS-1/2: 500 km
  - ASCAT-A/B/C: 2x550 km

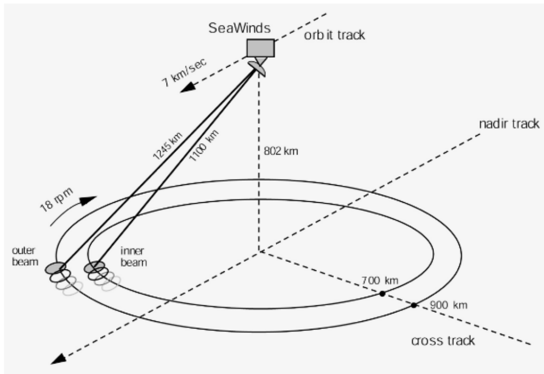


Ku-band scatterometers (Rotating pencil beam)

Used on US, Japanese, Indian and Chinese platforms:

- ✓ NSCAT, QuikSCAT, SeaWinds by NASA (and Japan)
- ✓ Oceansat, ScatSAT by ISRO
- ✓ Haiyang-2A/B/C/D by China
- ✓ RapidSCAT on the ISS

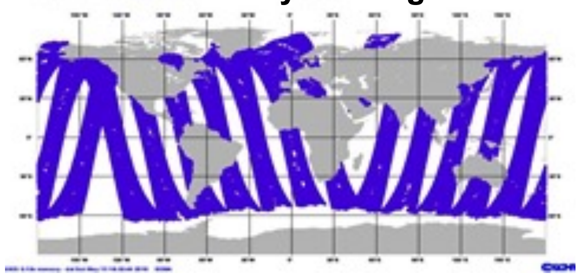
  - f~13.5 GHz (λ ~ 2.1 cm)
  - Two rotating pencil-beams (4 look angles)



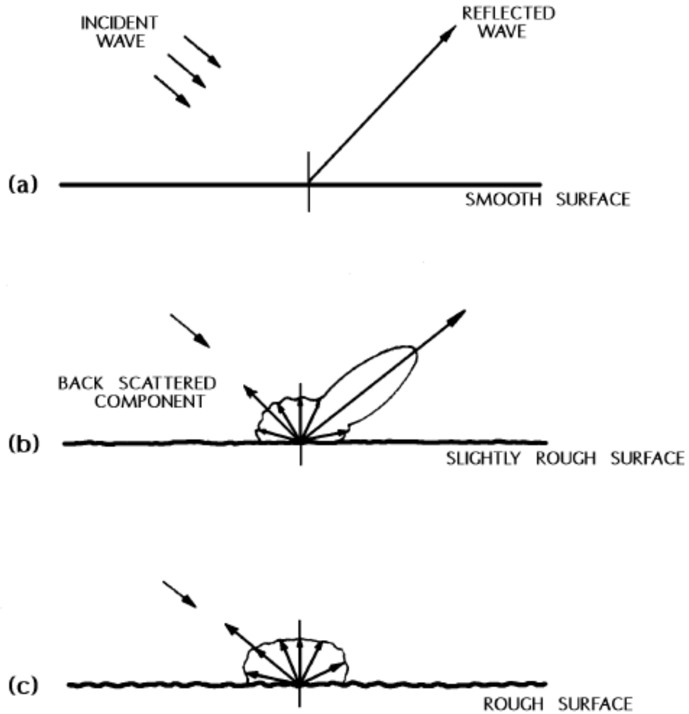
Pros and **cons:**

- ✓ Up to four wind solutions (rank-1 most often the correct one)
- ✓ Broad swath (1,800 km)
- ✓ Affected by rain
- ✓ Problems regarding wind direction:
  - azimuth diversity not good in centre of swath
  - outer 200 km only sensed by one beam.

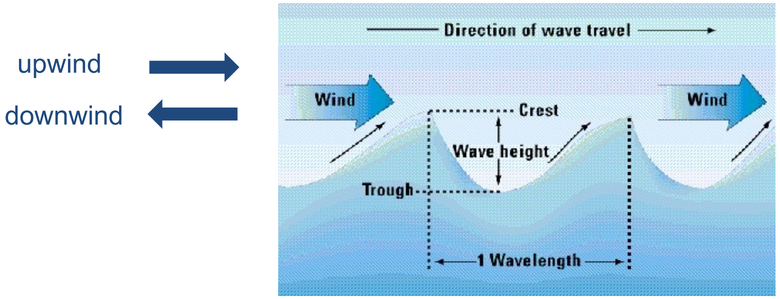
OSCAT daily coverage

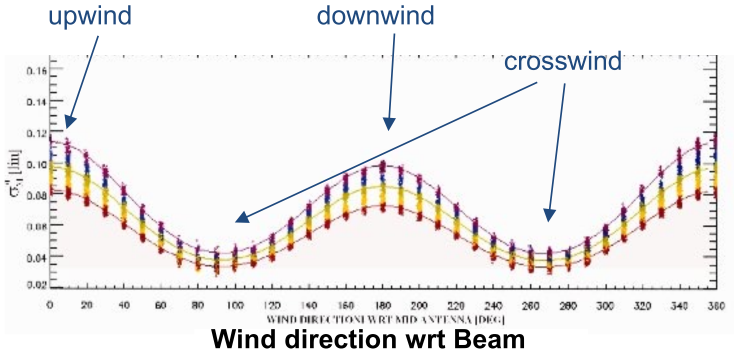


Dependency of the backscatter on... Wind speed



Dependency of the backscatter on... Wind direction



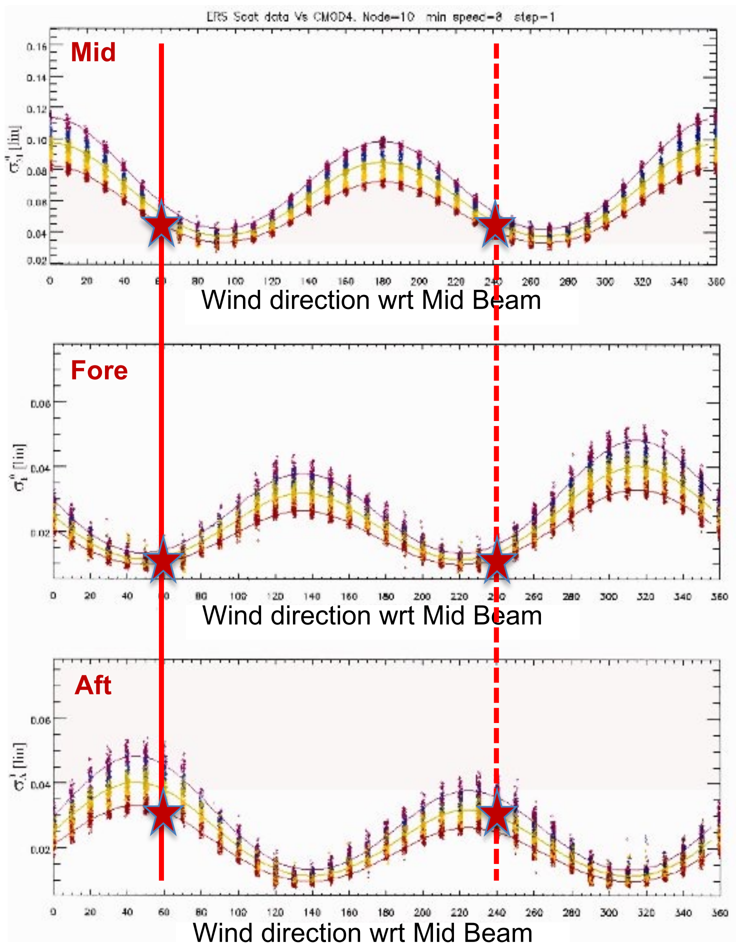


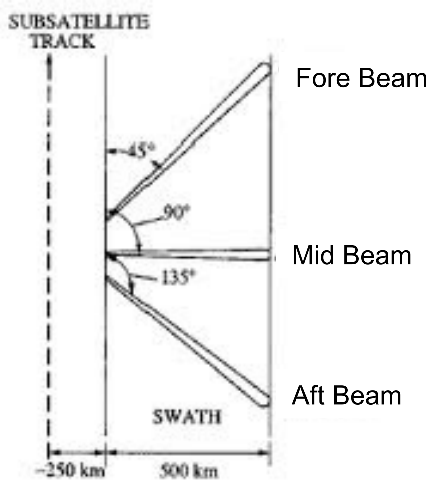
- Asymmetry Upwind – Downwind (particularly small)
- Asymmetry Upwind – Crosswind

Using multiple observations from different azimuth angles improves the accuracy of the derived wind direction

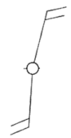
Dependency of the backscatter on... Wind direction

Backscatter response depends on the relative angle between the pulse and capillary wave direction (wind direction)



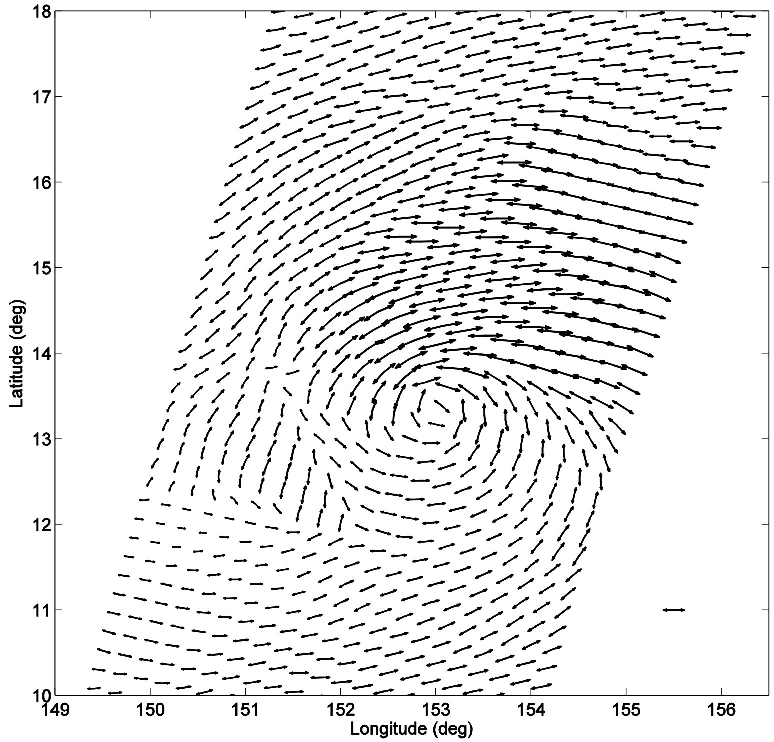


C-band SCAT geometry

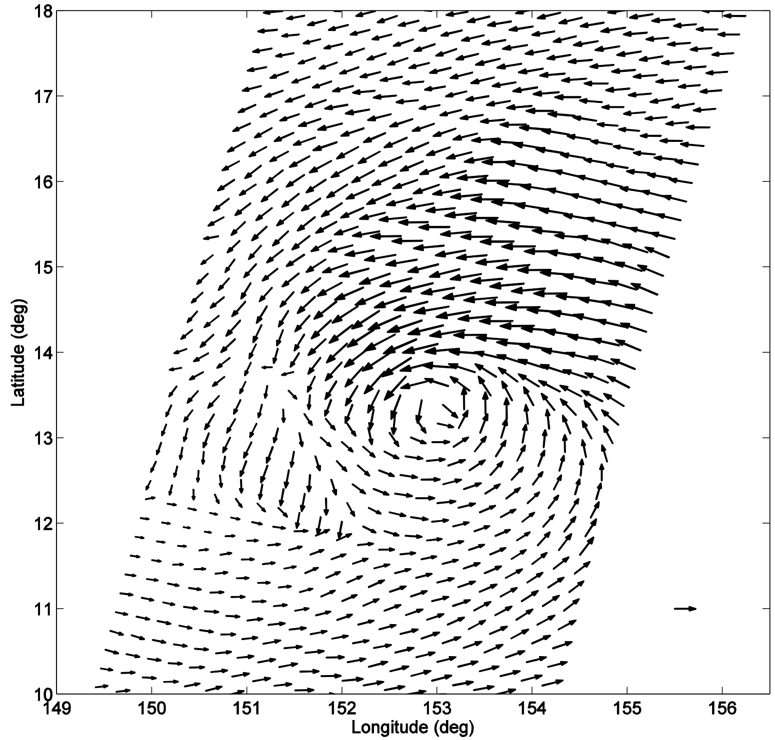


Wind Direction Ambiguity removal

- Measurements affected by noise
- Each wind vector cell has usually two possible solutions for wind direction and speed.
- The correct solution is determined by using NWP forecasts and wind field spatial patterns.



Ambiguities provided



Ambiguities selected

How can we relate backscatter to wind speed and direction?

Measurements sensitive to the **ocean-surface roughness** due to capillary gravity waves generated by local wind conditions (**surface stress**)

The relationship is determined empirically

- Ideally collocate with *surface stress* observations
- In practice with buoy and 10m model winds

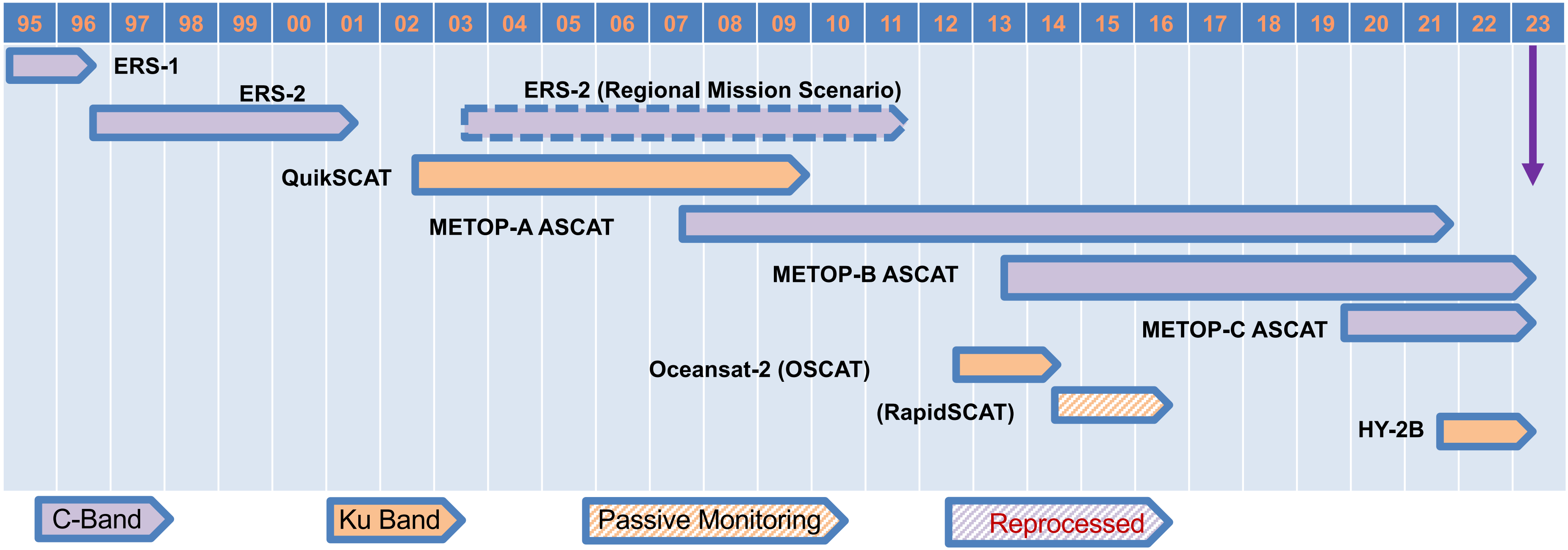
$$\sigma_0 = GMF(U_{10N}, \phi, \theta, p, \lambda)$$

U_{10N}: equivalent neutral wind speed
ϕ : wind direction w.r.t. beam pointing
θ : incidence angle
p : radar beam polarization
λ : microwave wavelength

Geophysical model functions (GMF) families

- C-band: CMOD (currently CMOD5.N in IFS)
- Ku-band: NSCAT, QSCAT

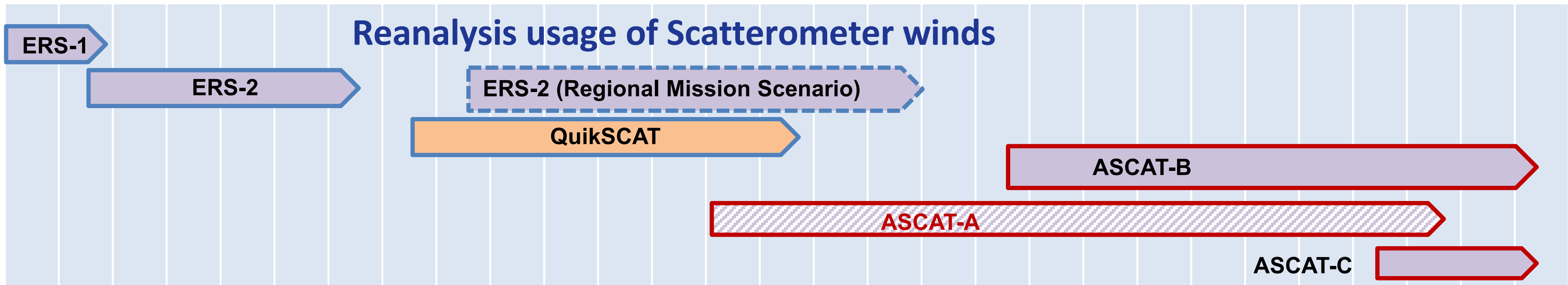
Operational usage of scatterometer winds at ECMWF



