

# Introduction to the parameterization of Land Surfaces processes in coupled Earth System Modelling

Souhail Boussetta

Land modelling Team, Earth System Modelling Section

[Souhail.Boussetta@ecmwf.int](mailto:Souhail.Boussetta@ecmwf.int), Office 1114

Thanks to Gabriele Arduini and all the land modelling team,  
and all previous contributors to the land modeling processes



ECMWF TC2025 - PA - Land Surface

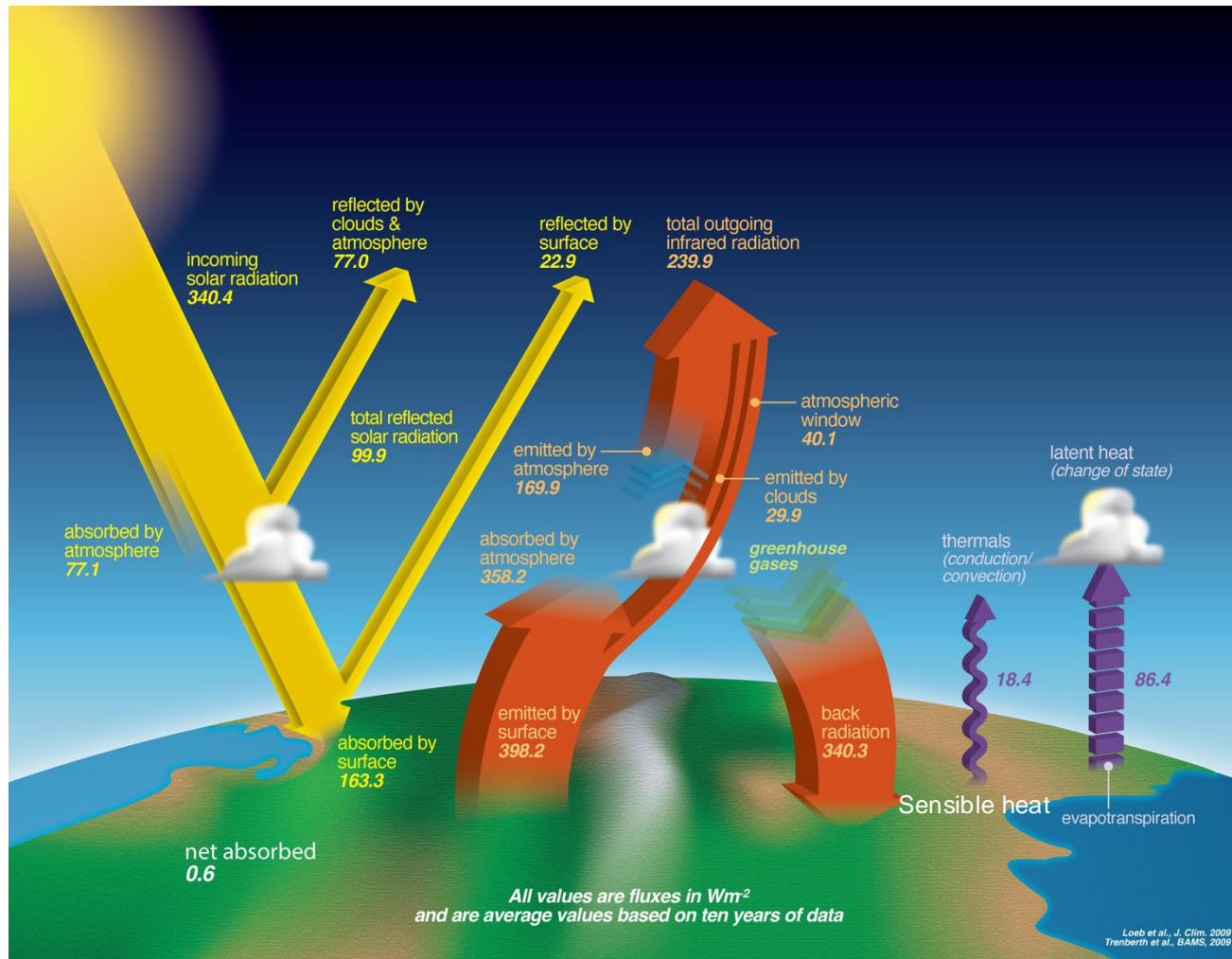


© ECMWF March 4, 2025

# Outline

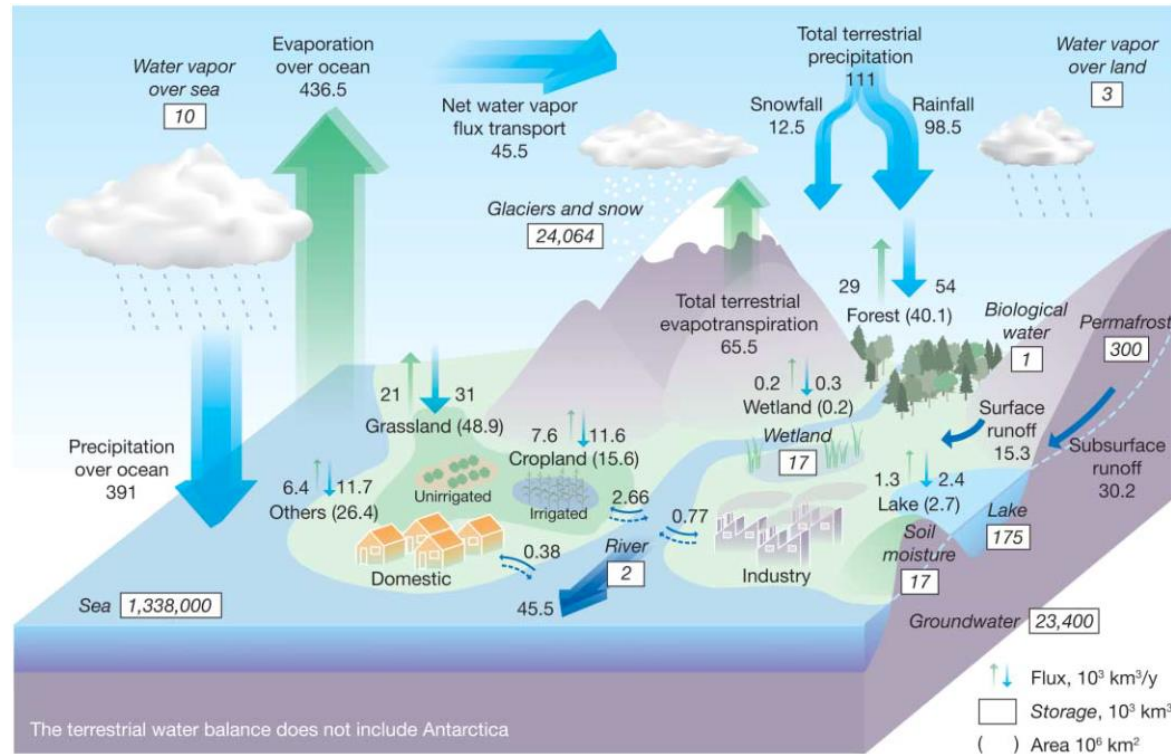
- Land-surface processes in the earth system:  
why is this important for Numerical Weather Predictions?
- Representing land-surface heterogeneities for NWP applications:
  - Coupling with the atmosphere
  - Surface energy balance and surface (skin) temperature
  - Energy and water budgets in the soil
  - Inland water bodies (lakes)
- Perspectives

# The land surface in the earth energy budget



- The surface is a physical boundary for atmospheric processes
- amount of energy reflected: surface albedo
- Conduction into the soil: amount of energy conducted and absorbed in the soil
- Emitted by the surface: surface temperature and surface emissivity
- Surface sensible heat flux
- Evapotranspiration (latent heat) flux

# The land-surface in the earth global water cycle



**Fig. 1.** Global hydrological fluxes ( $1000 \text{ km}^3/\text{year}$ ) and storages ( $1000 \text{ km}^3$ ) with natural and anthropogenic cycles are synthesized from various sources (1, 3–5). Big vertical arrows show total annual precipitation and evapotranspiration over land and ocean ( $1000 \text{ km}^3/\text{year}$ ), which include annual

precipitation and evapotranspiration in major landscapes ( $1000 \text{ km}^3/\text{year}$ ) presented by small vertical arrows; parentheses indicate area ( $\text{million km}^2$ ). The direct groundwater discharge, which is estimated to be about 10% of total river discharge globally (6), is included in river discharge.

