Role of Ocean Surface Waves in Modulating Tropical Cyclone Intensity

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Motivation

Tropical cyclone (TC) stands as one of the most damaging natural hazards in the world, posing threats to lives and property in coastal areas through strong winds, heavy rain, storm surges and landslides upon landfall. Despite significant progress in forecasting the TC tracks, improving the forecast of TC intensity remains a significant challenge. A TC can be conceptualized as a Carnot engine that draws energy from the warm ocean through the flux exchanges at the air-sea interface. Theoretical and observational evidence points to the potential critical role of ocean surface waves in modulating the air-sea fluxes, suggesting these wavy surfaces may help shape the intricate feedback mechanisms between TCs and the underlying ocean. However, these processes are not well-represented in numerical models.

Modelling Approach



FIO-AOW (Zhao et al., 2017; 2022; 2024) is a fully coupled Atmosphere-Ocean-Wave model. It consists of atmosphere component WRF, the 3rd generation ocean surface wave model MASNUM, circulation and ocean components POM&ROMS. These three components are integrated through a coupler known as Ccommunity Coupler. The code of FIO-AOW are assessable publicly https://github.com/Biao-Zhao/FIO-AOW.

Wave Coupling

The atmospheric and oceanic models in FIO-AOW are tightly coupled through ocean surface waves. Currently, the following wave-related physical processes have been considered to represent the physical interplay between air and ocean, including:

1.Sea-state-dependent air-sea momentum flux

A new air-sea momentum flux scheme considering both wave state (Janssen, 1989) and saturation (decrease) effect at high winds is introduced by

$$\begin{split} z_{o} &= \alpha_{ch} \frac{u_{*}^{2}}{g} + \frac{0.11\nu}{u_{*}} \\ \alpha_{ch} &= \frac{1}{\sqrt{1 - \frac{\vec{\tau}_{in}}{\vec{\tau}_{a}}}} \begin{cases} exp(-0.005755U_{10}^{2} + 0.3225U_{10} - 8.956) & 30 < U_{10} \le 55 \\ 0.00002 & U_{10} > 55 \end{cases} \\ \vec{\tau}_{in} &= \rho_{w}g \int_{0}^{2\pi} \int_{0}^{\infty} \frac{\vec{k}}{\omega} S_{in}(\omega, \theta) \, d\omega d\theta \end{split}$$

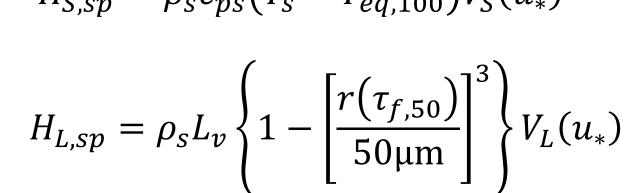
2. Sea spray-mediated air-sea enthalpy flux

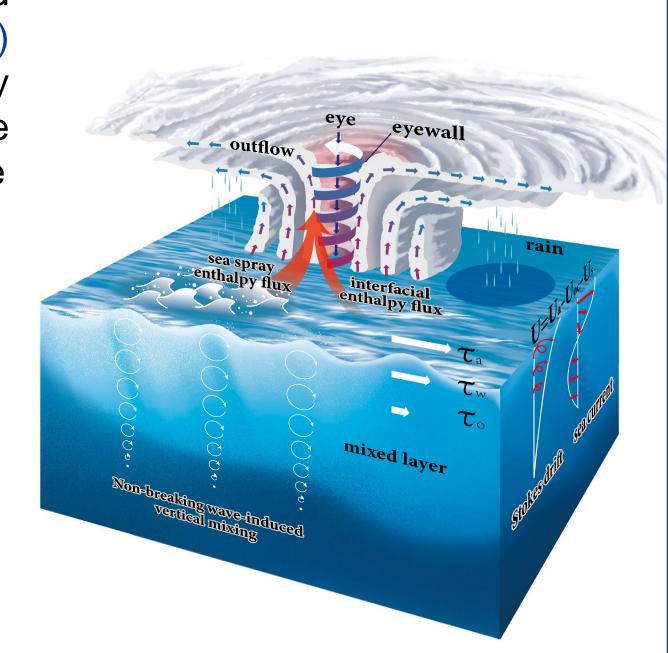
The sea spray-mediated sensible and latent fluxes (Andreas et al., 2015) have been added to the boundary layer interfacial heat fluxes in the atmosphere model boundary layer like

