

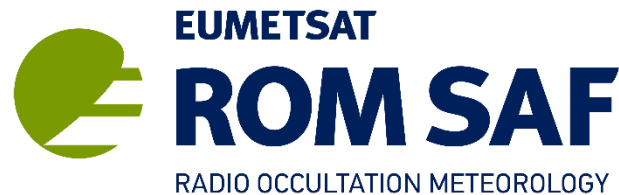
# **GNSS radio occultation lecture 2**

## **Impact/ some applications**

Sean Healy

(**sean.healy@ecmwf.int**)

NWP SAF lecture 2, May 17, 2023



<http://www.romsaf.org/>

# Acknowledge

- Hans Gleisner, DMI and ROM SAF
- Mark Ringer, Hadley Centre, Met Office
- Estel Cardellach, IEEC and ROM SAF
- Niels Bormann, ECMWF
- Daisuke Hotta, JMA
- Katrin Lonitz, ECMWF and ROM SAF

# **Presentations on GNSS-RO applications from ROM SAF/IROWG (2019), IROWG-8 (2021) and OPAC/IROWG (2022)**

- See presentations at:
  - <https://www.romsaf.org/romsaf-irowg-2019/en/content/21/program-agenda-by-day>
  - <https://cpaess.ucar.edu/events/8409/agenda>
  - <https://opacirowg2022.uni-graz.at/en/scientific-programme/>

# Outline

**Aim:** provide an overview of **some** GNSS-RO applications

Recap from lecture 1

- *GNSS-RO information content, key characteristics - ‘core region’, etc.*

*As you might expect, many applications are related to these characteristics in the core region*

- **GNSS-RO impact in NWP systems**
- **Key observation climate reanalyses**
- Climate monitoring
- Retrieving hydrometeor water content
- Summary

*EMAIL me for space weather/surface pressure information/PBL/radiosonde bias correction.*

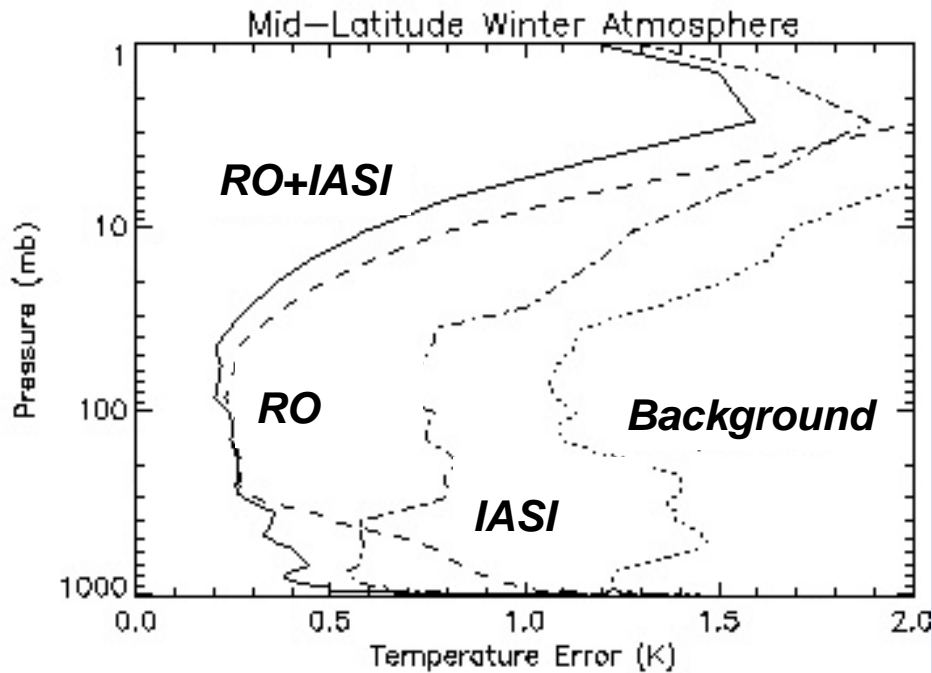
# Recap

- **All satellite measurements have strengths and weaknesses.**  
The aim is to construct a robust *global observing system* with a good balance of the different types, **given their distinct characteristics and information content.**
- **GNSS-RO measurements are useful because they complement satellite radiances**
  - Assimilation without bias correction (an *anchor measurement*)
  - Good vertical resolution
- The information content is largest in the “**core region**”, between 7-35 km, **and I will demonstrate** we see a large NWP impact on upper-tropospheric and lower/middle stratospheric temperatures.

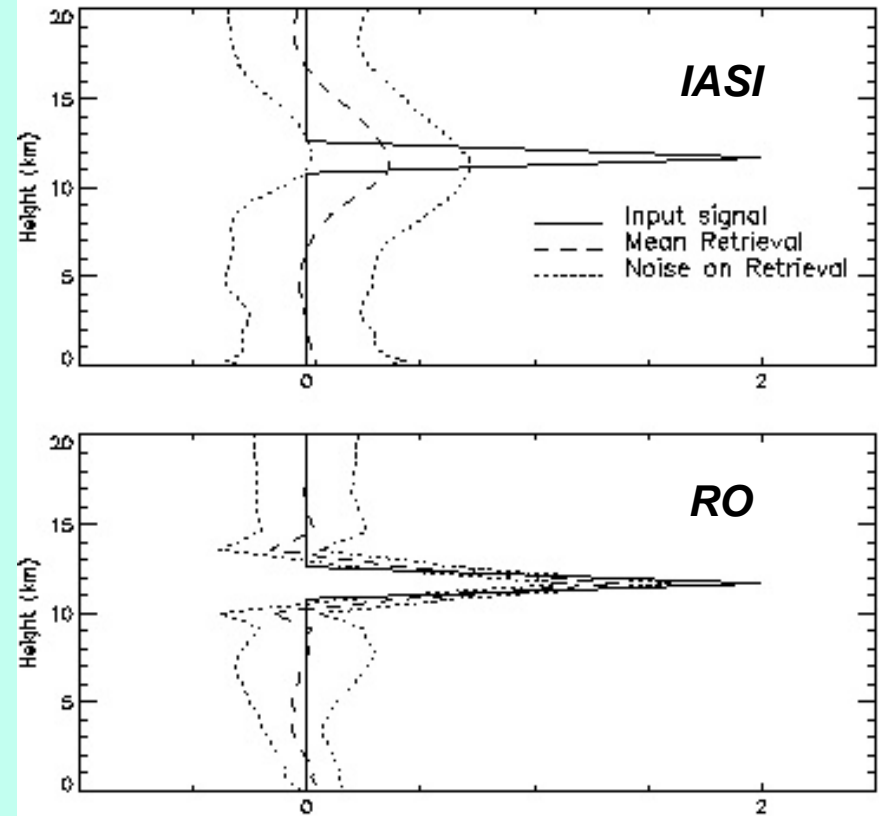
# GNSS-RO and IASI: 1DVAR simulations

Collard and Healy 2003,  
QJRMS:

*Expected retrieval uncertainty:*



*Power to resolve a peak-shaped error in background: Averaging Kernel.*



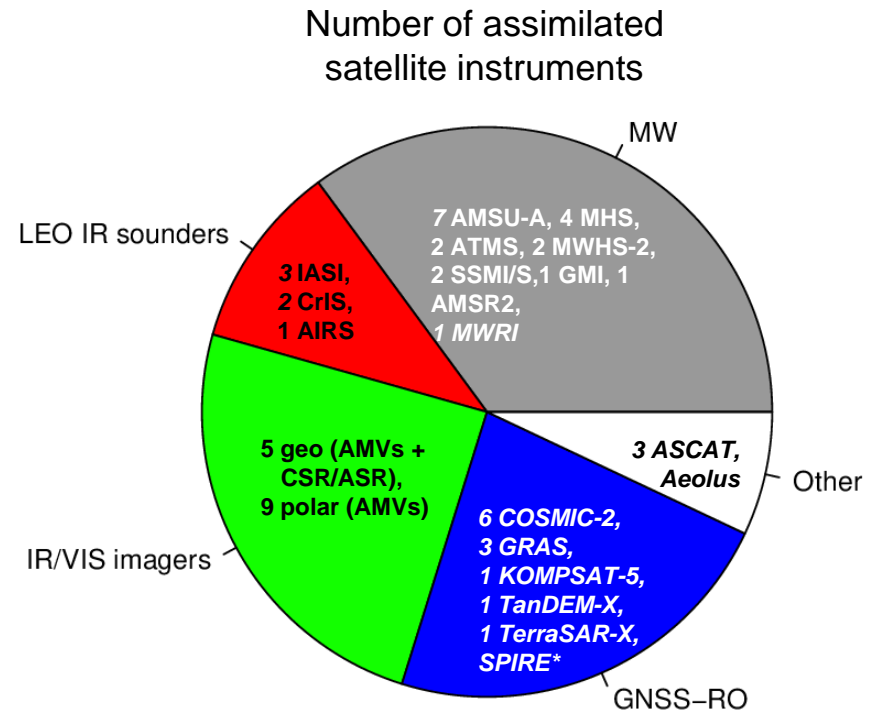
# Impact of various observing systems at ECMWF

Provided by Niels Bormann – 2021 annual seminar

[https://events.ecmwf.int/event/217/contributions/2049/attachments/1397/2509/AS2021\\_Bormann.pdf](https://events.ecmwf.int/event/217/contributions/2049/attachments/1397/2509/AS2021_Bormann.pdf)

# Observing system experiments

- *Periods, 6 months in total:*
  - 5 Sept – 2 Nov 2020
  - 1 Jan – 28 Feb 2021
  - 1 May – 30 June 2021(each + 4 days spin-up prior)
- Denial experiments compared to a full system for:
  - *Conventional in-situ observations*
  - *MW radiances*
  - *IR sounders from LEO*
  - *IR/VIS imagers (AMVs + IR radiances)*
  - *GNSS-RO*
- *Resolution:  $T_{CO}$  399 (~25 km)*
- *Background error from operational system*

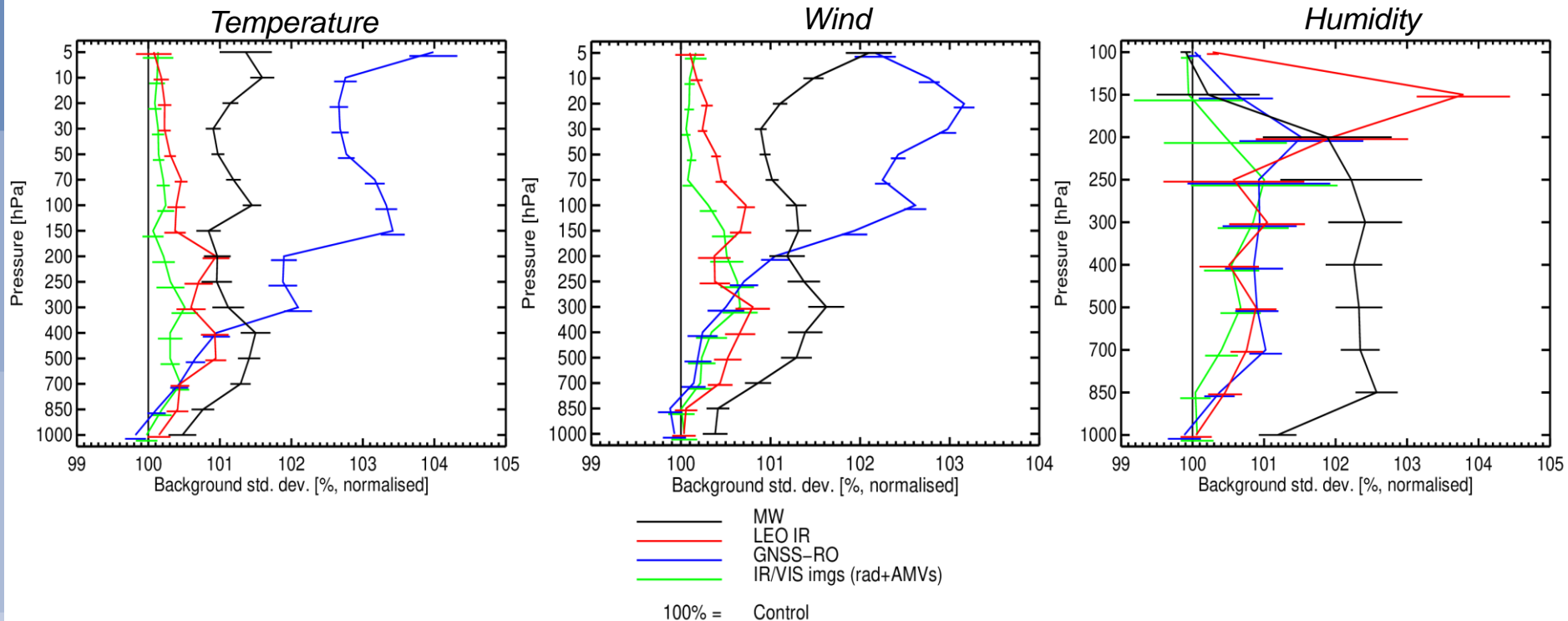


\* SPIRE: Sept 2020 only



# Short-range impact evaluated against in-situ observations: Stdev(o-b)

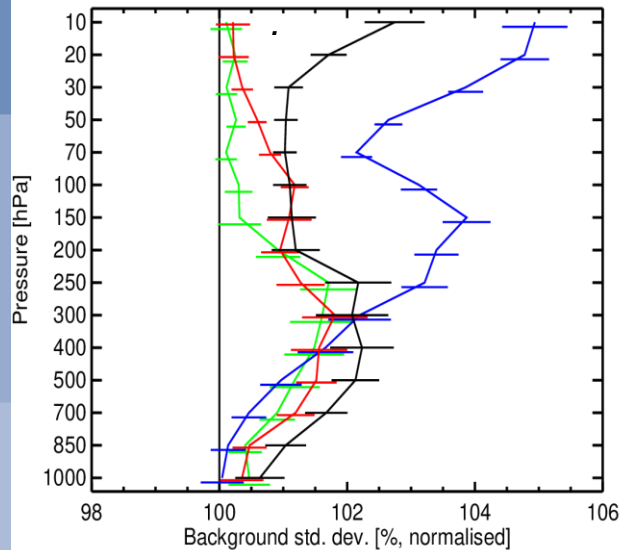
*Global, 3 periods combined*



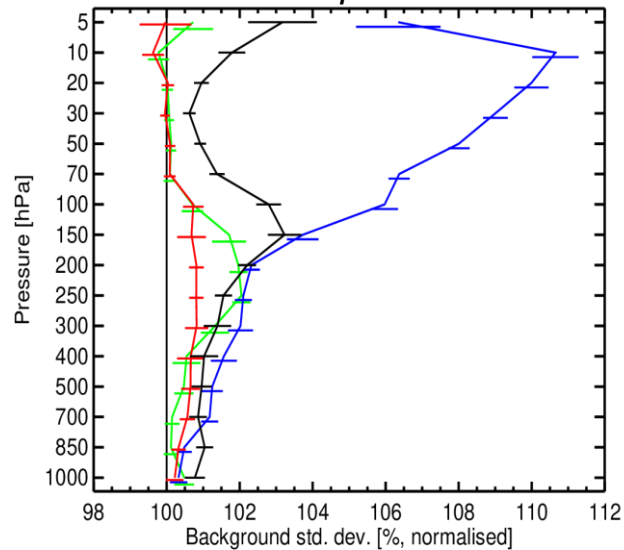
# Short-range impact evaluated against in-situ observations: **wind**

3 periods combined

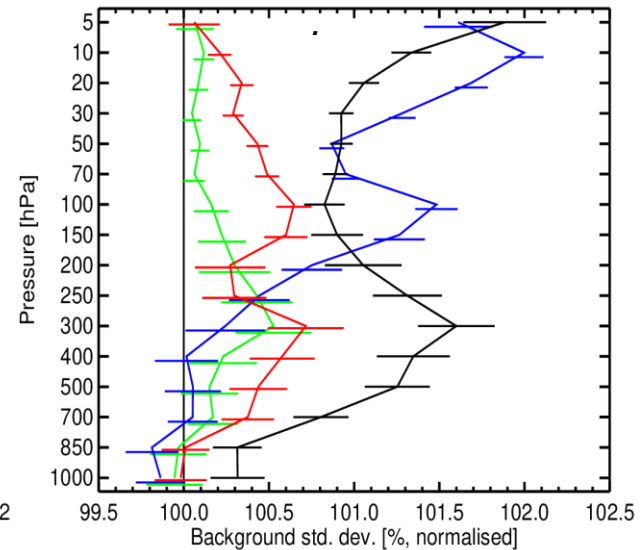
S.Hem



Tropics

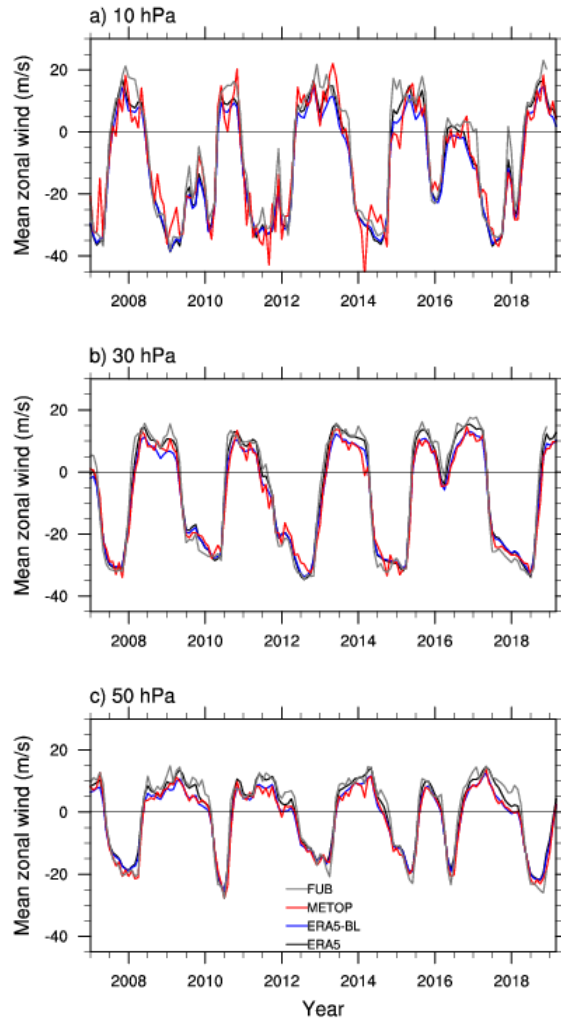


N.Hem



- MW
  - LEO IR
  - GNSS-RO
  - IR/VIS imgs (rad+AMVs)
- 100% = Control

# ASIDE: Zonally averaged zonal winds retrieved from a ROM SAF monthly mean GNSS-RO geopotential climatology



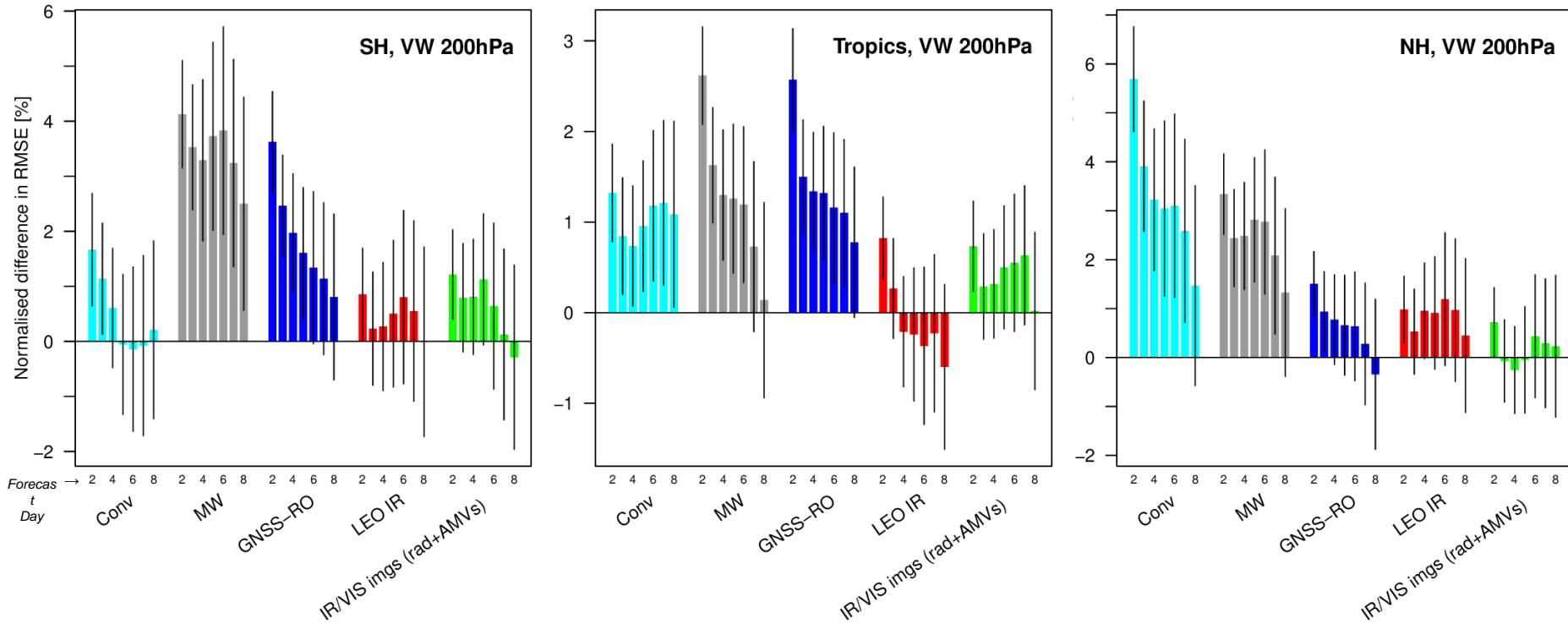
Compute the “balanced” GNSS-RO zonal winds from the second derivative of the zonally averaged geopotential height

$$\bar{U} \approx -\frac{1}{\beta} \frac{\partial^2 \bar{\phi}}{\partial y^2}$$
$$\beta = \frac{2\Omega}{a}$$