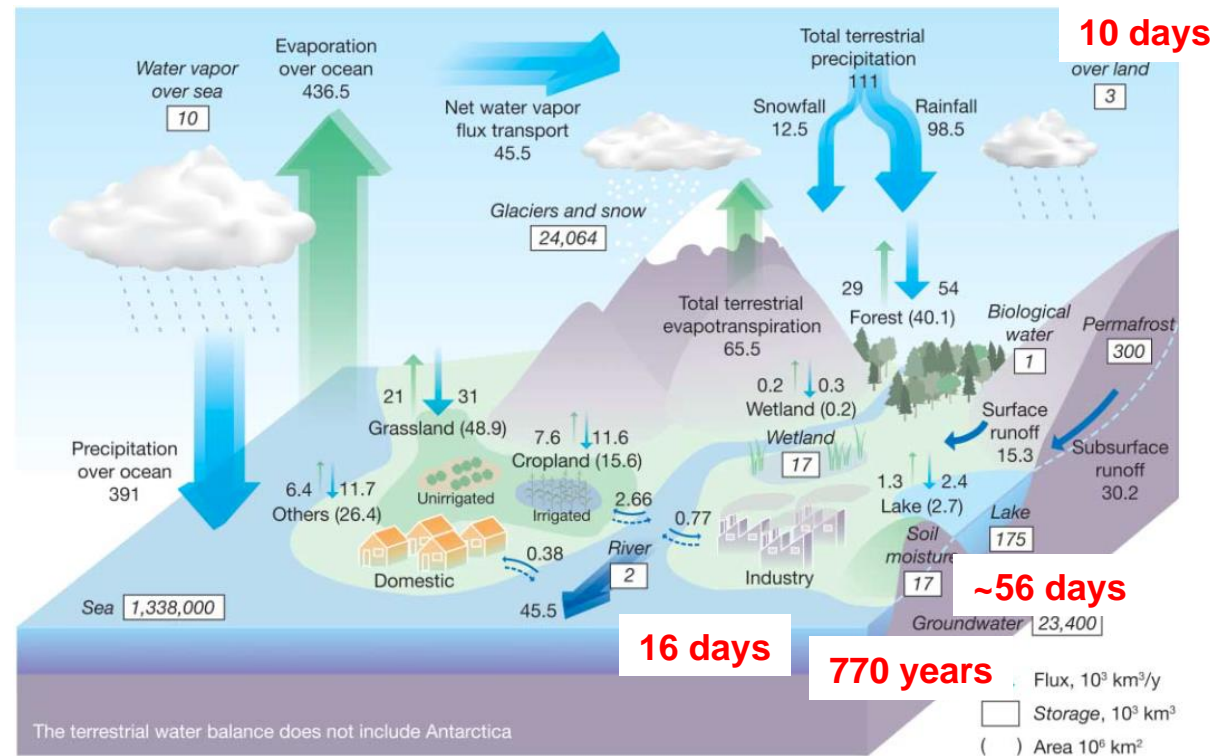
**Total land precipitation:**  
60% evapotranspiration  
40% runoff

**Total land evapotranspiration:**  
44% Forests  
32% Grassland  
12% Cropland  
12% Others  
} 88% via vegetation !

Oki and Kanae, Science 2006. Fluxes / Storage are estimates from different sources

[https://www.researchgate.net/publication/6856186\\_Global\\_Hydrological\\_Cycles\\_and\\_World\\_Water\\_Resources](https://www.researchgate.net/publication/6856186_Global_Hydrological_Cycles_and_World_Water_Resources)

# Earth Global Water Cycle – Reservoirs and timescales



**Total land precipitation:**  
 60% evapotranspiration  
 40% runoff

**Total land evapotranspiration:**  
 44% Forests  
 32% Grassland  
 12% Cropland  
 12% Others

**88% via vegetation !**

**Residence time = Reservoir size / flux**  
 Atmosphere : 10 days  
 Rivers : 16 days  
 Soil moisture : 56 days  
 Groundwater: 770 years

