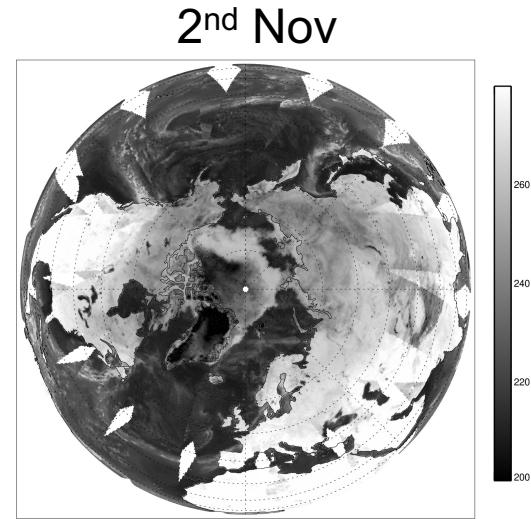
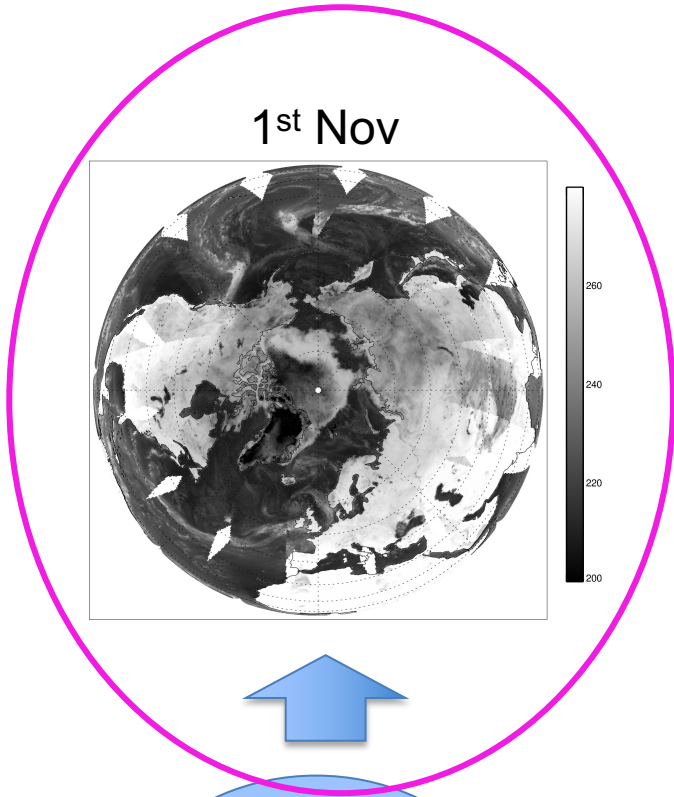


Microwave observations (part 2): cloud and precipitation; applications

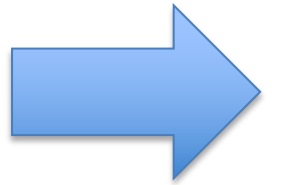
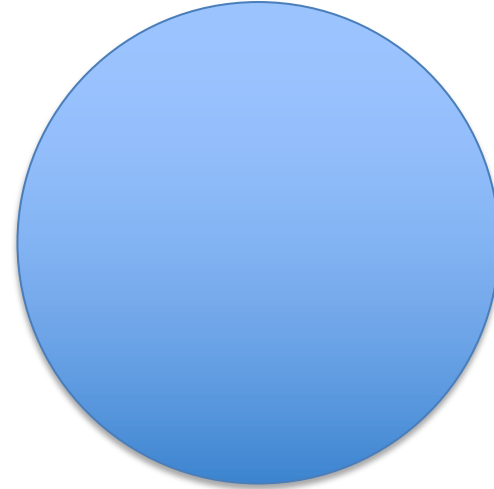
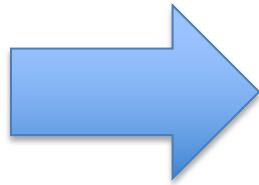
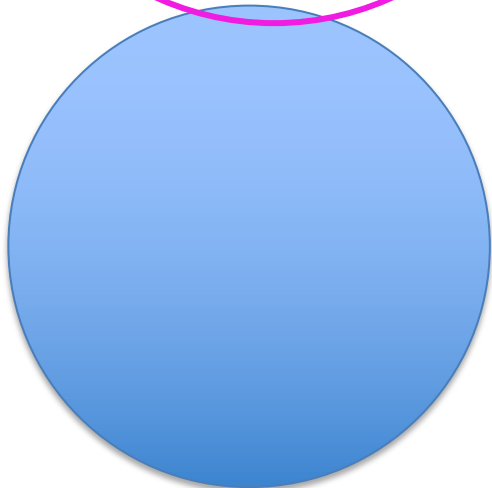
Alan Geer

EUMETSAT/ECMWF NWP-SAF satellite data assimilation training course, 15 – 19 May 2023

Observations

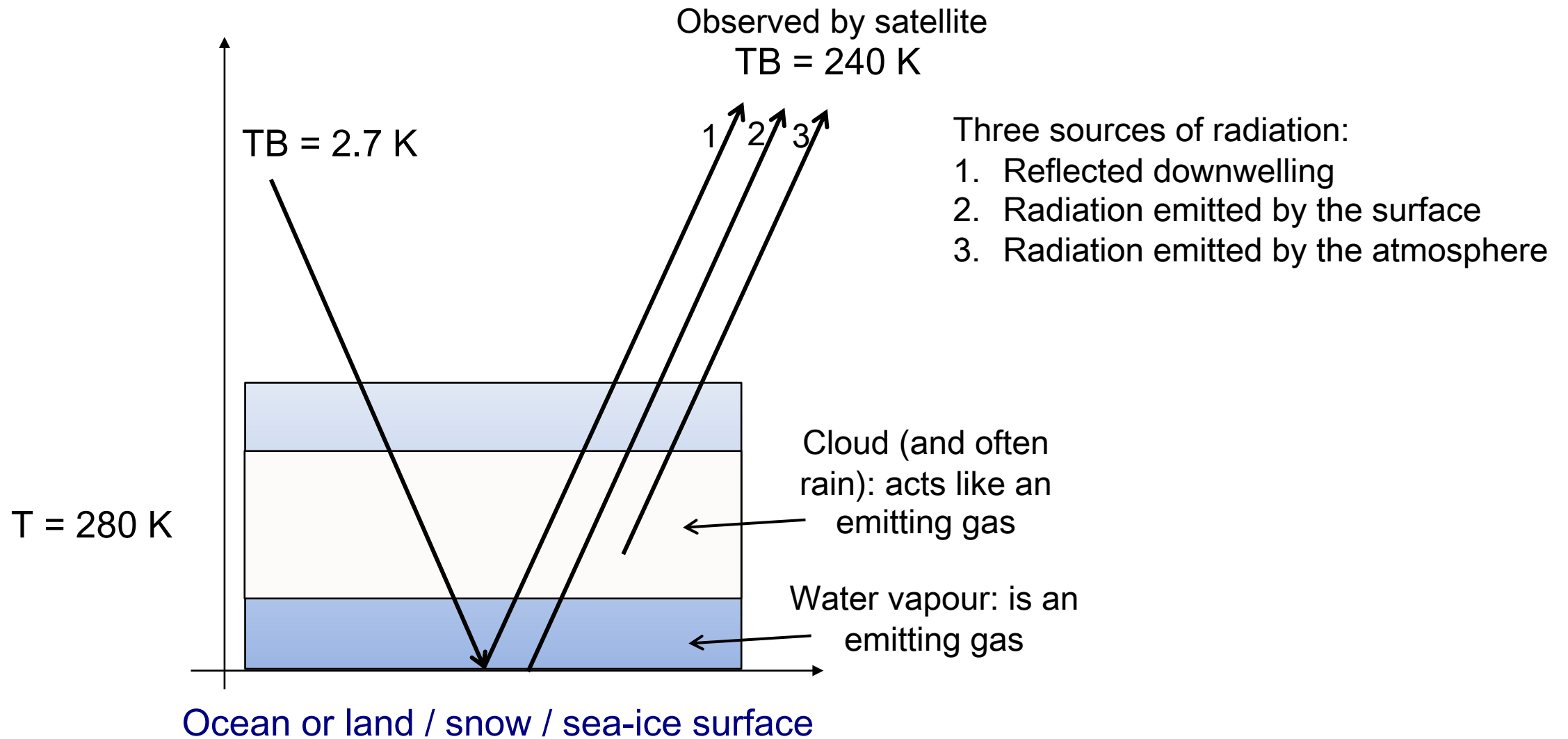


Earth system model:
geophysical state and its
forward propagation in time

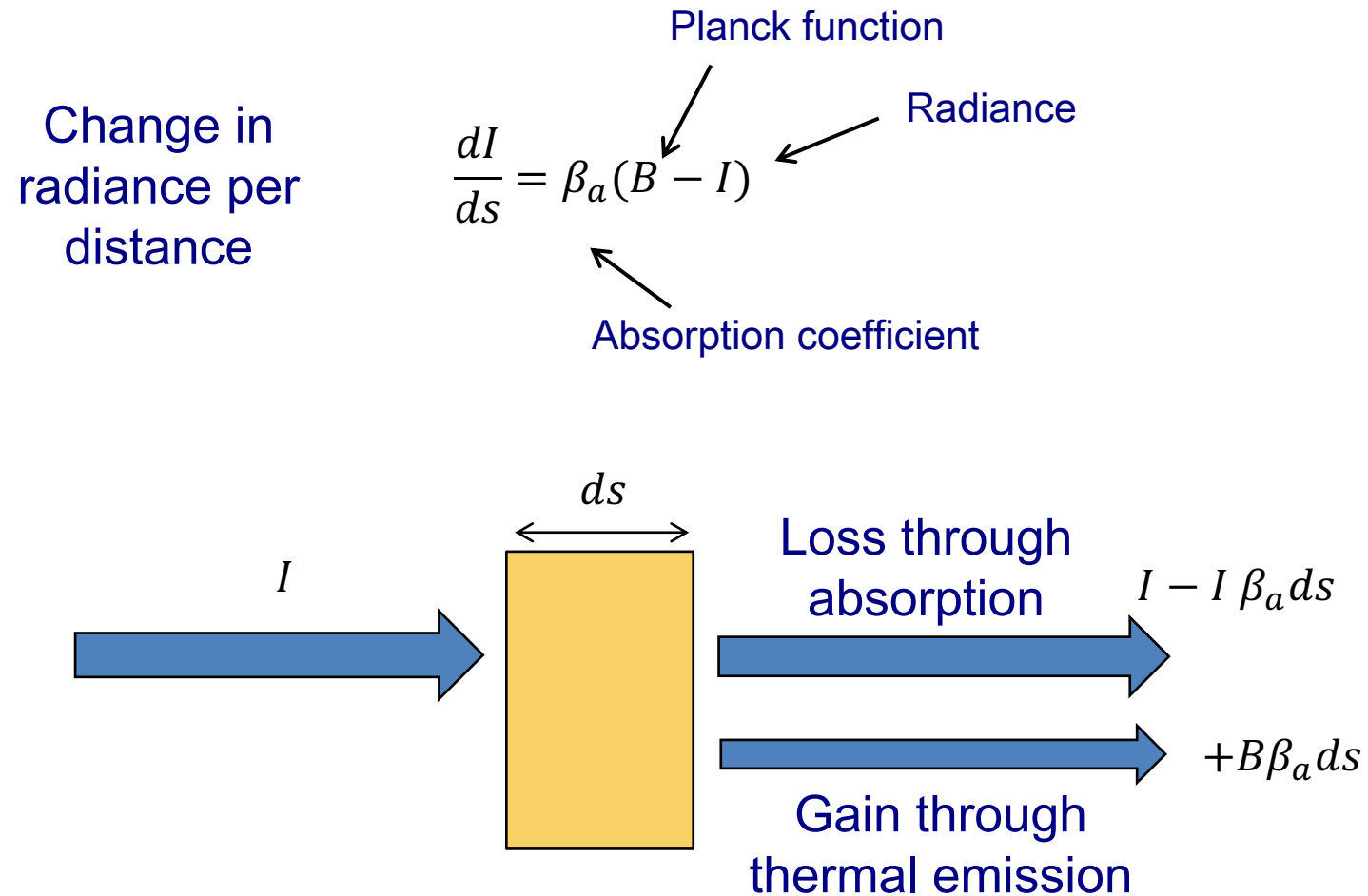


Scattering radiative transfer

Radiative transfer: window channels (ignoring scattering)



Schwarzchild's equation

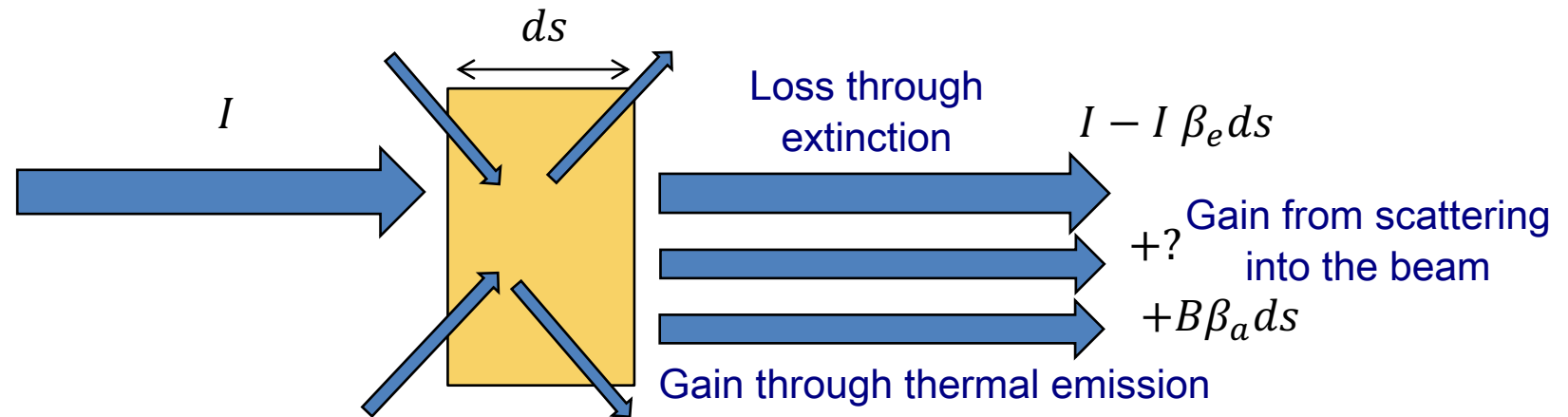


Adding scattering

Extinction coefficient

$$\beta_e = \beta_a + \beta_s$$

Scattering coefficient (describing the amount of scattering out of the beam)



Change in coordinates: optical depth

Change in optical depth
 $d\tau$ in a non-scattering
atmosphere

$$d\tau = -\beta_a ds$$

Change in optical depth
 $d\tau$ including extinction by
scattering

$$d\tau = -(\beta_a + \beta_s) ds = -\beta_e ds$$

The full scattering radiative transfer equation

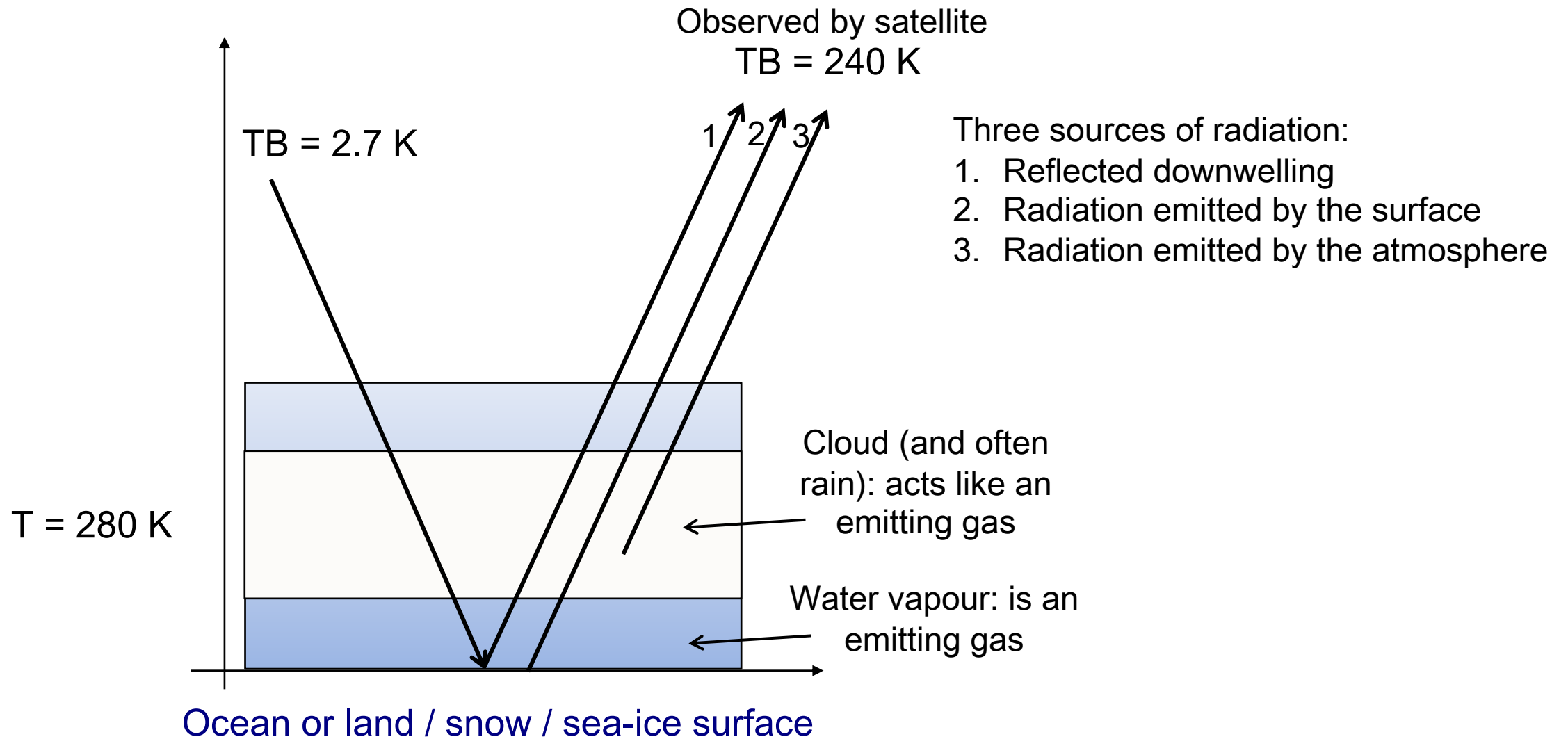
$$\frac{dI(\hat{\Omega})}{d\tau} = I(\hat{\Omega}) - (1 - \omega_s)B - \frac{\omega_s}{4\pi} \int_{4\pi} P(\hat{\Omega}, \hat{\Omega}') I(\hat{\Omega}') d\hat{\Omega}'$$