*FUB is the Free University Berlin radiosonde zonal wind climatology at Singapore.*

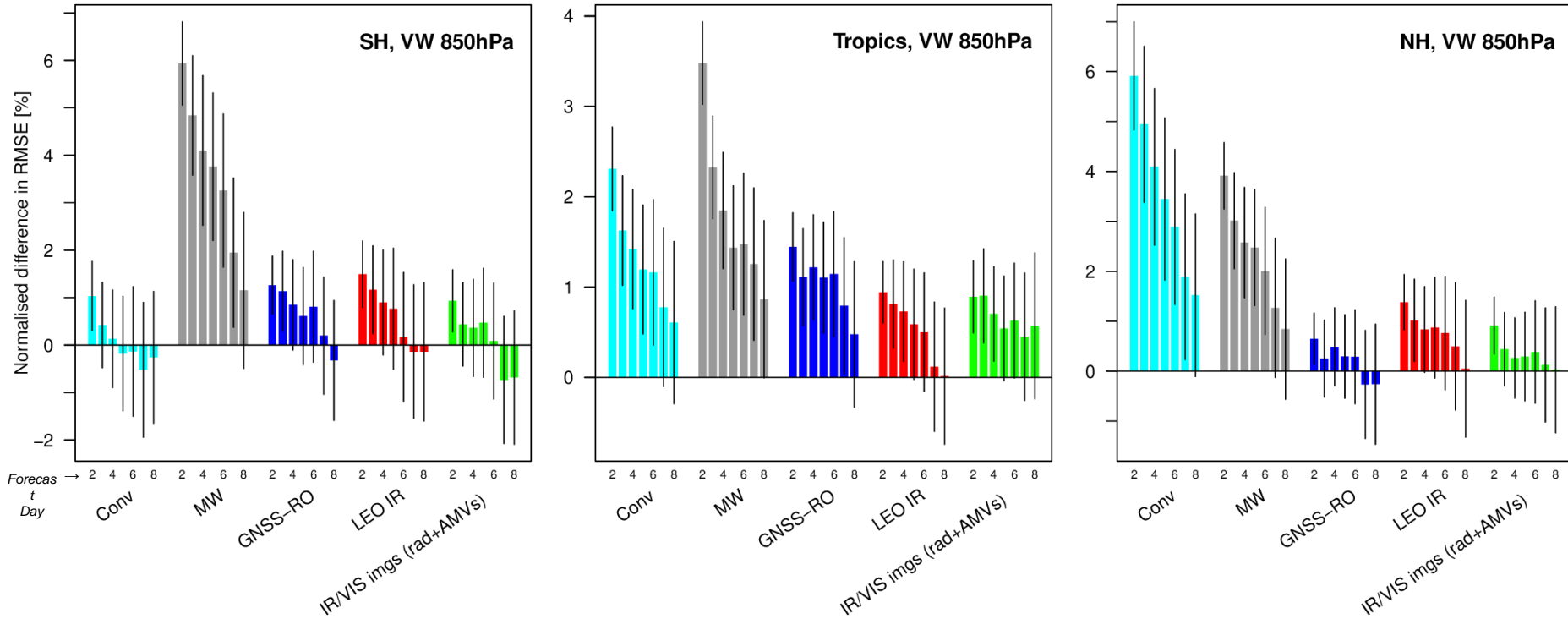
# Forecast impact, day 2-8: Wind at 200 hPa

Verified against operational analyses, 3 periods combined



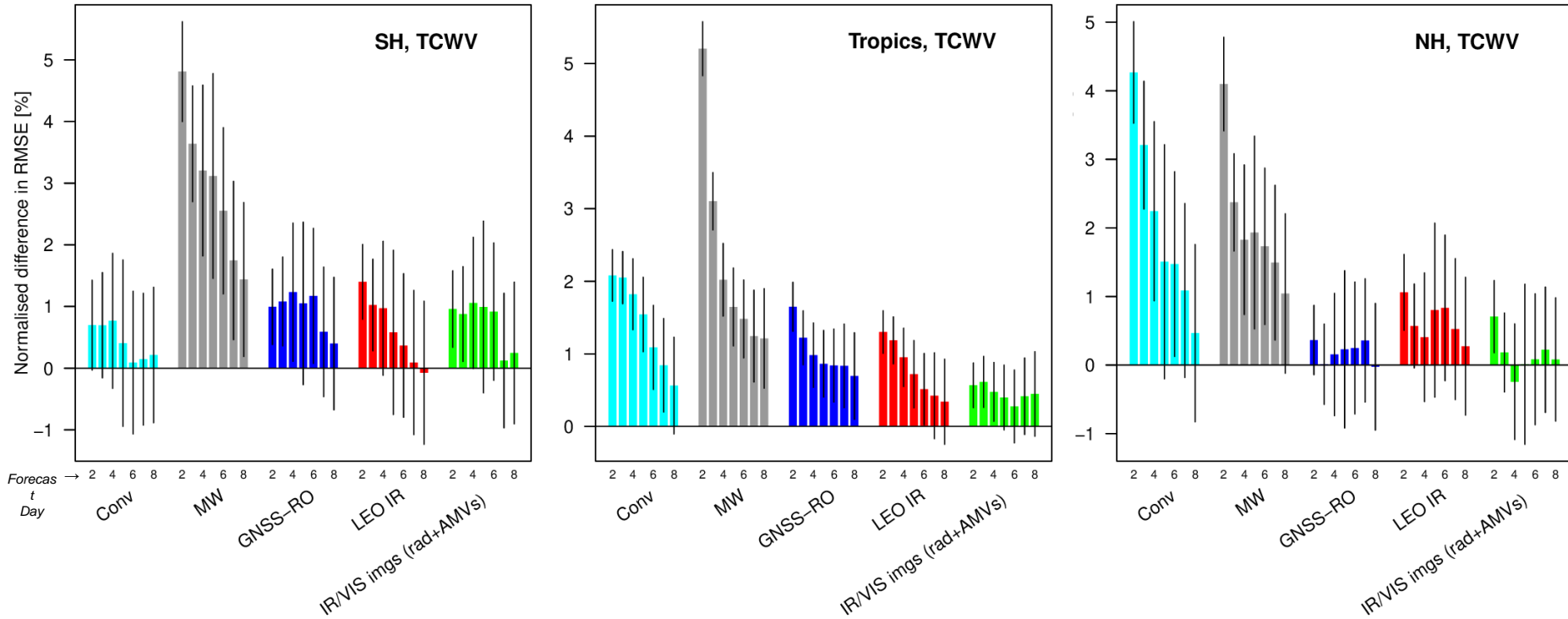
# Forecast impact, day 2-8: Wind at 850 hPa

Verified against operational analyses, 3 periods combined



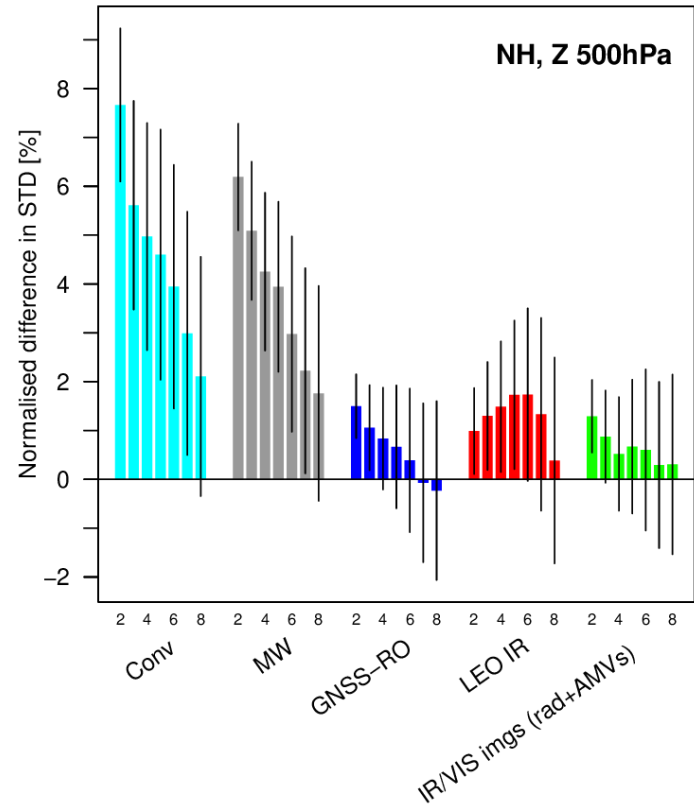
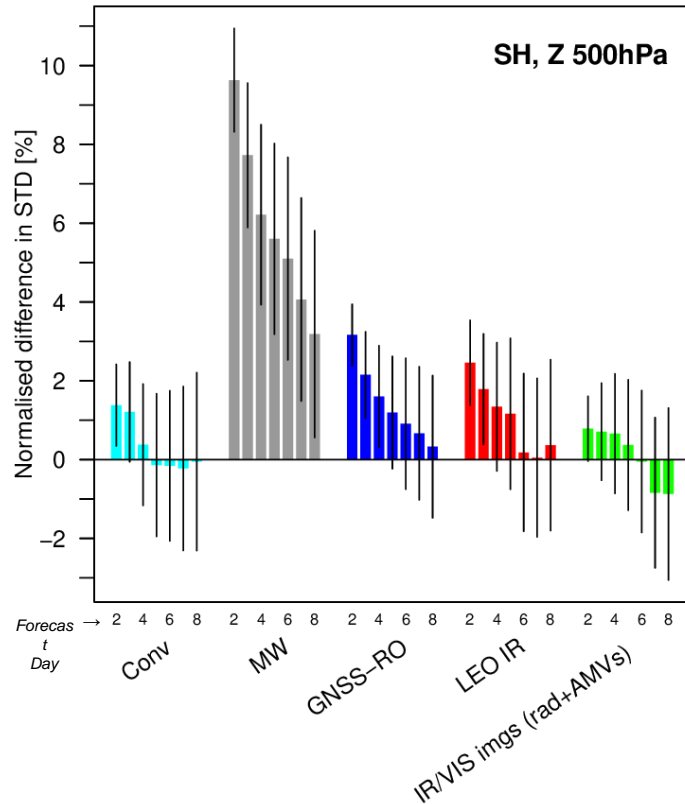
# Forecast impact, day 2-8: Total column water vapour

Verified against operational analyses, 3 periods combined



# Forecast impact, day 2-8: 500 hPa geopotential

Verified against operational analyses, 3 periods combined



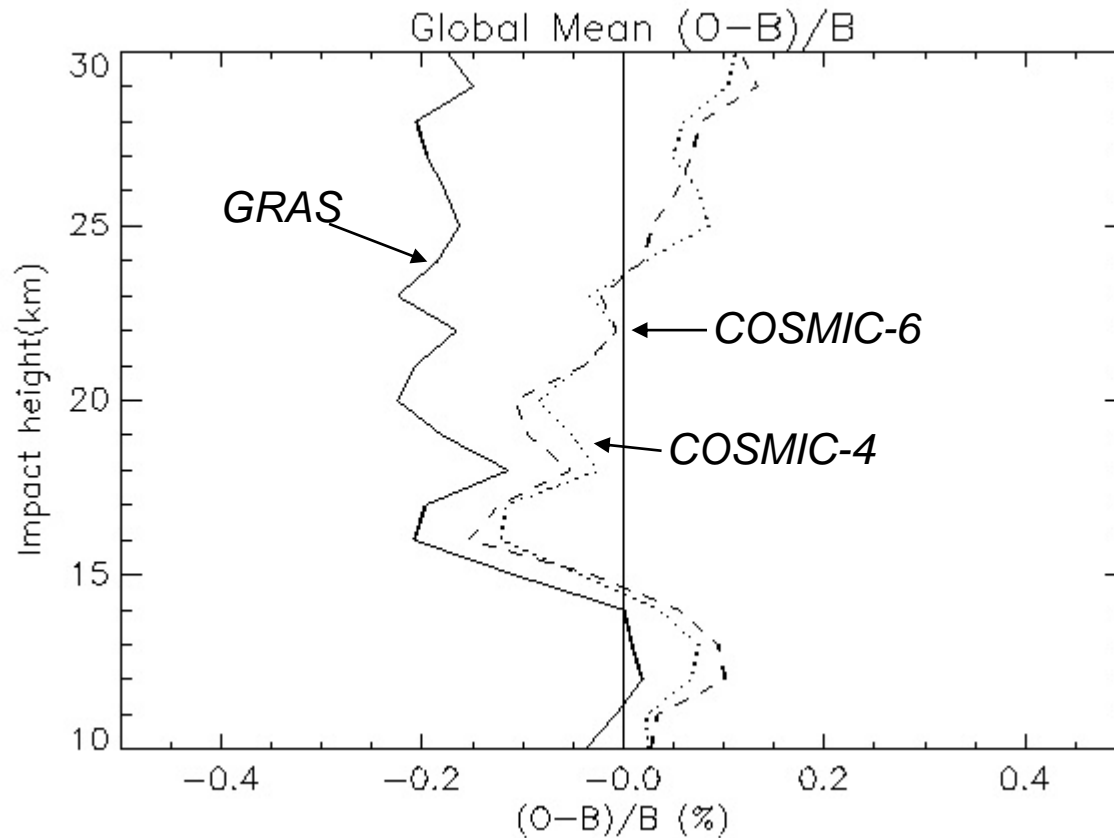
# Climate reanalysis applications



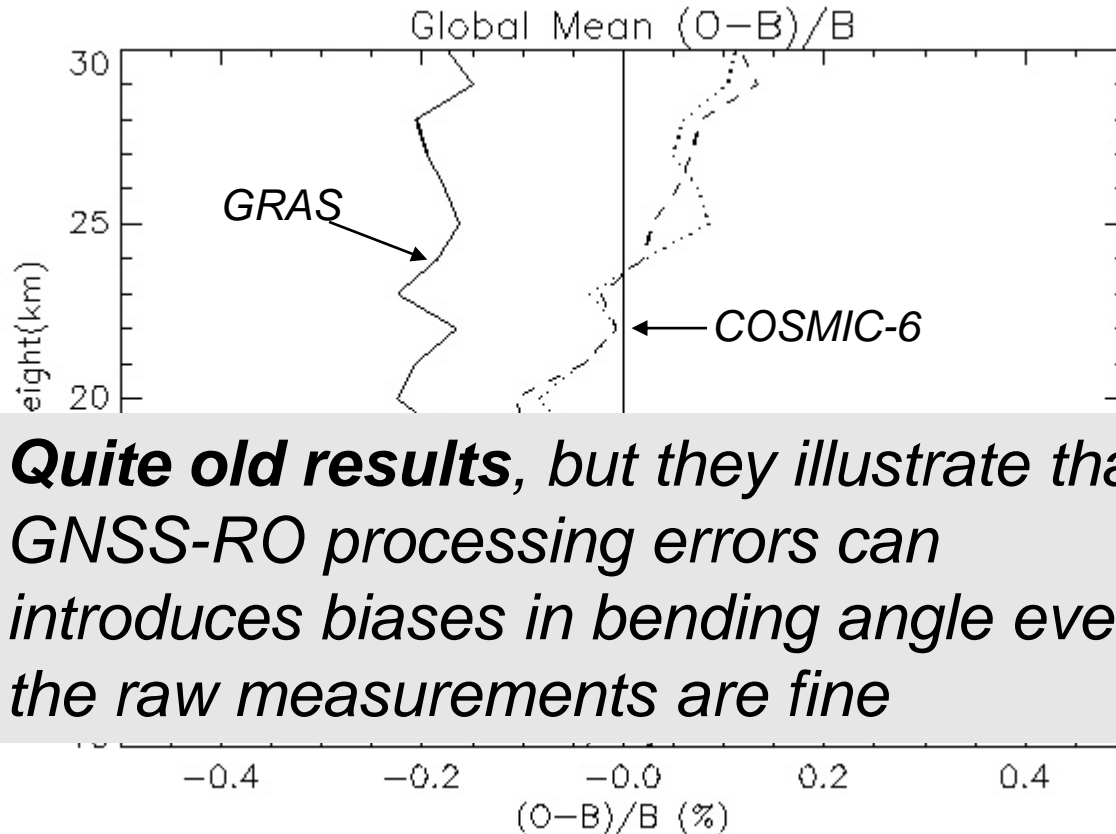
# Climate reanalysis applications

- **We have only had significant quantities of GNSS-RO since 2006 with the introduction of COSMIC**
- **Claim:** GNSS-RO measurements should not be biased.
  - It should be possible to introduce data from new instruments without long overlap periods for calibration.
  - No discontinuities in time-series as a result of interchange of GNSS-RO instruments.
- **Bending angle time series derived from the ERA-Interim reanalysis were used to investigate this claim**