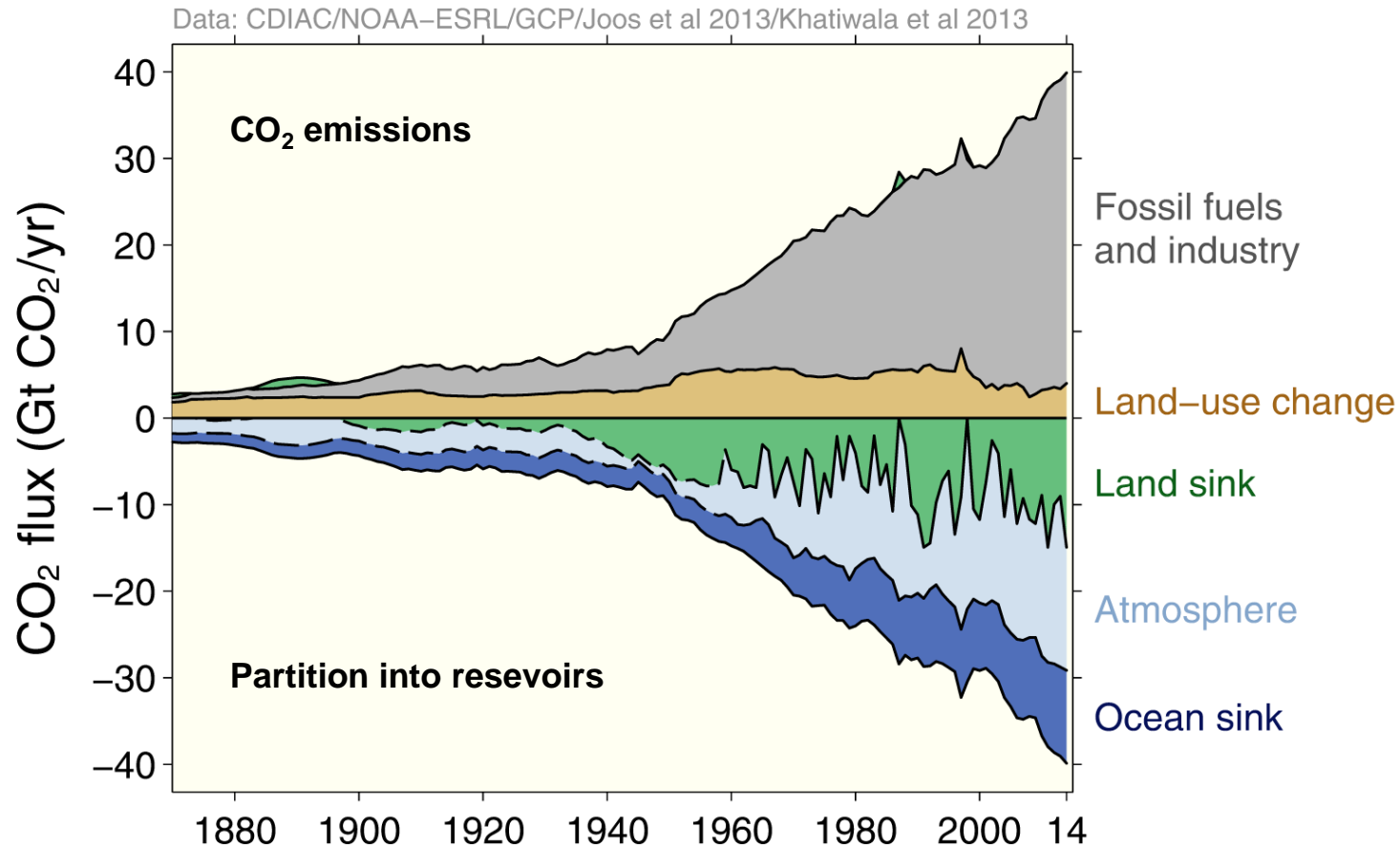
**Fig. 1.** Global hydrological fluxes ( $1000 \text{ km}^3/\text{year}$ ) and storages ( $1000 \text{ km}^3$ ) with natural and anthropogenic cycles are synthesized from various sources (1, 3–5). Big vertical arrows show total annual precipitation and evapotranspiration over land and ocean ( $1000 \text{ km}^3/\text{year}$ ), which include annual

precipitation and evapotranspiration in major landscapes ( $1000 \text{ km}^3/\text{year}$ ) presented by small vertical arrows; parentheses indicate area ( $\text{million km}^2$ ). The direct groundwater discharge, which is estimated to be about 10% of total river discharge globally (6), is included in river discharge.