Under assessment: *HY-2C, HY-2D*

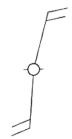
Waiting for data to evaluate: *Oceansat-3 (Nov '22), WindRad (July '21)*

Reanalysis usage of Scatterometer winds



Scatterometer assimilation strategy

	C-band	Ku-Band
Resolution	25 km	50 km
σ_0 bias correction	√	-
Wind Inversion	ECMWF	KNMI
Wind Speed bias correction	√	√
QC – Sea Ice check	√	√
Rain flag check	-	√
Thinning	100 km (*50 km in June 2023*)	-
Maximum wind speed assimilated	35 m/s	25 m/s
Assigned observation error	1.5 m/s (*2.25 m/s in June 2023*)	2 m/s
4D-Var	2 solutions	1 solution
Assimilated as 10m eq. neutral wind (U&V)	√	√

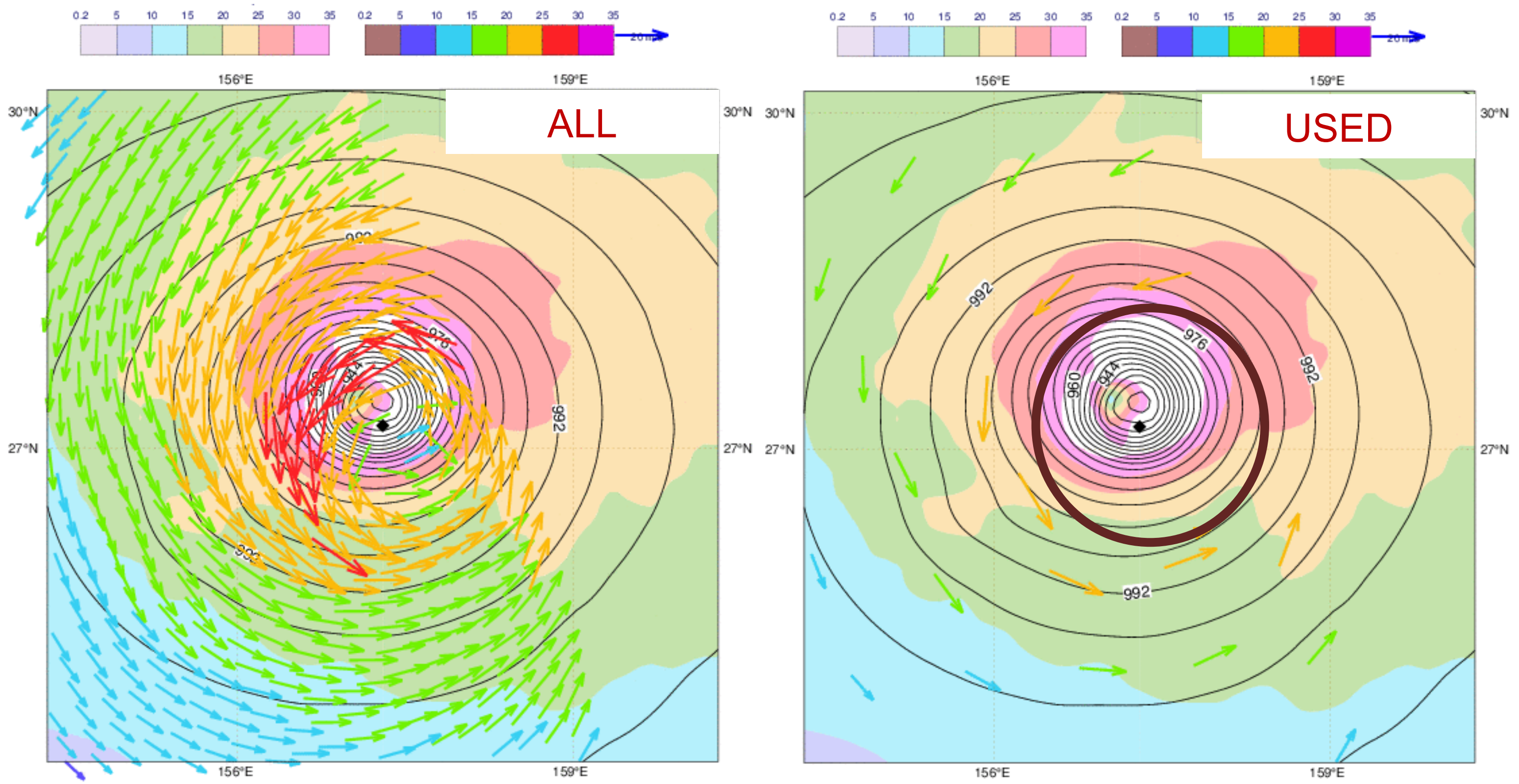


Why testing ASCAT reduced thinning?

ASCAT L2 products with 25km grid spacing are operationally assimilated with thinning=4 and OE=1.5 m/s

Fewer observations due to:
- Thinning
- VarQC

TC KILO – 2015090812 ASCAT Observations

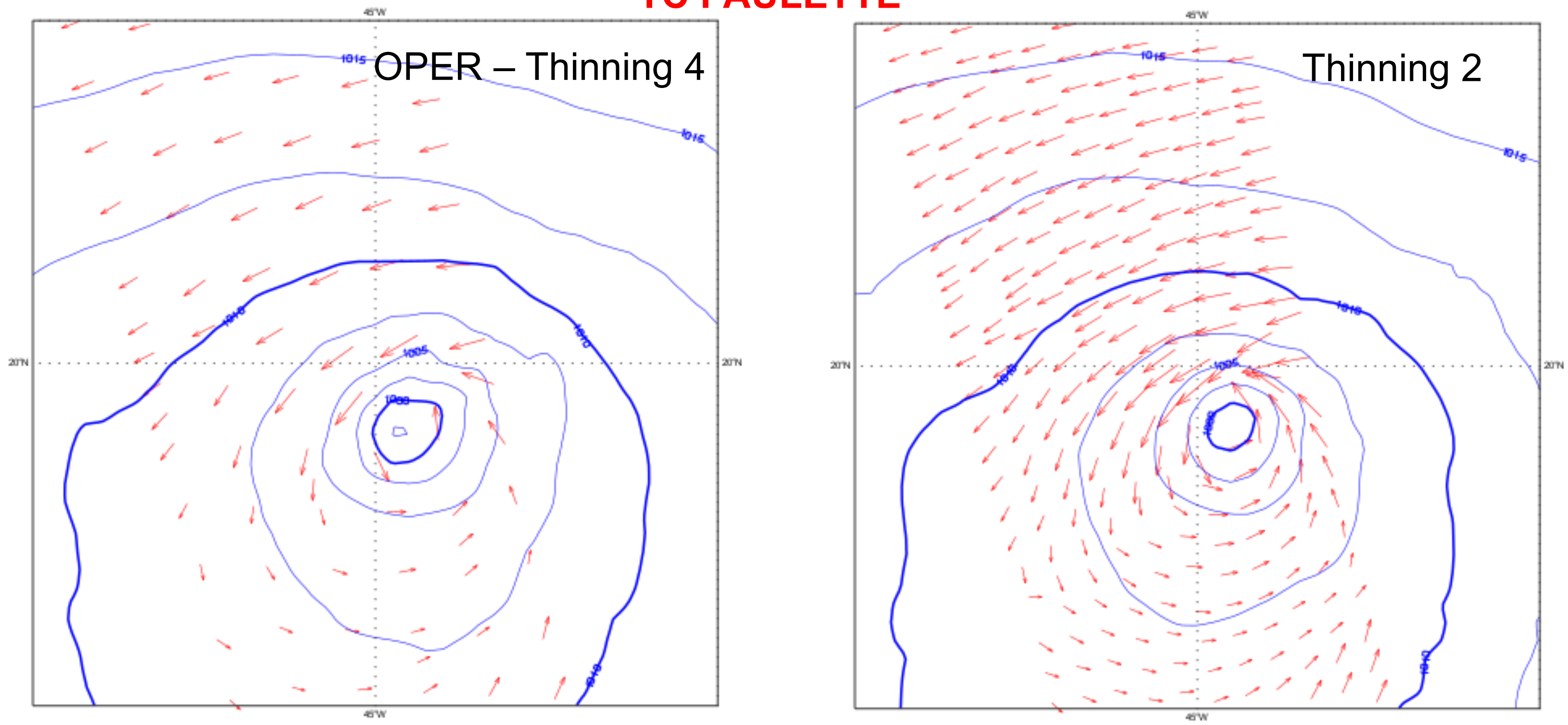


ASCAT reduced thinning

A reduced thinning would provide more information in more dynamic cases like TCs

When increasing the observations we need to inflate the OE (new OE=2.25 m/s)

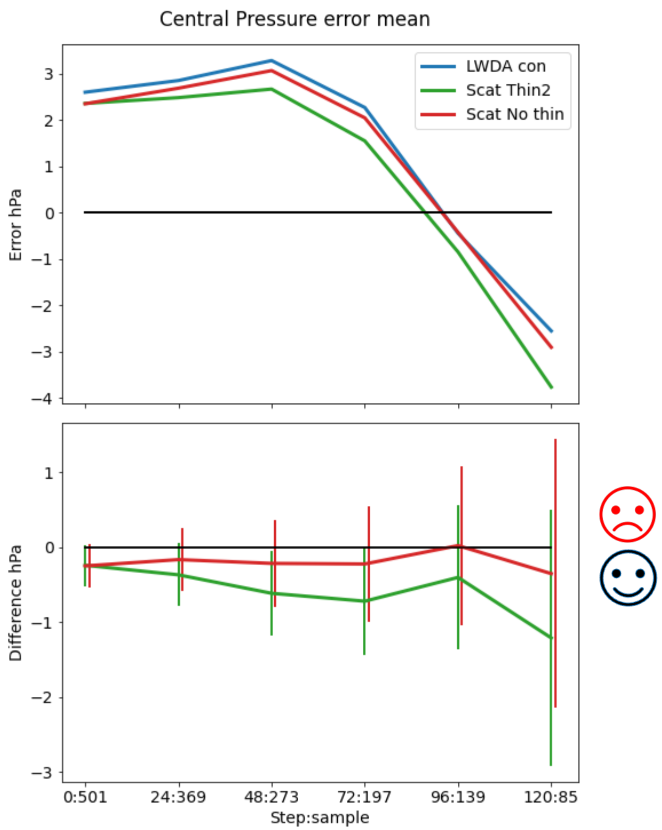
TC PAULETTE



Contour line = AN mslp

IMPACT ON TCs

3 months experiments (20190815 – 20191118) at Tco1279



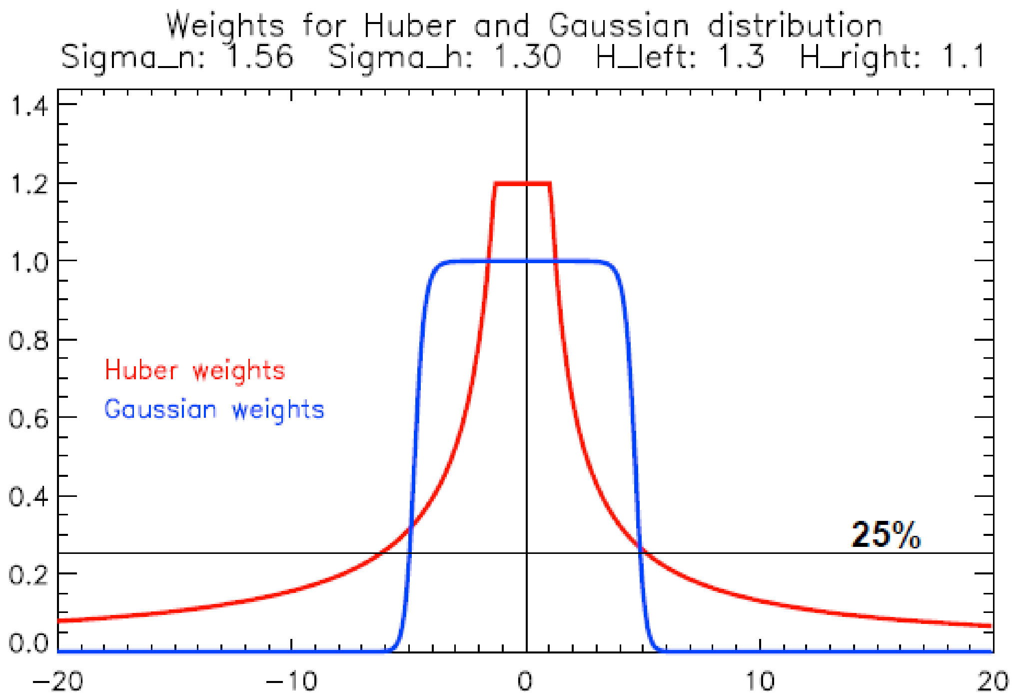
Global statistics over 3 months of:

- Slightly improved mean absolute error for position and intensity
- Reduced bias in intensity

VarQC

Observation weight: VarQC & Huber Norm

Comparing Observation weights:
Gaussian + flat (VarQC): more weight in the middle of the distribution
Huber Norm: more weight on the edges (to data with large departure)

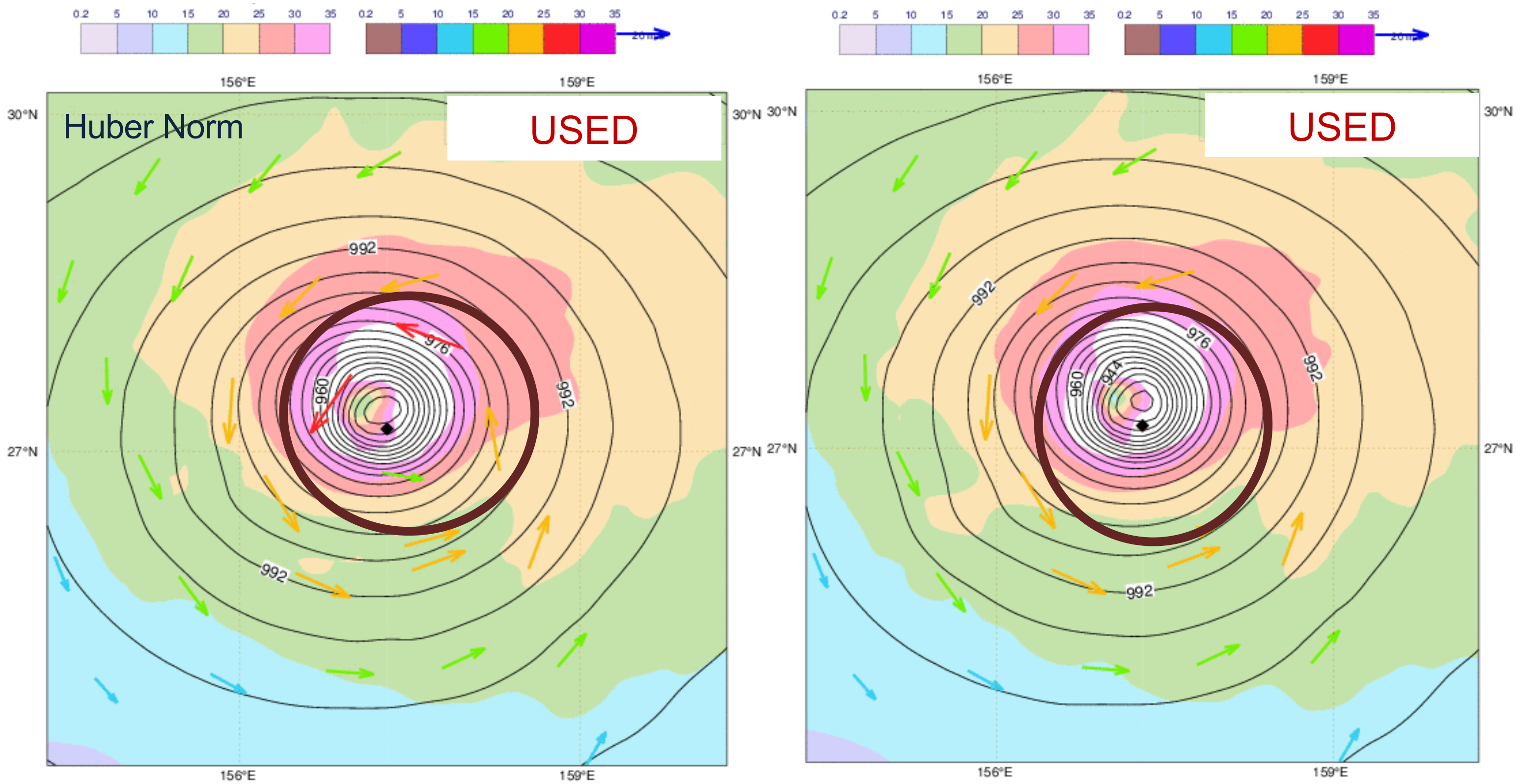


TC QC issues

TC KILO – 2015090812 ASCAT-A Observations

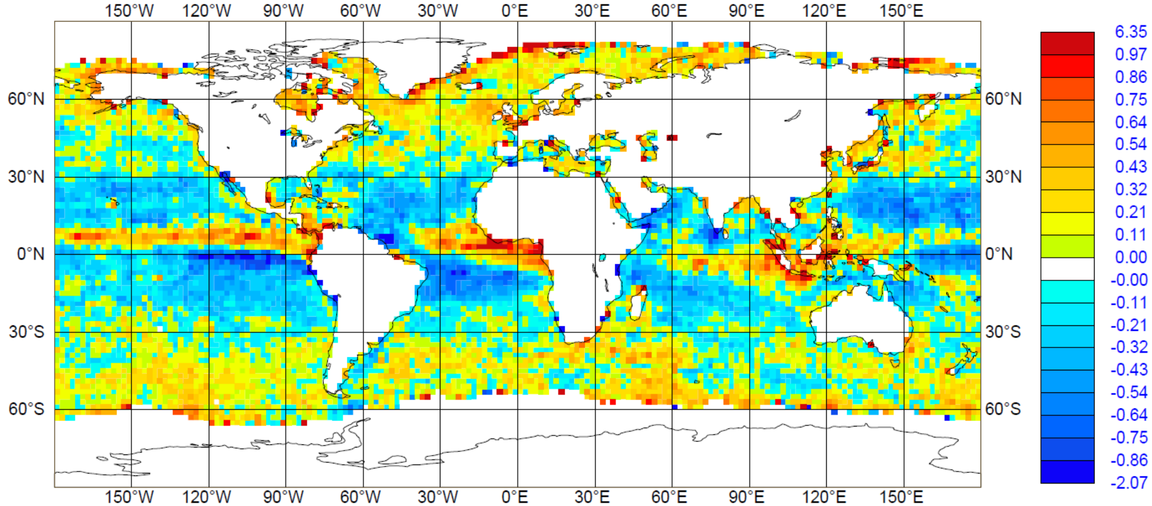
Less observations due to:

- Thinning
- VarQC



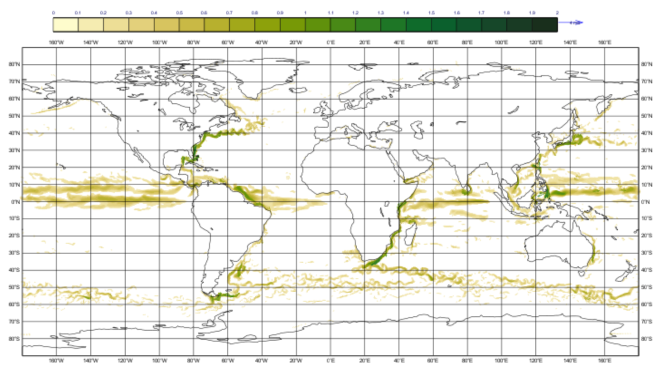
Background Departure

ASCAT-A Wind speed O-B

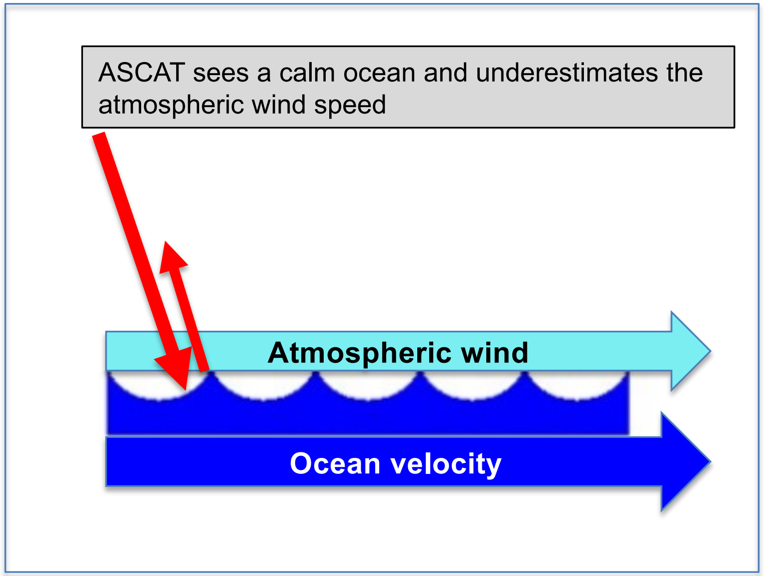


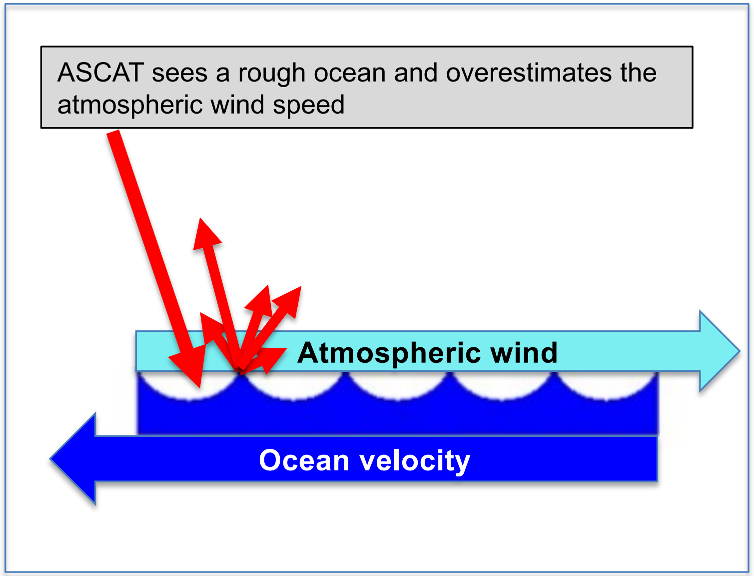
Wind speed bias in the Tropics: also due to Ocean Current?

Mean ocean currents from OCEAN5 from 20181101 to 20190129



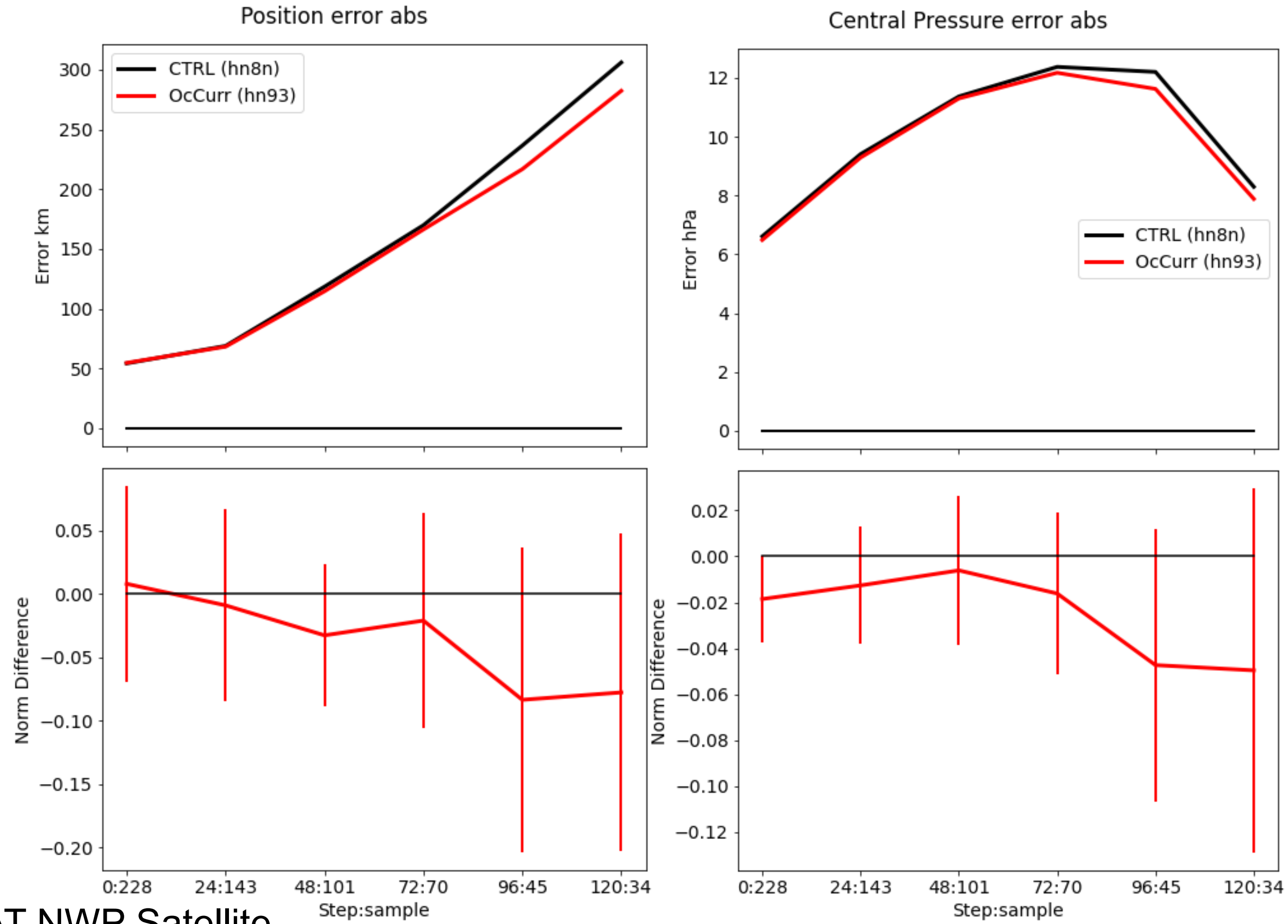
Scatterometer and Ocean Currents





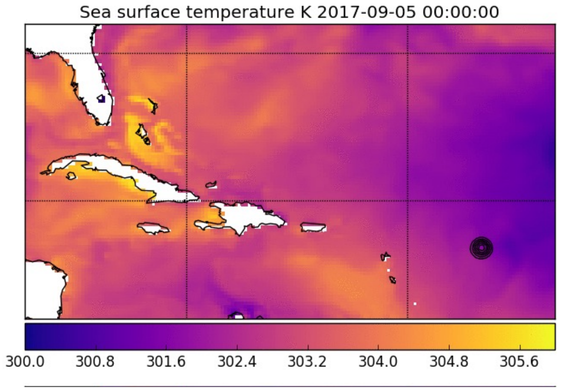
Scatterometer and Ocean Currents

Impact on TC - Preliminary results in weakly coupled DA



What about the impact of Scatterometer on the ocean?

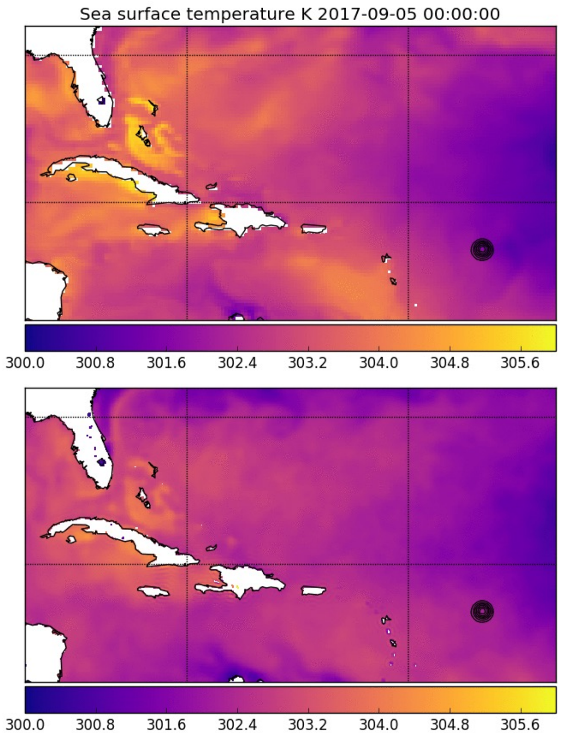
Coupled Data Assimilation (CDA)



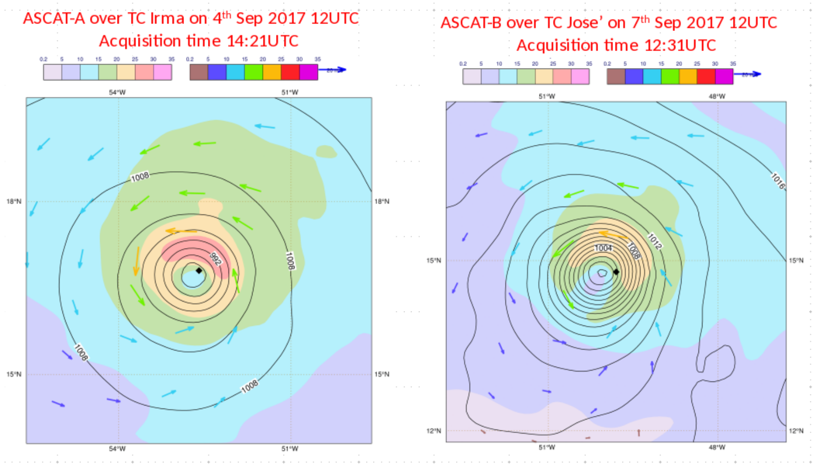
In the **coupled assimilation** the SST shows a clear and immediate impact on SST of the storm winds mixing the ocean (cold wake) and the storm's arrival in the Caribbean damping the usual pronounced diurnal cycle in the SST

Irma/Jose with ocean – atmosphere DA coupling

Coupled Data Assimilation (CDA)



What is the role of ASCAT (and JASON) in the coupled data assimilation during Irma and Jose?



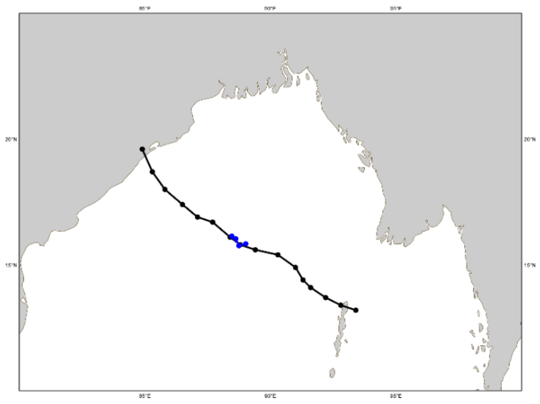
In CDA ASCAT gives SST information below Tropical Cyclones

Quantifying heat exchange between the storm and ocean surface is an important factor in predicting the intensification / de-intensification of Tropical Cyclones.

ASCAT sees through the cloud and rain (IR/MW cannot) and informs the coupled analysis of the surface roughening below the storm, in turn influencing the ocean mixing and thus the SST !