$$H_{L,T} = H_L + H_{L,Sp}$$

$$H_{S,Sp} = \rho_S c_{pS} (T_S - T_{eq,100}) V_S(u_*)$$

 $H_{S,T} = H_S + H_{S,sp}$





3. Non-breaking wave-induced vertical mixing

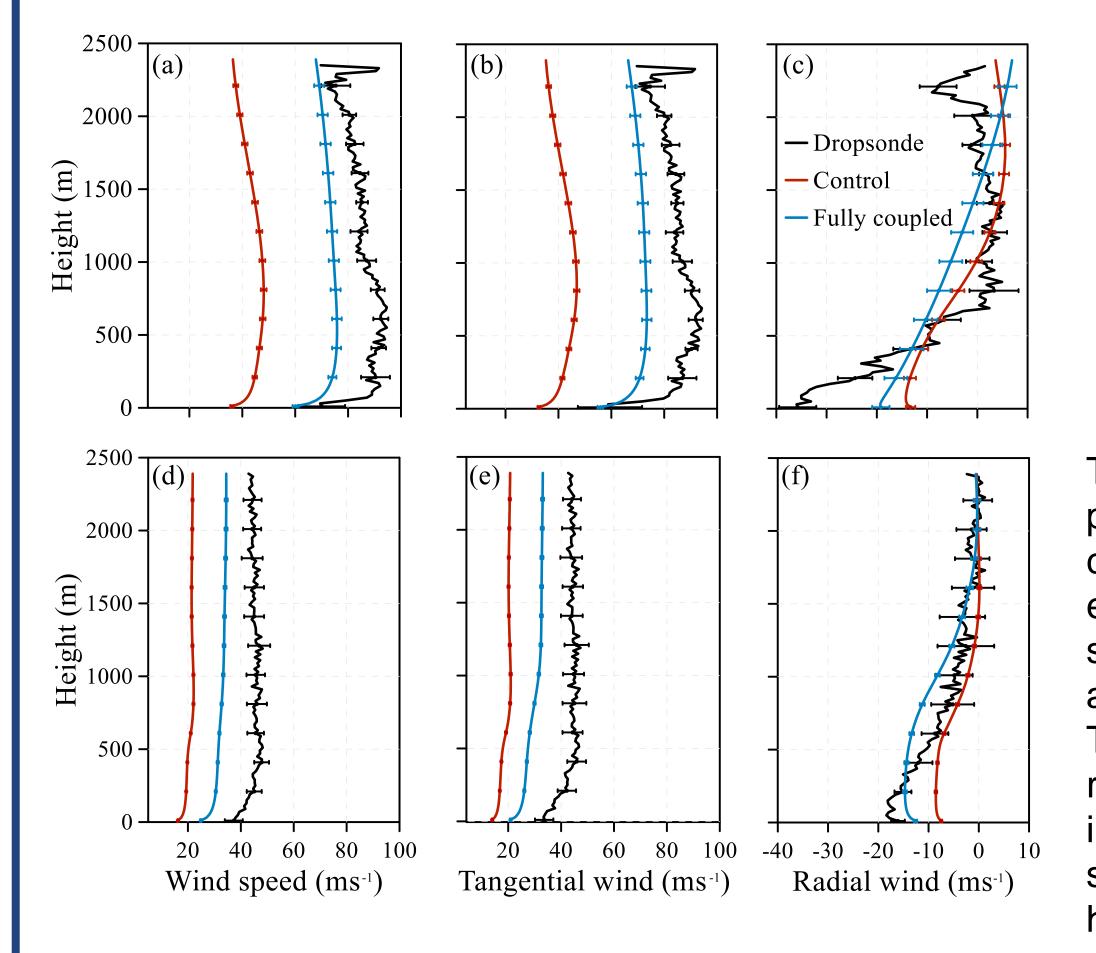
Turbulence generated by non-breaking waves (Qiao et al, 2010) is calculated by integrating wave spectrum as follows