Direction vector $\hat{\Omega}$
 Integral over all solid angle $\int_{4\pi}$
 Change in optical depth $\frac{dI(\hat{\Omega})}{d\tau}$
 Single scattering albedo ω_s
 Scattering phase function $P(\hat{\Omega}, \hat{\Omega}')$

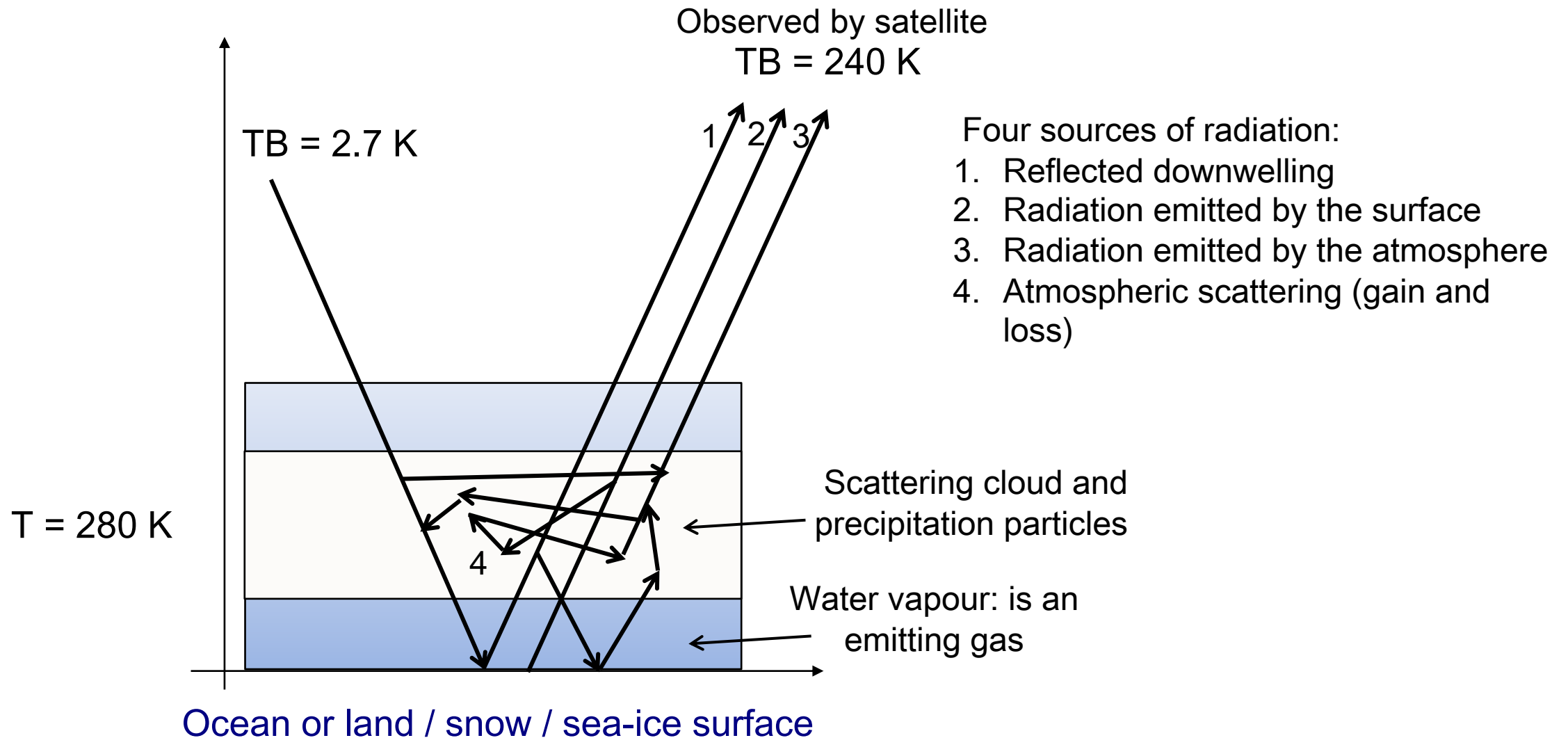
$$\omega_s = \frac{\beta_s}{(\beta_a + \beta_s)}$$

- Without scattering, just integrate this equation along the path travelled by the radiation (Tony's first lecture)
- With scattering, this can be complex to solve:
 $I(\hat{\Omega})$, the radiance in one direction, depends on radiance from all other directions: $I(\hat{\Omega}')$
 and all levels depend on each other

Radiative transfer: window channels (ignoring scattering)

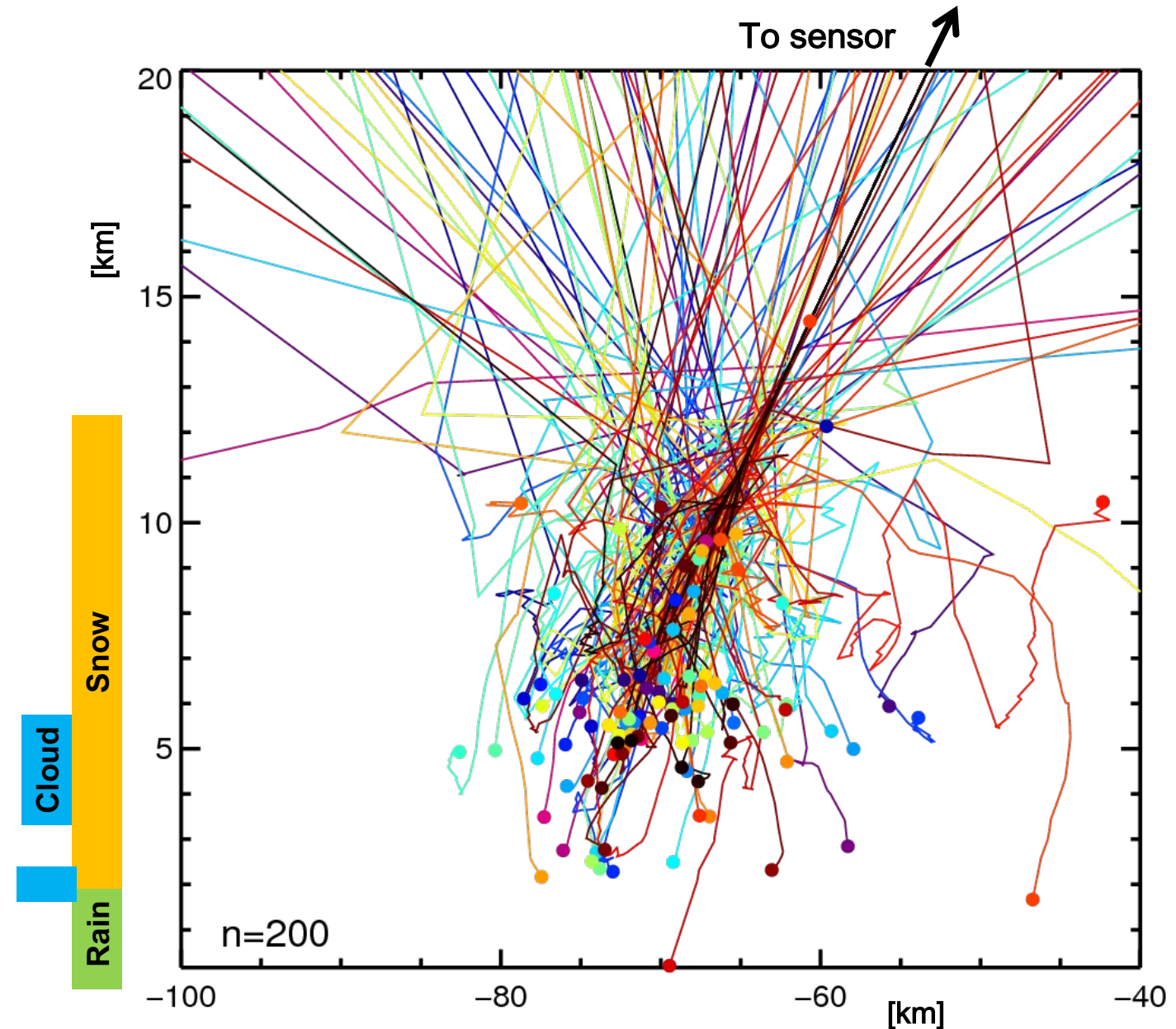


Radiative transfer: window channels (with scattering)



Strong scattering at 91 GHz

Reverse Monte-Carlo radiative transfer solver



The full scattering radiative transfer equation

Radiative transfer solver

$$\frac{dI(\hat{\Omega})}{d\tau} = I(\hat{\Omega}) - (1 - \omega_s)B - \frac{\omega_s}{4\pi} \int_{4\pi} P(\hat{\Omega}, \hat{\Omega}') I(\hat{\Omega}') d\hat{\Omega}'$$

Direction vector $\hat{\Omega}$
 Integral over all solid angle $\int_{4\pi}$
 Change in optical depth $d\tau$
 Single scattering albedo ω_s
 Scattering phase function $P(\hat{\Omega}, \hat{\Omega}')$

$$\omega_s = \frac{\beta_s}{(\beta_a + \beta_s)}$$

+ Surface description (water, sea ice, snow, soil)

Optical properties (gas and hydrometeor), temperature

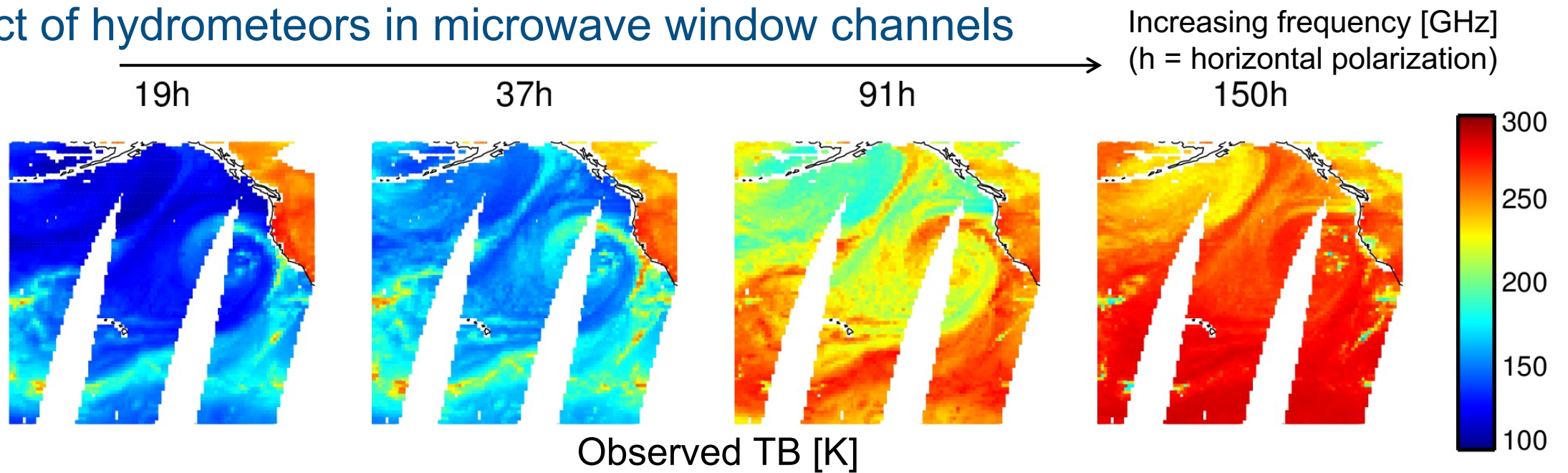
Gas spectroscopy

Single-particle scattering models

Atmospheric profiles (temperature, pressure, water vapour, hydrometeors)

Cloud effects in observations

Effect of hydrometeors in microwave window channels



Effect of hydrometeors in microwave window channels

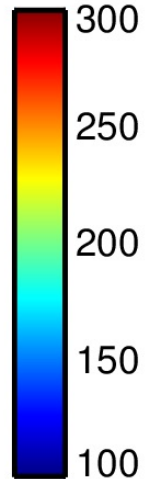
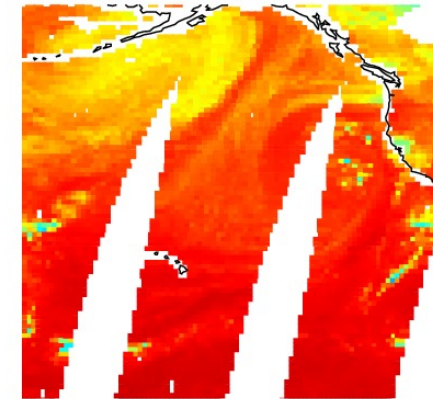
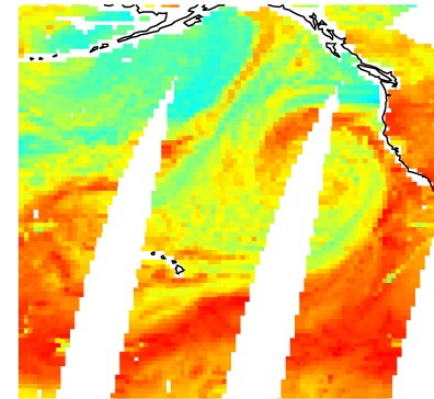
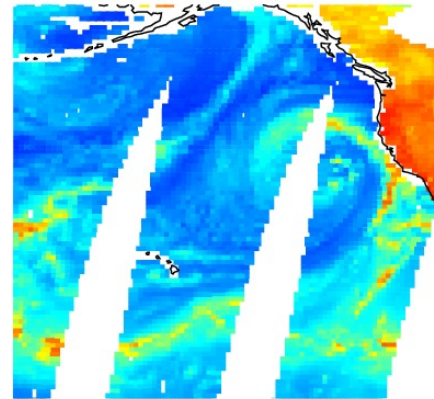
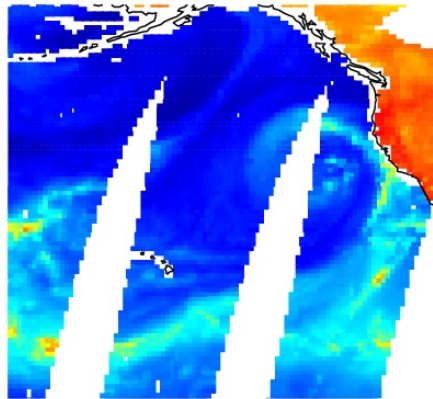
Increasing frequency [GHz]
(h = horizontal polarization)