# Global bending angle (o-b)/b departure statistics from ECMWF operations for Aug.20 to Sept. 20, 2009



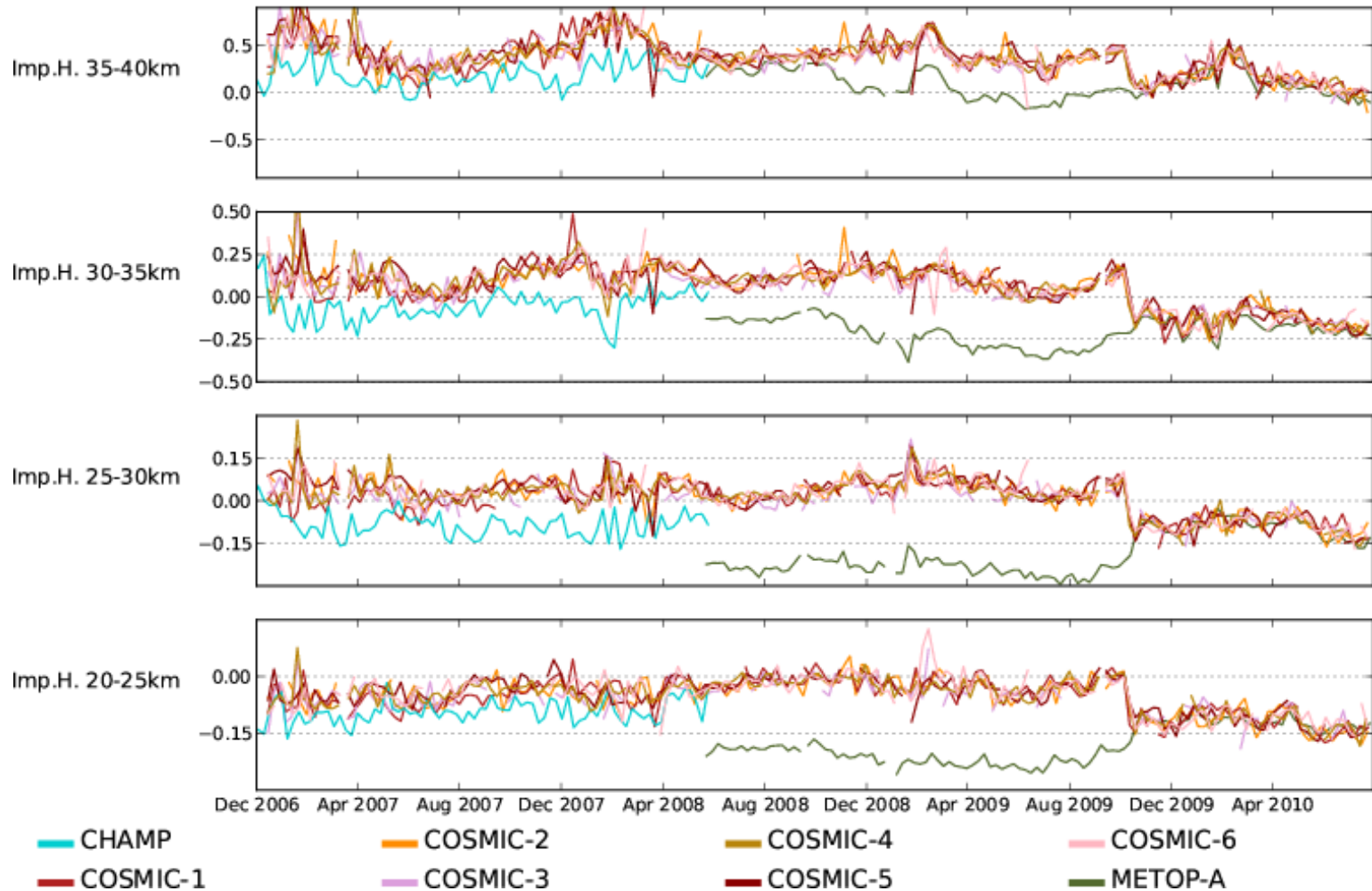
# Global bending angle $(o-b)/b$ departure statistics from ECMWF operations for Aug.20 to Sept. 20, 2009



***Quite old results, but they illustrate that GNSS-RO processing errors can introduce biases in bending angle even if the raw measurements are fine***

# Consistency of GNSS-RO bending angles (ERA-Interim Reanalysis, Paul Poli)

ERA-Interim daily Obs minus Background statistics GPSRO B.A. (percent) N.Hem. (20N-90N)



# GNSS-RO and the bias correction of radiances

- “Bias correction schemes for satellite radiances need to be grounded by a reference.” The reference measurements are often called “*anchor measurements*”
- The assimilation of GNSS-RO anchors the bias corrections we apply to radiances
- We can illustrate this by plotting how the bias corrections applied to radiances change with/without GNSS-RO

# VarBC is used at ECMWF

Dee, QJRMS (2007), 131, pp 3323-3343

- Bias corrected radiances are assimilated.

$$\tilde{\mathbf{y}} = \mathbf{y} - \mathbf{b}(\boldsymbol{\beta}, \mathbf{x})$$

$$\mathbf{b}(\boldsymbol{\beta}, \mathbf{x}) = \sum_i \beta_i \mathbf{p}(\mathbf{x})$$

$$J(\mathbf{x}, \boldsymbol{\beta}) = (\mathbf{x}_b - \mathbf{x})^T \mathbf{B}_x^{-1} (\mathbf{x}_b - \mathbf{x})$$

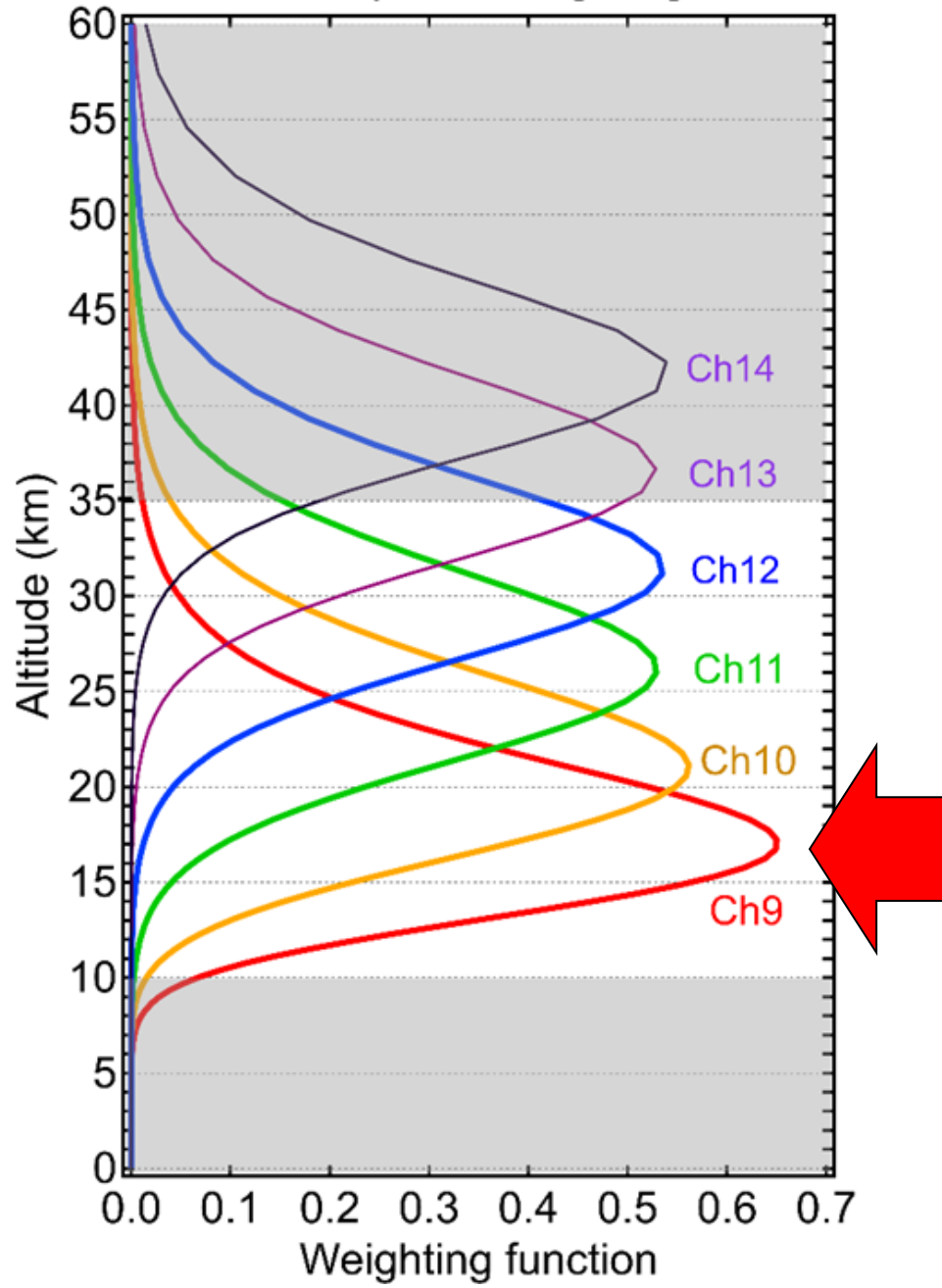
$$+ (\boldsymbol{\beta}_b - \boldsymbol{\beta})^T \mathbf{B}_\beta^{-1} (\boldsymbol{\beta}_b - \boldsymbol{\beta}) +$$

$$(\mathbf{y} - \mathbf{b}(\boldsymbol{\beta}, \mathbf{x}) - H(\mathbf{x}))^T \mathbf{R}^{-1} (\mathbf{y} - \mathbf{b}(\boldsymbol{\beta}, \mathbf{x}) - H(\mathbf{x}))$$

*In the 4D-Var, we minimize an augmented cost function, where the bias coefficients are estimated.*

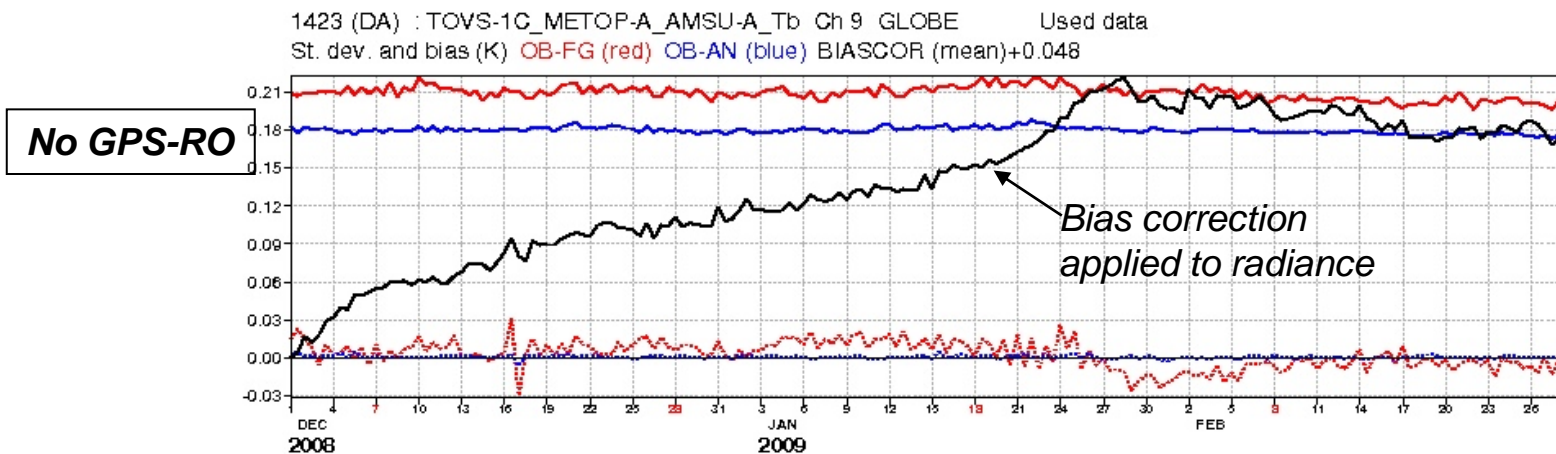
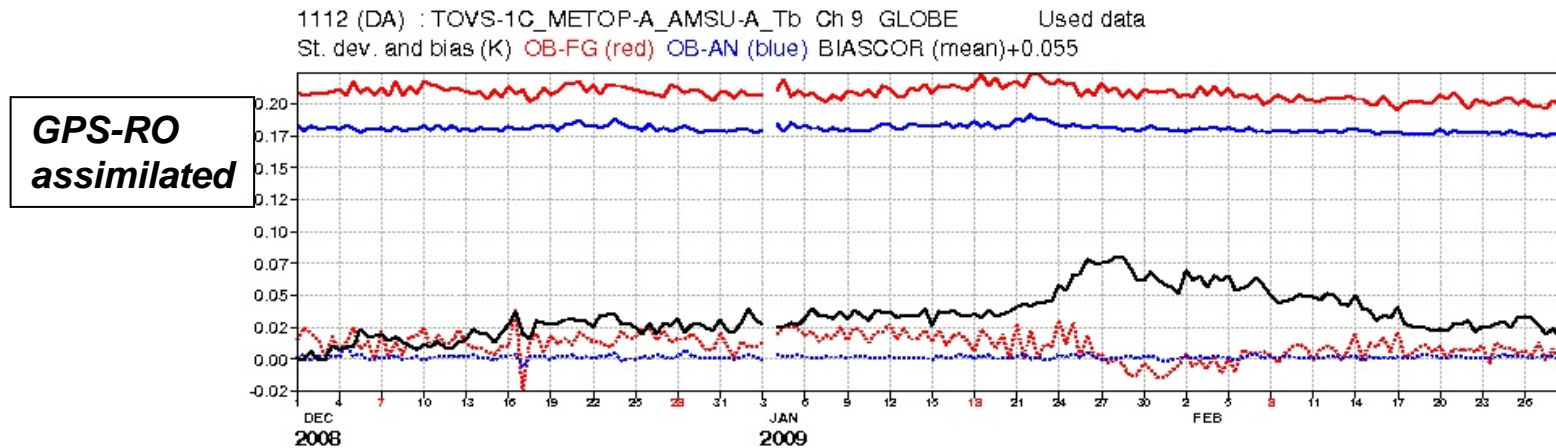
- **VarBC assumes an unbiased model.**

AMSU stratospheric weighting functions



# Experiment removing GPS-RO from ERA-Interim (Dec. 08, Jan-Feb 09)

- Impact on bias correction. E.g., globally averaged **Metop-A, AMSU-A channel 9 bias correction**.

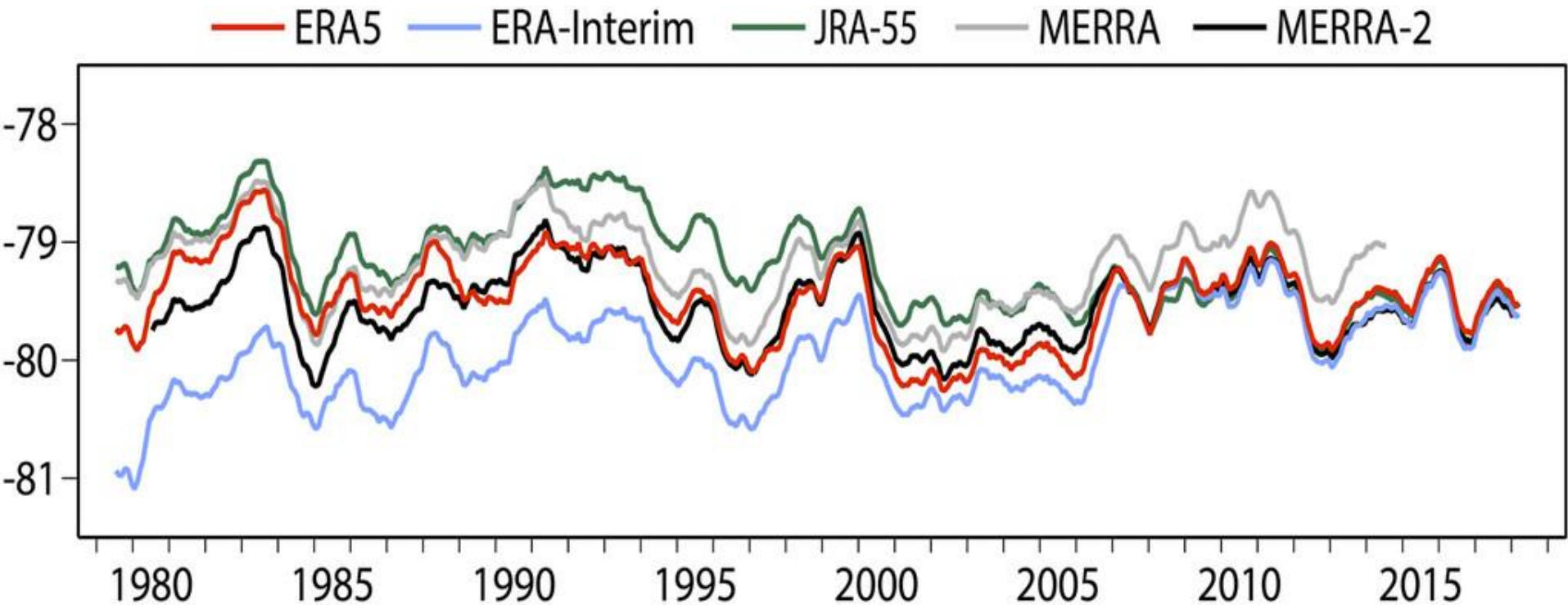




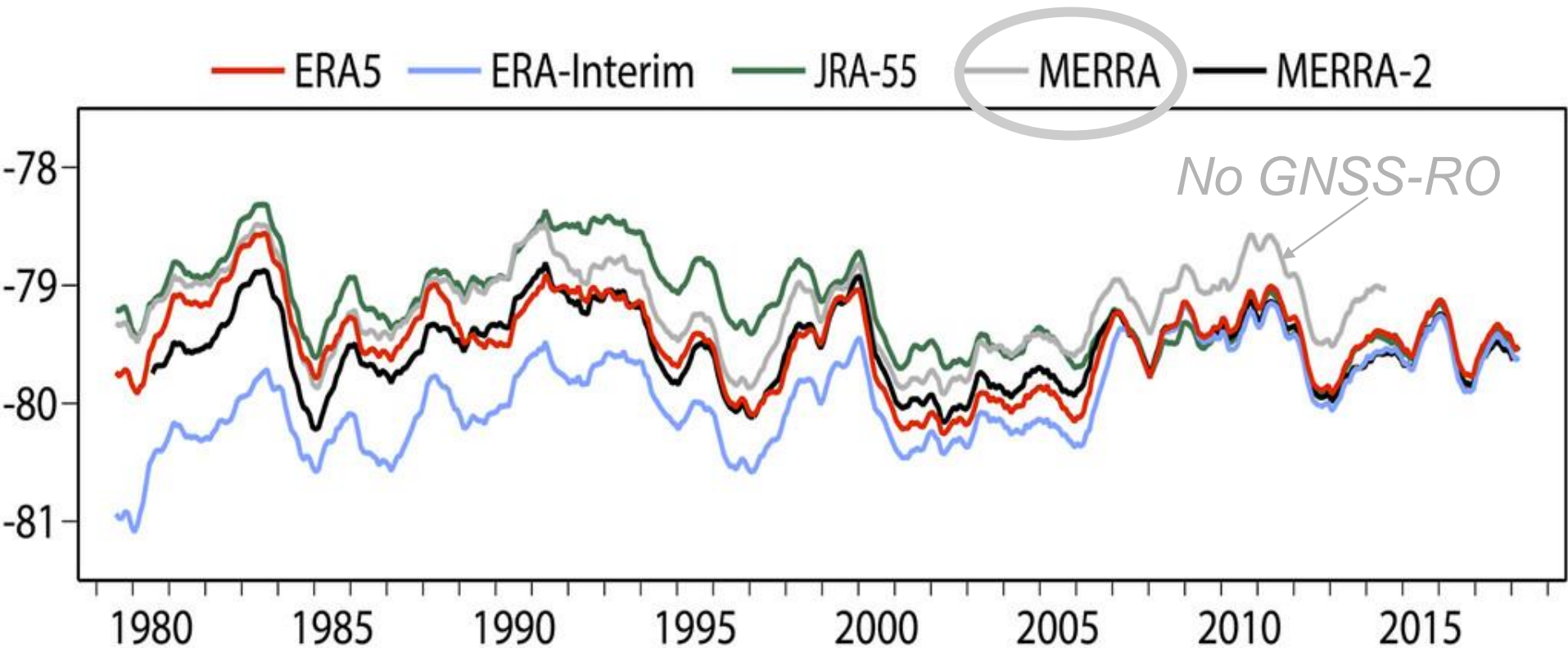
**GPS-RO have improved the consistency between climate reanalyses in the upper-troposphere and lower/middle stratosphere since 2006 with the introduction of COSMIC**