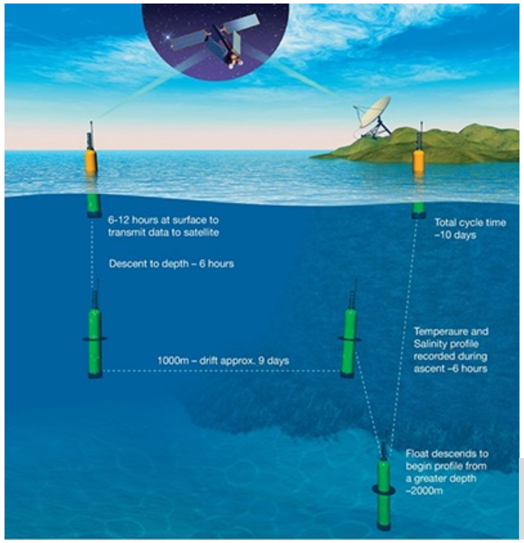
Impact of scatterometer winds …on the ocean parameters

Coupled Data Assimilation (CDA)

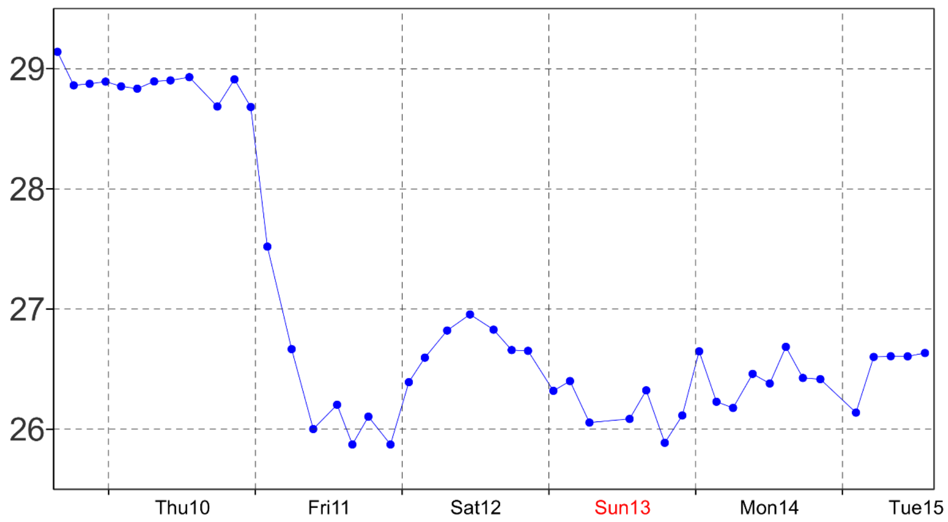


Focus on a specific weather event:

- TC Phailin
- Bay of Bengal
- formed on the 4th October 2013
- Argo probe with high-frequency measurements



Temperature measurements at 40-meter depth

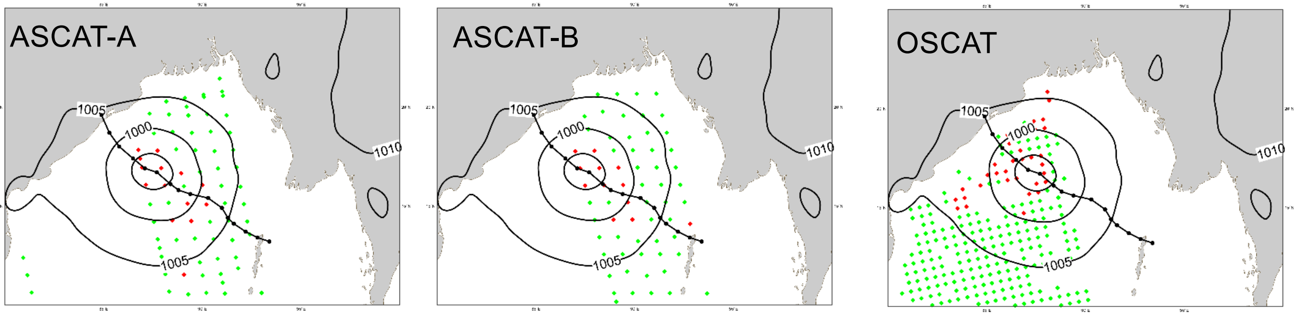


Impact of scatterometer surface wind data in the ECMWF coupled assimilation system
P. Laloyaux, J-N Thépaut and D. Dee. MWR, 2016

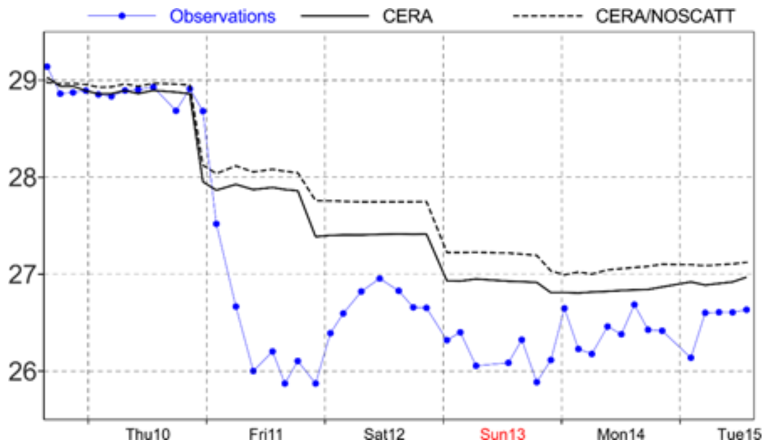
Impact of scatterometer winds …on the ocean parameters

TC Phailin

Wind measurements from scatterometers (ascending pass, 11 October 2013)

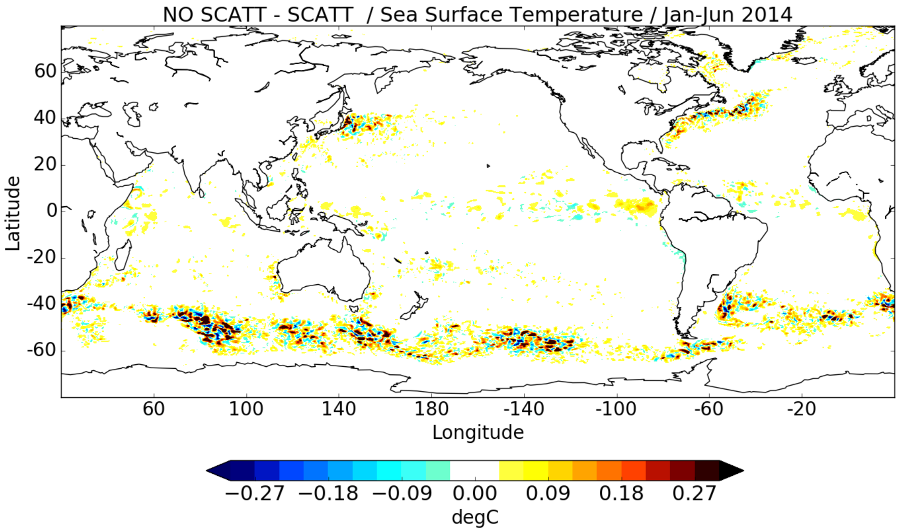


Ocean temperature analysis at 40-meter depth (scatterometer data are assimilated)

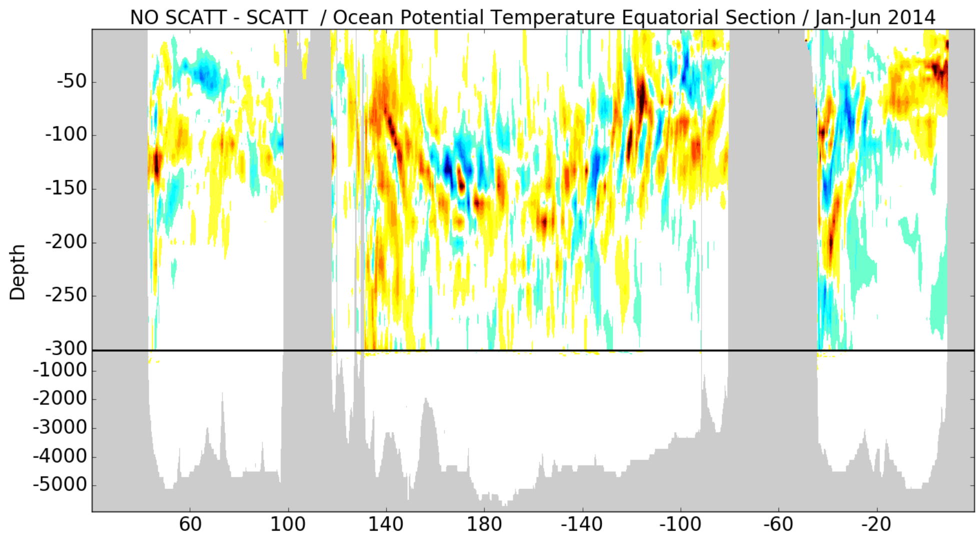


Coupled analysis with Scatterometer winds is closer to the observations with a stronger cold wake

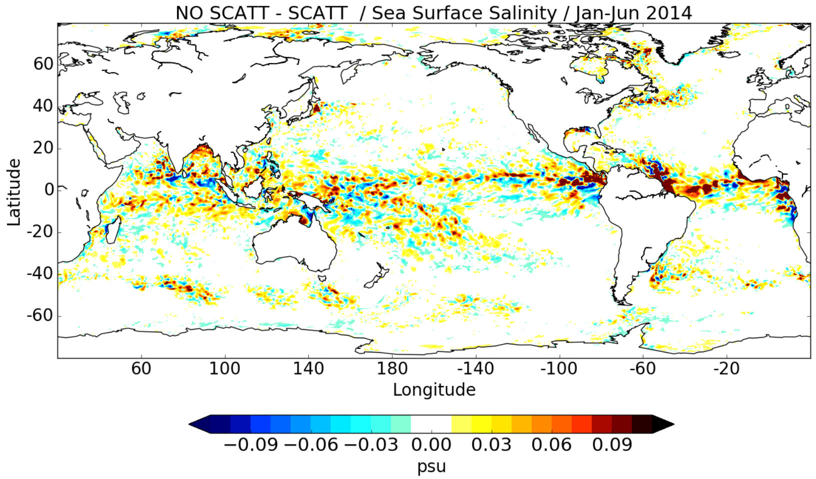
Impact of Scatterometer on Ocean Temperature



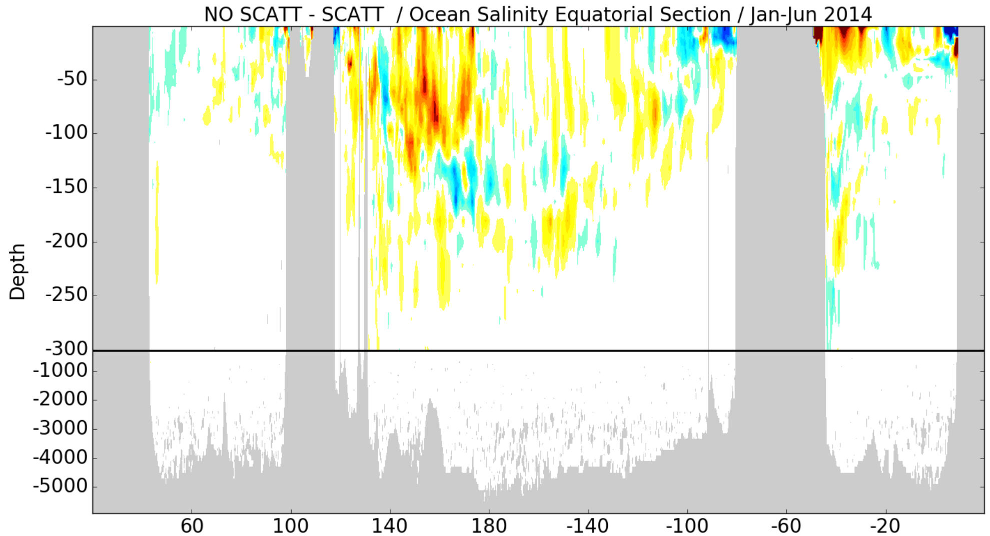
No SCATT – SCATT



Impact of Scatterometer on Ocean Salinity



No SCATT – SCATT



Scatterometer Concluding remarks

Scatterometer observations widely used in NWP

- Ocean wind vectors
- Positive impact on analysis and the forecast
- Global scale and extreme events
- Impact on Atmospheric, Ocean and Wave model

ECMWF has a long experience with scatterometry

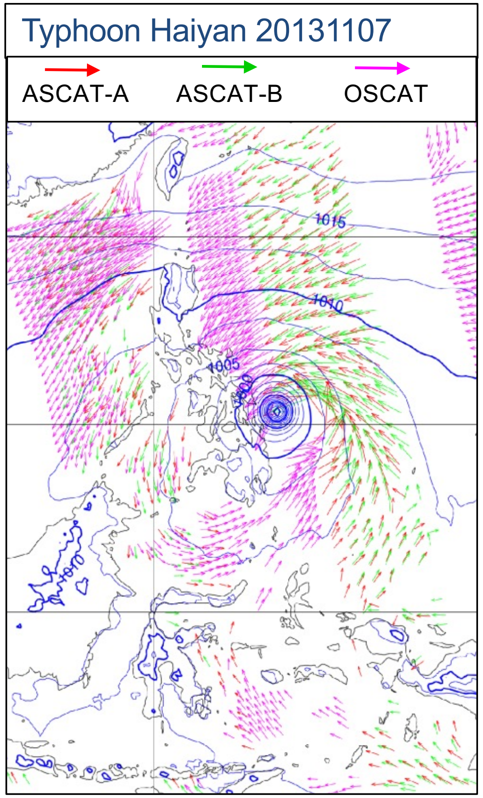
- Available continuously from 1991 onwards:
- GMF development
- Monitoring, validation, assimilation, re-calibration

On-going efforts to improve usage and impact

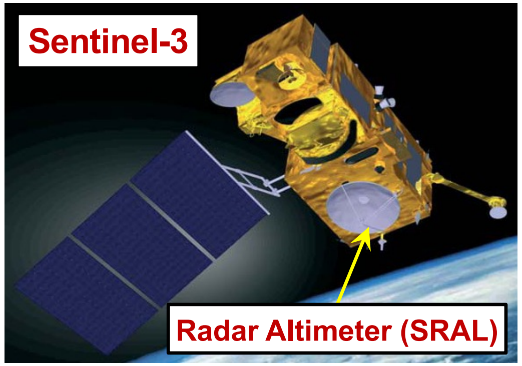
- Improve QC
- Adapt observation errors
- Include dependency from other geophysical quantities (i.e. Ocean Currents)
- Currently testing the assimilation of the backscatter rather than the wind

Use in the Reanalysis

- ERS1/2 and QuikSCAT in ERA-Interim
- ASCAT-A, ASCAT-B, ASCAT-C and HY-2B products used in ERA5

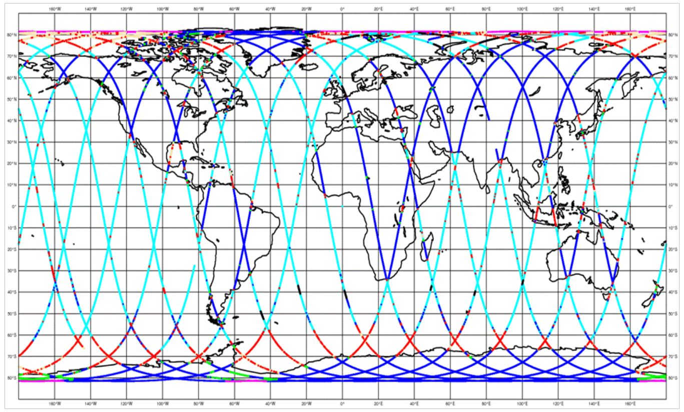


Radar Altimeters

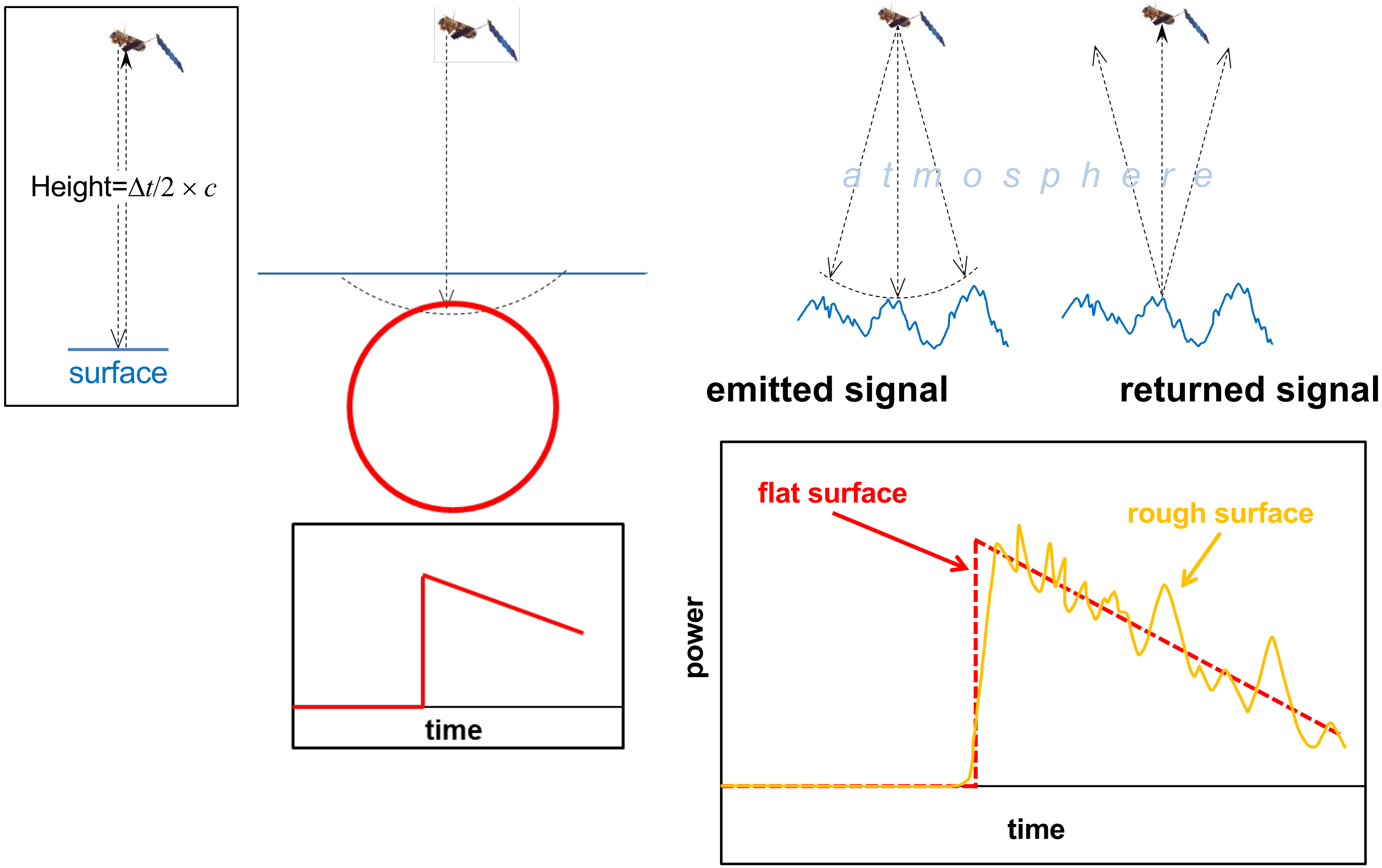


- ✓ Radar altimeter is a nadir looking instrument.
- ✓ Specular reflection.
- ✓ Electromagnetic wave bands used in altimeters:
  - Primary:
    - Ku-band (~ 2.5 cm) – ERS-1/2, Envisat, Jason-1/2/3, Sentinel-3/6
    - Ka-band (~ 0.8 cm) – SARAL/AltiKa (only example)
  - Secondary:
    - C-band (~ 5.5 cm) – Jason-1/2/3, Topex, Sentinel-3/6
    - S-band (~ 9.0 cm) – Envisat
- ✓ Main parameters measured by an altimeter:
  - Significant wave height (*wave model*)
  - Wind speed (*used for verification*)
  - Sea surface height (*ocean model*)