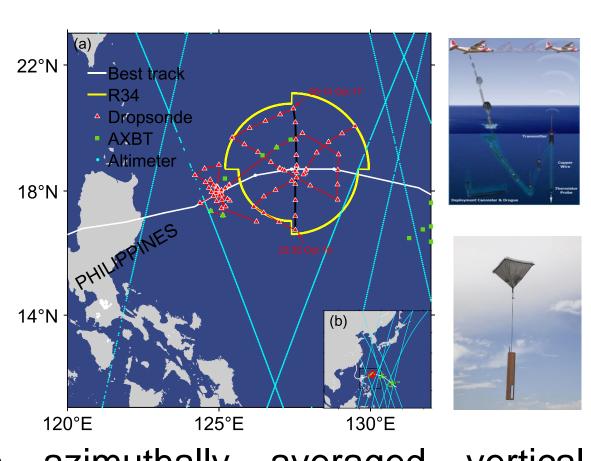
$$B_{v} = \alpha \int E(\vec{k}) \exp\{2kz\} d\vec{k} \frac{\partial}{\partial z} \left(\iint_{\vec{k}} \omega^{2} E(\vec{k}) \exp\{2kz\} d\vec{k} \right)^{\frac{1}{2}}$$

This coefficient is added to the vertical viscosity and diffusivity of ocean model by

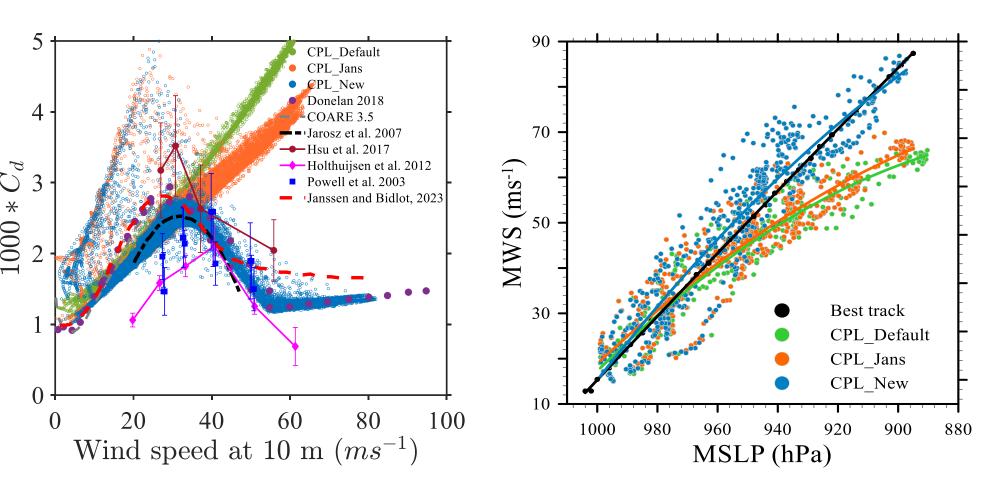
$$K_m = K_{mc} + B_v$$
, $K_h = K_{hc} + B_v$

Tropical Cyclone Intensity



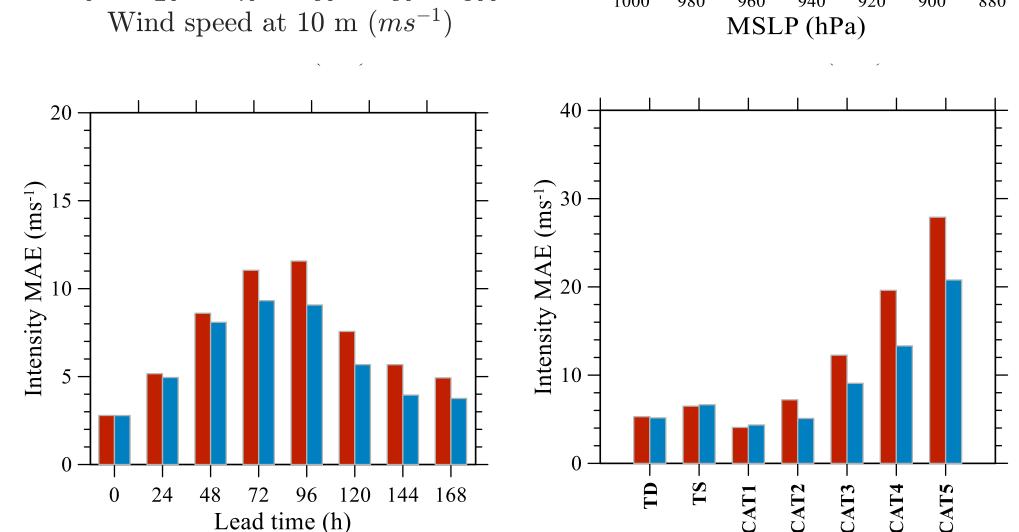


The azimuthally averaged vertical profiles of simulated wind speeds are compared with dropsonde in the eyewall (0.75≤r/RMW≤1.25, RMW stands for Radius of Maximum Wind) and outer core (3≤r/RMW≤5) regions. The introduction of surface wave related physical processes leads to improved wind speeds and TC structure in terms of boundary layer height and inflow strength.