19h

37h

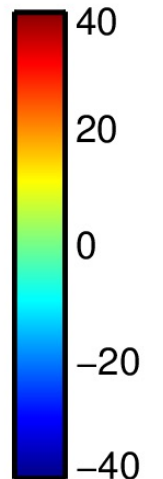
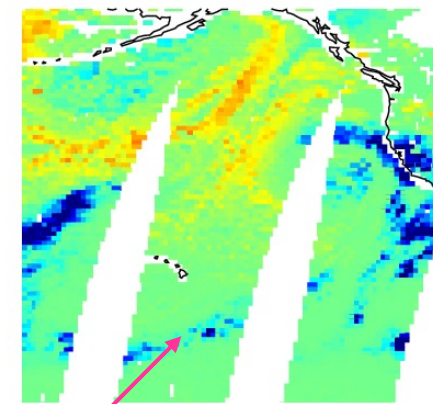
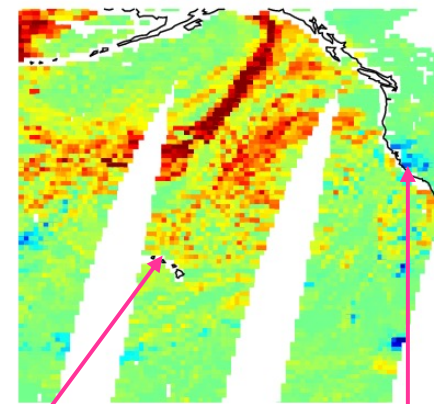
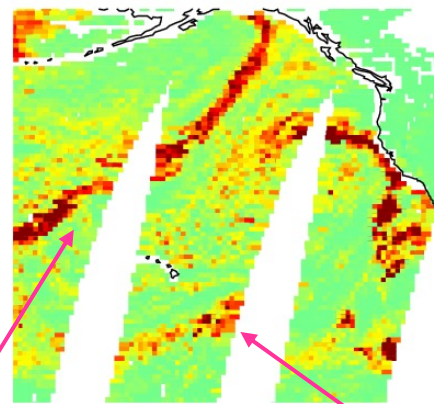
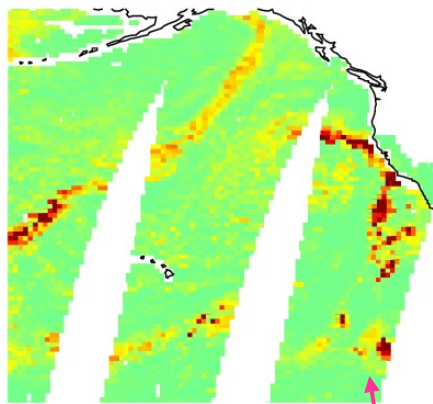
91h

150h



Observed TB [K]

Hydrometeor effect: observed TB – Simulated clear-sky TB [K]

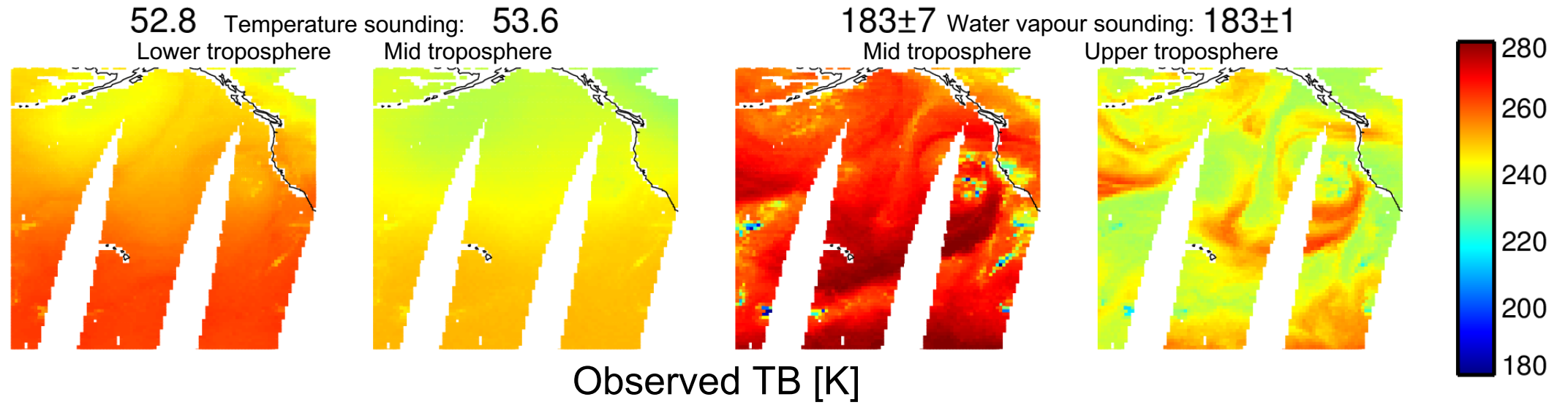


Rain (absorption, increases TB)

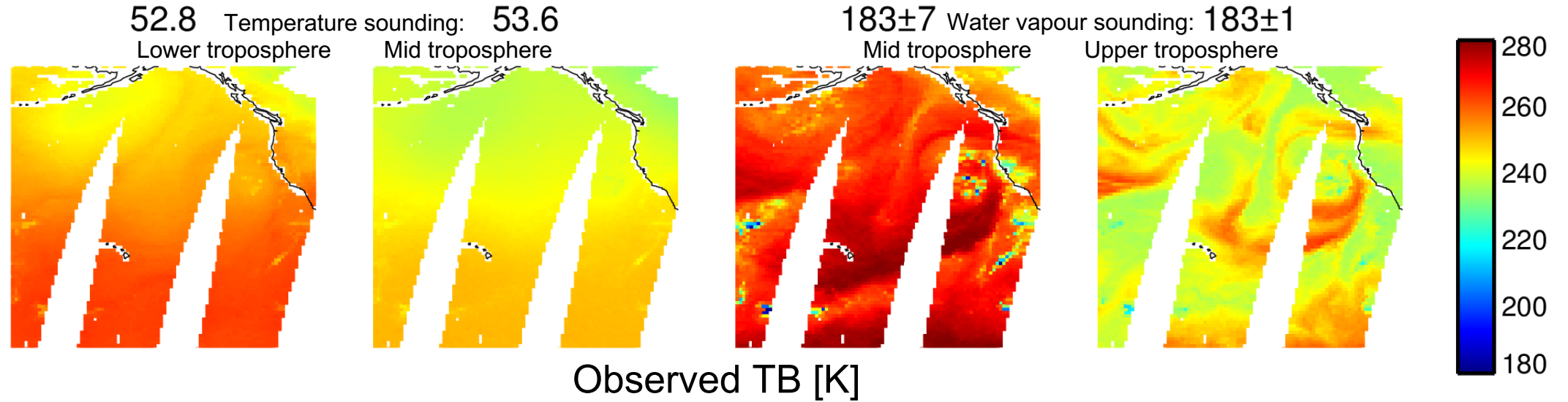
Cloud (absorption, increases TB)

Snow/graupel/hail (scattering, decreases TB)

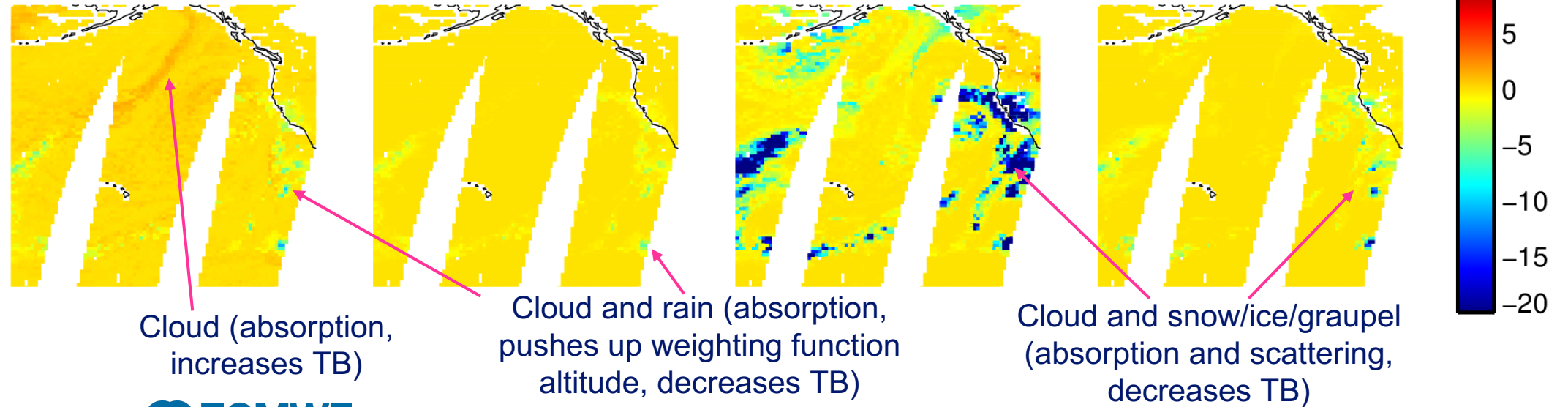
Effect of hydrometeors in microwave sounding channels



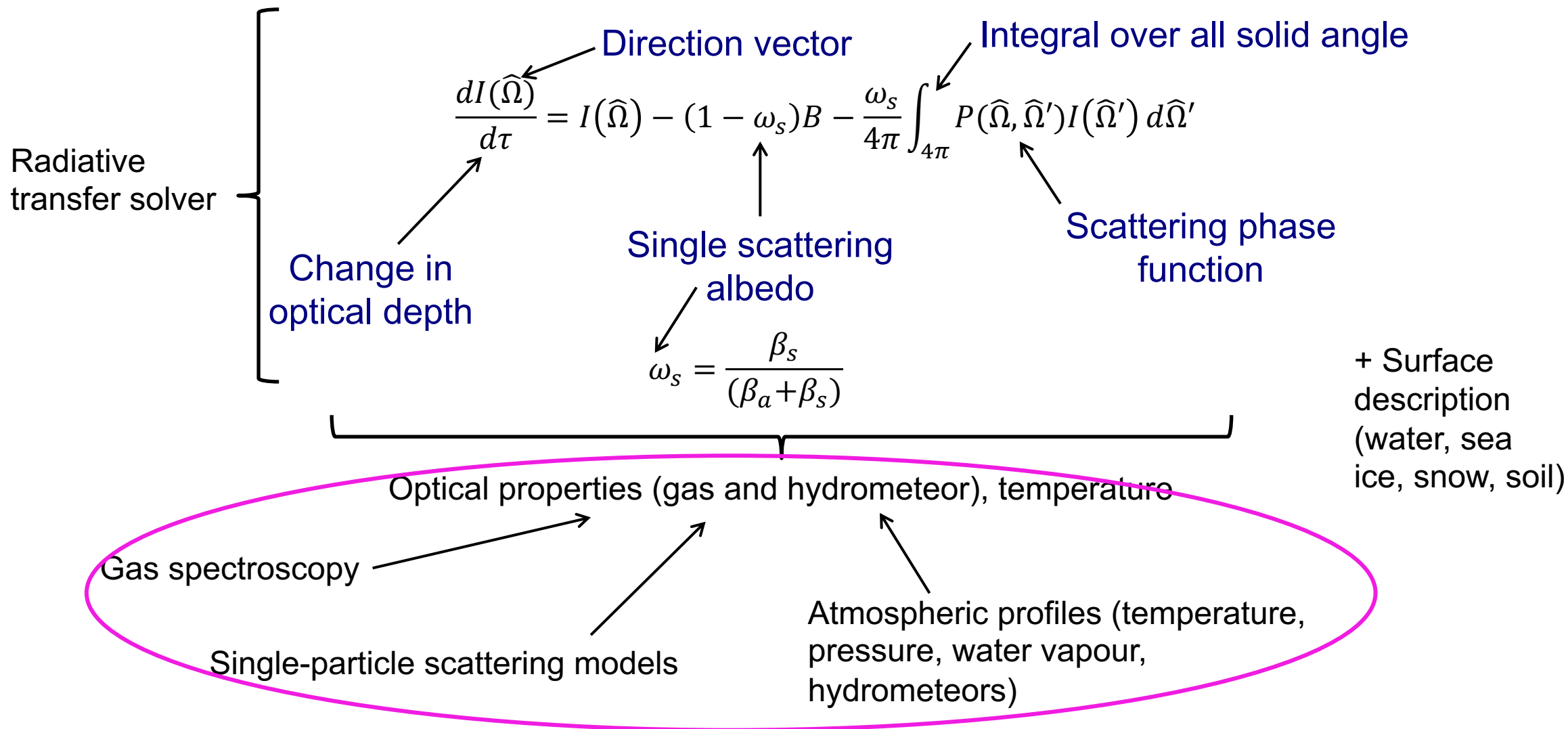
Effect of hydrometeors in microwave sounding channels



Hydrometeor effect: observed TB – Simulated clear-sky TB [K]