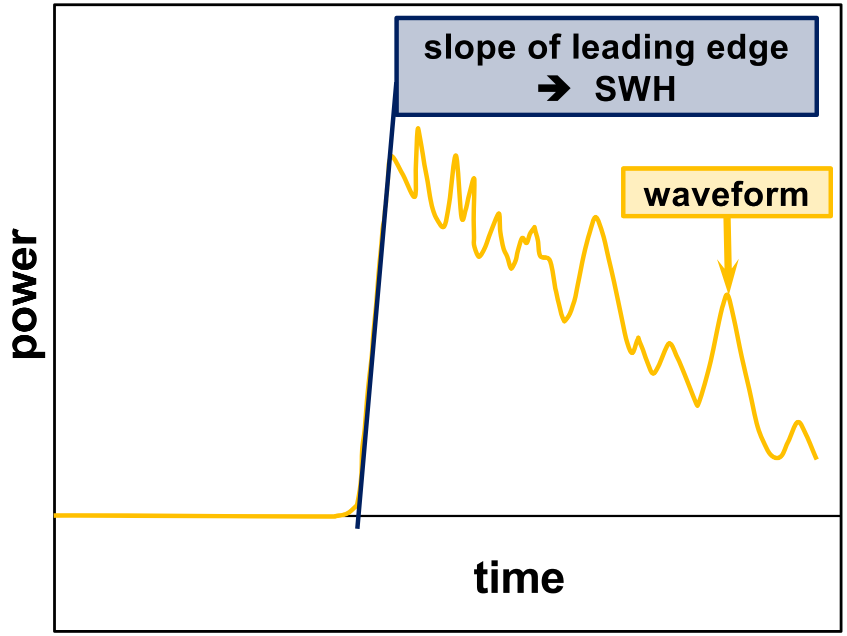
Envisat 1-day coverage



How Altimeter Works

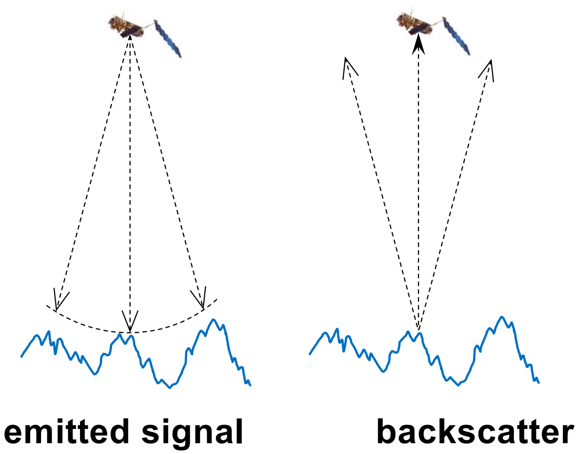


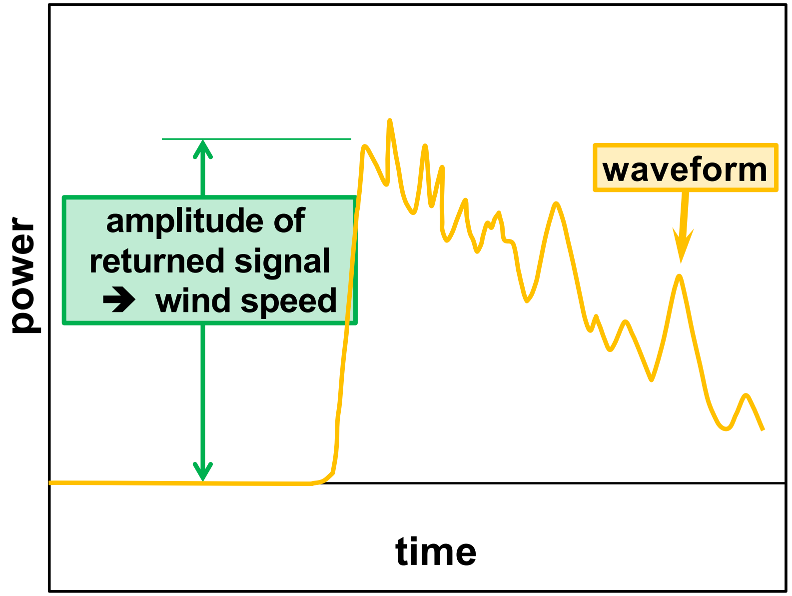
Significant Wave Height (SWH)



- ✓ SWH is the mean height of highest 1/3 of the surface ocean waves
- ✓ Higher SWH → smaller slope of waveform leading edge
- ✓ Errors are mainly due to waveform retracking (algorithm) and instrument characterisation

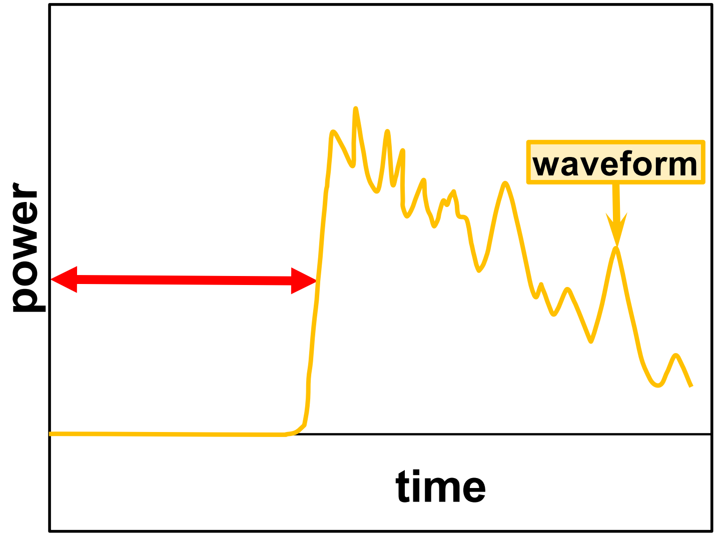
Surface wind speed





- ✓ Backscatter is related to water surface Mean Square Slope (MSS)
- ✓ MSS can be related to wind speed
- ✓ Stronger wind → higher MSS → smaller backscatter
- ✓ Errors are mainly due to algorithm assumptions, waveform retracking (algorithm), unaccounted-for attenuation & backscatter.

Sea Surface Height



- ✓ Time delay → sea surface height
- ✓ Radar signal attenuation due to the atmosphere is caused by:
  - Water vapour impact: ~ 10's cm.
  - Dry air impact: ~ 2.0 m

  Correction made using radiometer and model data

Operational Assimilation of SWH (wave model)

Sentinel-3A/B (Q2, 2021), CFOSAT-SWIM (2022), Sentinel-6 (2022).

Assimilation method for SWH data:

- ✓ Data are subjected to a quality control process (inc. super-obbing).
- ✓ Bias correction is applied.
- ✓ Simple optimum interpolation (OI) scheme on SWH.
- ✓ The SWH analysis increments ➔ wave spectrum adjustments…

Instrument sinergies

Altimeter SWH data available from five satellites – nice synergy!
Plot shows random error reduction of SWH compared to model only.

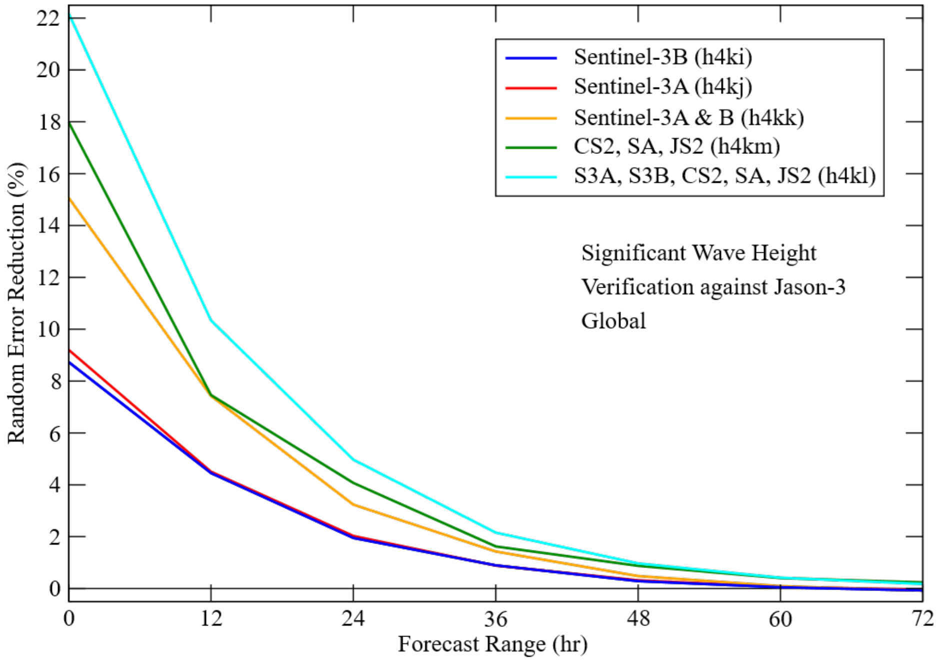
All the five altimeter instruments listed below

Cryosat-2 (CS2)+
SARAL AltiKa (SA)+
Jason-2 (J2)

Sentinel-3A&B

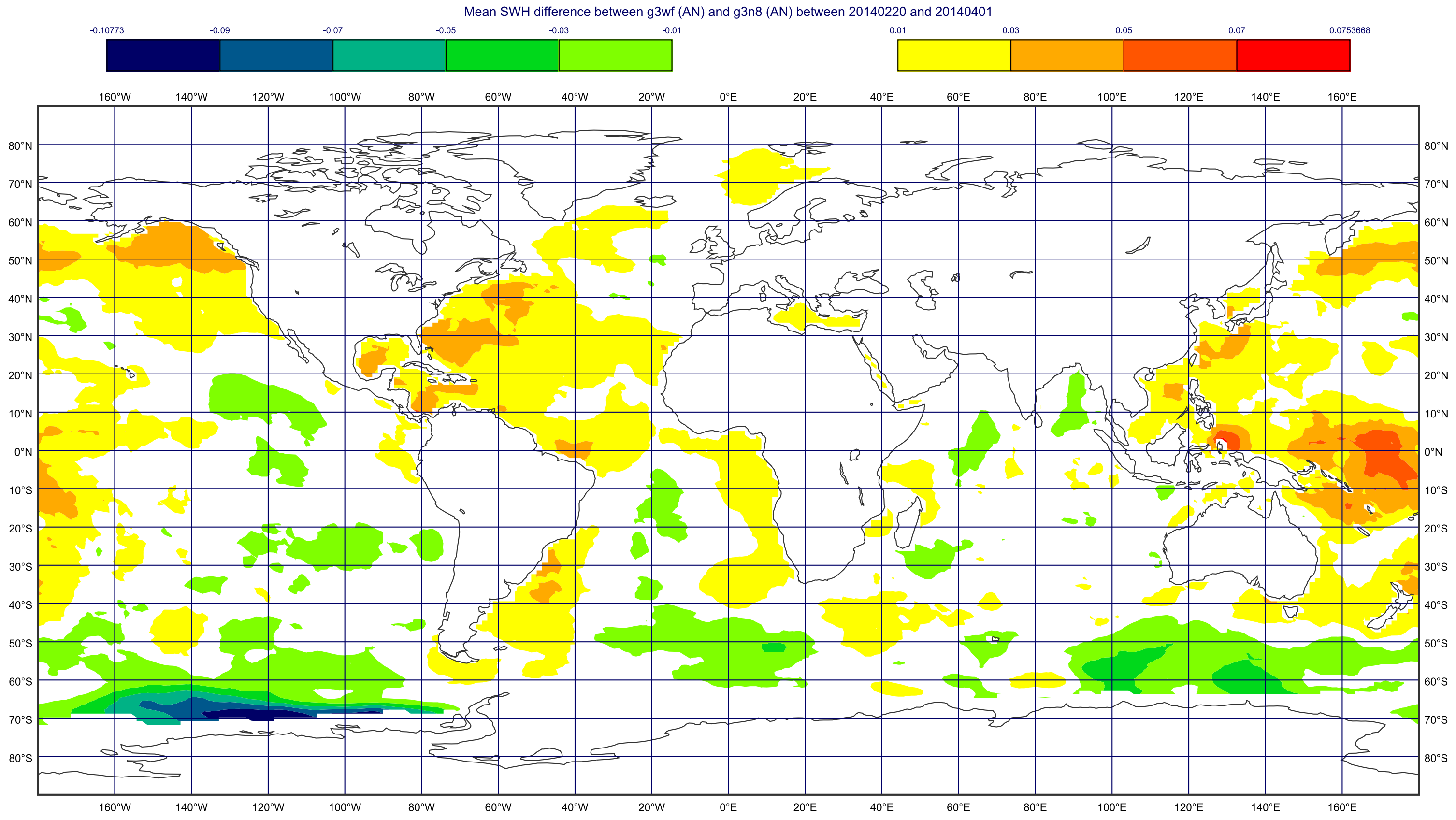
Sentinel-3A

Sentinel-3B

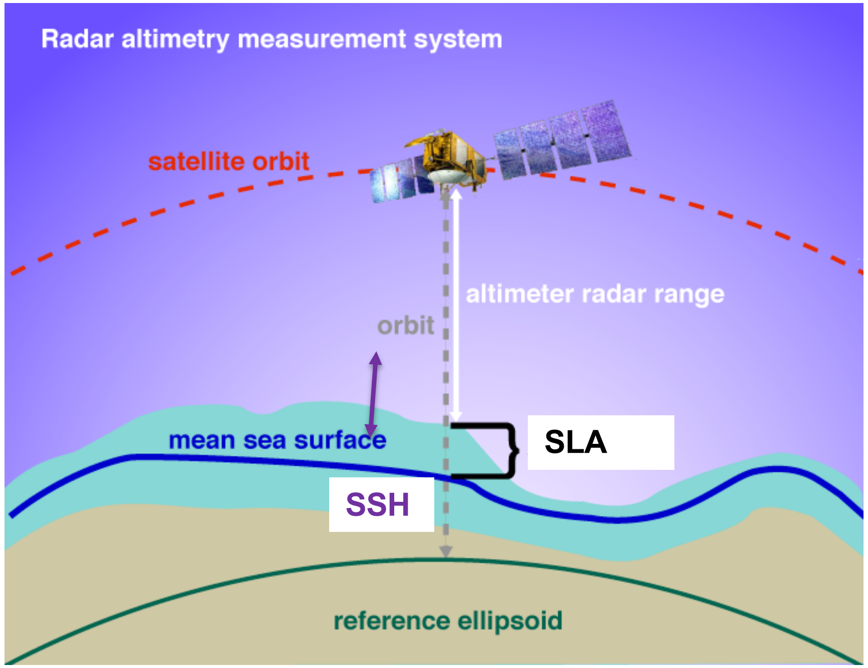


Impact of SWH assimilation

Impact of one additional altimeter on the SWH analysis [CS & J2] - [J2 only]



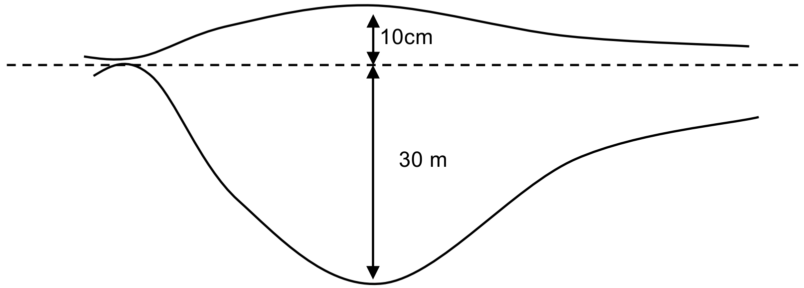
Altimeter data in the Ocean Analysis System



The altimeter measures the range which can be used to determine Sea Surface Height (SSH)

SLA = Sea Level Anomaly

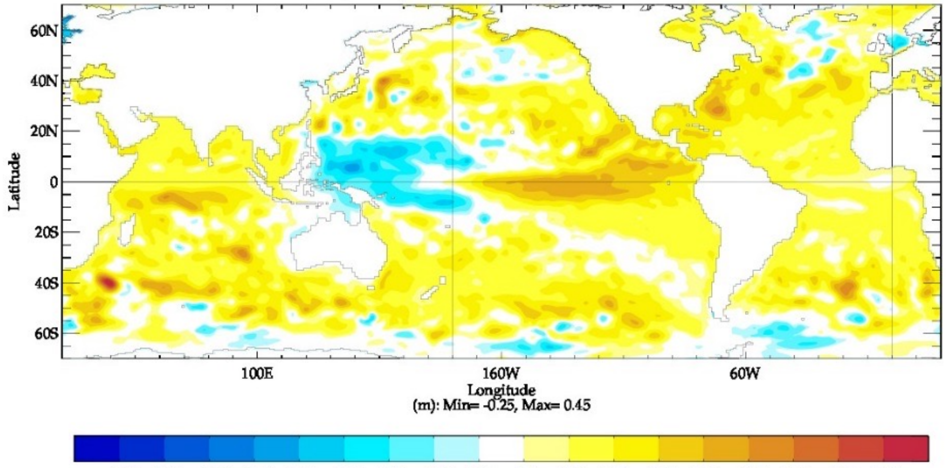
Assimilated in the ocean model



From sea level observation it is possible to infer information on the vertical density structure

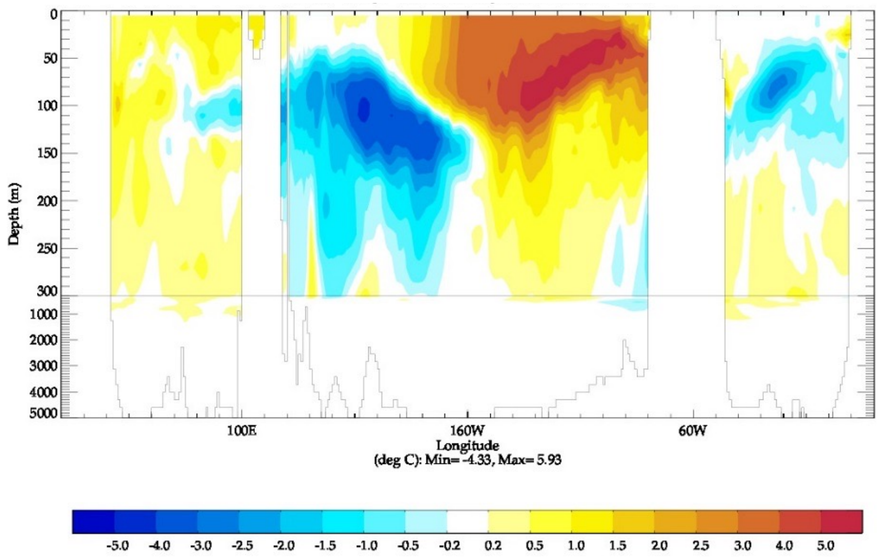
Why do we need SLA?