Compare ERA-Interim, JRA-55, MERRA, MERRA2, ERA5 reanalysis

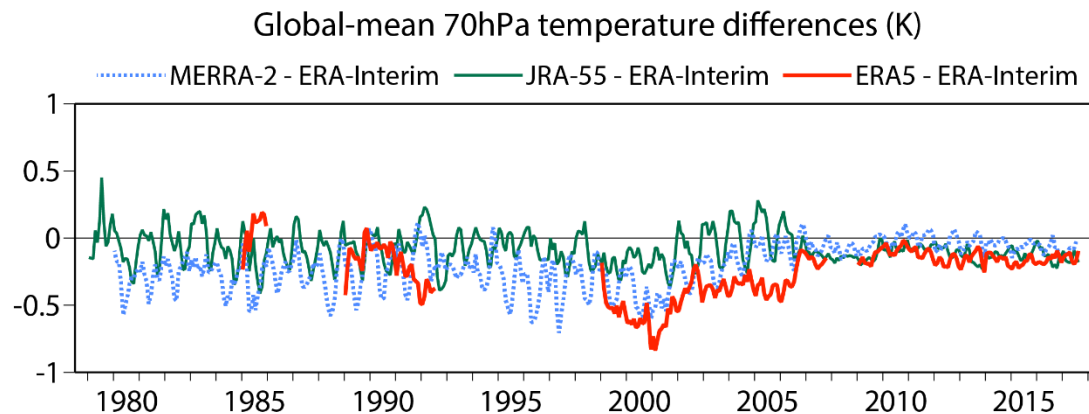
Twelve-month running mean temperature ( $^{\circ}\text{C}$ ) at 100 hPa averaged over the tropics ( $20^{\circ}\text{S}$  to  $20^{\circ}\text{N}$ ) from five global reanalyses.



Twelve-month running mean temperature ( $^{\circ}\text{C}$ ) at 100 hPa averaged over the tropics ( $20^{\circ}\text{S}$  to  $20^{\circ}\text{N}$ ) from five global reanalyses.



# Lower stratospheric global temperature bias in ERA5 (corrected in ERA5.1)



*The version of the assimilating model used for ERA5 has a larger cold bias in the lower stratosphere than the version used for ERA-Interim.*

*The cold bias is controlled by assimilating GPSRO data.*

*Radiosonde data exert a less-effective control on bias in ERA5 than they do in ERA-Interim.*

*See ERA 5.1 Tech Memo 859*

*<https://www.ecmwf.int/en/publications/technical-memoranda>*

# An indirect impact of GPS-RO on stratospheric humidity in reanalyses

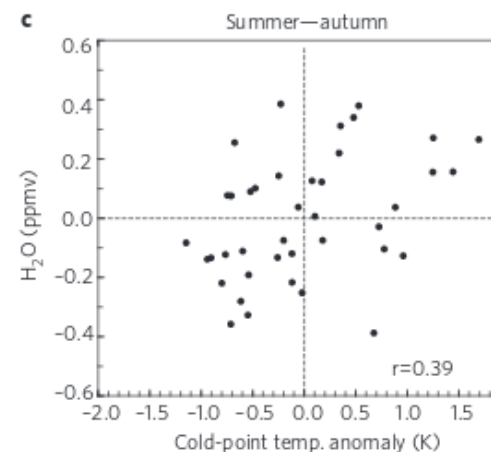
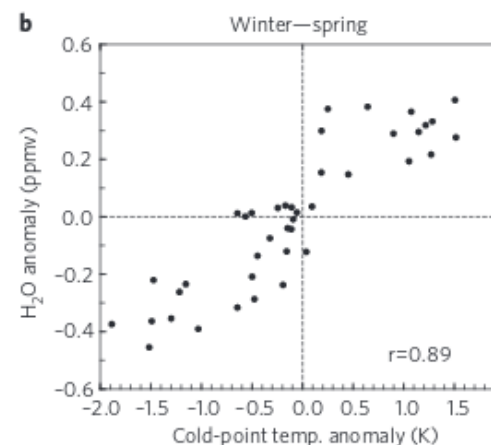
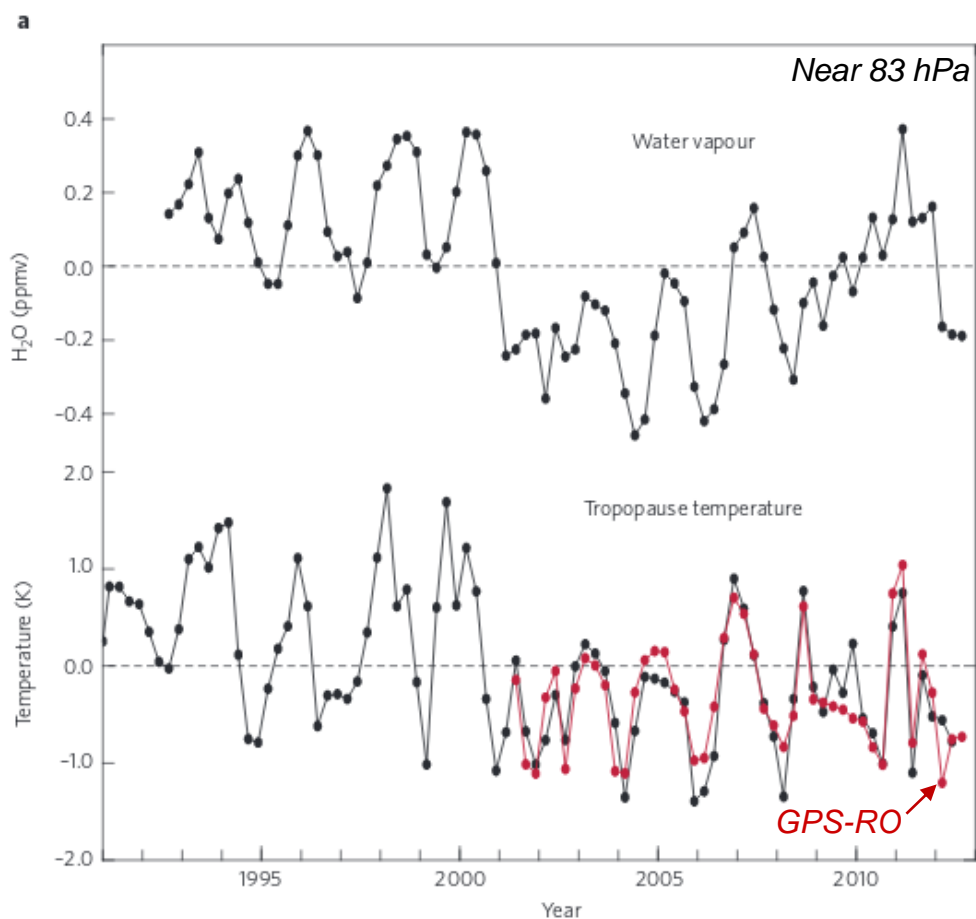
- Recall from lecture 1, that the stratospheric humidity is set to 0 in the “classical” temperature retrieval:

$$N = 10^6 (n - 1)$$
$$= \frac{c_1 P}{T} + \frac{c_2 P}{1 - \frac{P}{P_0}}$$

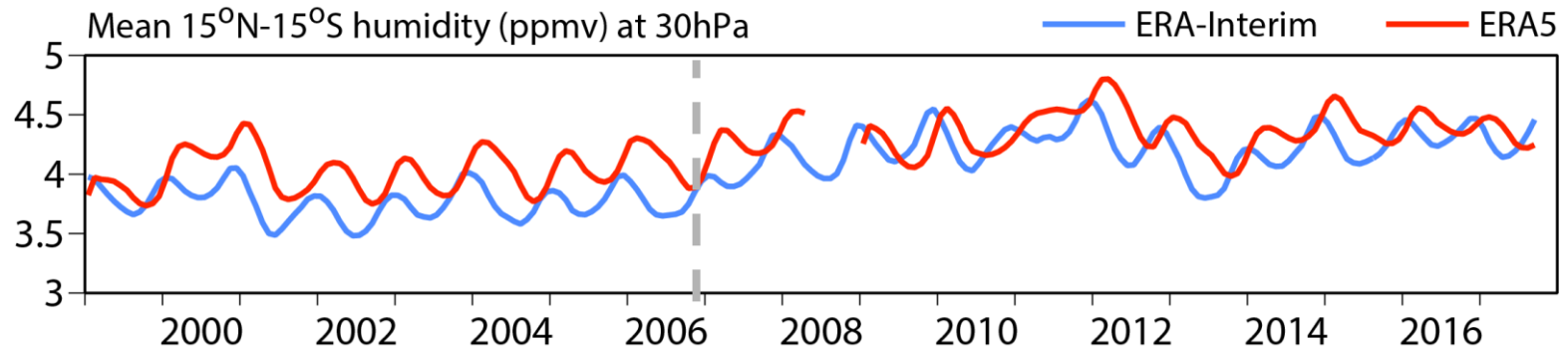
- This is reasonable because the contribution to the refractivity in the stratosphere from humidity is negligible. **The GPS-RO alone does not provide information about stratospheric humidity.**
- However, air enters the stratosphere primarily in the tropics (*The Brewer-Dobson Circulation*). The composition of the air is determined by the tropical tropopause layer (TTL).
- The air passing through the TTL is dehydrated at the cold point tropopause, leading to the extreme dryness in the stratosphere.

# Physical processes in the tropical tropopause layer and their roles in a changing climate

William J. Randel<sup>1\*</sup> and Eric J. Jensen<sup>2</sup>



# Tropical stratospheric humidity from ERA5/ERA-Interim



COSMIC is active in ERA-Int in Dec 2006.

COSMIC warms the tropical tropopause.

⇒ Moister stratosphere.

⇒ Gradient  $d(Q_s)/d(T_{cp}) \sim 0.5$  ppmv/K

See also ERA 5.1 Tech Memo 85

<https://www.ecmwf.int/en/publications/technical-memoranda>

# Comparing Q reanalyses in the stratosphere

Atmos. Chem. Phys., 17, 12743–12778, 2017  
 https://doi.org/10.5194/acp-17-12743-2017  
 © Author(s) 2017. This work is distributed under  
 the Creative Commons Attribution 3.0 License.



S. M. Davis et al.: Reanalysis UTLS water vapor and ozone assessment

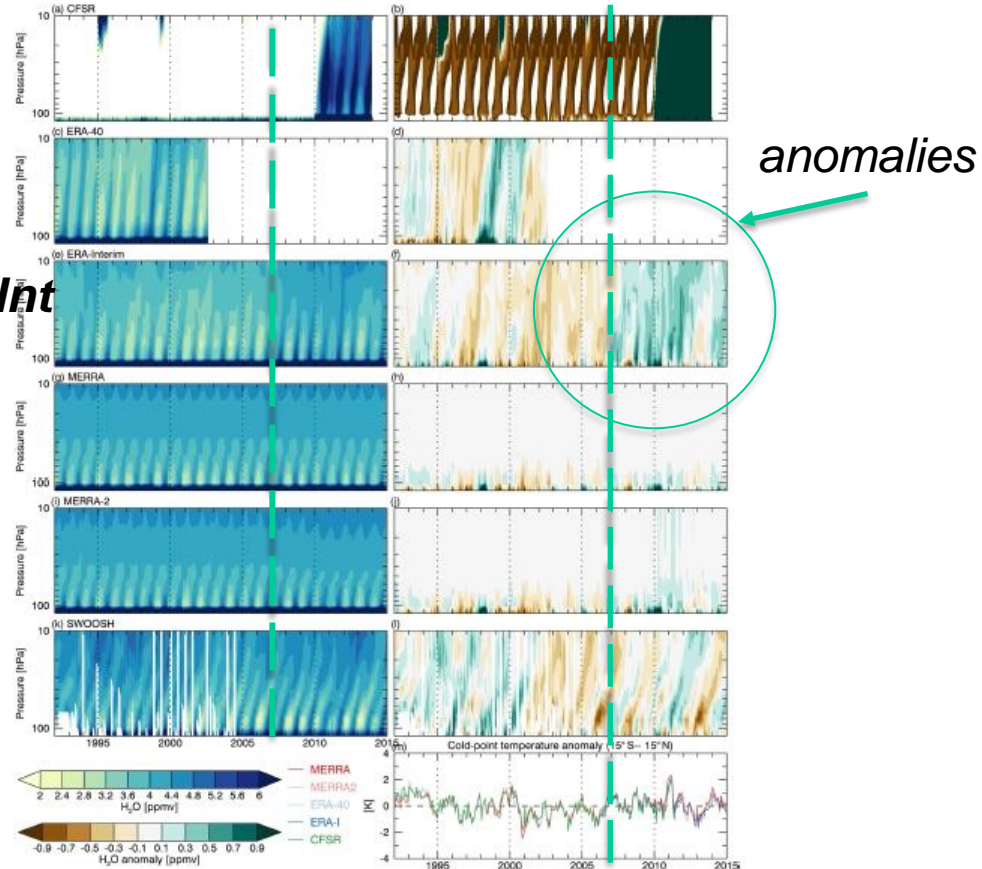
12769

## Assessment of upper tropospheric and stratospheric water vapor and ozone in reanalyses as part of S-RIP

Sean M. Davis<sup>1,2</sup>, Michaela I. Hegglin<sup>3</sup>, Masatomo Fujiwara<sup>4</sup>, Rossana Dragani<sup>5</sup>, Yayoi Harada<sup>6</sup>,  
 Chiaki Kobayashi<sup>6,7</sup>, Craig Long<sup>8</sup>, Gloria L. Manney<sup>9,10</sup>, Eric R. Nash<sup>11</sup>, Gerald L. Potter<sup>12</sup>, Susann Tegtmeier<sup>13</sup>,  
 Tao Wang<sup>14</sup>, Krzysztof Wargan<sup>11,15</sup>, and Jonathon S. Wright<sup>16</sup>

**ERA-Int**

*Note the SPARC community emphasise that reanalysis stratospheric humidity values should be used with caution, although ERA-Interim is described as “... surprisingly reasonable ...”*



**Figure 15.** The tropical tape recorder signal as represented in reanalyses and the SWOOSH merged satellite product, defined as the height-time evolution of water vapor averaged over the 15°S–15°N tropical band. Both absolute values (a) and anomalies relative to the mean water vapor seasonal cycle at each level (b) are shown. Anomalies are computed separately for each data set. Monthly mean anomalies in tropical (15°S–15°N) cold-point tropopause temperatures calculated from 6 h data on the native vertical resolution of each reanalysis model are shown for context (m).



# Climate monitoring applications

*GNSS-RO is becoming more important for climate monitoring, as the observation time-series lengthens*

***But which variables should we monitor?***

*Bending angles or more geophysical quantities?  
Recall,*

*'Satellites do not measure temperature, ...'*

## Recall basic GPS-RO processing chain:

- Excess phase delays.
- Doppler shift.
- Bending angle.
- Refractivity.
- Pressure/Temp. Geopotential height.

# The RoTrends Project

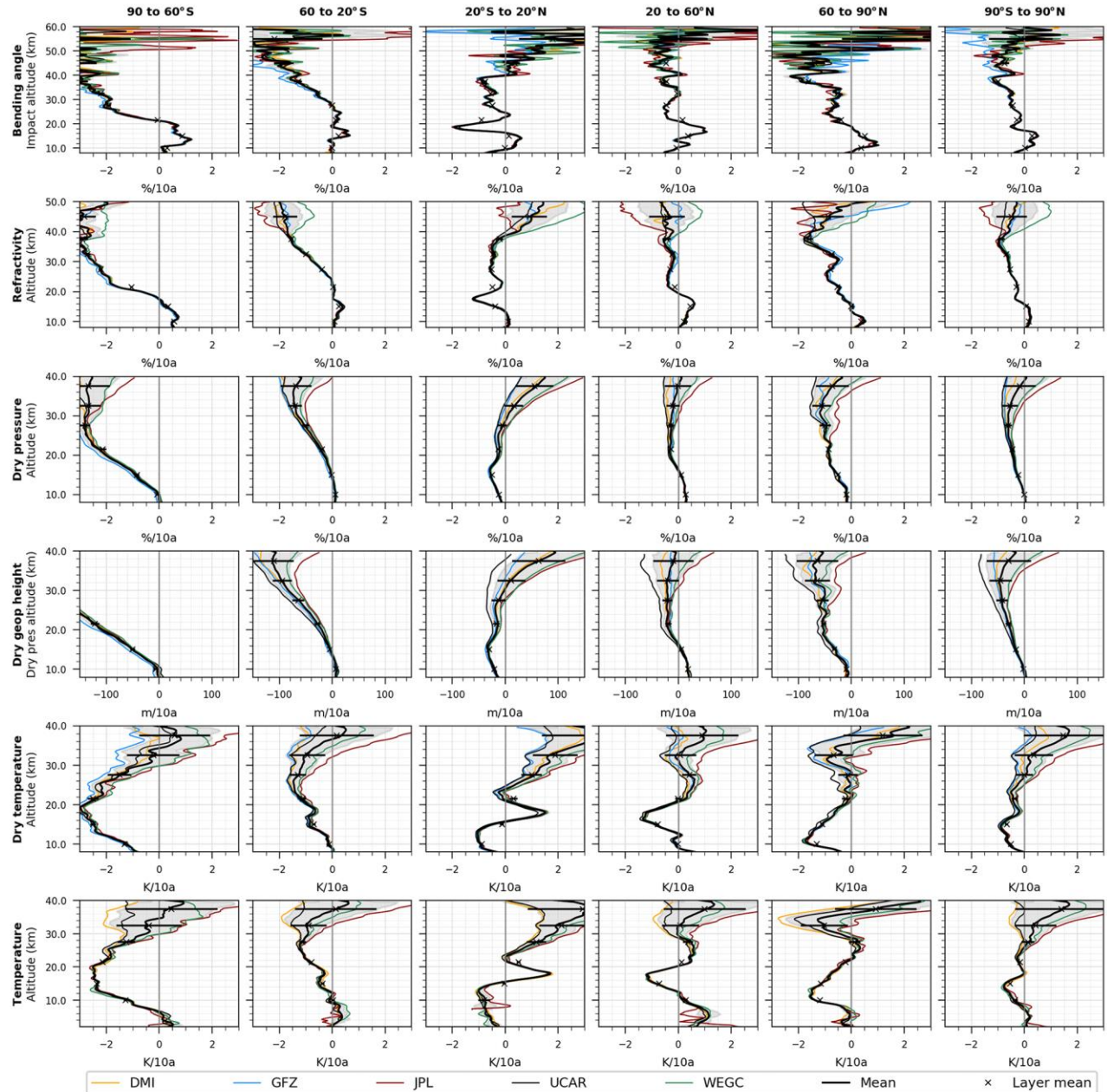
## *ROtrends* collaboration

RO community started comparison of different processing centres in 2007 (*ROtrends*).