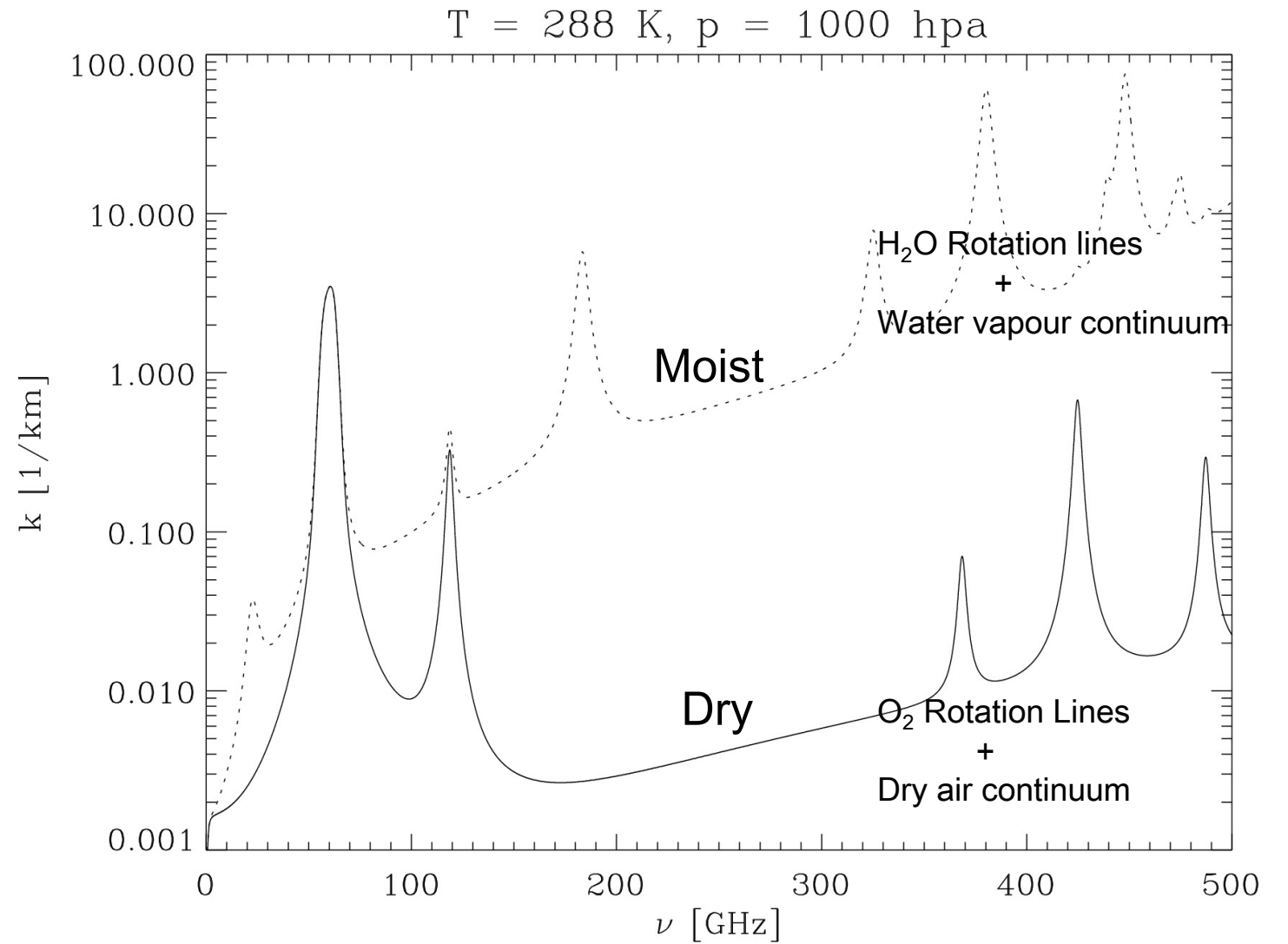


The full scattering radiative transfer equation

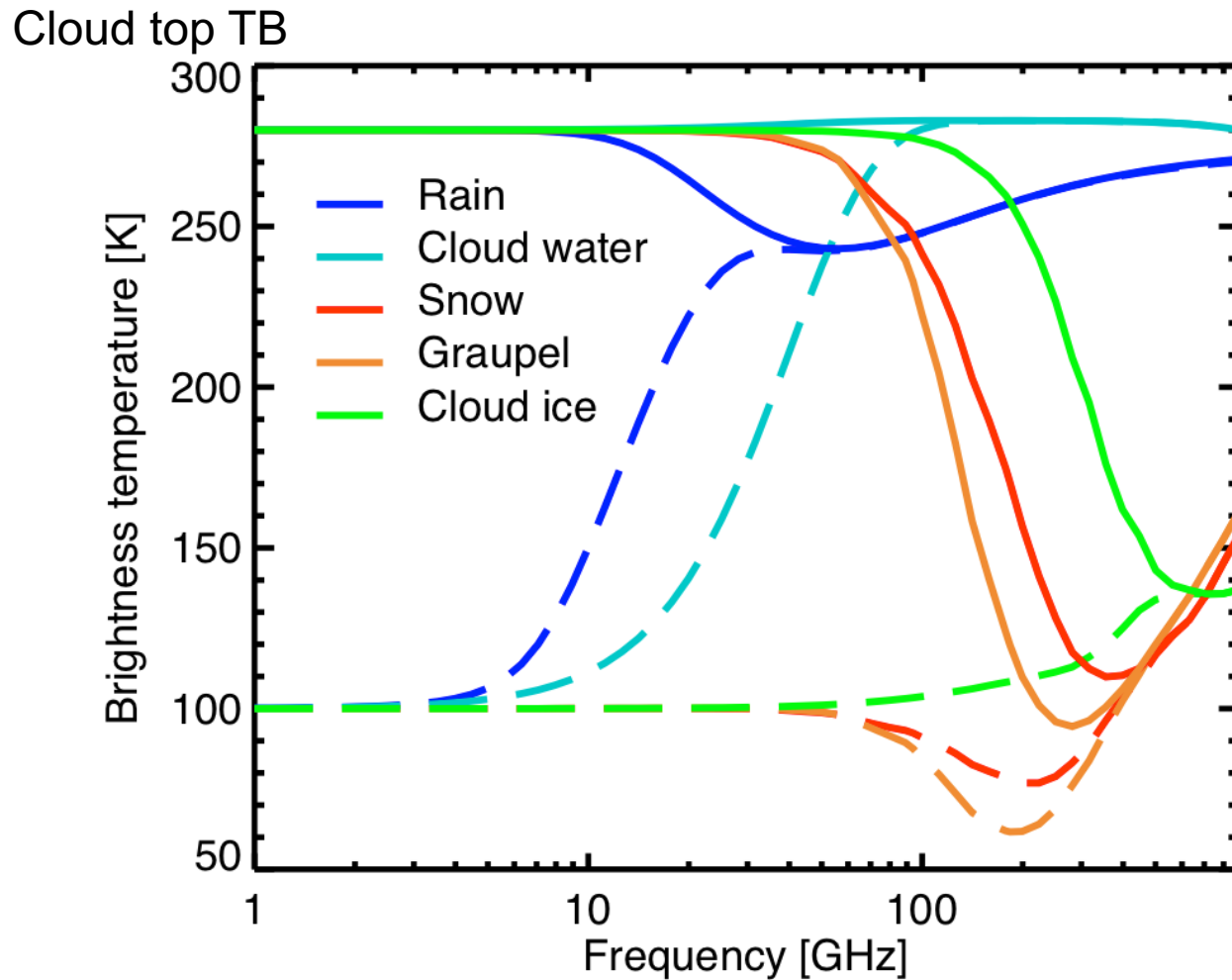
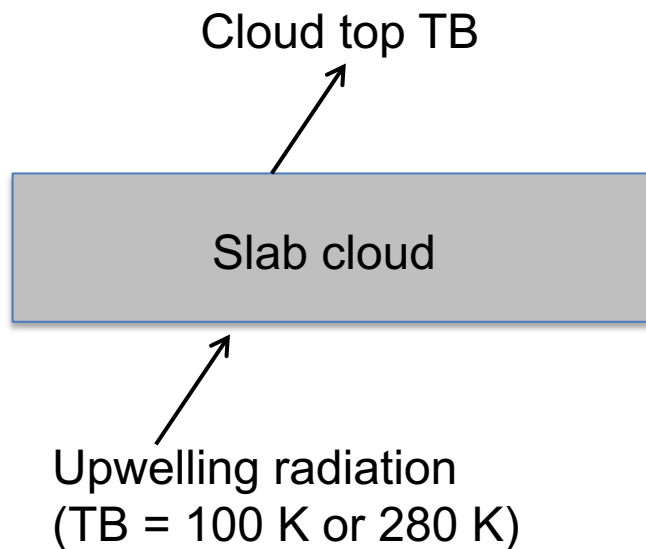


Gas absorption: the microwave spectrum

Absorption coefficient β_a [1/km]



Cloud and precipitation optical properties: the microwave spectrum

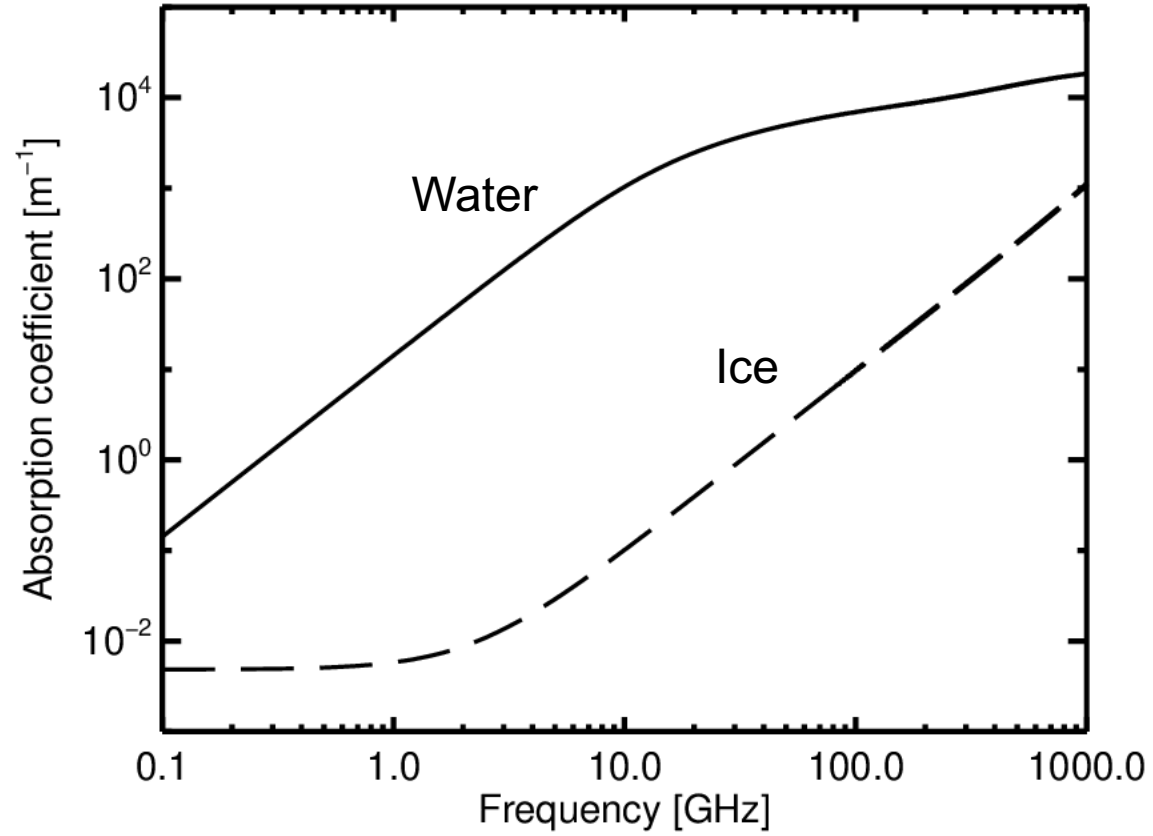


Slab cloud at 283K above a 280K surface (solid)

Slab cloud at 283K above a 100K surface (dashed)

Geer et al. (2021, GMD, Bulk hydrometeor optical properties for microwave and sub-millimetre radiative transfer in RTTOV-SCATT v13.0)

Absorption in pure water or ice



Effect of hydrometeors – particles

- 30 GHz frequency ↔ 10mm wavelength (λ)

Size parameter

$$x = \frac{2\pi r}{\lambda}$$

$x \ll 1$: Rayleigh scattering

$x \sim 1$: Mie sphere, discrete dipole approximation, etc.

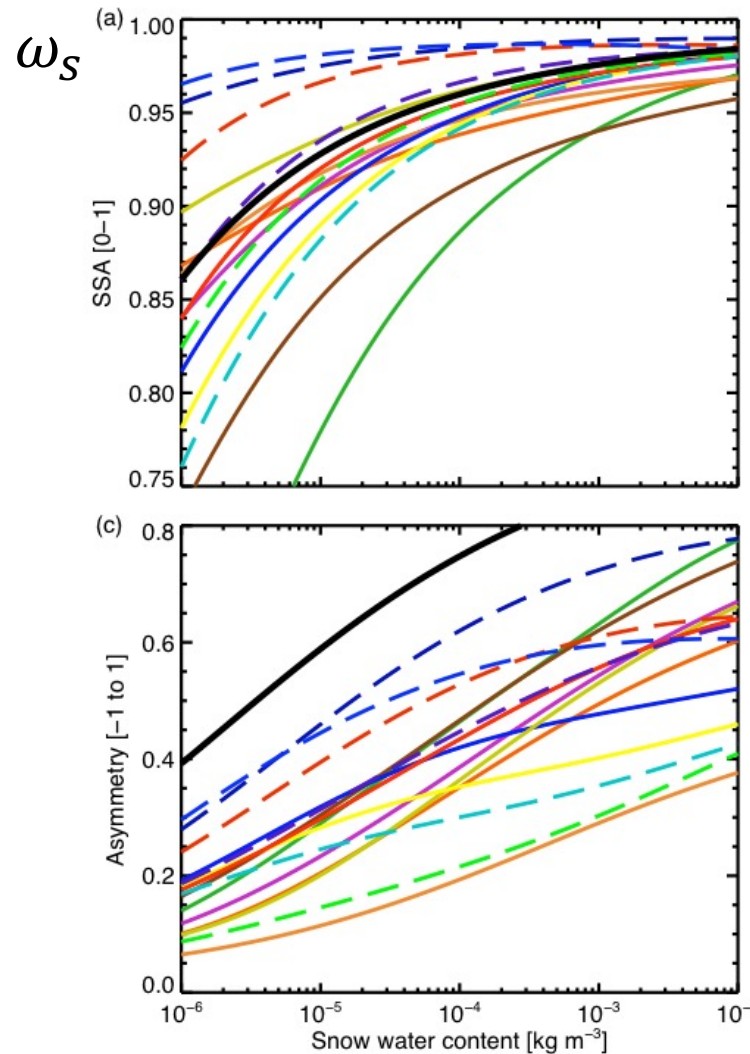
$x \gg 1$: Geometric optics

Particle type	Size range, r	Size parameter, x
Cloud droplets	5 – 50 μm	0.003 – 0.03
Drizzle	~100 μm	0.06
Rain drop	0.1 – 3 mm	0.06 – 1.8
Ice crystals	10 – 100 μm	0.006 – 0.06
Snow	1 – 10 mm	0.6 – 6
Hailstone	~10 mm	6

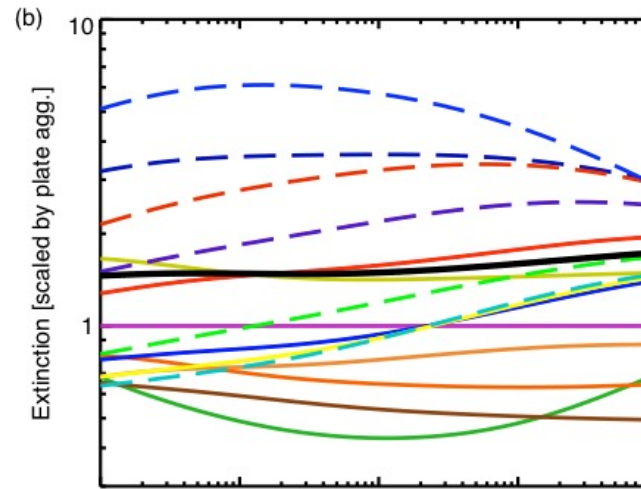
- Effect of particles on radiation is a function of the particle shape and structure, size relative to the radiation, and composition (complex refractive index / permittivity)
- Bulk effect of particles is an integral over the particle size distribution (PSD)

Optical properties of hydrometeors in RTTOV-SCATT: at 183 GHz

Lookup tables for snow hydrometeors as a function of snow water content



- Mie
- ICON hail
- 8-column aggregate
- ICON cloud ice
- Column type 1
- Gem graupel
- Flat 3-bullet rosette
- Plate type 1



$$\beta_e = \beta_a + \beta_s$$

Bulk extinction coefficient scaled relative to a large plate aggregate

$$g = \int_{-1}^1 \cos \theta P(\theta) d\cos \theta$$

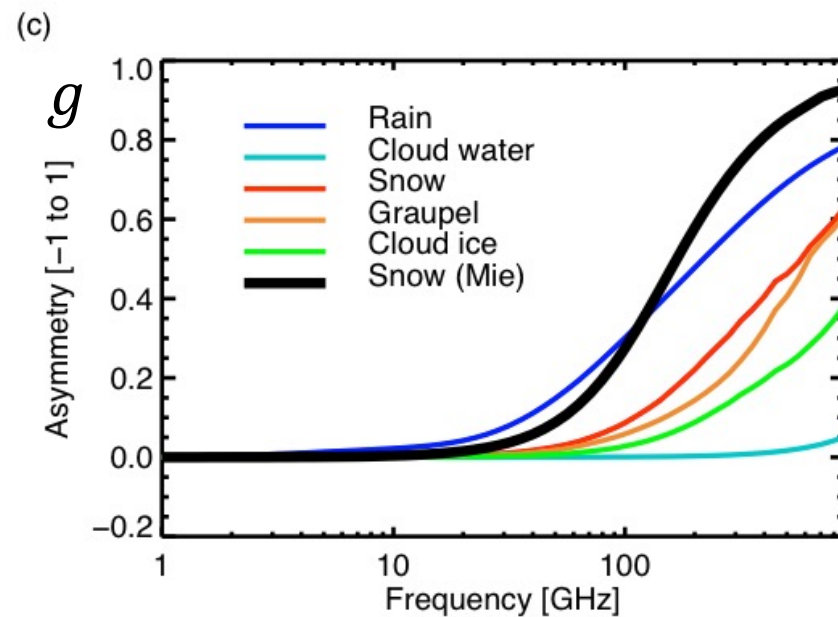
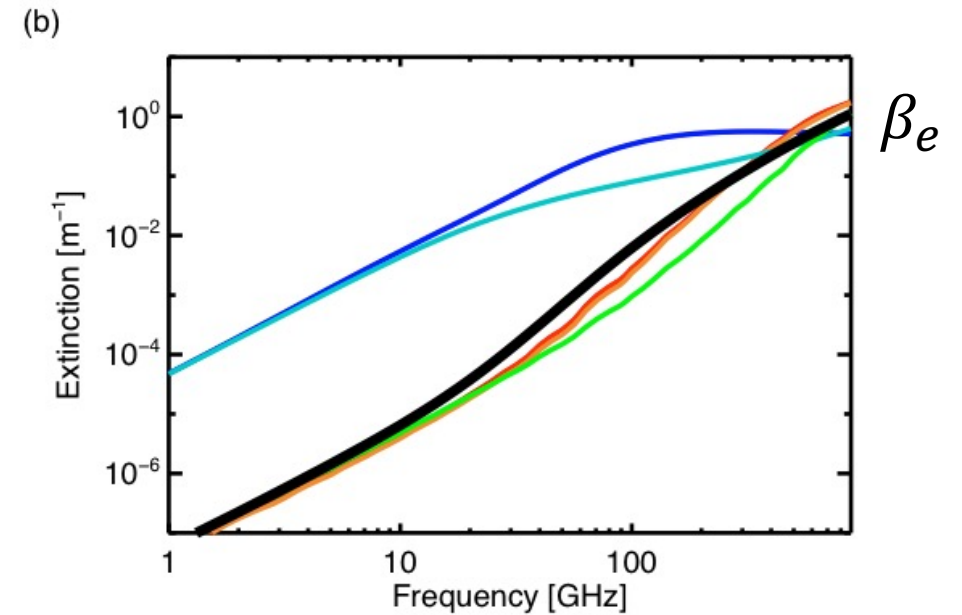
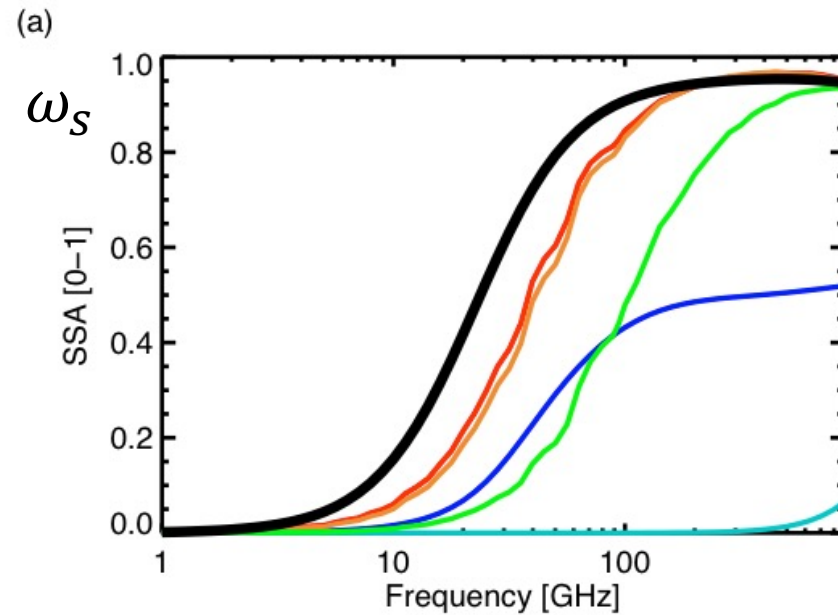
Bulk asymmetry parameter = average over the (already azimuthally averaged) scattering phase function $P(\theta)$

- Perpendicular 4-bullet rosette
- Large block aggregate
- 6-bullet rosette
- Sector snowflake
- Large plate aggregate
- ICON snow
- Large column aggregate
- Evans snow aggregate