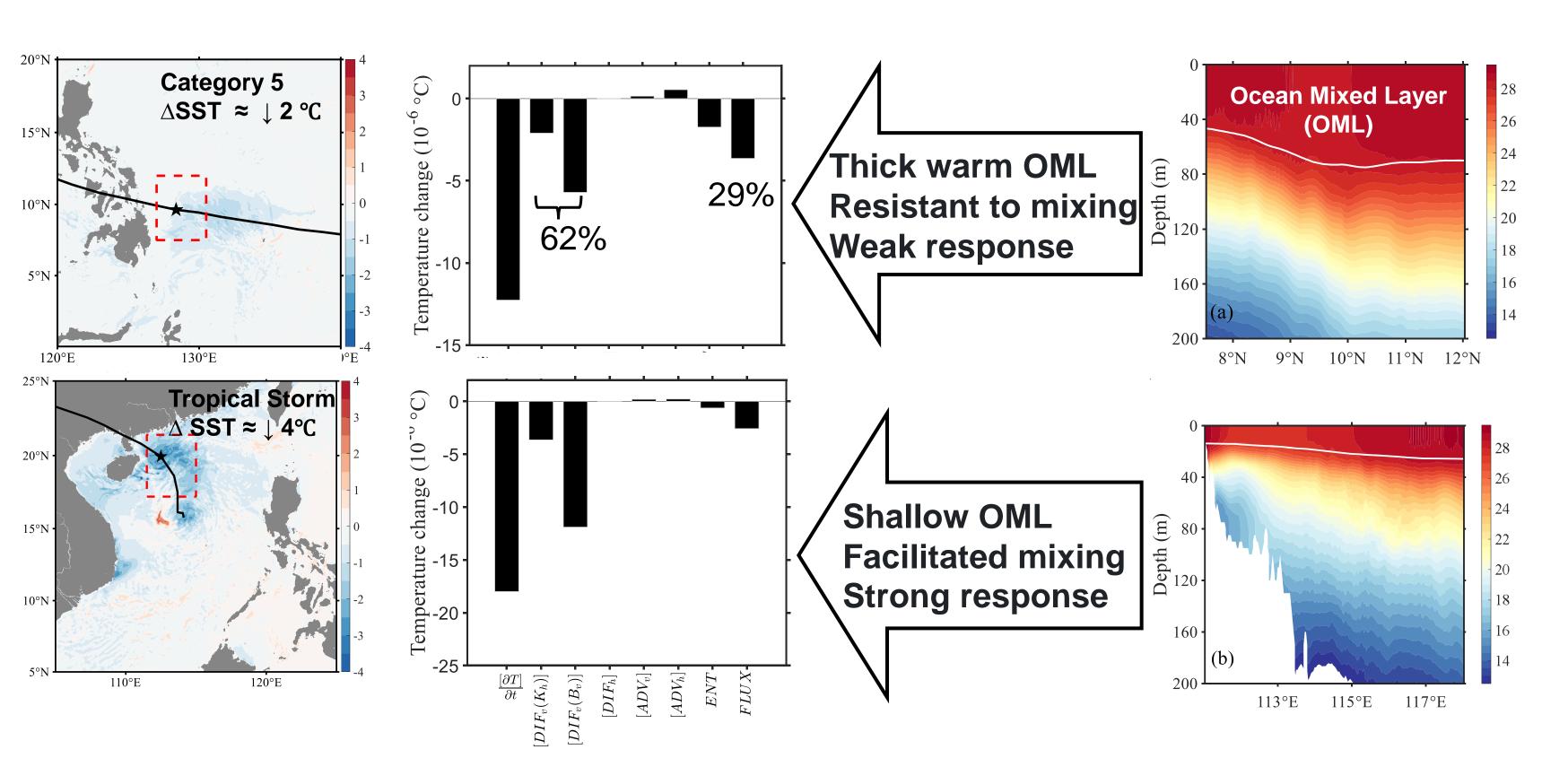


Saffir-Simpson



TCs of 2013 in western North Pacific are retrospectively simulated. Results suggest that the surface wave related physical processes play an important role in TC intensity and wind-pressure relationship. Wind-pressure relationship is very sensitive to the airsea momentum flux parameterization at high winds

Ocean Response



References

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