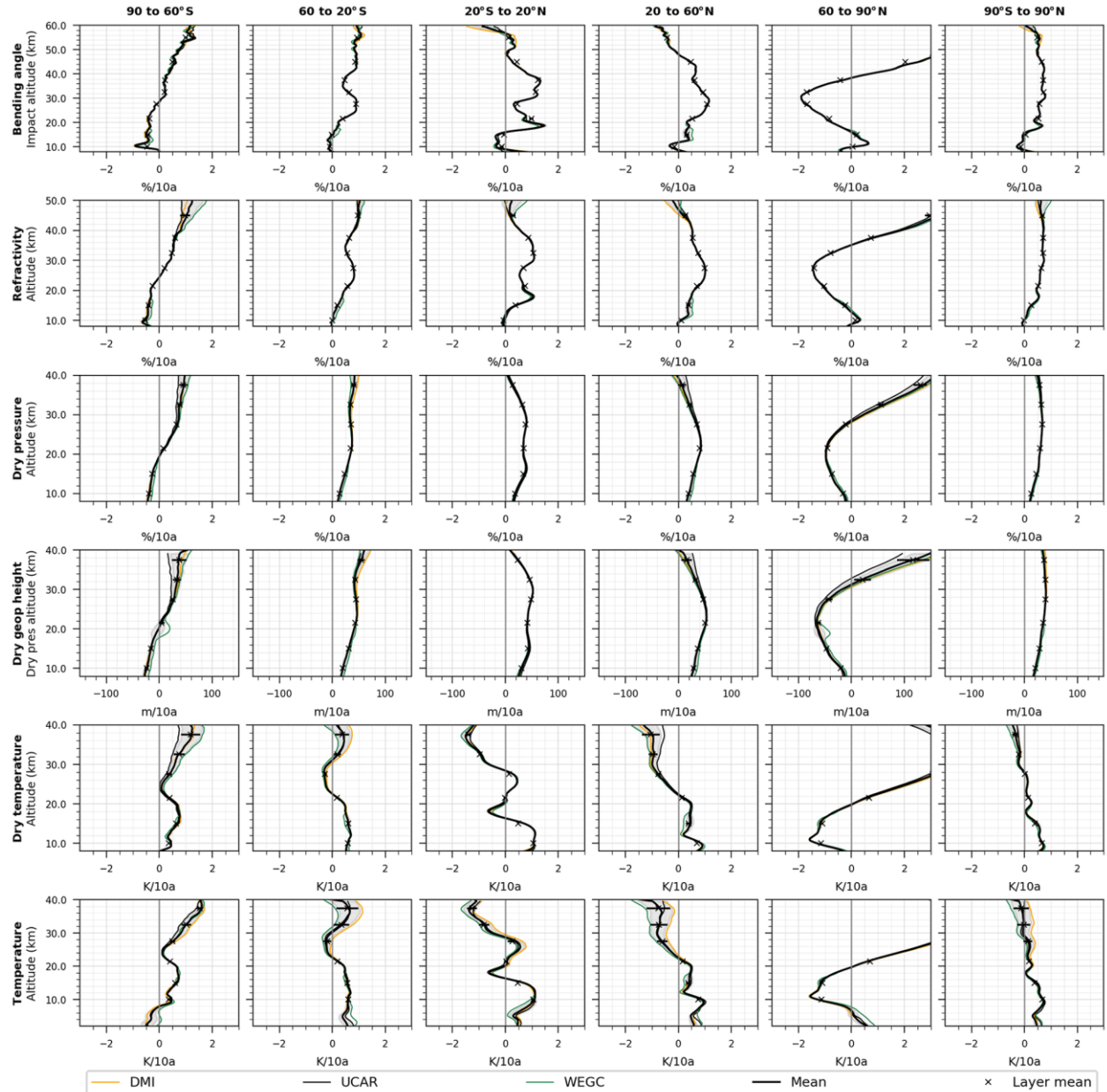
Main aim is to *validate* RO as a climate benchmark, identifying the impact of processing assumptions (*structural uncertainty*).

- *ROtrends* partners: DMI, JPL, GFZ, UCAR, WEGC, and EUMETSAT
- Common focus on CHAMP data, Aug 2001 to Sep 2008
- Aiming at improved understanding of ***structural uncertainty***, whilst still keeping the algorithm/software development independent
- Some recent results described in *Steiner et al. [2020]*
  - <https://amt.copernicus.org/articles/13/2547/2020/>

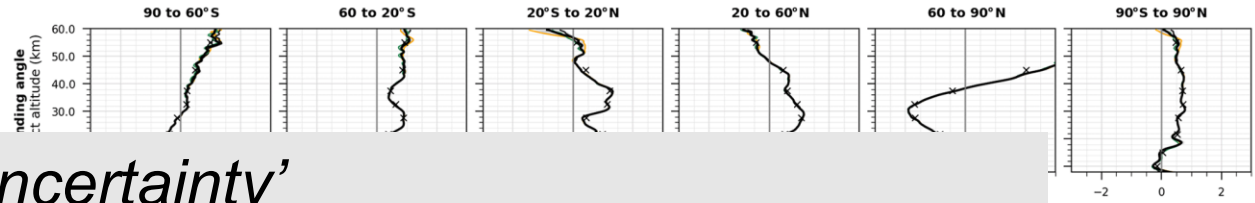
# Steiner et al, Figure 8 (Champ)



# Steiner et al, Figure 10 (Metop GRAS)



# Steiner et al, Figure 10 (Metop GRAS)



*'Structural uncertainty'*

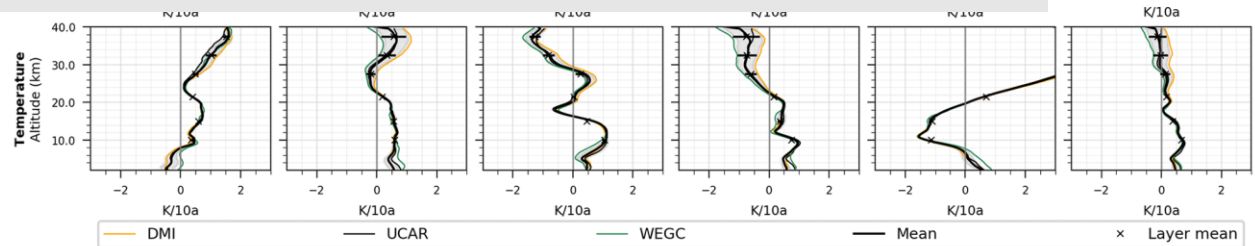
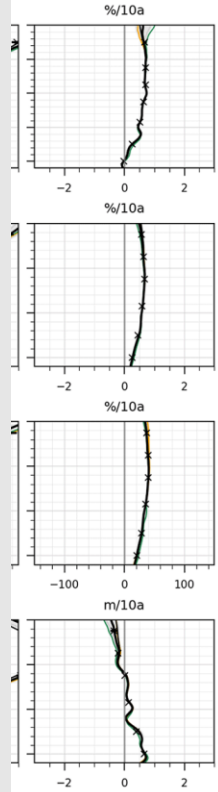
*The trends from the centres diverge as we move to more geophysical parameters because of different assumptions made in the processing*

*Illustrates the reliance on a-priori!*

*'Satellites do not measure temperature ...'*

Solid black  
mean

Gray = stan  
mean



— DMI — UCAR — WEGC — Mean × Layer mean

# Bending angle for climate monitoring

## Simulation study using the Hadley Centre climate model

### **Simulation studies to assess:**

- potential of GNSS-RO for detecting climate trends
- what variable should we monitor
- information content of GNSS-RO in relation to other sensors

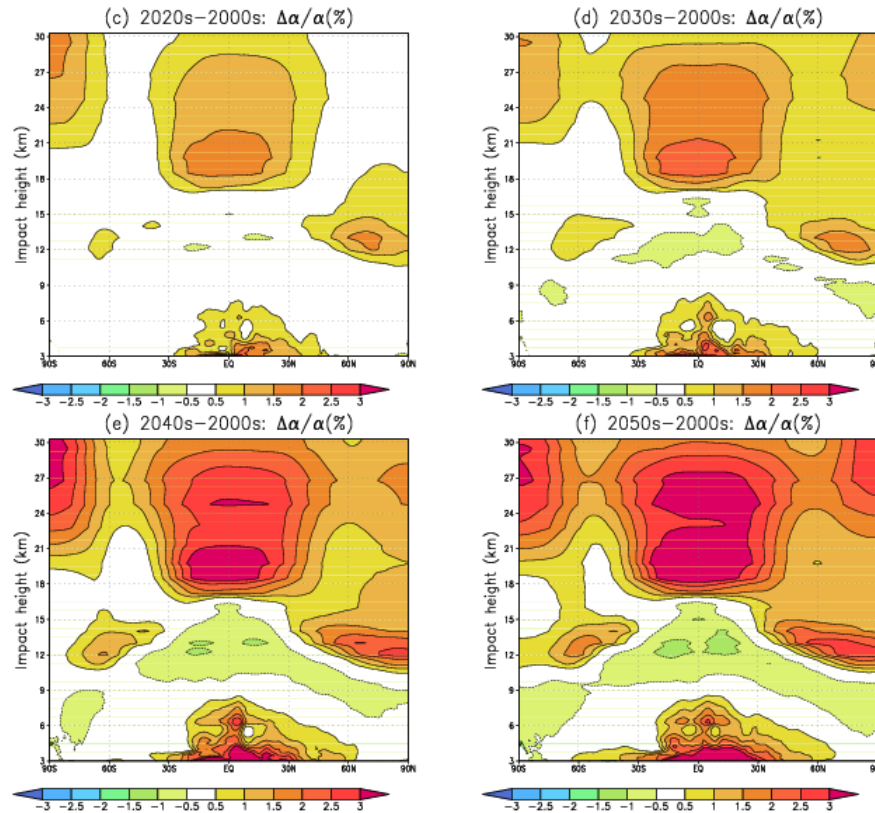
### **Simulations use:**

- Met Office Hadley Centre coupled climate model (HadGEM1)
- Climate change scenario (A1B) for 2000 – 2100
- Forward modelling of the GNSS-RO bending angles

*Provided by Mark Ringer (Hadley Centre)*

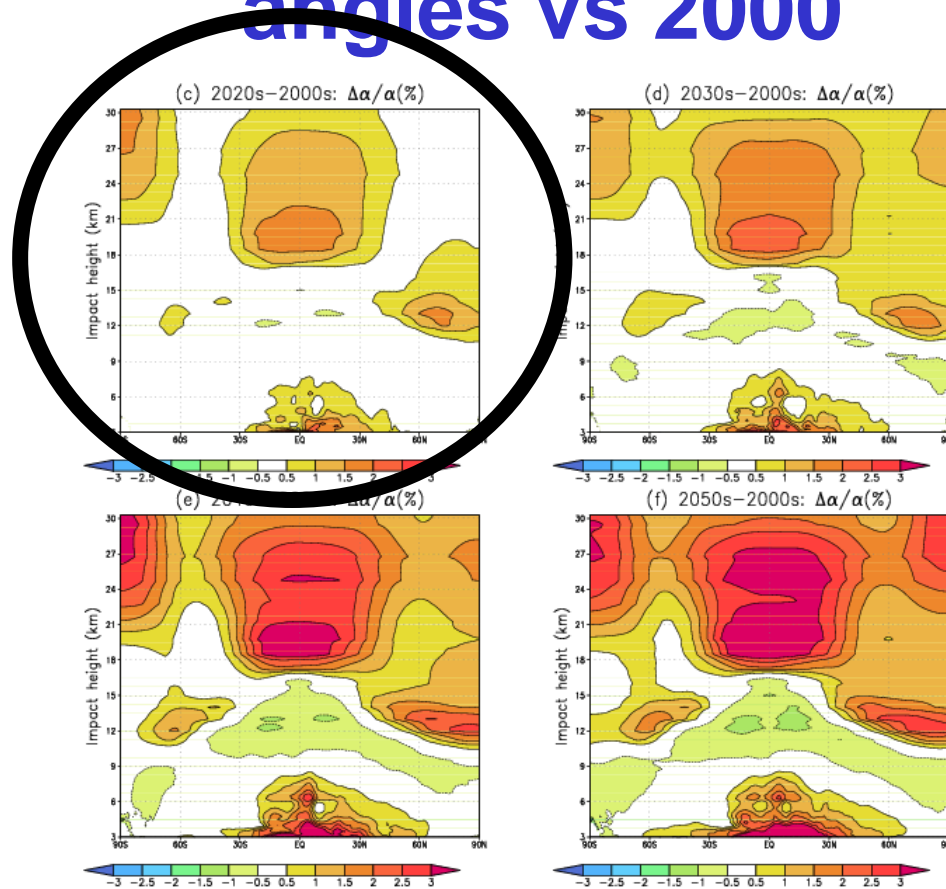


# Change in zonally averaged bending angles vs 2000



[https://www.romsaf.org/Publications/articles/2007\\_ringer-healy\\_GL032462.pdf](https://www.romsaf.org/Publications/articles/2007_ringer-healy_GL032462.pdf)

# Change in zonally averaged bending angles vs 2000

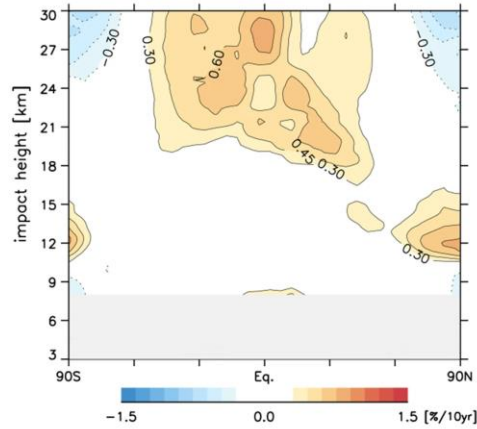


[https://www.romsaf.org/Publications/articles/2007\\_ringer-healy\\_GL032462.pdf](https://www.romsaf.org/Publications/articles/2007_ringer-healy_GL032462.pdf)

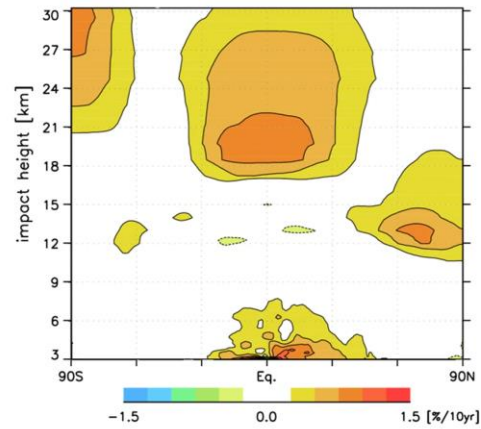
# ROM SAF

# Simulated in 2006

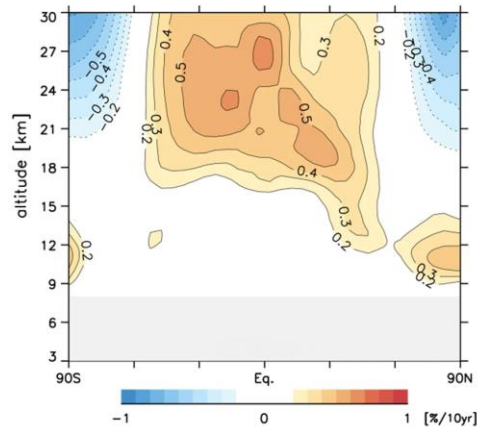
(a) RO bending angle trends



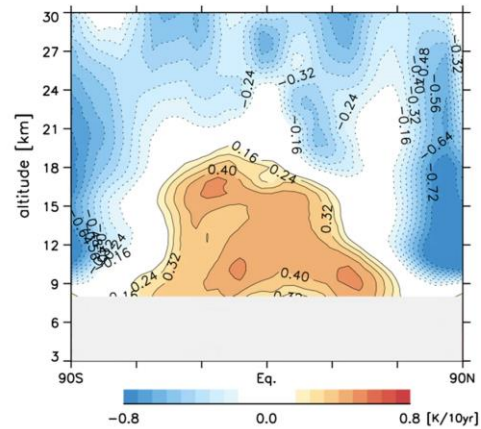
(b) HadGEM1 bending angle trends



(c) RO refractivity trends



(d) RO temperature trends

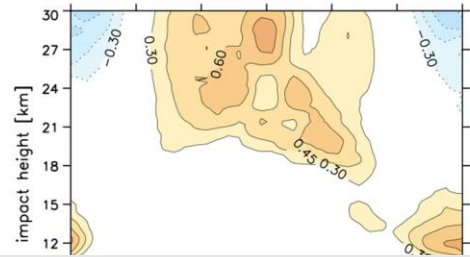


<https://www.nature.com/articles/s41612-022-00229-7>

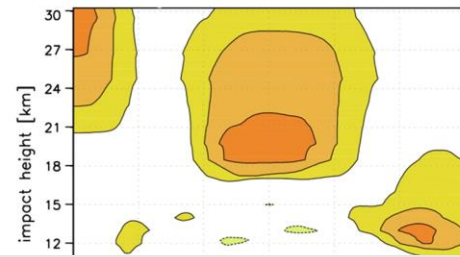
# ROM SAF

# Simulated in 2006

(a) RO bending angle trends

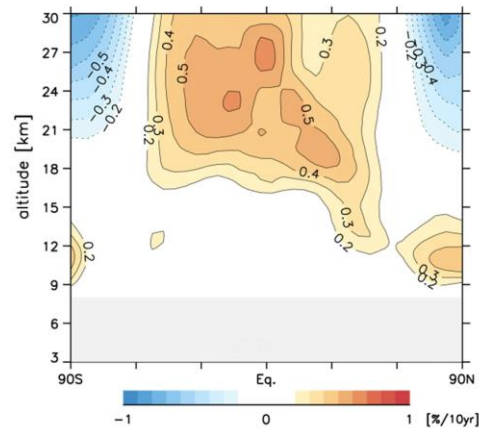


(b) HadGEM1 bending angle trends

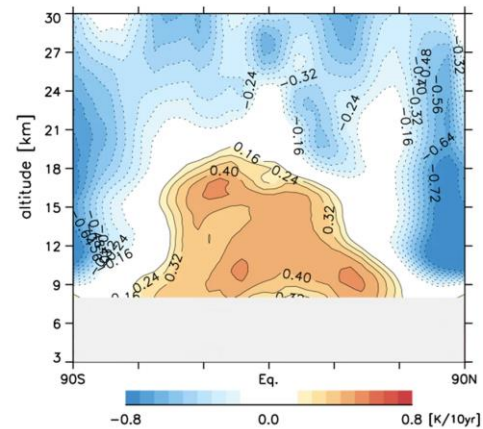


*Similar pattern – about (3/4) of magnitude*

(c) RO refractivity trends



(d) RO temperature trends



# Problem with monitoring bending angles

- More difficult to interpret than geophysical quantities.
- Most climate related work looks at temperature/geopotential heights.