DDA non-spherical particles from the ARTS scattering database

Optical properties of hydrometeors in RTTOV-SCATT: across frequencies

Lookup tables for snow hydrometeors of water content 10^{-4} kg/m^3 as a function of frequency



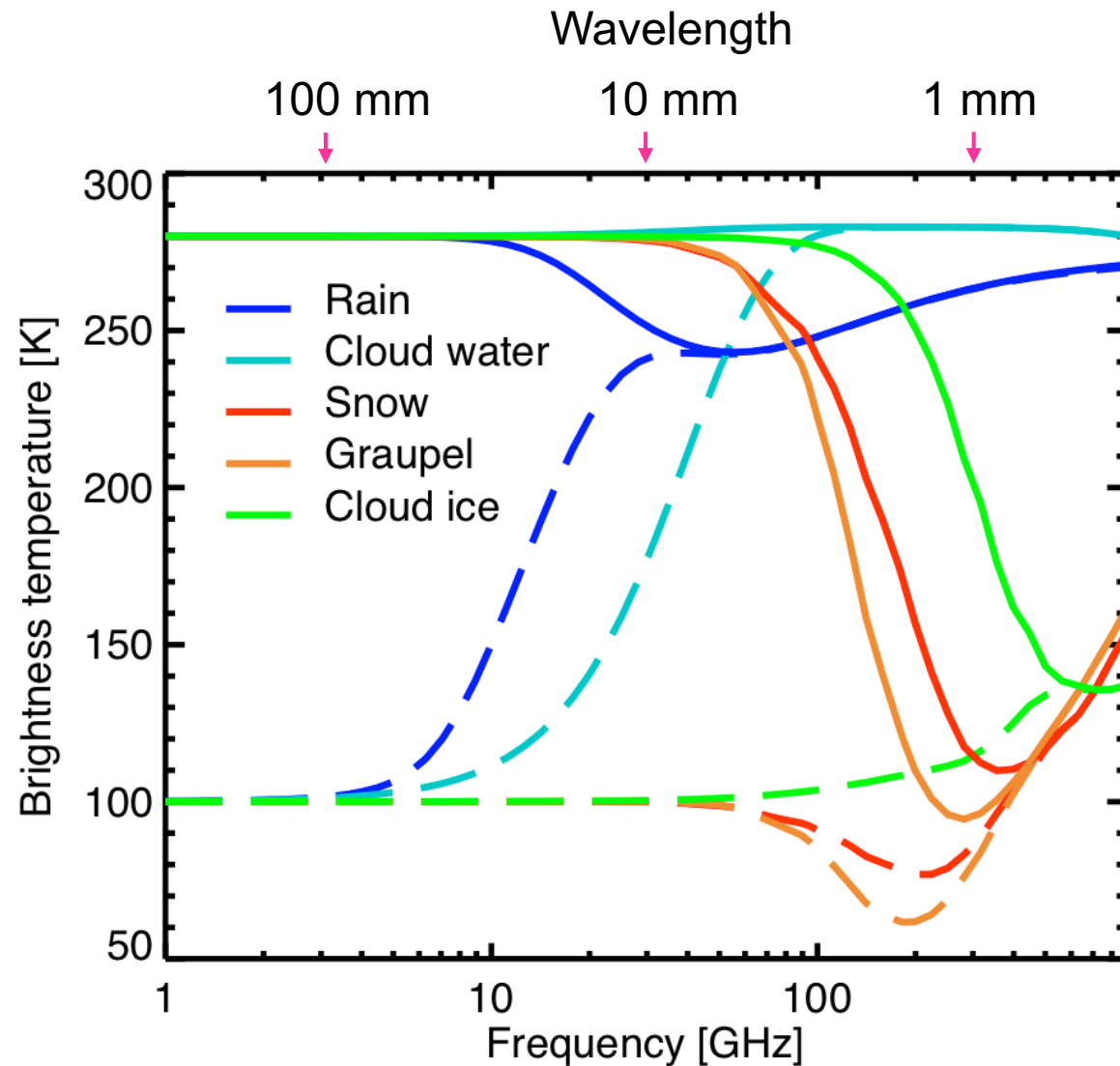
(d)

Hydrometeor placeholder	Scattering type	Particle shape	PSD	MGD parameters			
				N_0	μ	Λ	γ
Rain	Mie	sphere	MGD	8×10^6	0	free	1
Snow	ARTS	large plate aggregate	F07 T	-	-	-	-
Graupel	ARTS	column	F07 T	-	-	-	-
Cloud water	Mie	sphere	MGD	free	2	2.13×10^5	1
Cloud ice	ARTS	large column aggregate	MGD	free	0	1×10^4	1

Geer et al. (2021, GMD)



Cloud and precipitation optical properties: the microwave spectrum

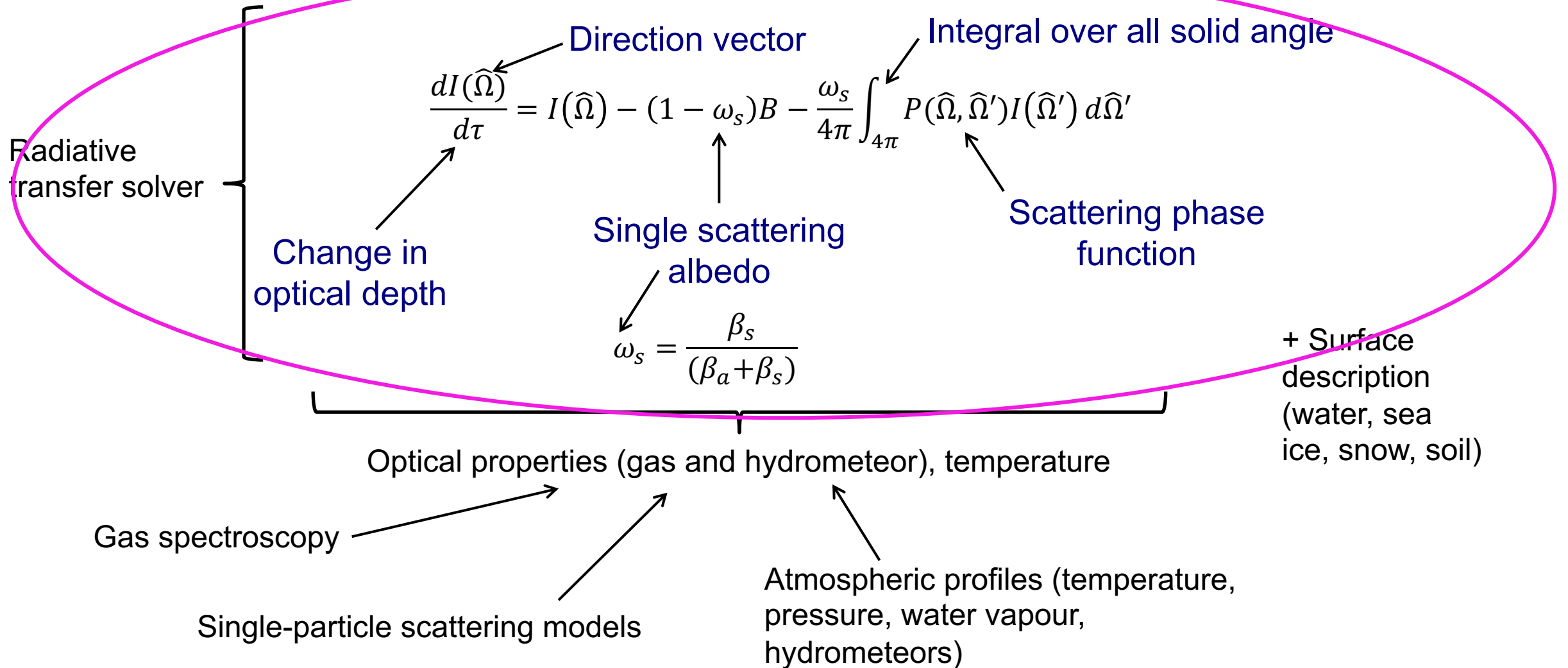


Slab cloud at 283K above a 280K surface (solid)

Slab cloud at 283K above a 100K surface (dashed)


Geer et al. (2021, GMD)

The full scattering radiative transfer equation



Fast approximate solver: RTTOV- SCATT

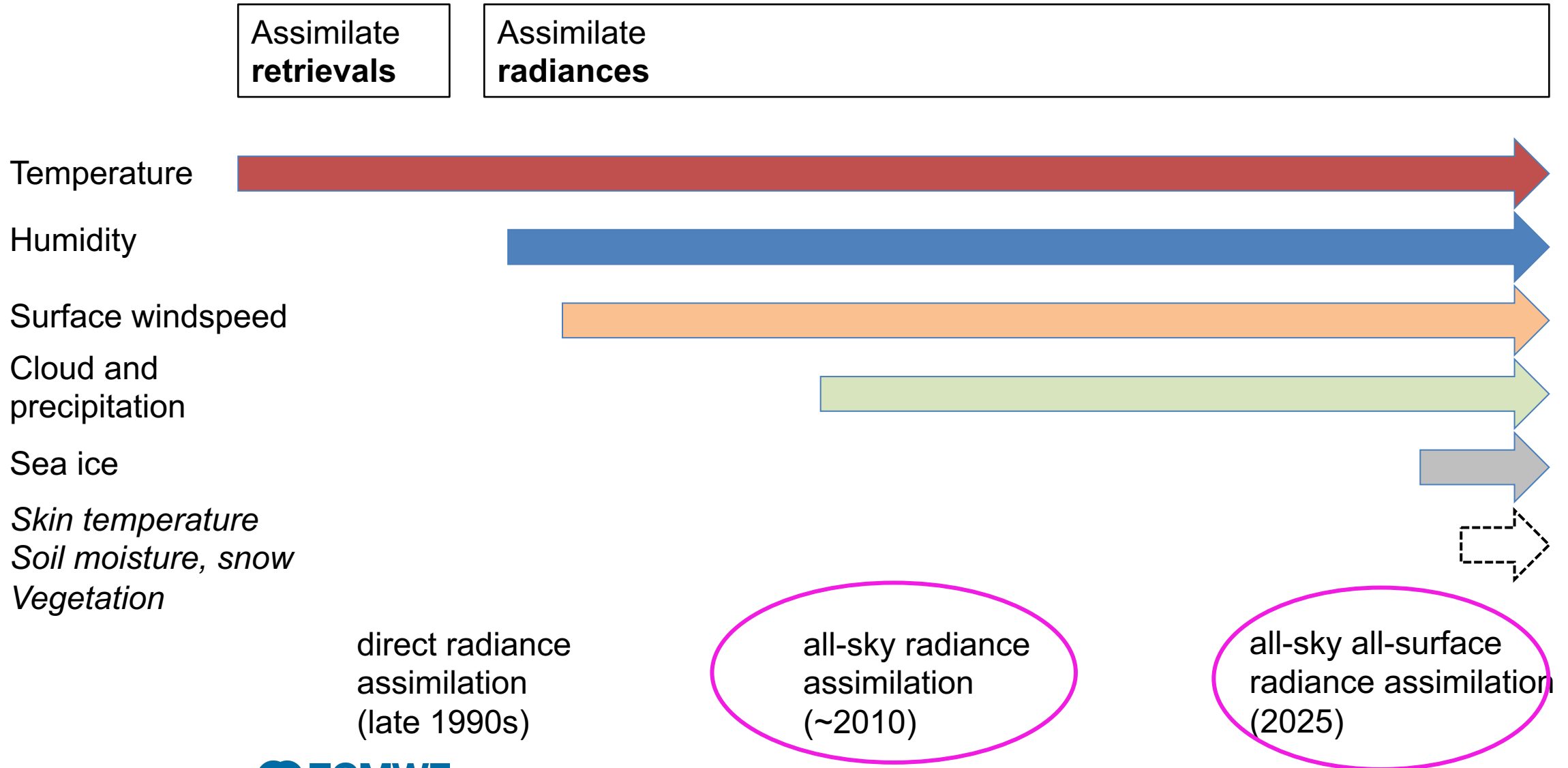
- For each hydrometeor type (e.g. cloud, rain etc.), compute single-species bulk optical properties as a function of mixing ratio
- Sum up single-species optical properties along with gas optical properties
- Assume a simplified scattering phase-function
- Assume no azimuth dependence of the radiance field
- Assume radiance is a simple function of zenith angle (theta): $I = I_0 + I_1 \cos \theta$ (Eddington approximation, similar to the two-stream approach)
- delta-scaling (treat sharp peaks in the scattering phase function as no scattering at all)
- Set up coupled equations describing radiance all levels simultaneously; solve using matrix algebra for the approximate radiance field I at each level
- Integrate the radiative transfer equation along the line of sight, using the delta-Eddington solution to provide the 'source terms'

$$\frac{dI(\hat{\Omega})}{d\tau} = I(\hat{\Omega}) - (1 - \omega_s)B - \frac{\omega_s}{4\pi} \int_{4\pi} P(\hat{\Omega}, \hat{\Omega}') I(\hat{\Omega}') d\hat{\Omega}'$$


Applications

(depending on time available)

Rough timeline of satellite microwave data assimilation in 'atmospheric' DA

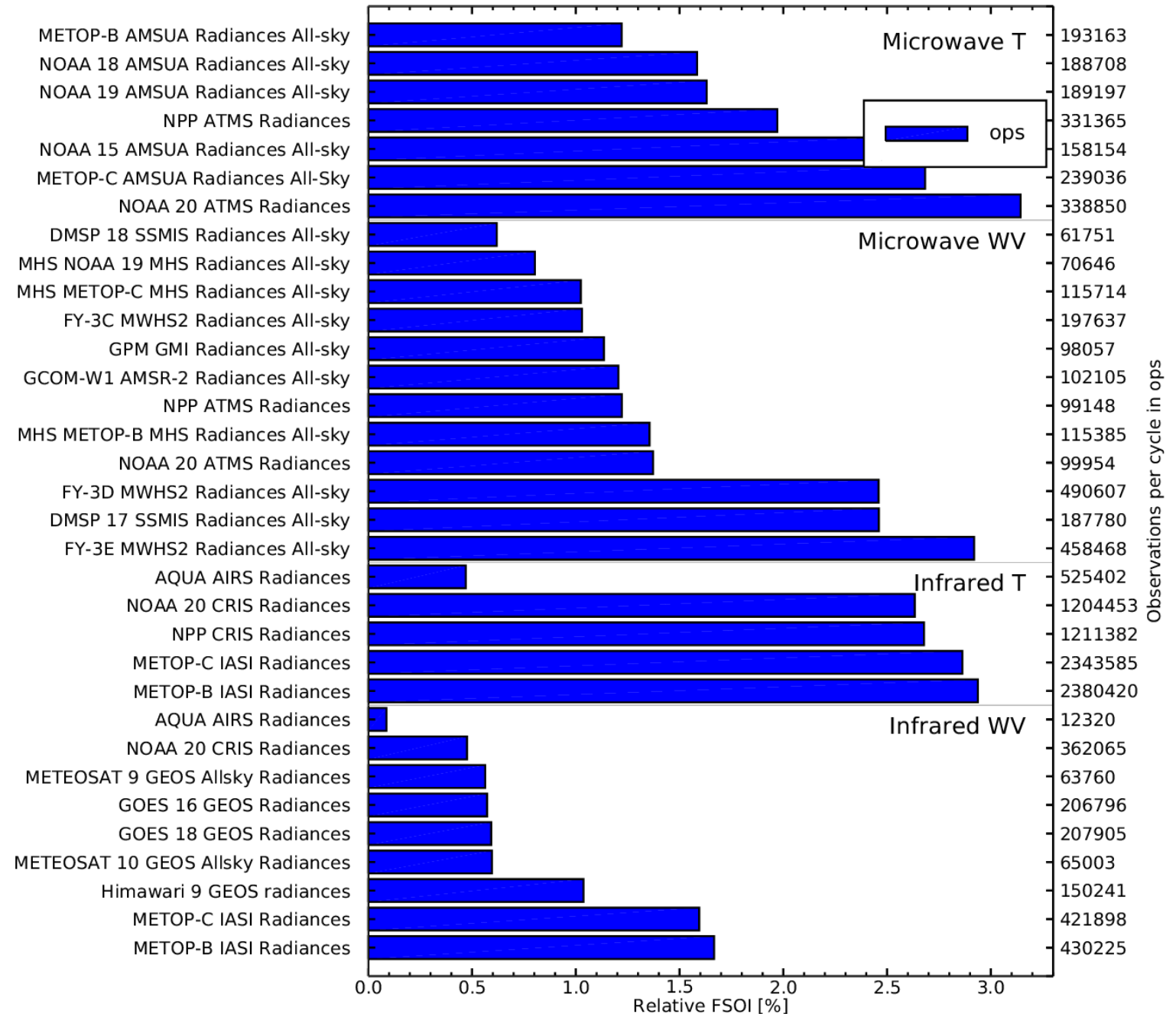


All-sky assimilation

Clear-sky or all-sky?