SL Anomalies Feb 2016

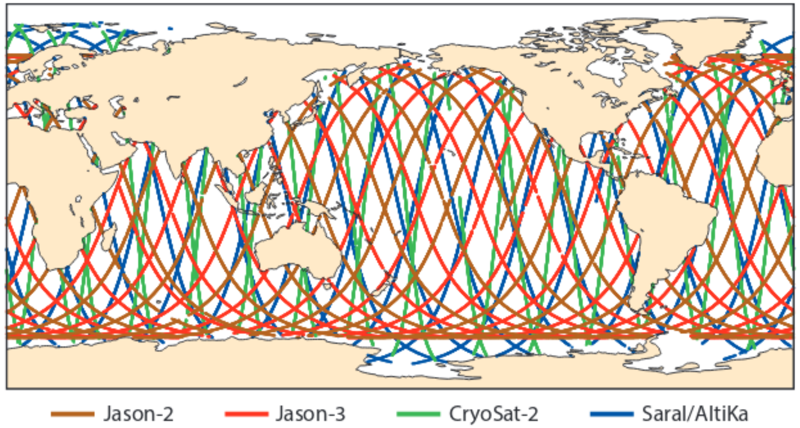


El Nino 2015/2016

Lon/Depth Temp anom Feb 2016



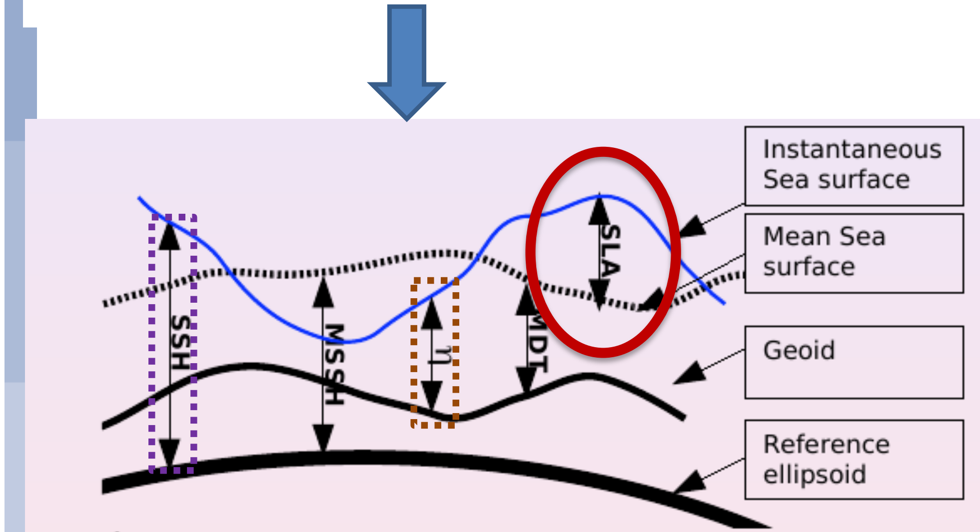
From Sea Surface Height to Sea Level Anomaly



Altimeter measures SSH (respect reference ellipsoide)
Model represents η (SSH referred to the Geoid)

SSH-Geoid= η

Geoid was poorly known (until recent years) and changes in time



Alternative: Assimilate Sea Level Anomalies (SLA) with respect to a time period

Obs: SSH anomalies = SSH-MSSH = Obs SLA
Mod: η anomalies = η – MDT = Mod SLA

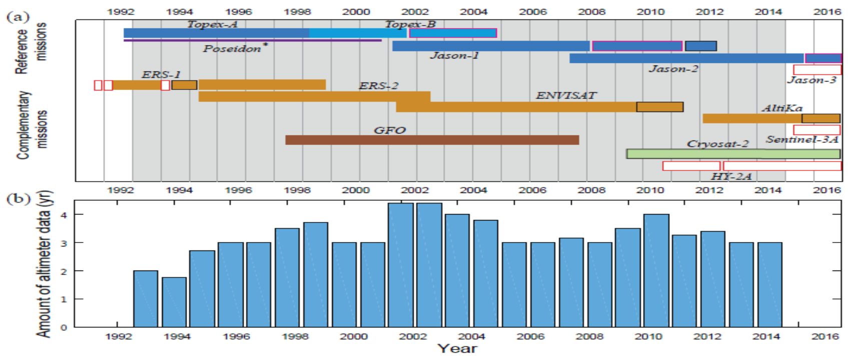
Where: MSSH= Temporal Mean SSH ;
MDT = Temporal Mean of model SL (Mean Dynamic Topography)
MSSH – Geoid = MDT

Assimilation of SLA as implemented in OCEAN5

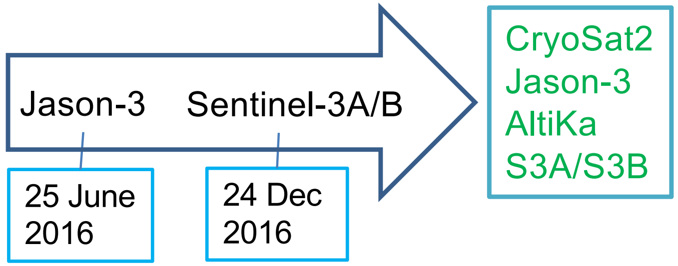
- Assimilate L3 along-track SLA data to constrain regional sea-level changes
- Assimilate L4 gridded MSLA maps to constrain global mean sea-level changes (via freshwater balance)
- Assimilate SLA with a model MDT (Mean Dynamic Topography) approach (require pre-computation)
- Apply cut-off to remove all high-latitude SLA data, as well all near coast observations.

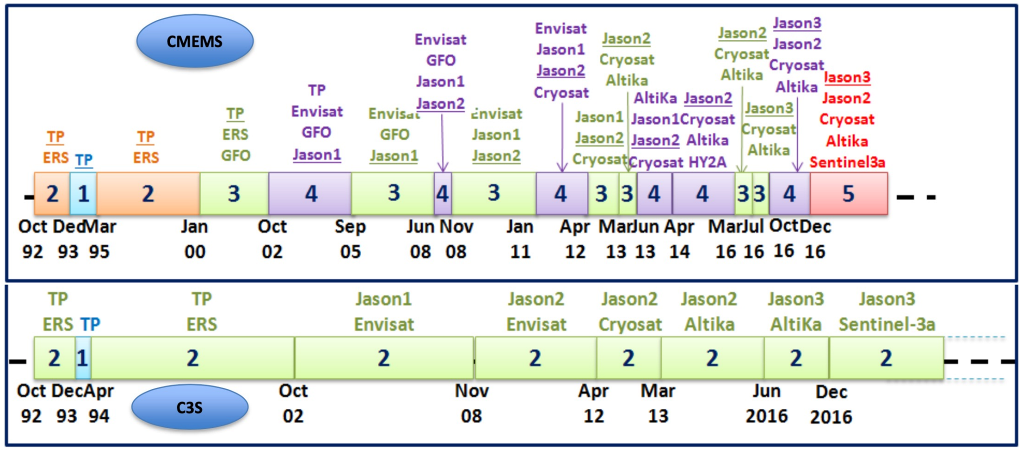
The effectiveness of SLA DA is mostly determined by elements listed above, but is also affected by bias correction settings; balance operator, BKG errors settings for other variables; consistency of the available SLA data; land freshwater input at the river mouths; treatment of freshwater constrain; etc

Available L3 along-track SLA data



Number of available SLA missions



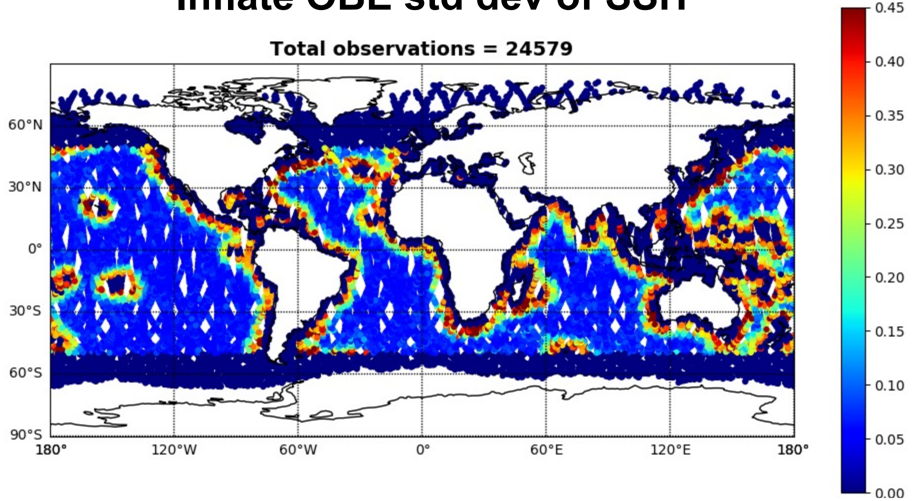


JF Legeais, et al., DUACS DT2018, poster

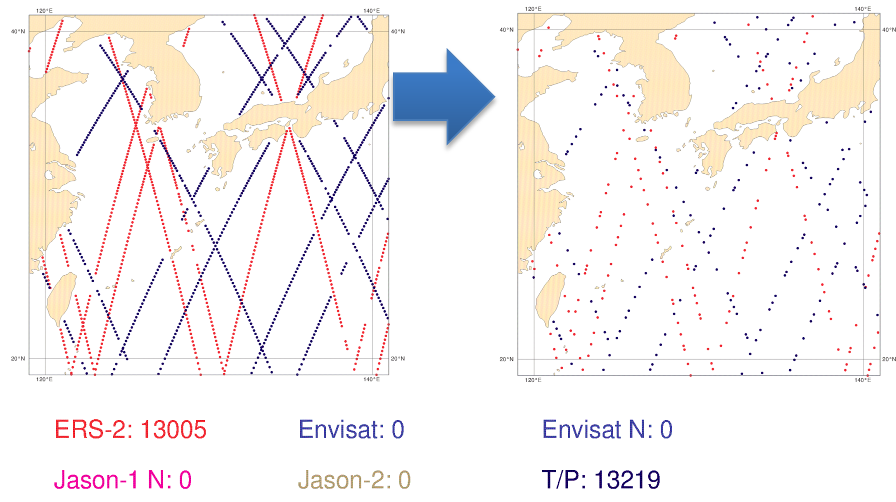
Assimilation of SLA: pre-process obs

- **The SLA along track data has very high spatial (9-14km) resolution for the operational ocean assimilation systems.**
    - Features in the data which the model can not represent
    - “Overfitting” to SLA obs
- **This can be dealt with in different ways:**
    - Inflate the observation error
    - Construction of “superobs” or thinning

Inflate OBE std dev of SSH



Thinning of SLA obs

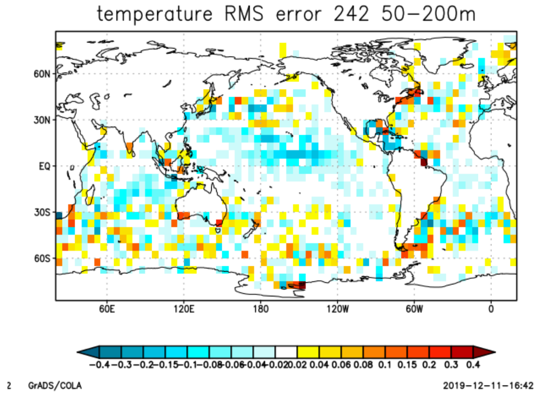


Assimilation of SLA: impact on the ocean state

Assimilation of SLA improves simulated ocean states

- Global mean sea-level changes
- Regional sea-level changes
- Subsurface temperature and salinity
- Large-scale ocean circulations

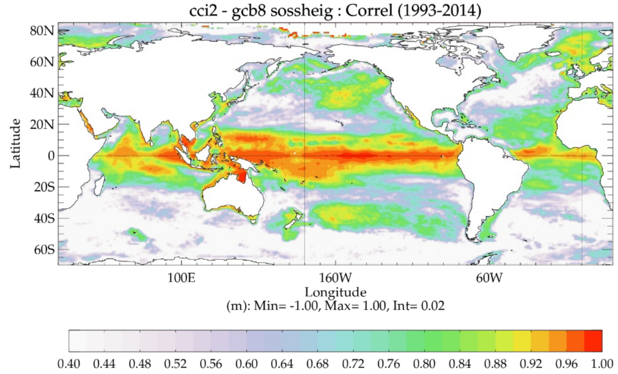
T ∆RMSE (O-B):
assim. SLA – not assim. SLA



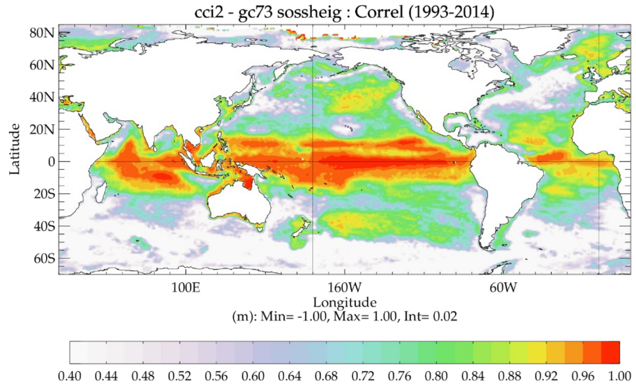
Temporal correlation (monthly) to AVISO data

ORAS5-NoAlti

Zuo et al., 2018

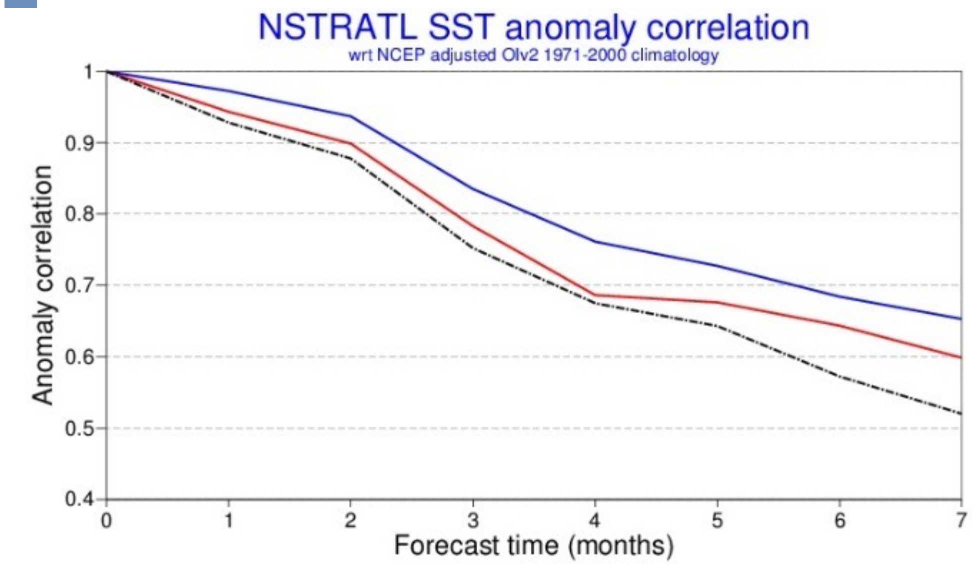


ORAS5



Impact of Altimeter SLA on Seasonal forecast

No Alt ocean ini
Seasonal Fc with Alt



Consistent Improvement of Skill over the Atlantic Ocean

Shown are Northern Subtropical Atlantic.
Potential for prediction of tropical cyclones.