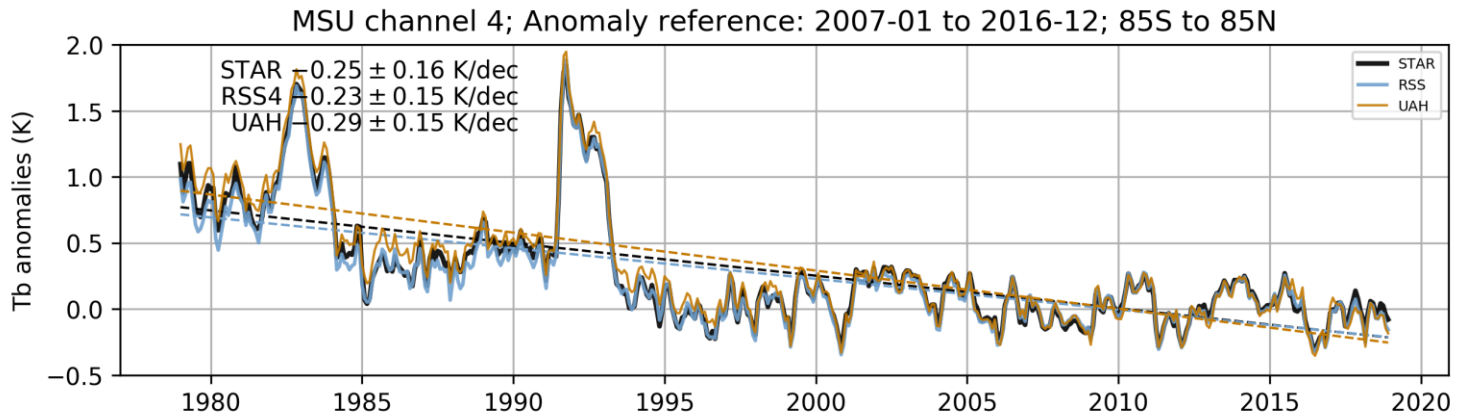
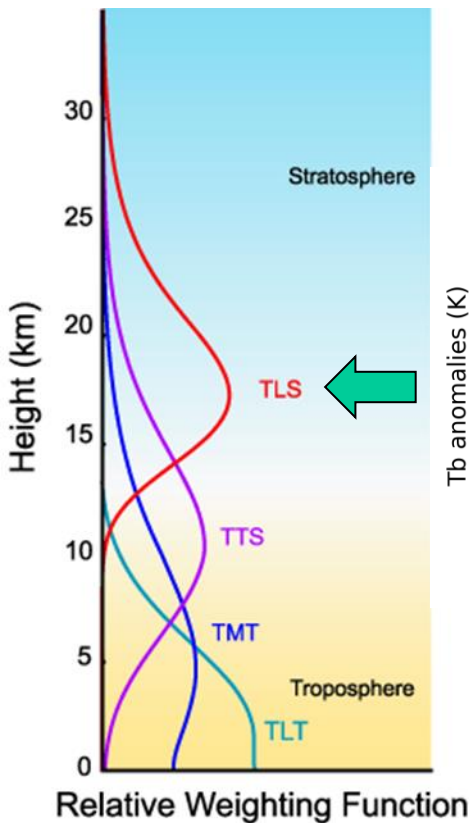
# Contribution to IPCC AR6

<https://www.ipcc.ch/assessment-report/ar6/>

*Compare trends retrieved from GNSS-RO with other observations?*

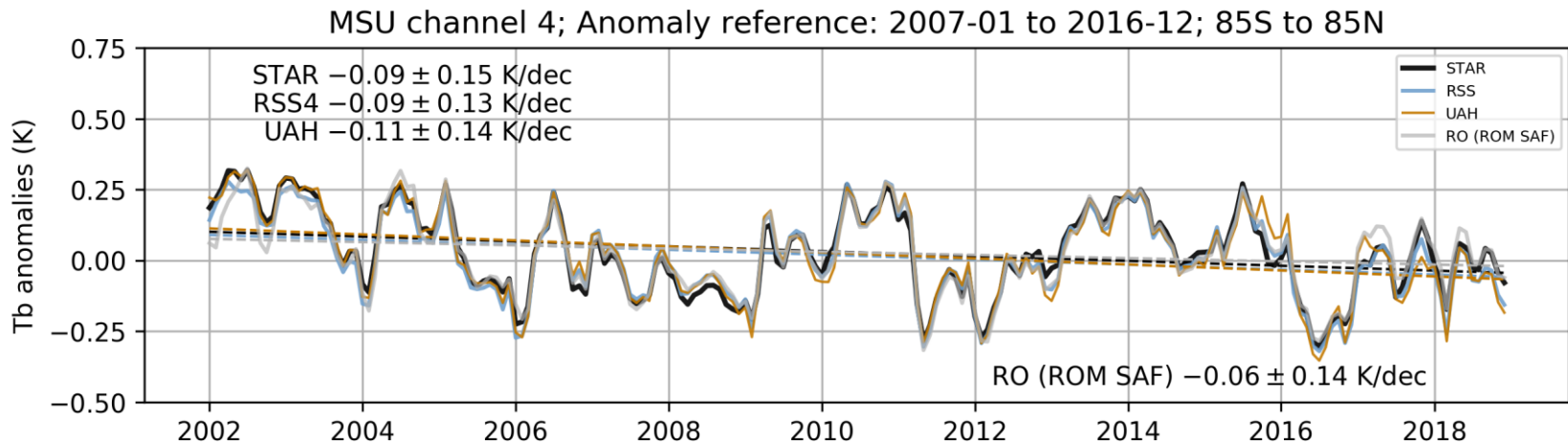
*How do temperature trends in the tropics vary with height? Do the climate models look reasonable?*

# Comparing GPS-RO with MSU radiances (Global)



*We can forward model GNSS-RO anomalies to MSU ch 4 (AMSU ch 9, TLS) brightness temperatures from 2002 (CHAMP).*

# MSU-4 and GPS-RO anomalies 2002-2018

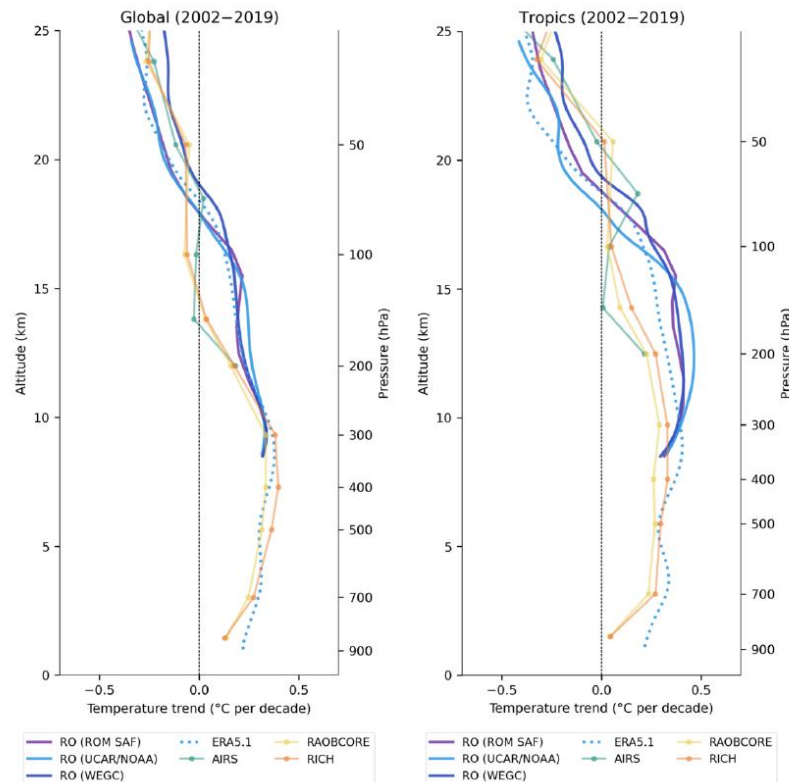


*Good consistency between MSU-4 and GNSS-RO from 2002.*



# Contribution to the IPCC AR6 report

– observed temperature trends in the tropical upper troposphere –



Decadal temperature trends as function of altitude from:

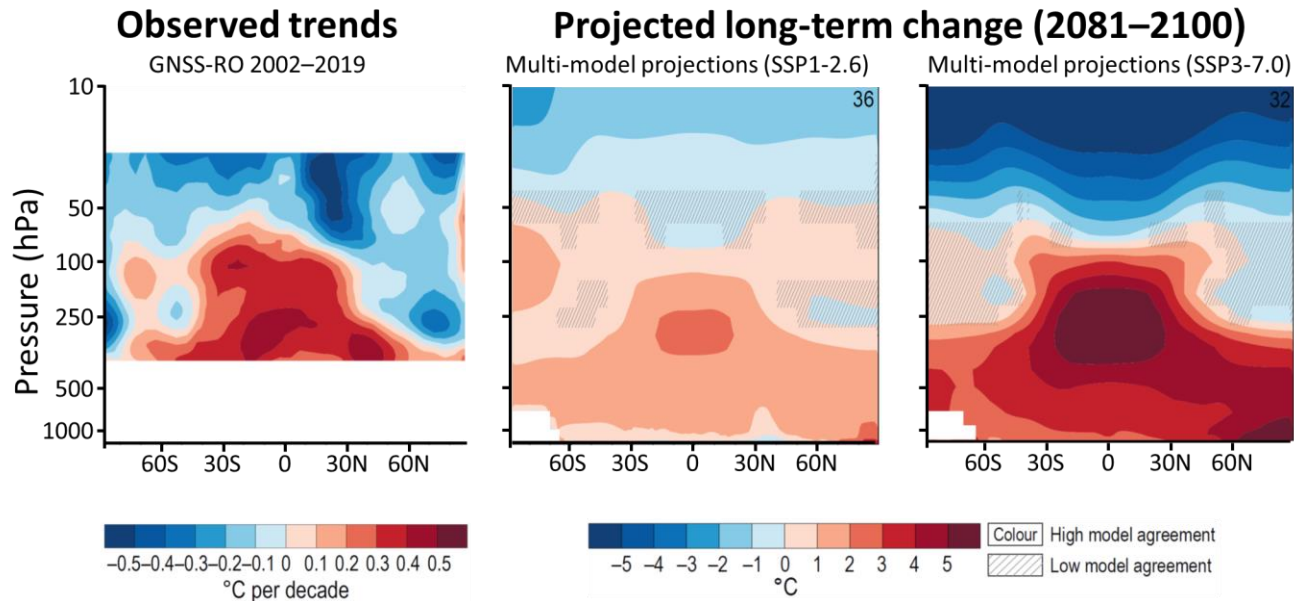
- RO data from: ROM SAF, UCAR/NOAA, Wegener Center, Graz (WEGC)
- Radiosonde datasets: RAOBCORE, RICH
- AIRS data
- ERA5 data

From ROM SAF VS40 report (Florian Ladstädter)

[https://www.romsaf.org/Publications/reports/romsaf\\_vs40\\_rep\\_v10.pdf](https://www.romsaf.org/Publications/reports/romsaf_vs40_rep_v10.pdf)

# Global upper air temperature trends

contribution to the IPCC AR6 WG1 report



From IPCC AR6 WG 1, Technical Summary

*Observed trends in ROM SAF RO data (left) in K/decade compared to projected temperature changes in CMIP6 models under a middle/low scenario (SSP1-2.6) and a middle-high scenario (SSP3-7.0).*

# Polarimetric RO (Slides, Estel Cardellach).

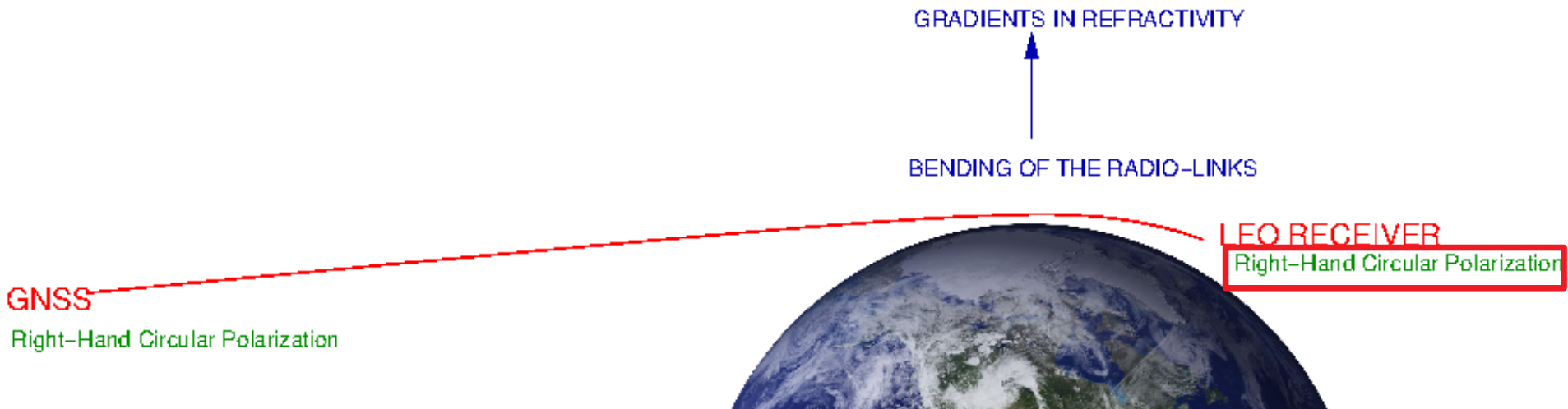
## Geophysical Research Letters

Research Letter |  Open Access |    

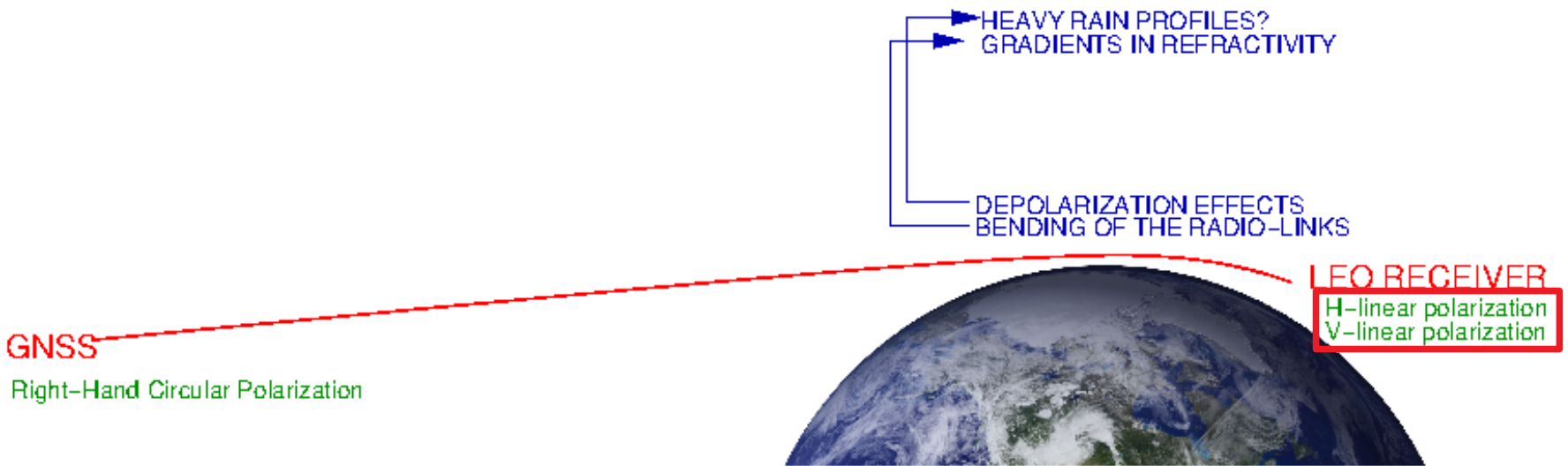
### Sensing Heavy Precipitation With GNSS Polarimetric Radio Occultations

E. Cardellach , S. Oliveras, A. Rius, S. Tomás, C. O. Ao, G. W. Franklin, B. A. Iijima, D. Kuang, T. K. Meehan, R. Padullés, M. de la Torre Juárez, F. J. Turk, D. C. Hunt, W. S. Schreiner ... [See all authors](#)

First published: 21 December 2018 | <https://doi.org/10.1029/2018GL080412>



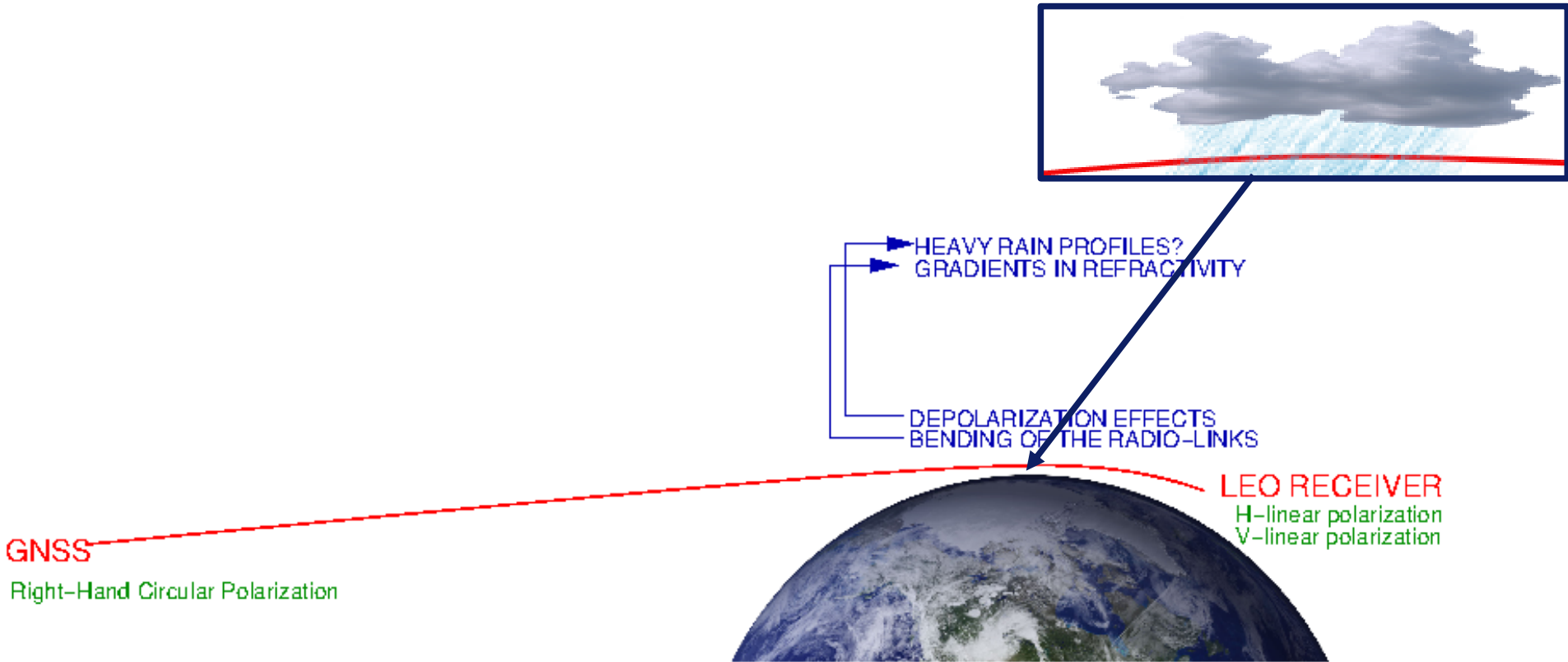
***'TYPICAL' GNSS RO PRODUCTS: VERTICAL PROFILES OF THERMODYNAMIC VARIABLES at the tangent point (typically temperature, pressure, humidity)***