- Clear-sky assimilation:
 - Remove any cloud-contaminated observations
 - Do not model the effect of cloud on brightness temperatures
 - Traditionally used for temperature sounding channels (e.g. AMSU-A channel 5)
 - Extract small signals of temperature forecast errors (order 0.1K) that would be swamped by errors from displaced clouds and precipitation (10-100K)
- All-sky assimilation
 - Model the effect of cloud and precipitation on the observations
 - Assimilate all data, whether clear, cloudy or precipitating
 - Initially developed for water-vapour sounding and imaging channels, but now also applied to temperature-sounding channels
 - Use the tracing mechanism of 4D-Var to infer the dynamical state from errors in the location/intensity of water vapour, cloud and precipitation
- Broadly, the clear-sky approach is now outdated at microwave frequencies
 - At ECMWF, all but a handful of microwave sensors are now assimilated in all-sky conditions
 - Broadly, going to all-sky assimilation doubles the impact of a sensor (ECMWF TM 741, 2014)

Impact (FSOI) of satellite radiances at ECMWF on short-range forecast, by sensor (100% = all obs)



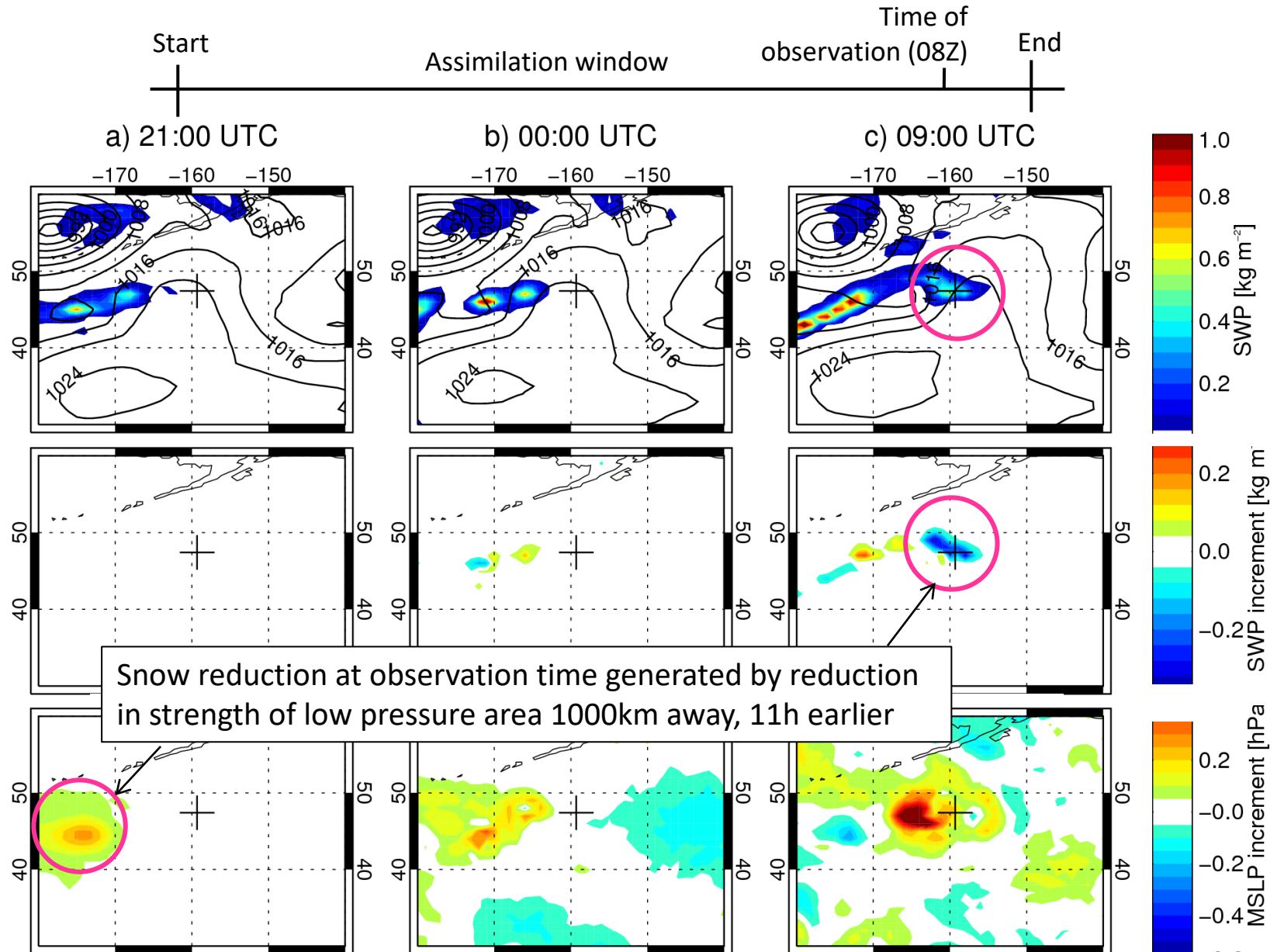
All-sky 'tracer' effect

Single 190
GHz
observation
assimilation
test case

MSLP and
snow
column
(FG)

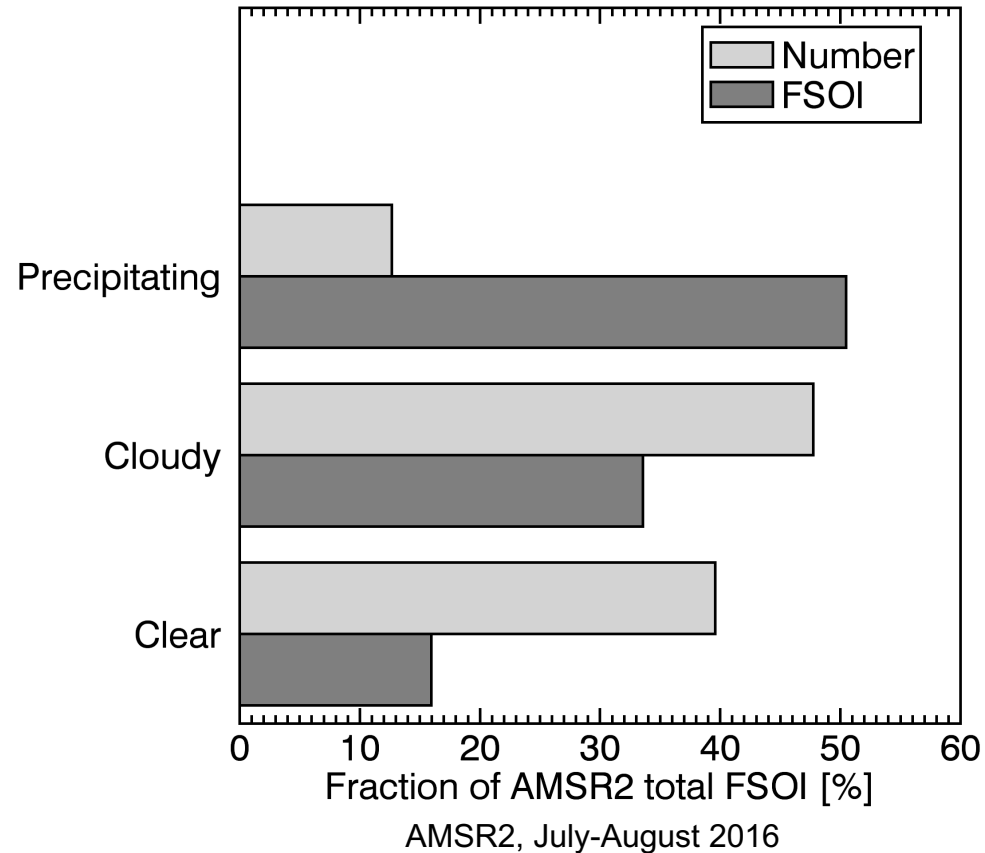
Snow
column
increment

MSLP
increment



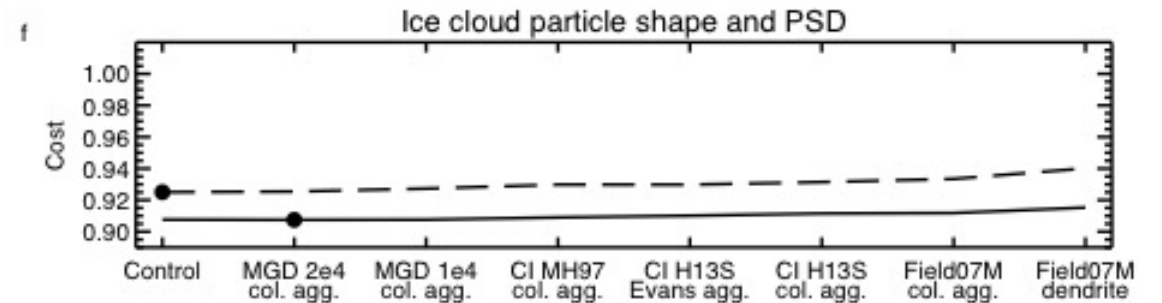
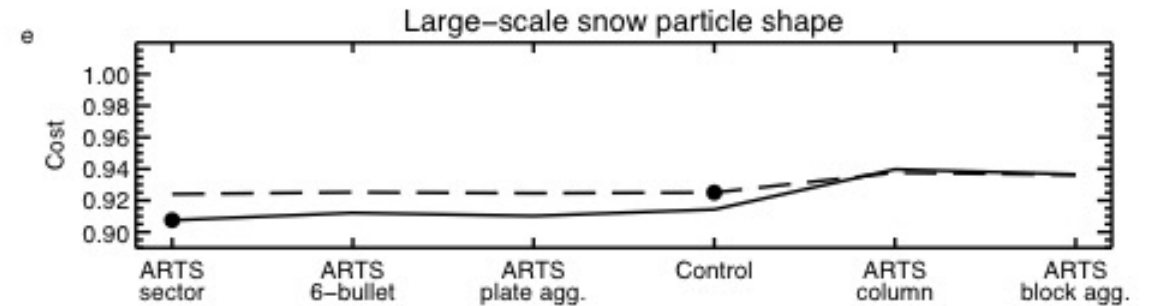
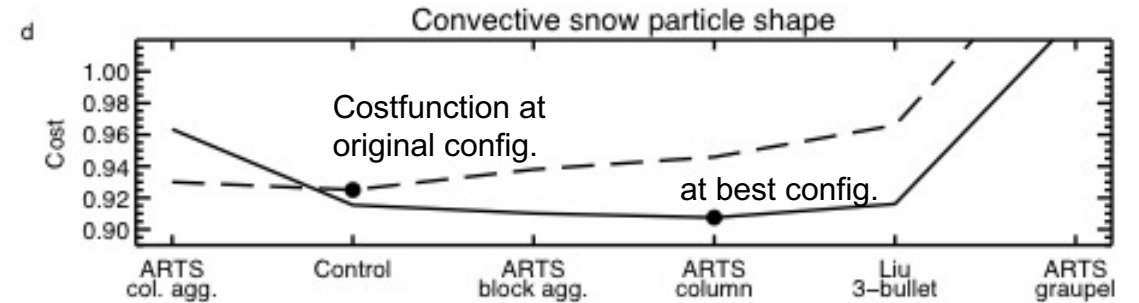
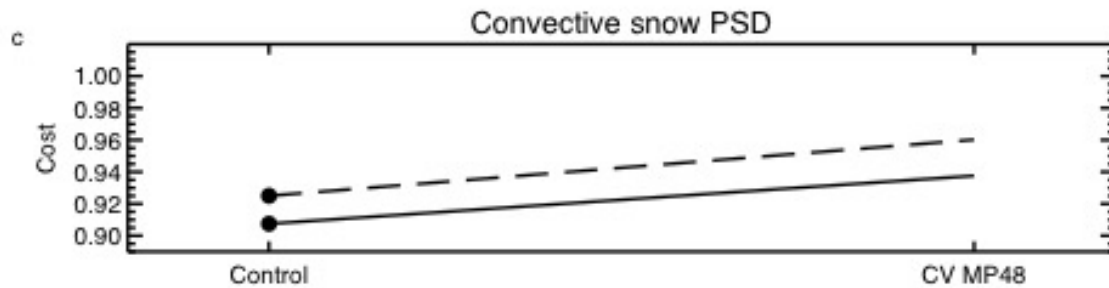
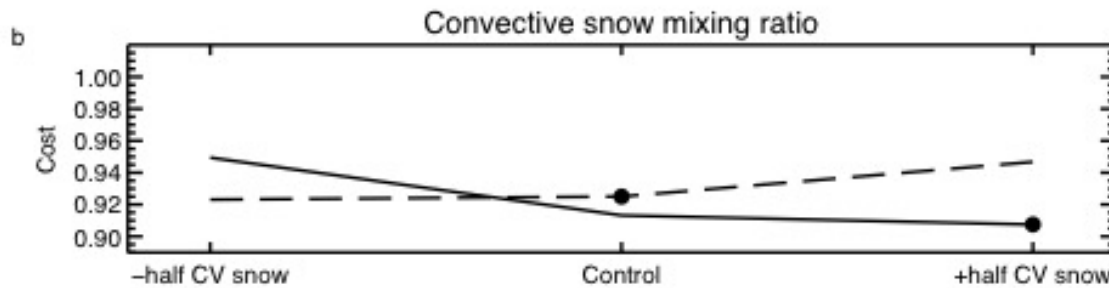
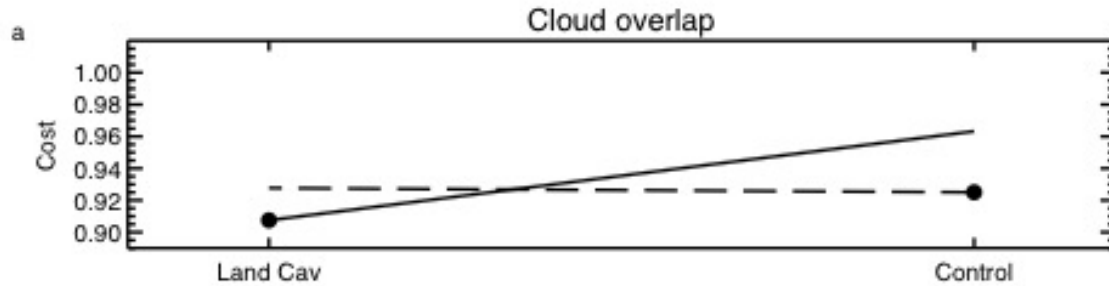
All-sky microwave imagers: a unique contribution from precipitation-affected observations

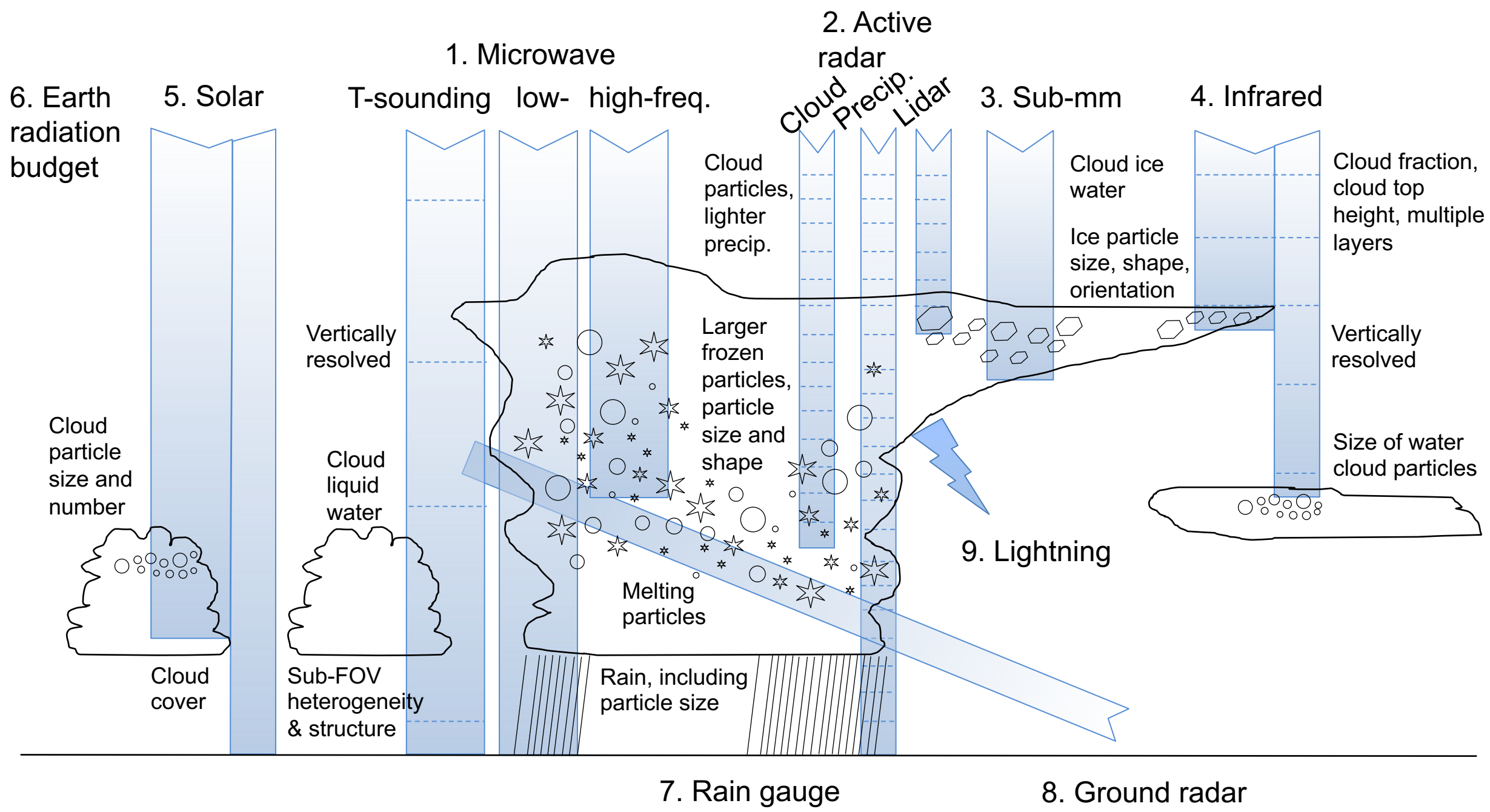
Microwave imagers give their largest forecast impact from a small fraction of precipitating scenes.