In other regions the impact is not so obvious and varies with season.

In the Pacific, skill is improved mainly in forecasts initialized during spring (important for ENSO onset).

Courtesy of Magdalena

Altimeter Concluding remarks

ECMWF has a long experience with altimetry in the wave model

- Available continuously from 1993 onwards:
  - ERS1/2, Envisat, Jason1/2/3, Cryosat, Saral, Sentinel-3…
  - Now with new missions: CFOSat-SWIM, Sentinel-6,…

Altimeter wind and wave data are used for:

- Data assimilation
- Error estimation
- Use in reanalyses (assimilation and validation)
- Long term assessments & climate studies
- Monitoring of model performance (inc. model resolution) & Assessment of model changes

Altimeter sea level anomaly:

- Use for assimilation and validation
- Significant impact for surface and sub-surface ocean
- Importance for reanalysis and climate studies
- Uncertainty from the ensemble members potentially used for model error

SMOS

SMOS (Soil Moisture and Ocean Salinity) was **launched in November 2009**

The SMOS synthetic antenna consists of **69 radiometer elements** operating at L-band (**frequency ~1.4 GHz, λ=21 cm**) and distributed along three equally spaced arms, resulting in a planar Y-shaped structure.

In **aperture synthesis radiometers**, a TB image is formed through Fourier synthesis from the cross correlations between simultaneous signals obtained from pairs of antenna elements.

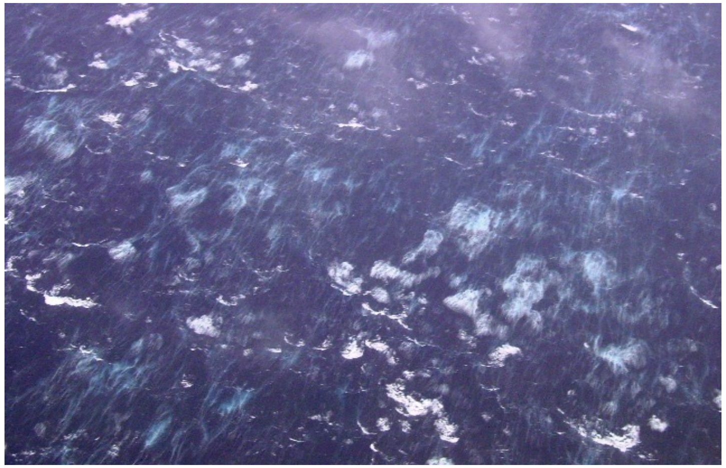
Multi-angular images of the brightness temperature are obtained over a large swath width (1200 km), with a **spatial resolution** varying within the swath **from 30 km to about 80 km**, and with a **revisit time of less than 3 days**.

SMOS wind data

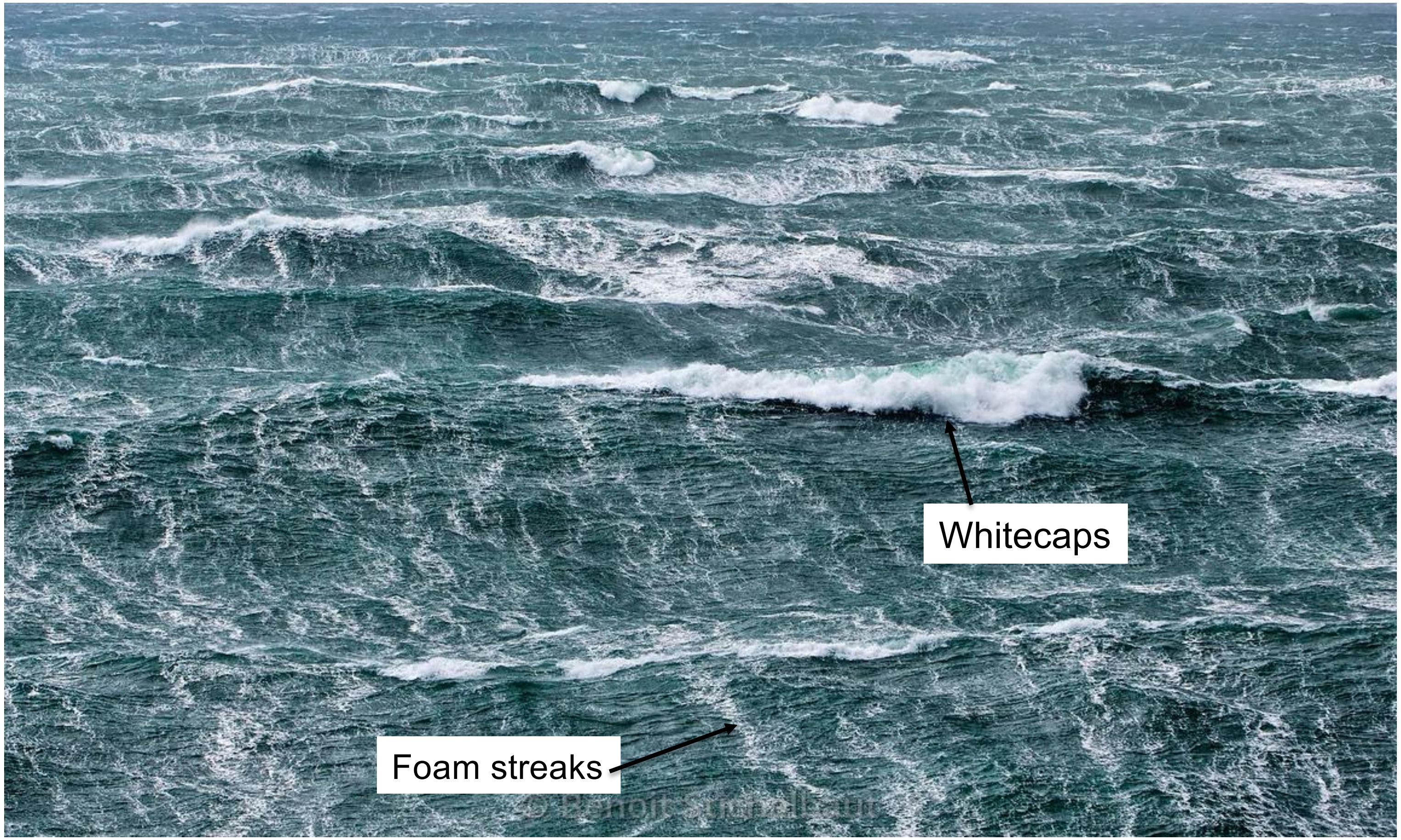
- ✓ L-band is less affected by rain, spray and atmospheric effects than higher mw frequencies (C-band, Ku-band)

- ✓ There is no saturation at high wind speed like for radars

- ✓ Sea foam, generated by breaking waves which mainly depends on surface wind strength and sea state development, increases the microwave ocean emissivity



Ocean-Atmosphere Interface in very High Wind speed conditions

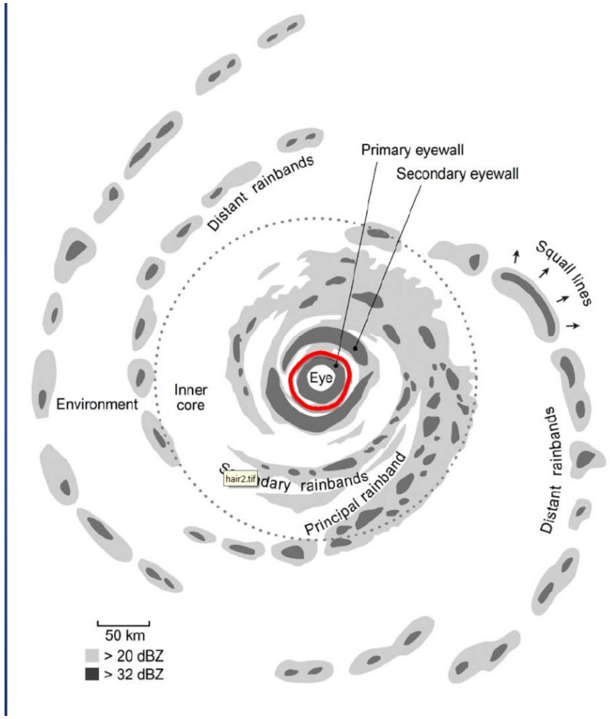


Most of the increased surface whitening at & above hurricane force (>33 m/s) is principally induced by the **increased streaks coverage**

Whitecap coverage is found ~constant above Hurricane force ~4 [Holthuijsen et al. JGR 2012]

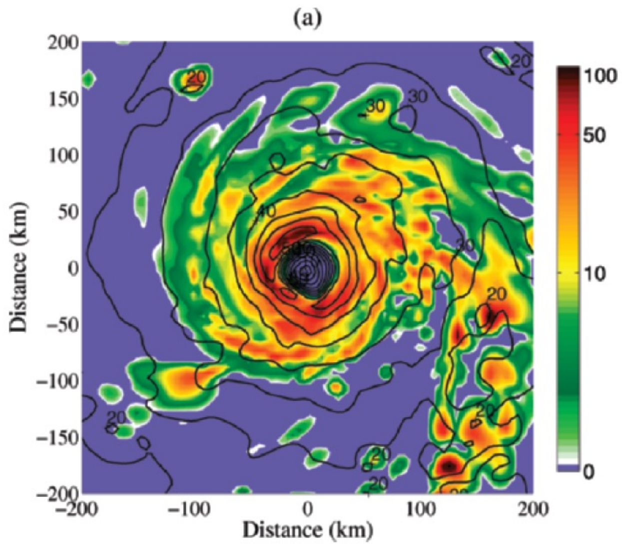
High winds in Hurricanes are very often associated with High rain rates

Rain Anatomy in a hurricane



Houze 2010

Rain rate [mm/h]

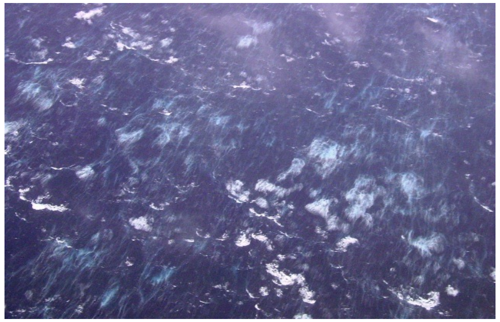


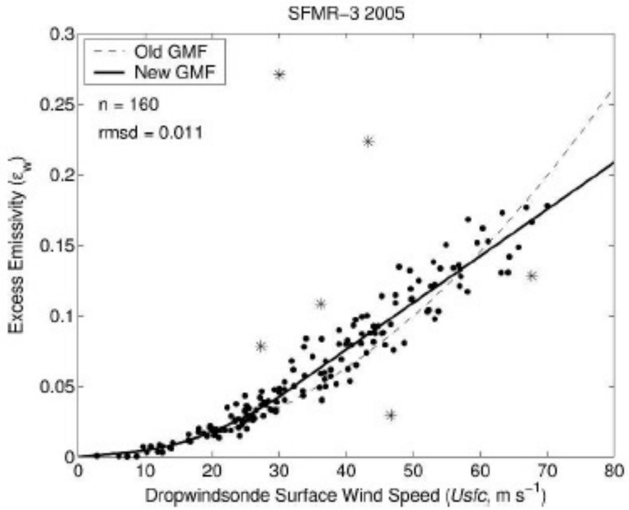
S.Shen and J. Tenerelli 2007

Because of the small ratio of raindrop size to the SMOS electromagnetic wavelength (~21 cm), scattering by rain is almost negligible at L-band, even at the high rain rates experienced in hurricanes.

Wind speed retrieval in extreme winds : SFMR

Increase of the microwave ocean emissivity
with wind speed ⇔ surface foam change impacts





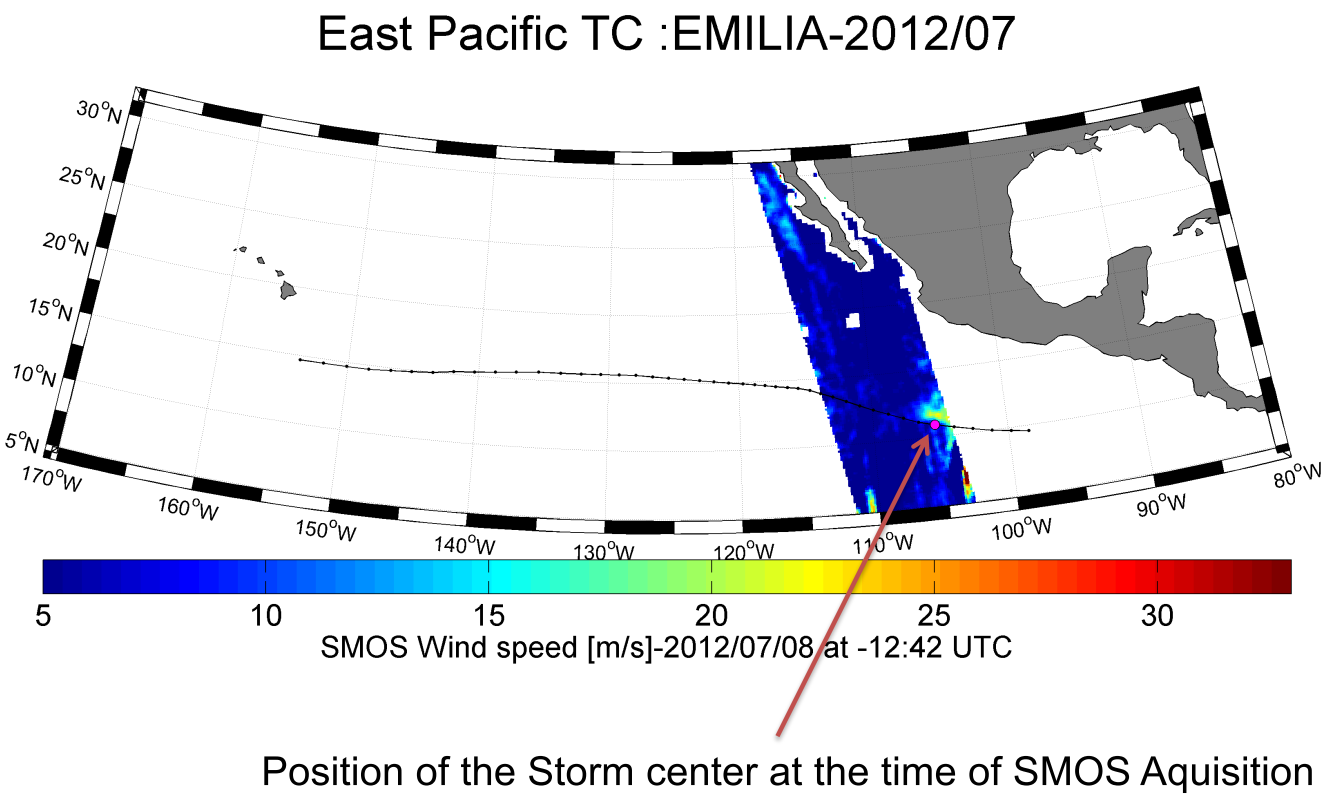
Hurricane hunter P-3

This information can be used to retrieve the surface wind speed in Hurricanes:

Principle of **the Step Frequency Microwave Radiometer** (SFMR)
C-band: => Use multi-frequency C-band channels to separate wind from rain effects
NOAA's primary airborne sensor for measuring Tropical Cyclone surface wind speeds since 30 year (Ulhorn et al., 2003, 2007).