***'NEW' GNSS-PRO PRODUCTS:***

*VERTICAL PROFILES OF THERMODYNAMIC VARIABLES (typically temperature, pressure, water vapor)*

***+ VERTICAL PROFILES OF INTENSE RAIN***



***'NEW' GNSS-PRO PRODUCTS:***

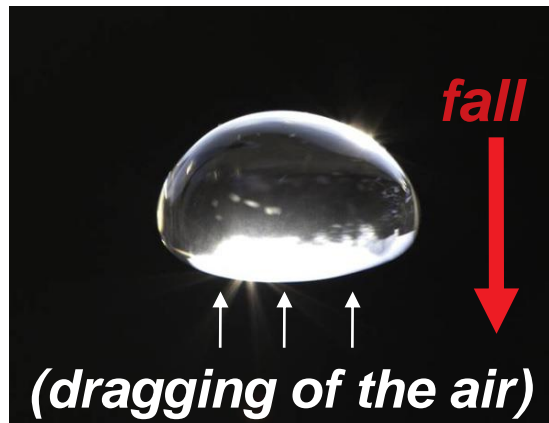
*VERTICAL PROFILES OF THERMODYNAMIC VARIABLES (typically temperature, pressure, water vapor)*

***+ VERTICAL PROFILES OF INTENSE RAIN***

***To understand this concept it is important to keep in mind that the big falling rain drops ARE NOT like this***



***but rather LIKE***

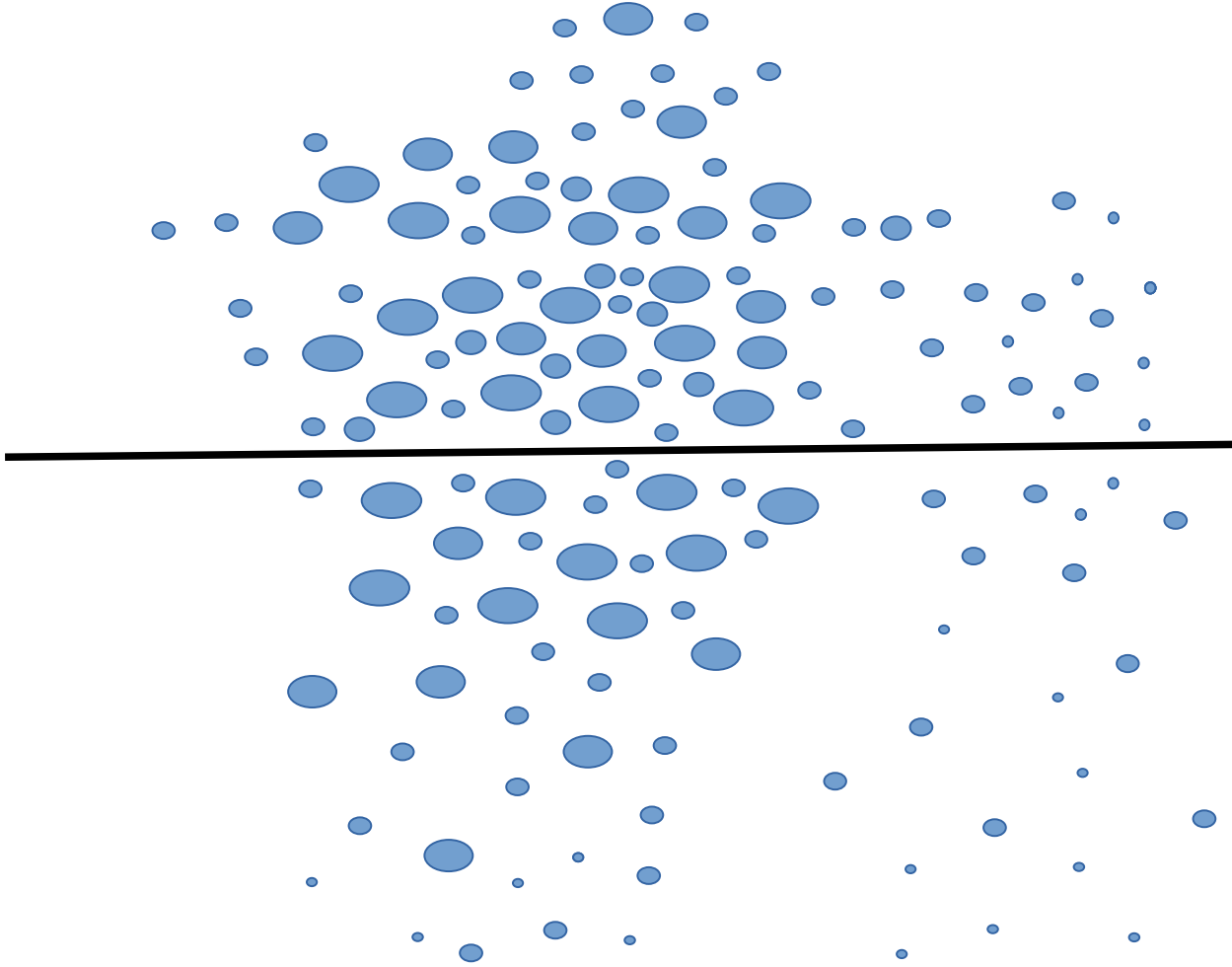


***Vertical dimension  
shorter than  
Horizontal  
dimension →  
different  
propagation delays***

*precipitation cell*

*GNSS*

*LEO*



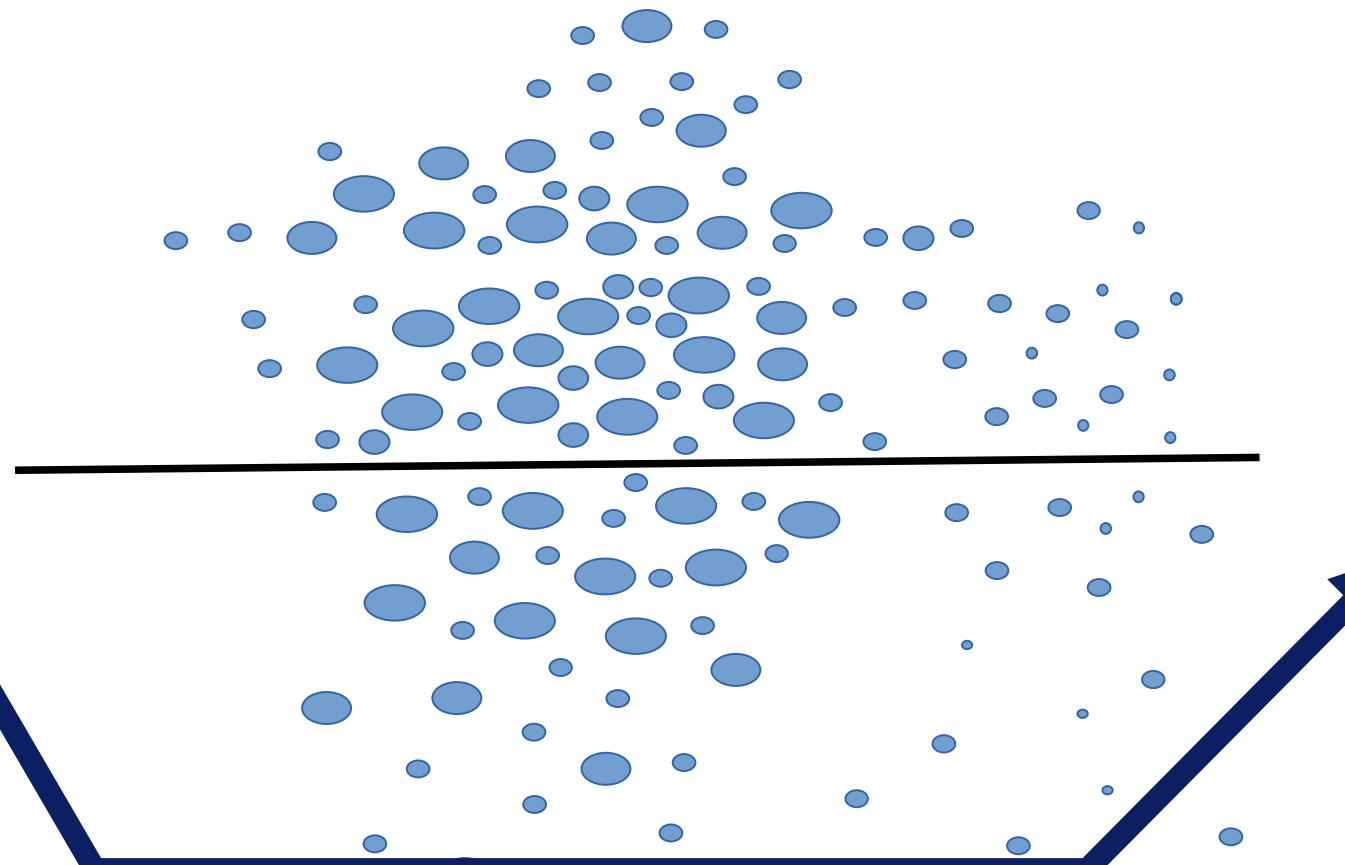


*precipitation cell*

*GNSS*

*LEO*

***Bistatic radar: transmitter and receiver at different locations***



*precipitation cell*

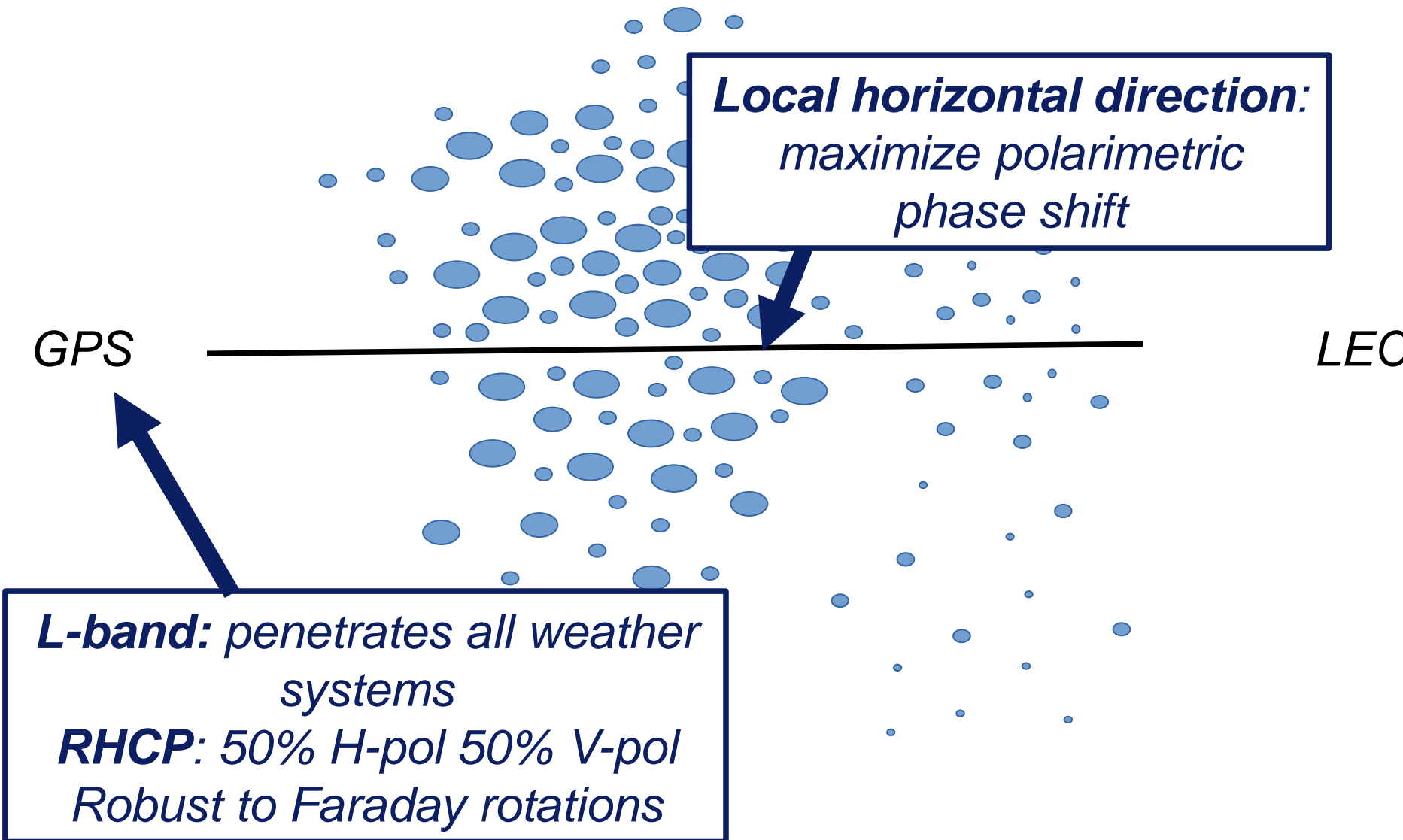
**Local horizontal direction:**  
*maximize polarimetric  
phase shift*

*GPS*

*LEO*

**L-band:** *penetrates all weather  
systems*

**RHCP:** *50% H-pol 50% V-pol  
Robust to Faraday rotations*



*precipitation cell*

$$\Delta\phi^{atm} = \int_L K_{dp}(l) dl$$

GNSS



LEO

***Kdp***

***Kdp***

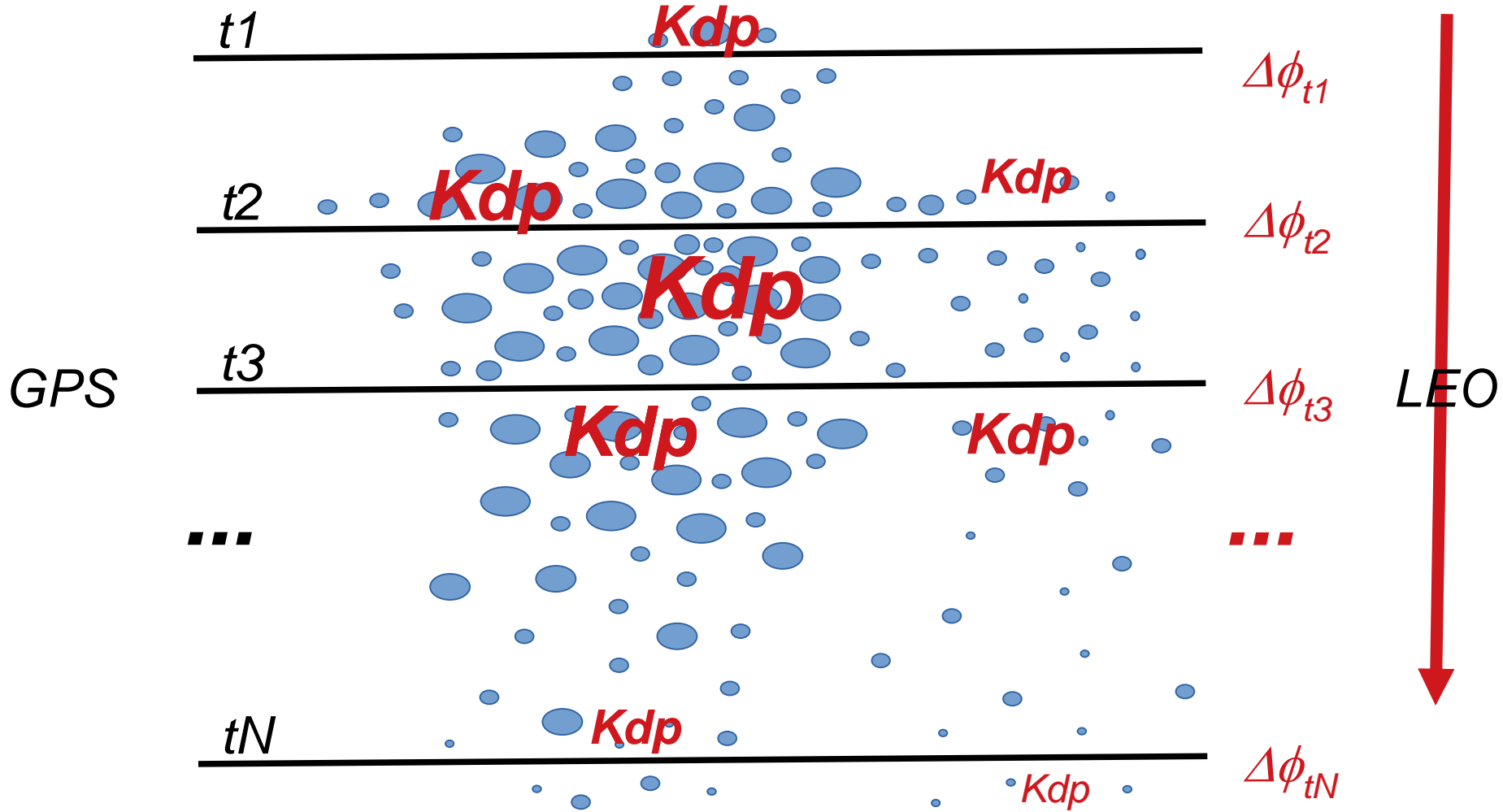
***Δφ***

**Observable:** horizontally integrated polarimetric phase shift (or polarimetric phase delay):

$$\Delta\phi = \phi_H - \phi_V$$

***Delay of H-pol longer than V-pol!***

*precipitation cell*



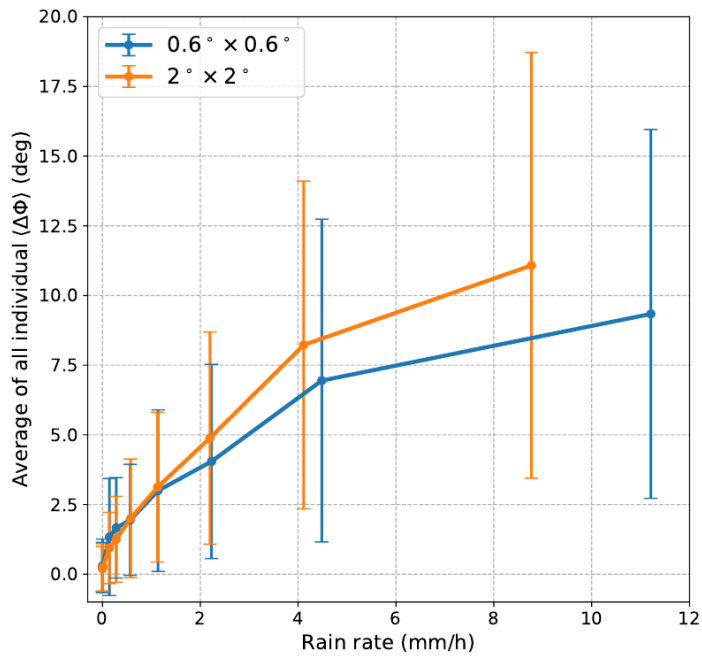
**Vertical scanning**

*Measurement concept being tested aboard the PAZ satellite (ROHP-PAZ experiment)*