Parameter estimation for 6 macro- and microphysical variables

Geer (2021, AMT): Physical characteristics of frozen hydrometeors inferred using parameter estimation





7. Rain gauge

8. Ground radar

All-sky all-surface assimilation

Information content: window (i.e. surface sensitive) channels

SSMIS F-17 channel 13 (19 GHz, v polarisation)
Observed TB, 3rd December 2014

Ocean waves, wind, skin temperature

Atmospheric
water vapour

Cloud and
precipitation

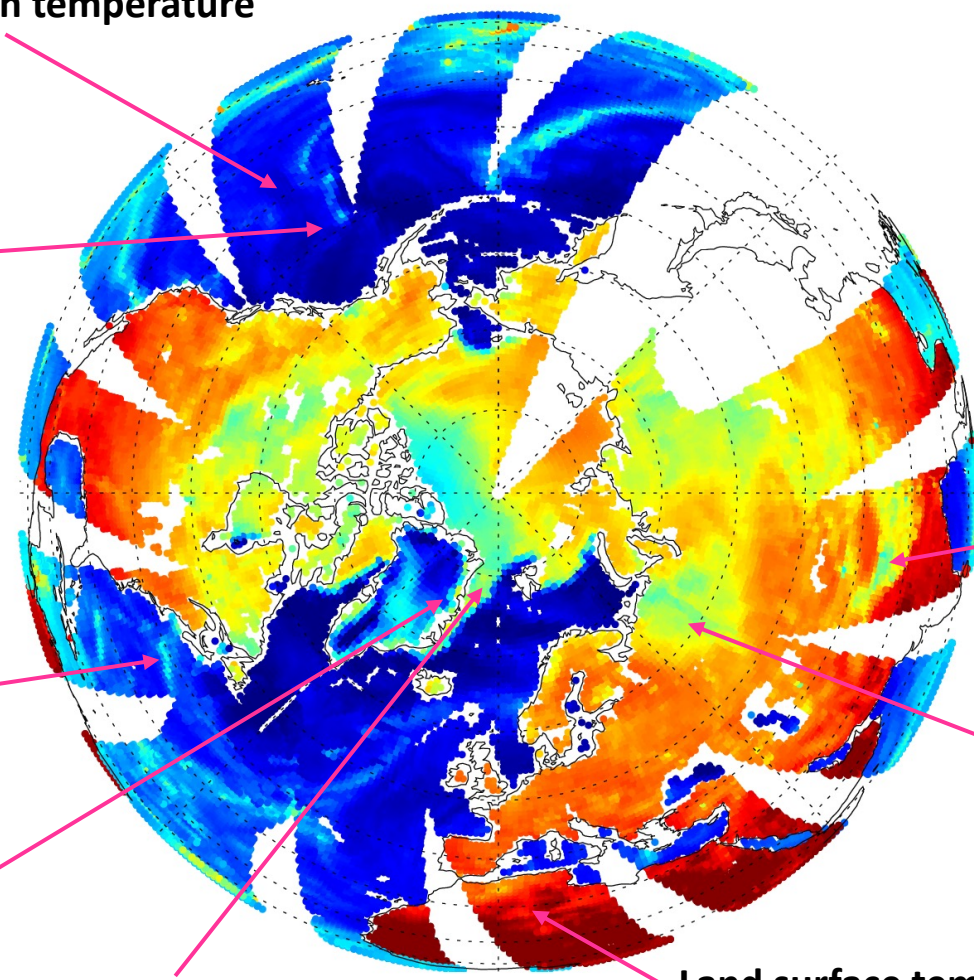
Special snow
and ice
conditions

Sea-ice

High altitude

Snow cover

Land surface temperature,
biomass, soil/rock, soil moisture

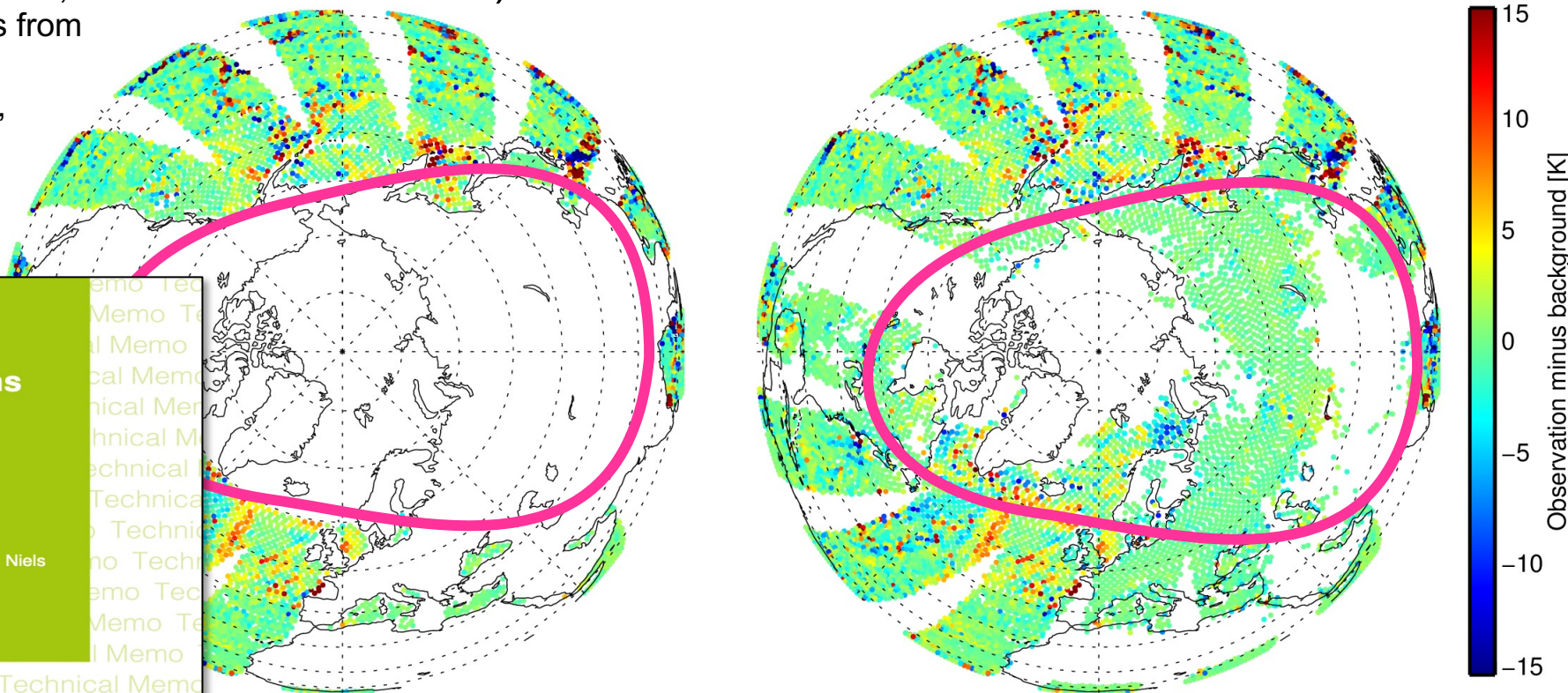


Developments for surface-sensitive microwave channels in cycle 48r1 (June 2023)

Active channel 10 (36.5 GHz, v-polarised) observations from AMSR2 during 00 UTC analysis cycle, 26th June, 2019

now (all-sky but not all-surface)

upgrade



Improved surface treatment for all-sky microwave observations

Alan J. Geer, Katrin Lonitz, David I. Duncan, Niels Bormann
(Research Department)
February 2022

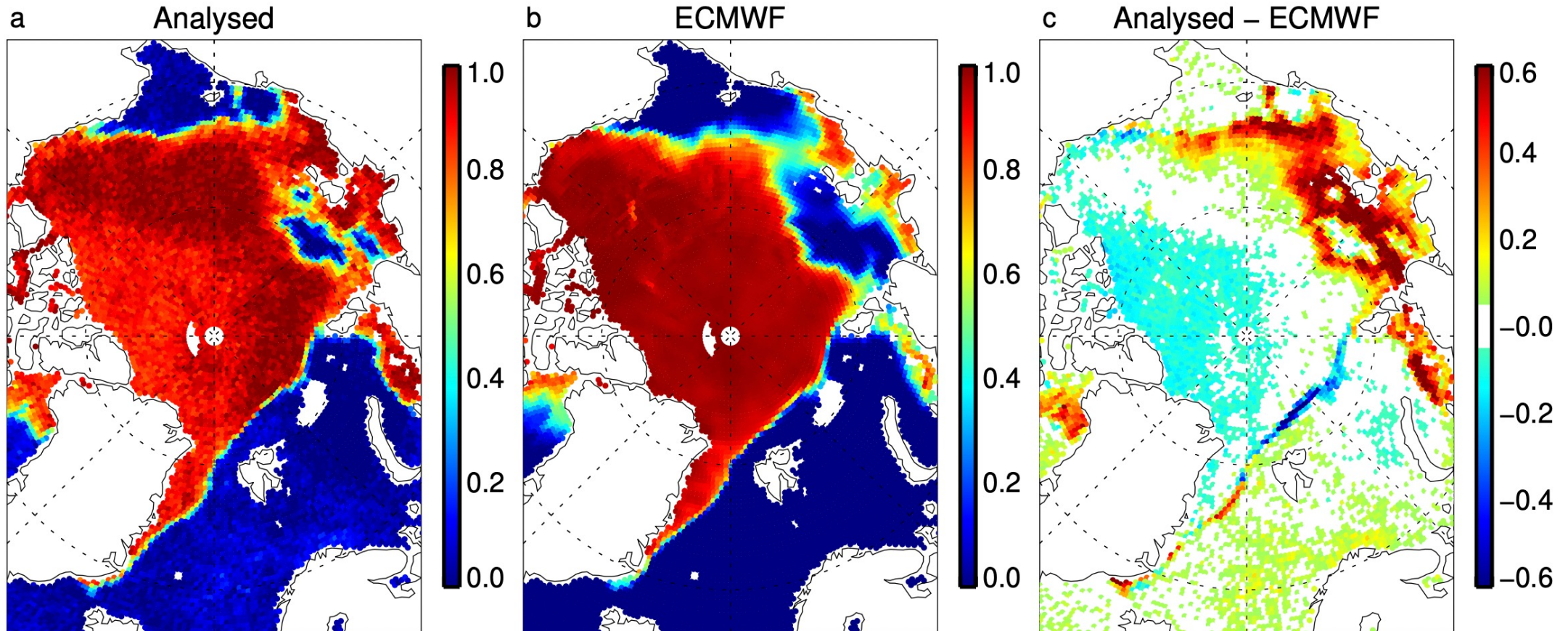
adding higher latitudes, land surfaces, mixed scenes (land – water)
(but excluding sea-ice, snow, high altitudes, desert soils)

For cycle 49r1 (June 2024?)– assimilation of microwave imagers over sea ice

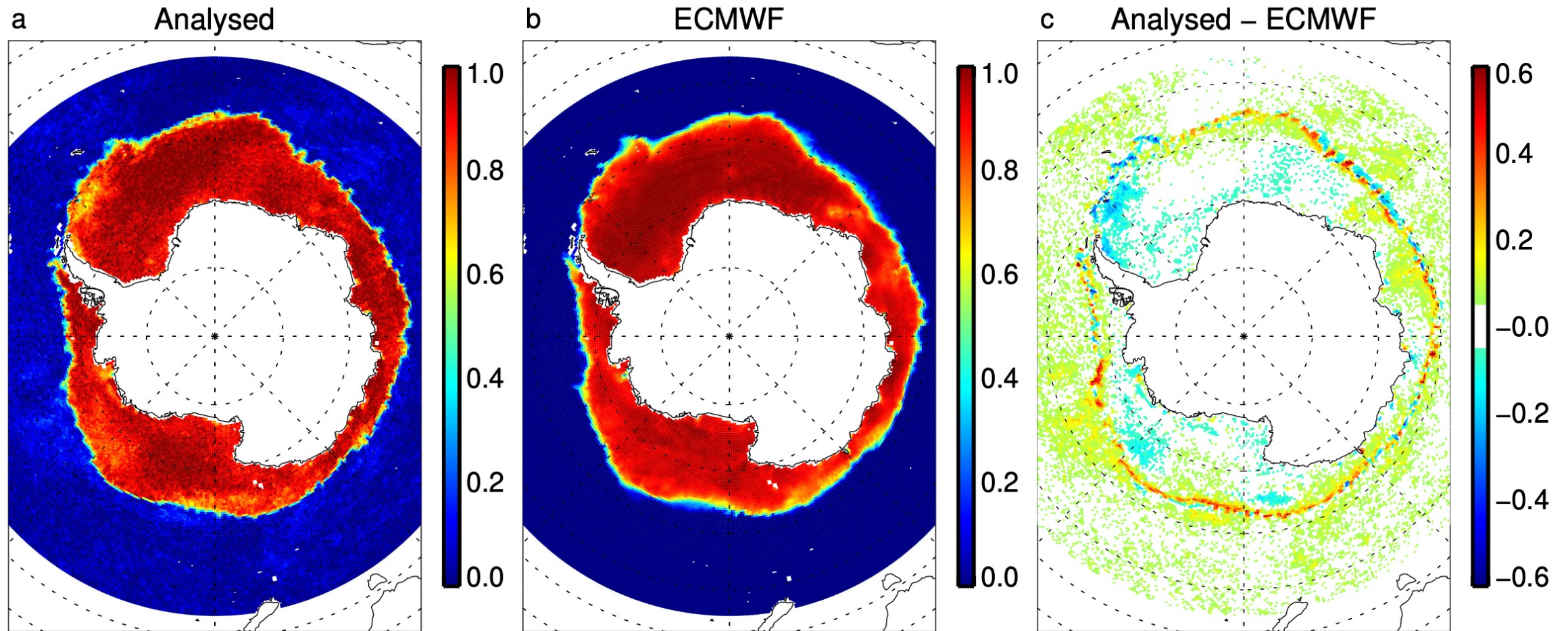
- Offline hybrid machine learning – data assimilation technique to create a surface emissivity model for sea ice, driven by a set of “empirical state” variables describing unknown aspects of the sea ice surface (grain size, ice age, wetness, etc.)
- In ECMWF 4D-Var – use the empirical surface emissivity model with an observation space “sink variable” to estimate the sea ice fraction and the empirical state.
 - Water vapour, cloud and precipitation signal above the sea ice is also used
 - Significant improvement in southern ocean forecasts (T, wind) from being able to assimilate observations over ocean but in close proximity to sea ice