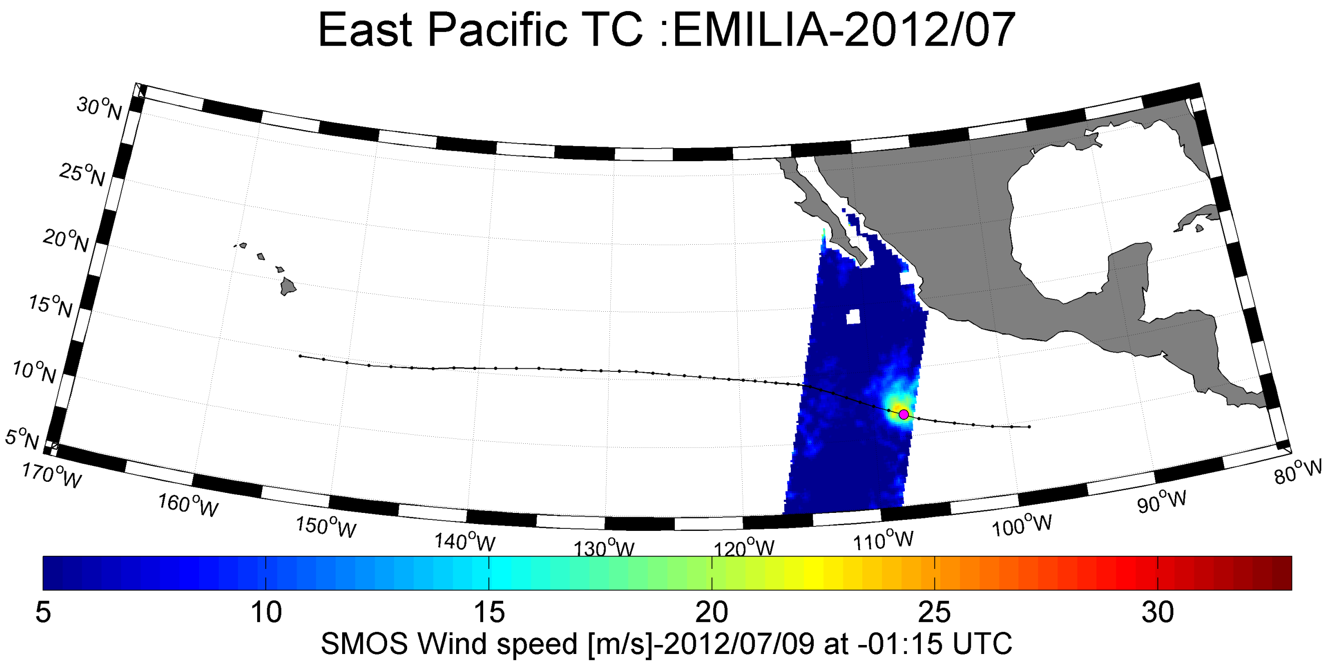
SMOS+STORM Evolution

Detect the useful TC & ETC events in SMOS data: Example of EMILIA



SMOS+STORM Evolution

Tasks 2: Detect the useful TC & ETC events in SMOS data: Example of EMILIA



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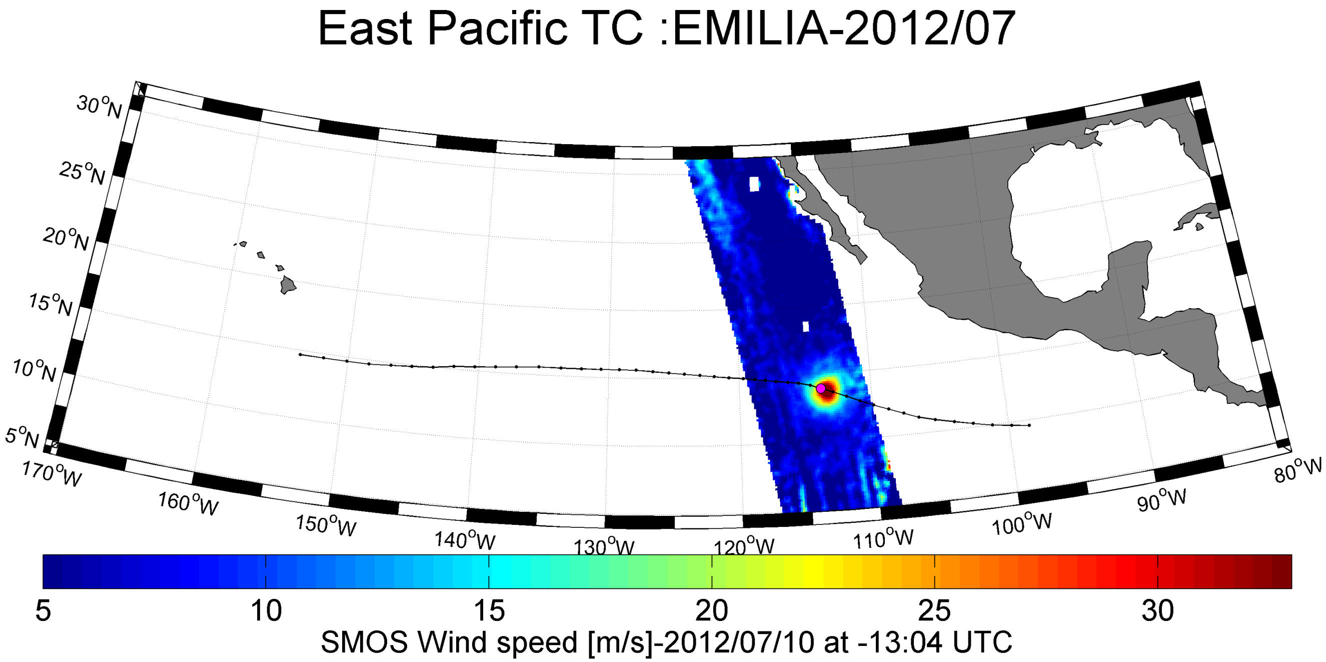

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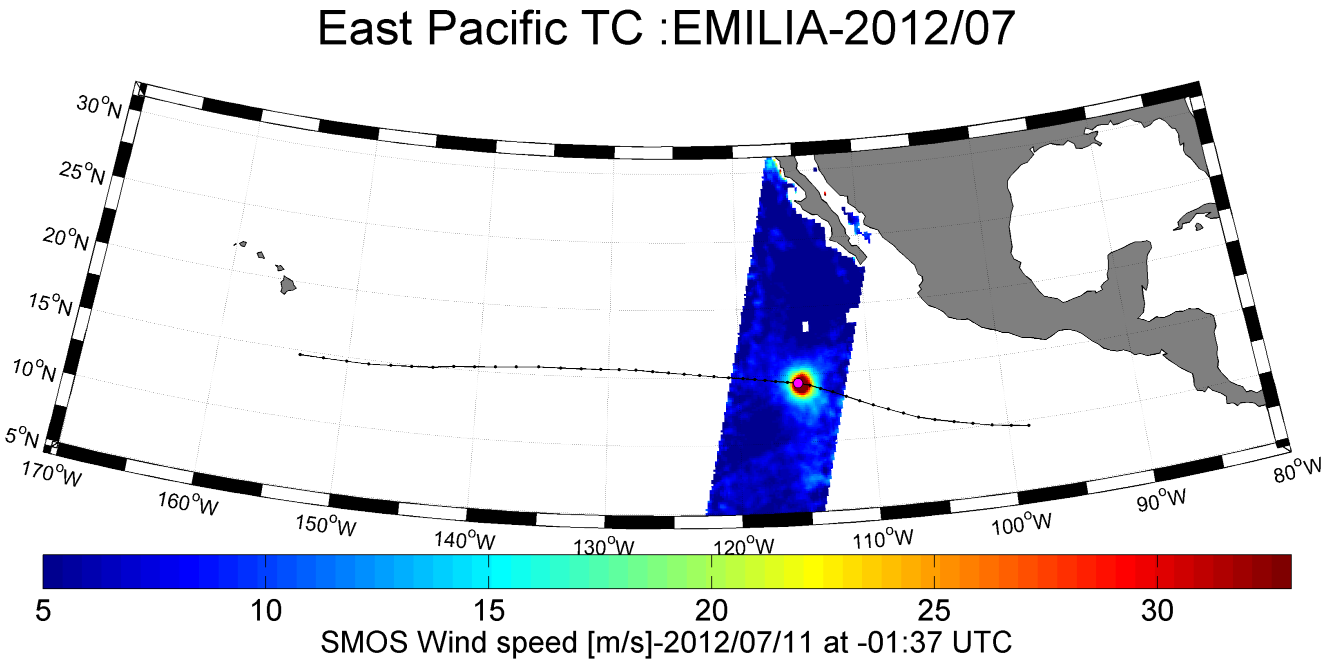

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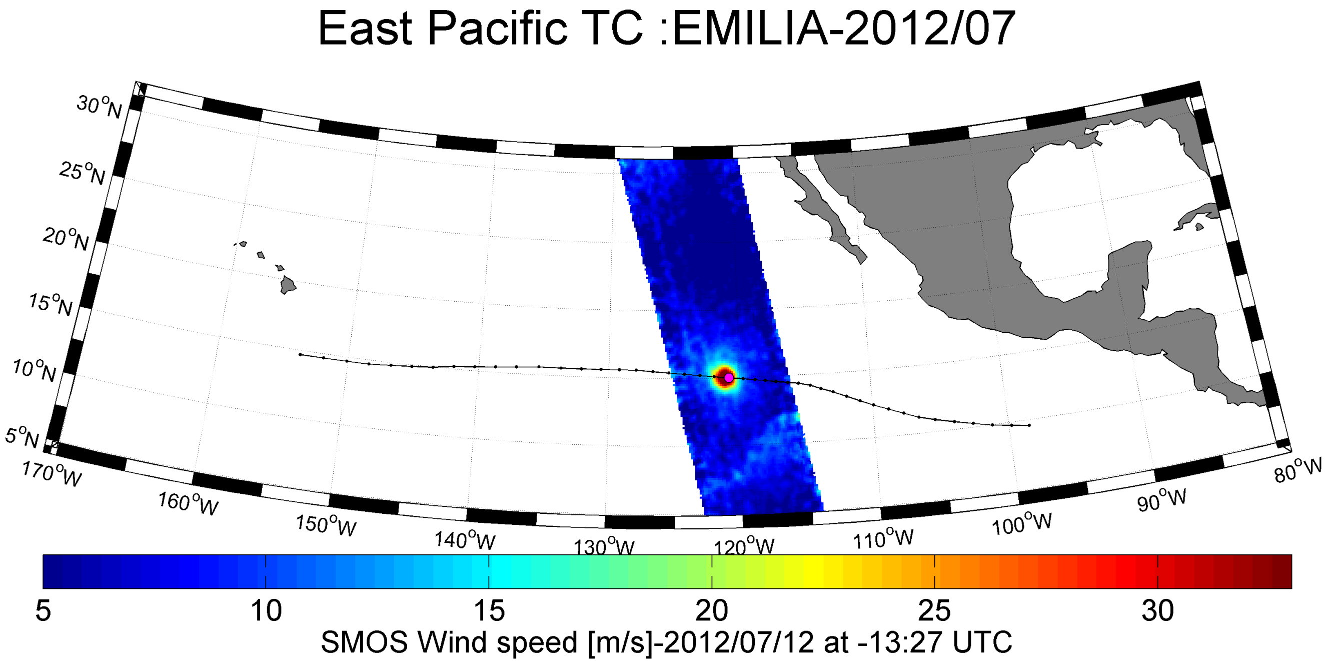

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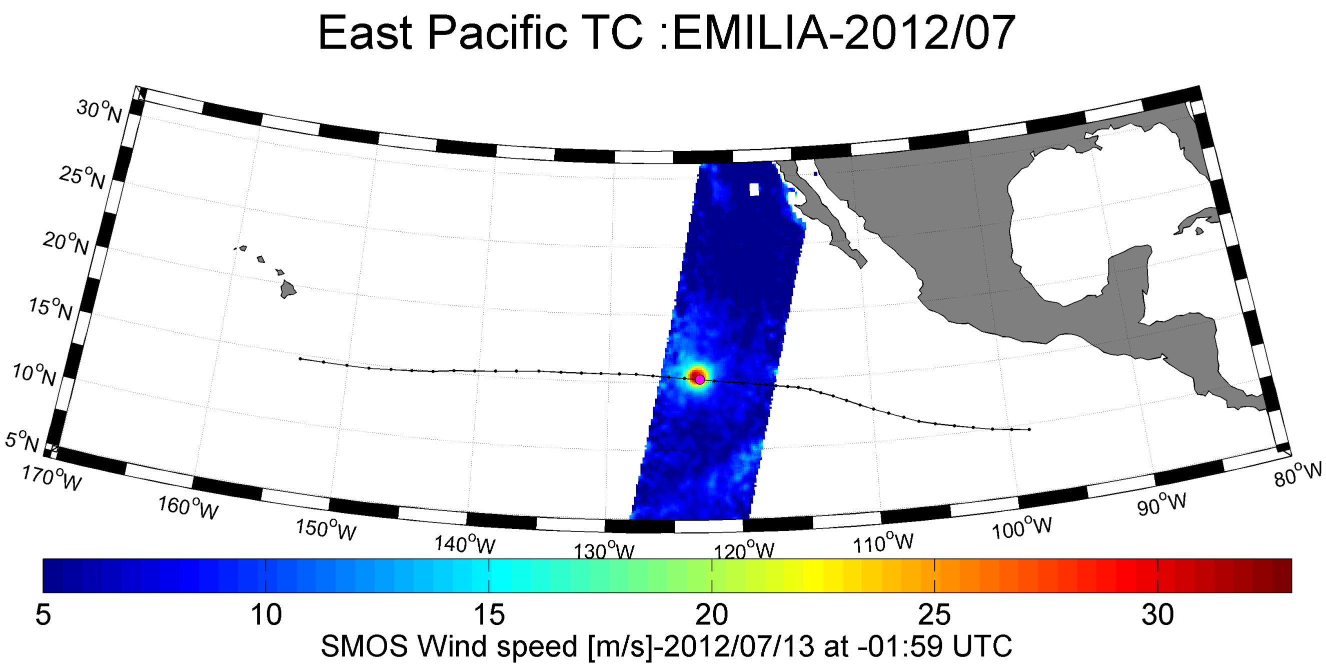
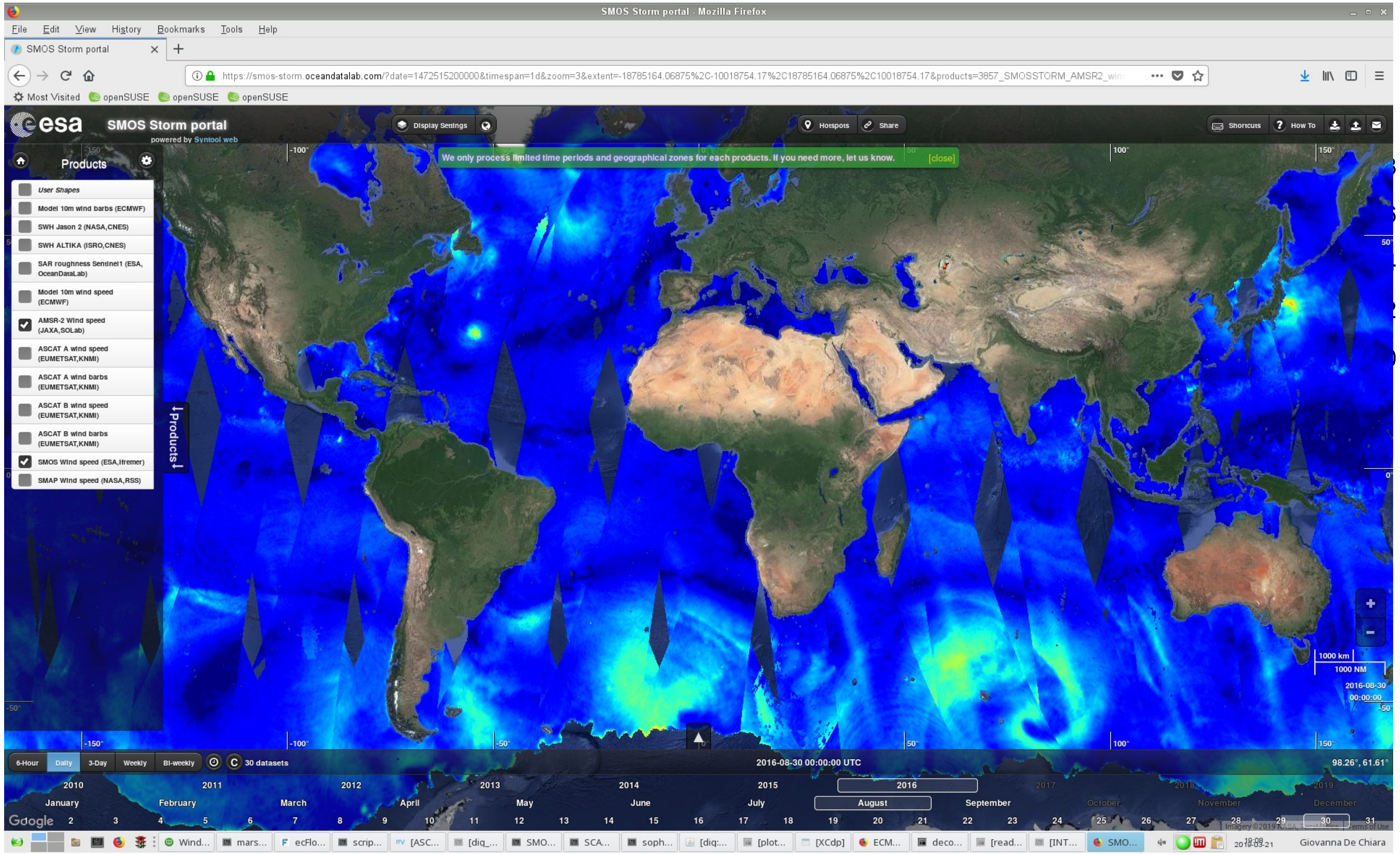


A view at the SMOS-STORM 2010-2015 TC database



https://smos-storm.oceandatalab.com

Summary

- The SMOS brightness temperature signal (ΔT_B) is clearly associated with the passage of Tropical Cyclones
- Correlations between L-band Tb increase with TC intensity from Cat 1 to Cat 5 was demonstrated
- L-band observations provide a first non-atmosphere corrupted view of the ocean surface in extreme conditions=> wind speed retrieval with ~5m/s accuracy
- A complete storm database as been generated for the SMOS mission archive:

TC & ETC 2010-now available at http://www.smosstorm.org/

Thanks!!!