*Sucessful launch on **February 22, 2018**, by SpaceX (Falcon9).*

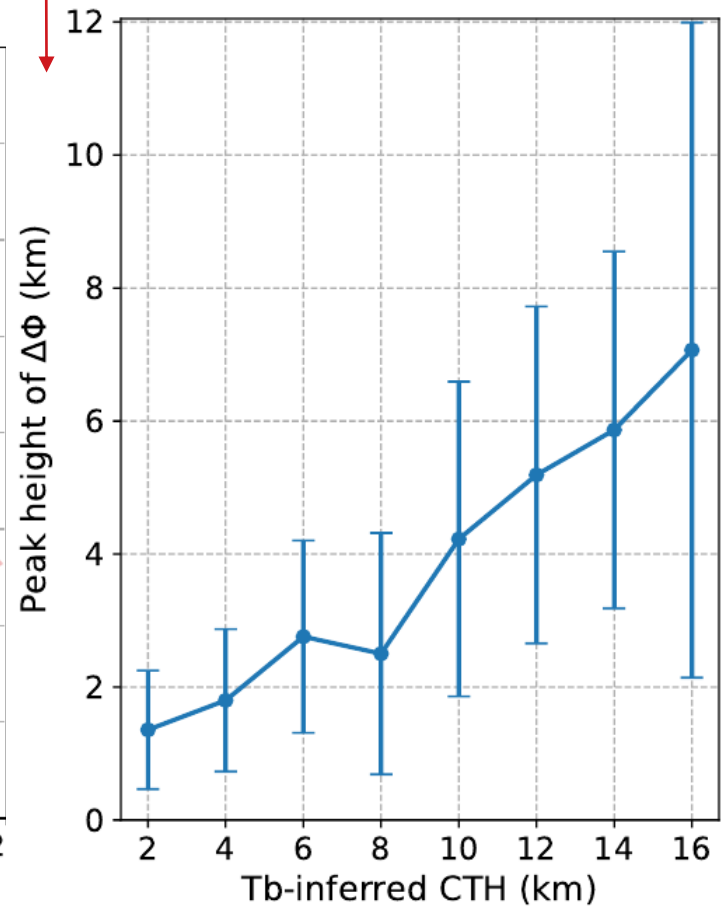
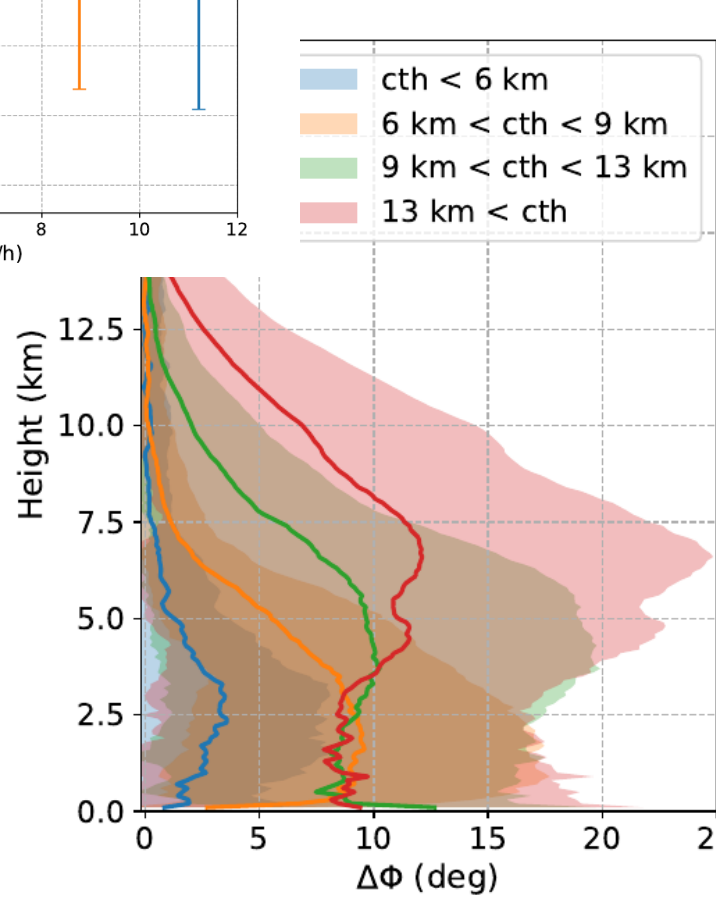
*GNSS RO experiment **activated on May 10, 2018**.*





*Rain rate vs polarimetric delay*

*polarimetric delay as function of altitude and Cloud Top Height (cth)*



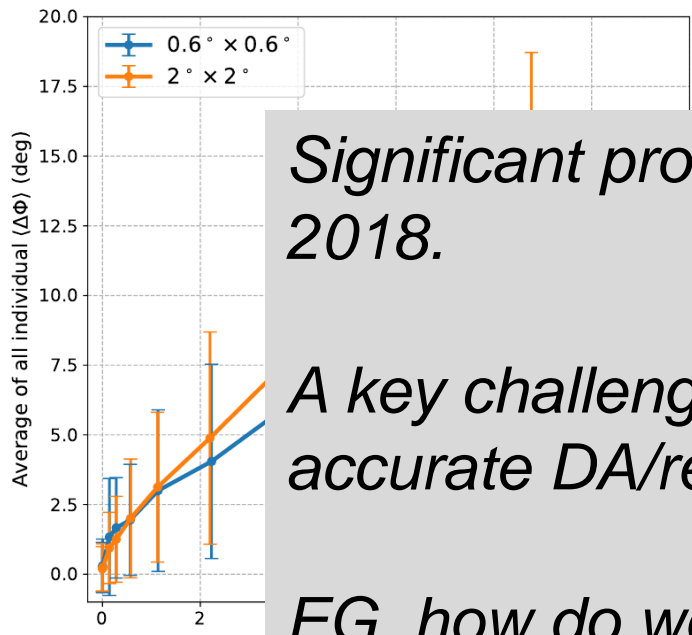
*Rain rate vs polarimetric delay*

*Significant progress since the launch of PAZ in 2018.*

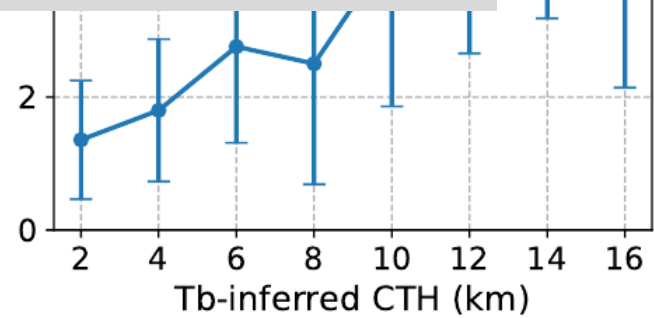
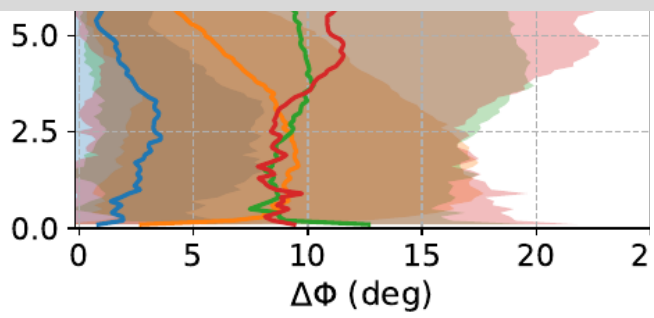
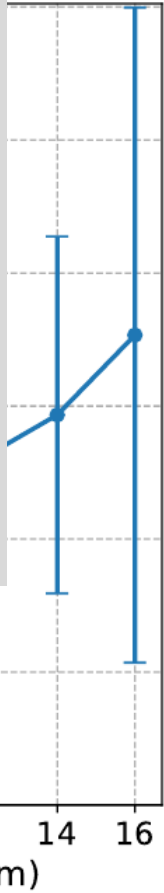
*A key challenge will be to demonstrate and accurate DA/retrieval approach.*

*EG, how do we distinguish between light rain over a long path or intense rain over a short path?*

*If we assimilated this data, modelling 2D aspects will be key*



*th)*



# PRO observable $\Phi_{DP}$ (differential phase shift) and its analogy to GNSS-RO bending angle

- PRO observable
- = integral of  $K_{DP}$  along the ray path:

$$\Phi_{DP} = \int_{GNSS}^{LEO} K_{DP}(s) ds$$

- RO observable
- = bending angle
- = integral of infinitesimal bending along the

ray path:

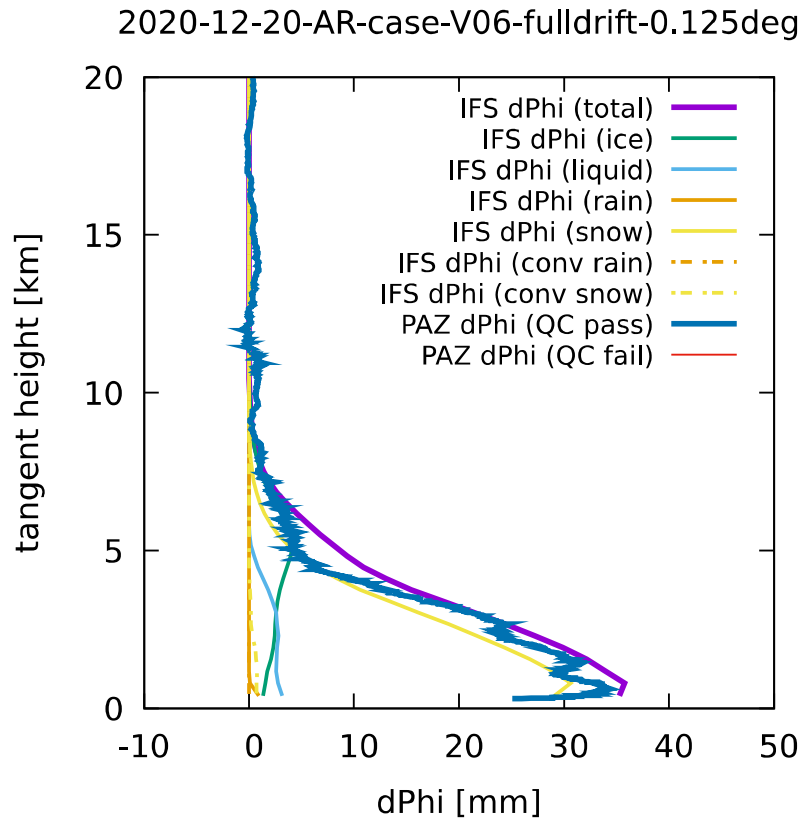
$$\alpha = \int_{GNSS}^{LEO} \left( \frac{d\alpha}{ds} \right) ds$$

→ Natural analogy between PRO and regular RO observable

→ can exploit the existing forward operator for RO bending angle

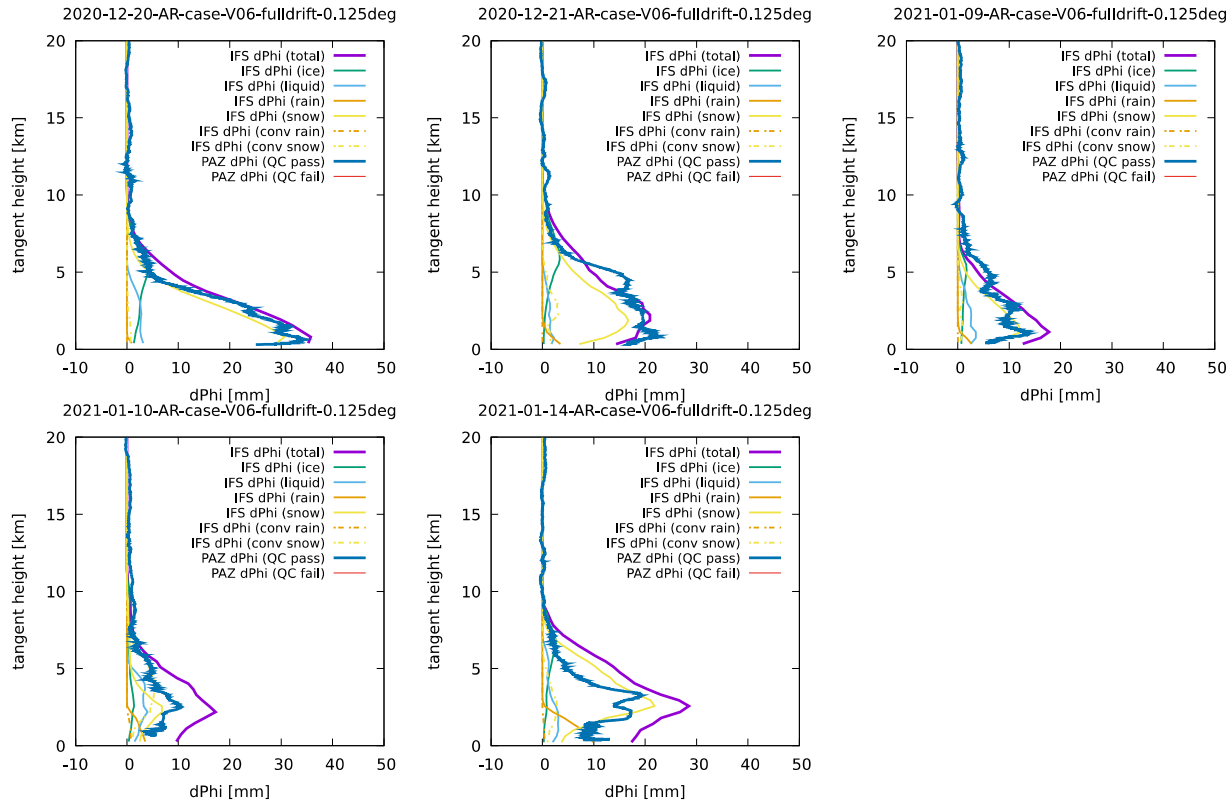


# Results: Overall agreement of simulated vs observed $\Phi_{DP}$



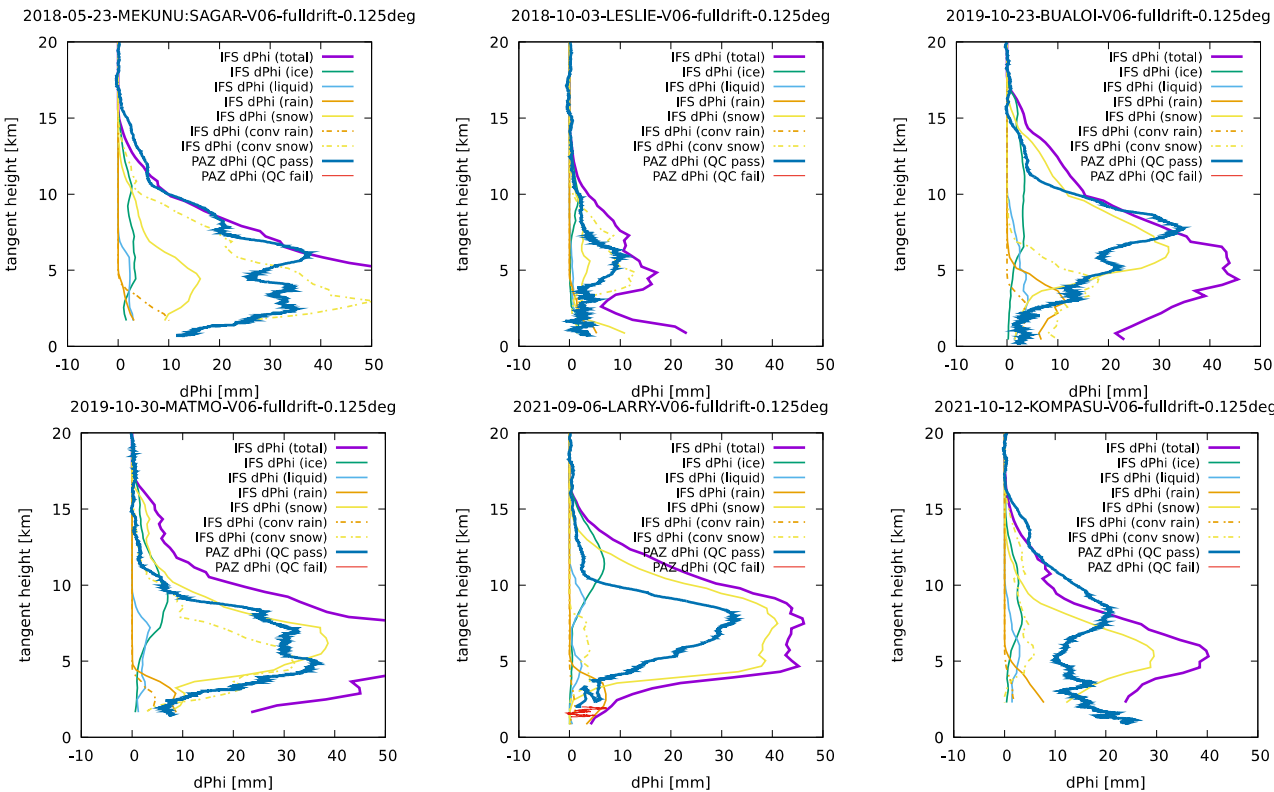
- Result for an Atmospheric River (AR) case
- Very good agreement between simulated total (blue) and observed (purple) profiles.
  - despite many simplifying assumptions!
- Resolved-scale snow (yellow solid) is the dominant contribution
  - which is a surprise, to be discussed later

# Results: Overall agreement of simulated vs observed $\Phi_{DP}$



- *Results for all 5 AR cases*
- *Very good agreement in all the cases, which is great!*
- *However....*

# Results: Overall agreement of simulated vs observed $\Phi_{DP}$



- Results for all TC cases show poor agreement between simulation and observation

- “Shape” of the profiles do not match

- Amplitude also systematically overestimated

- Why?

# Summary

- Given an overview of applications and pointed to published sources where possible.
- Recent impact on NWP performance
  - Impact on stratospheric winds in the tropics
- The GNSS-RO are now key observations for climate reanalyses and have led to improved consistency between reanalyses since 2006
- Climate monitoring with GNSS-RO is becoming increasingly important. Inclusion in the IPCC AR6 is an important step forward for the community
- Introduced the polarimetric RO concept, early work