Rapid freeze-up: sea ice fraction - 7th Nov 2020

IFS sea-ice patterns do eventually catch up with AMSR2 retrievals after 2-3 days

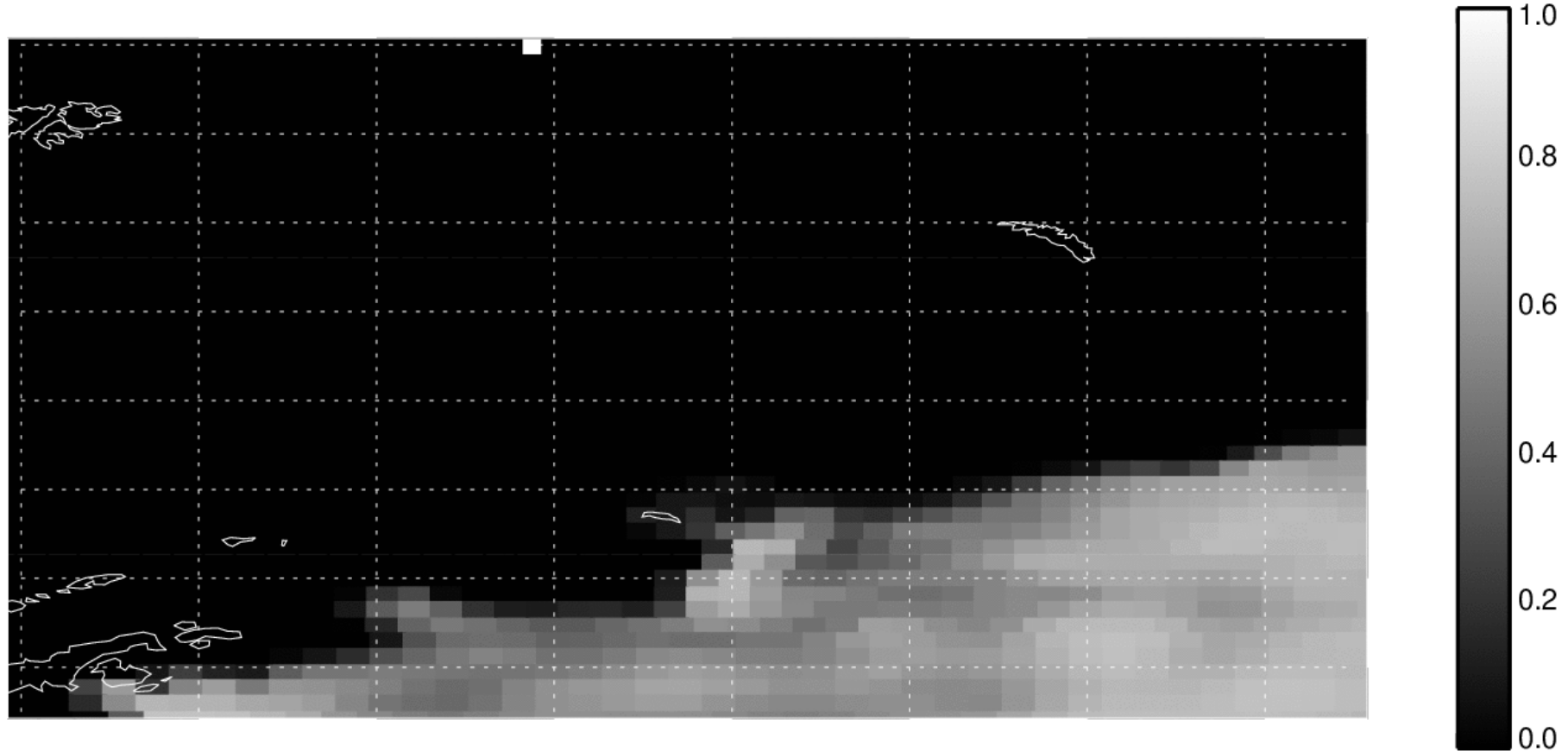


Antarctic example – sea ice fraction 7th August 2020



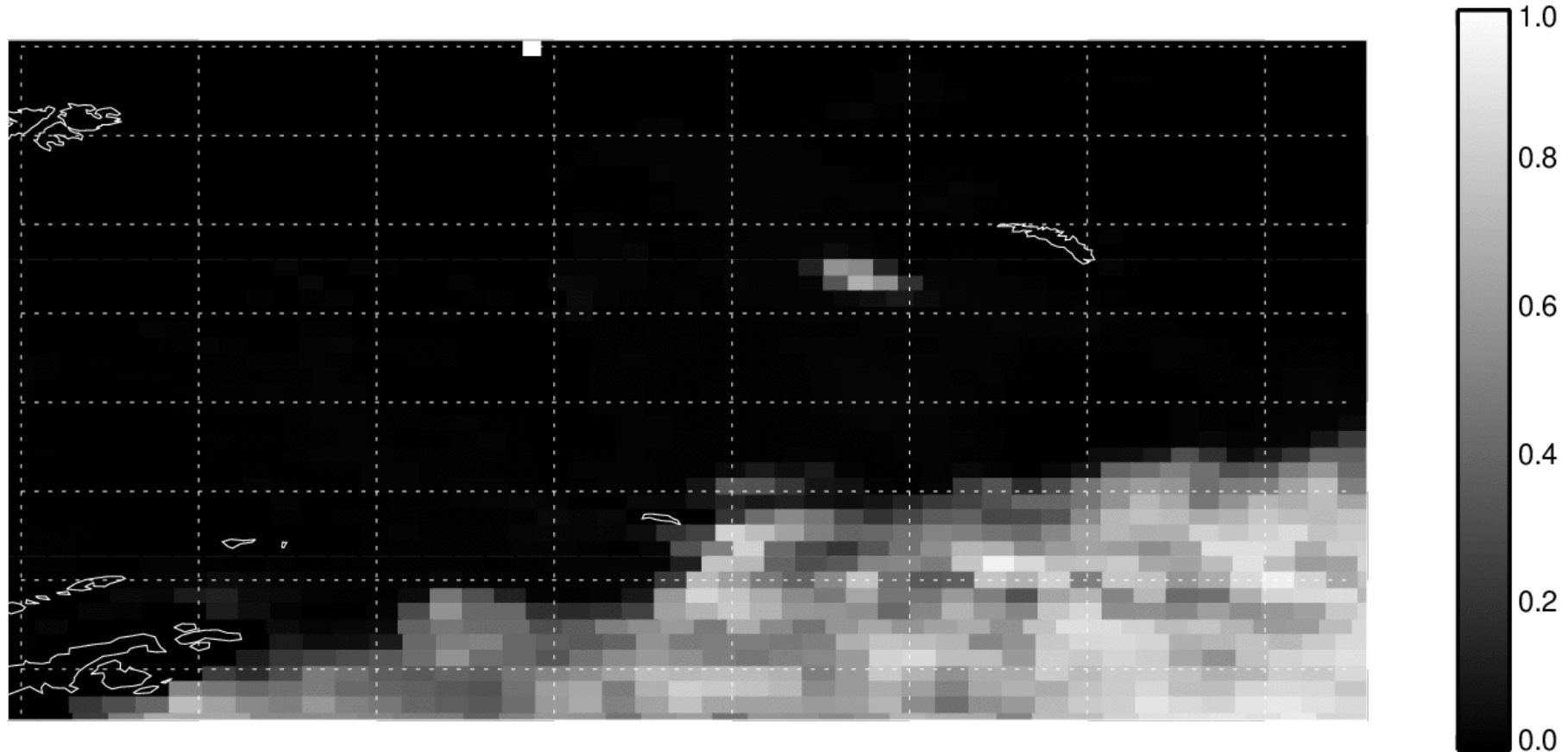
IFS sea ice concentration at AMSR2 locations

12Z 2-Dec-2020



AMSR2 sink variable sea ice concentration

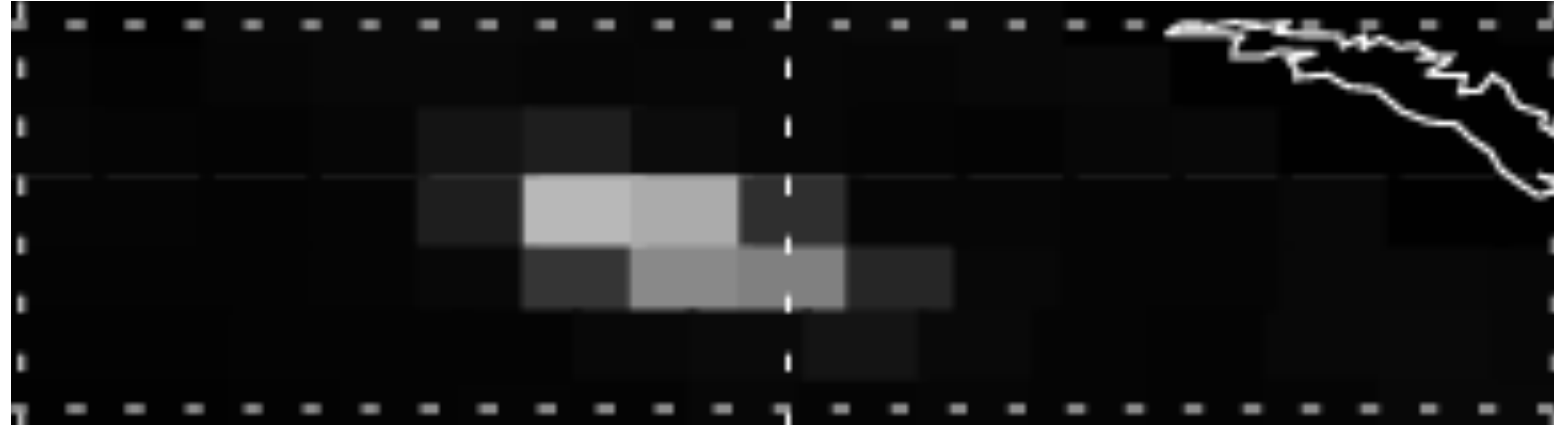
12Z 2-Dec-2020



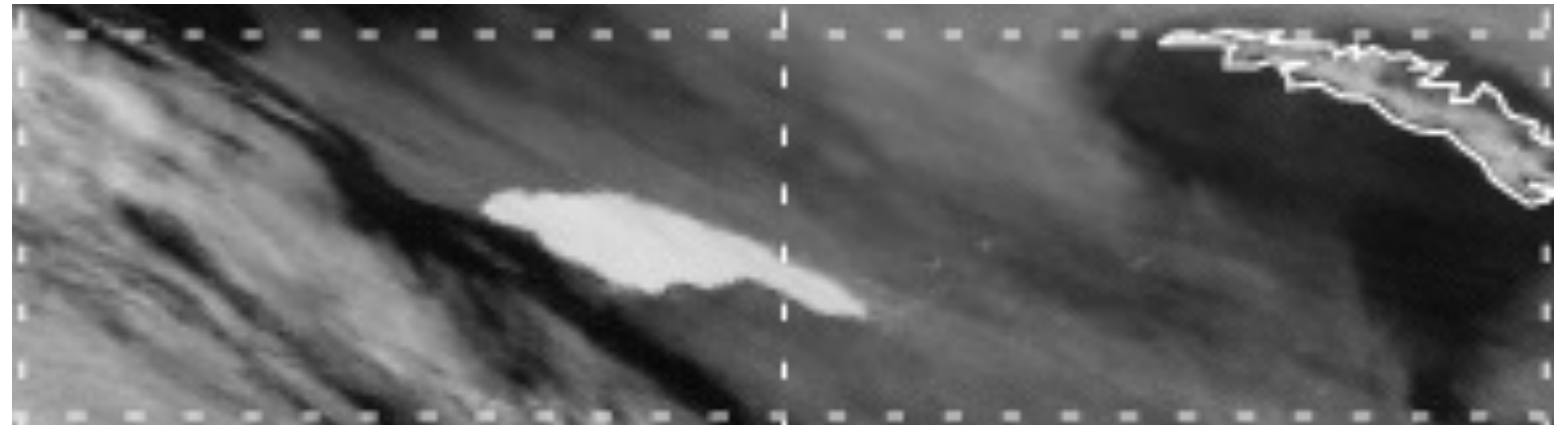
AMSR2 sea ice fraction vs OLCI image: A68A iceberg

AMSR2

1 pixel ~
40x40 km



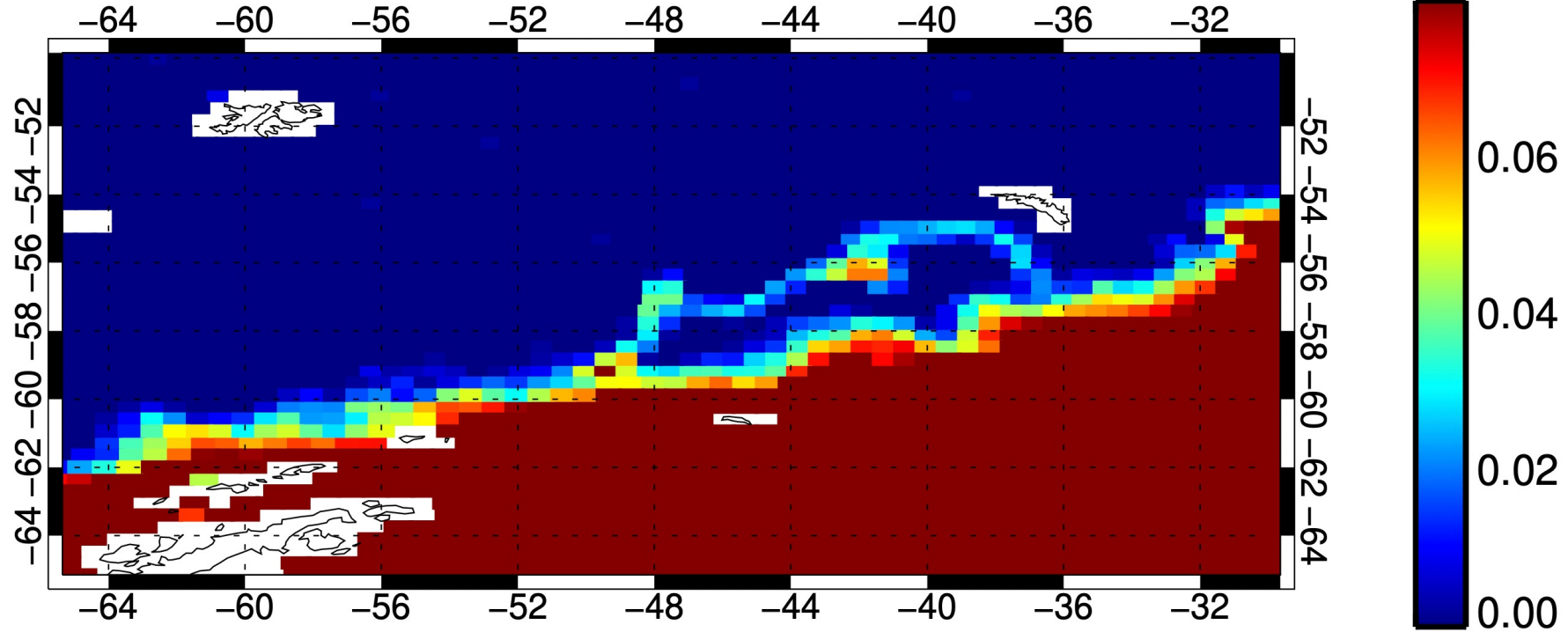
12Z 4th Dec 2020



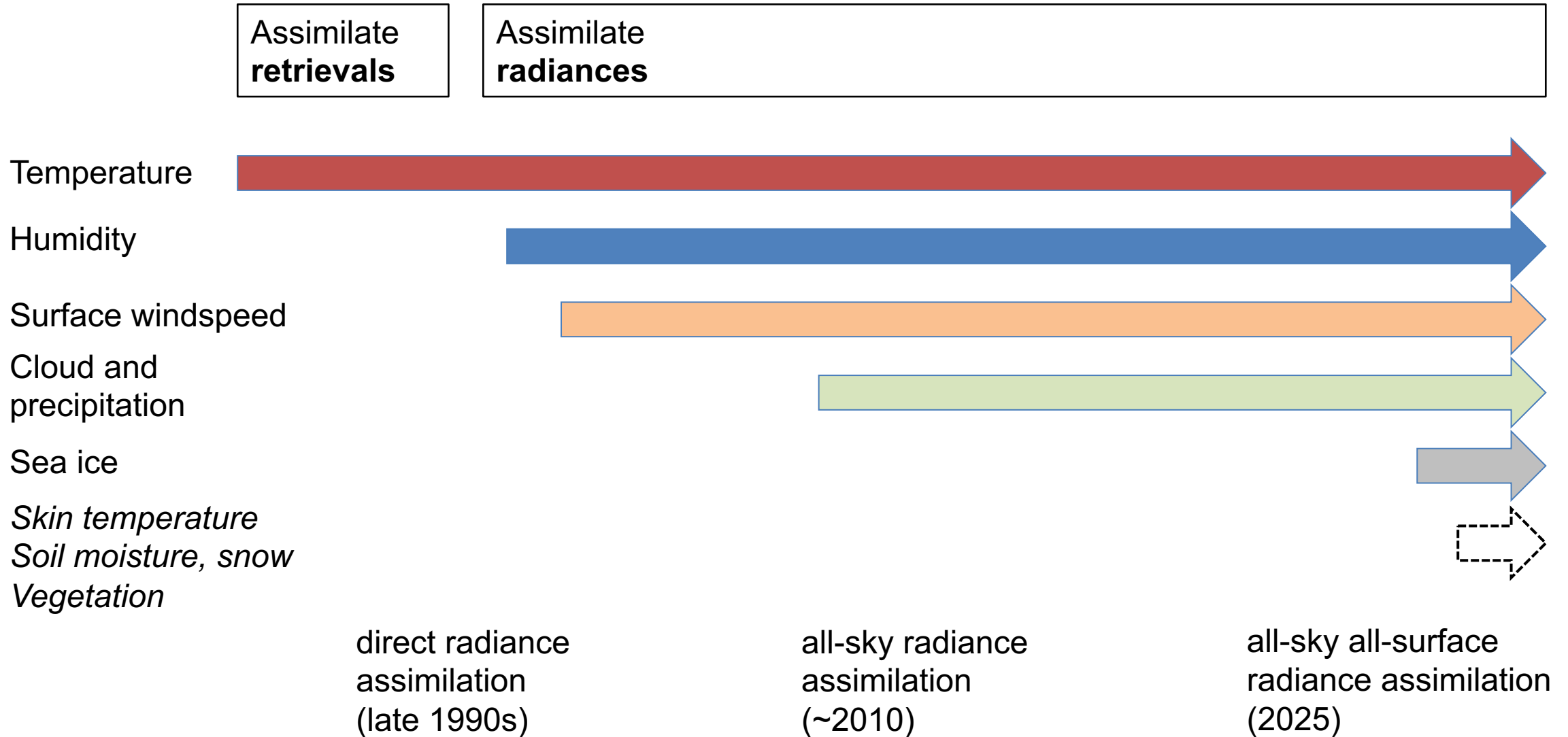
OLCI channel
10 (681 nm)

a

Mean sea ice concentration Sep – Dec 2020



Rough timeline of satellite microwave data assimilation in 'atmospheric' DA



Questions?