Oki and Kanae, Science 2006. Fluxes / Storage are estimates from different sources

[https://www.researchgate.net/publication/6856186\\_Global\\_Hydrological\\_Cycles\\_and\\_World\\_Water\\_Resources](https://www.researchgate.net/publication/6856186_Global_Hydrological_Cycles_and_World_Water_Resources)

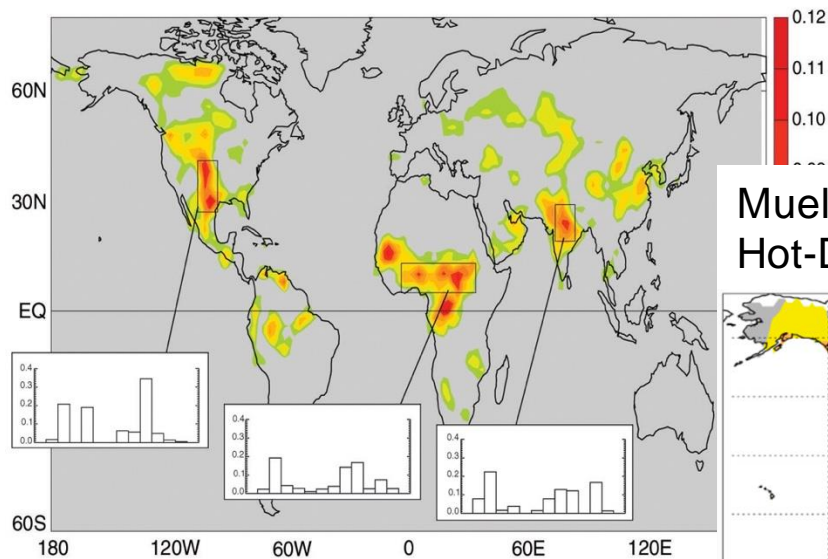
# Earth Global Carbon Cycle - Reservoirs



# Earth surface role, experimental evidence (soil moisture)

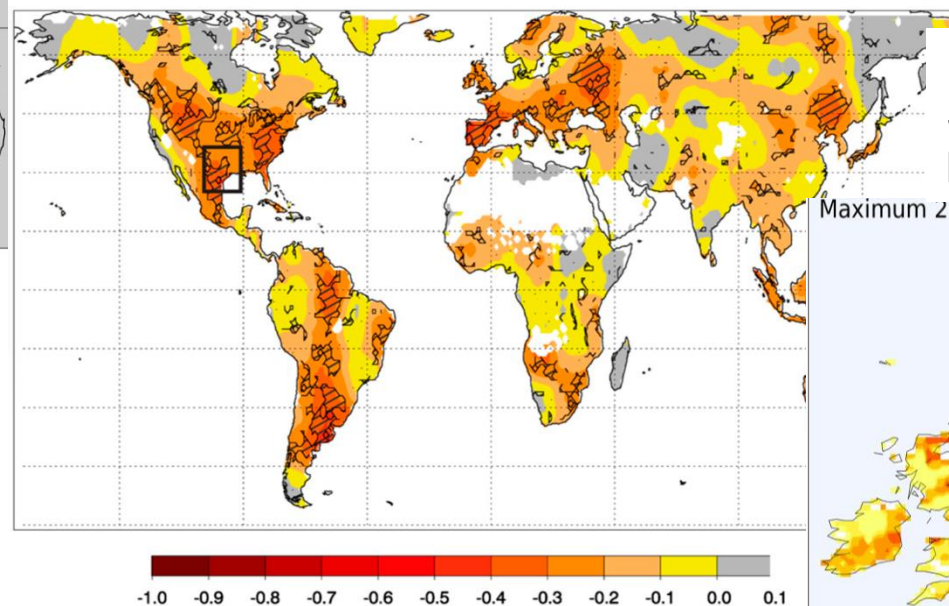
Koster et al. 2004 Science

Land-atmosphere coupling (SoilMoist-Precip feedback), JJA



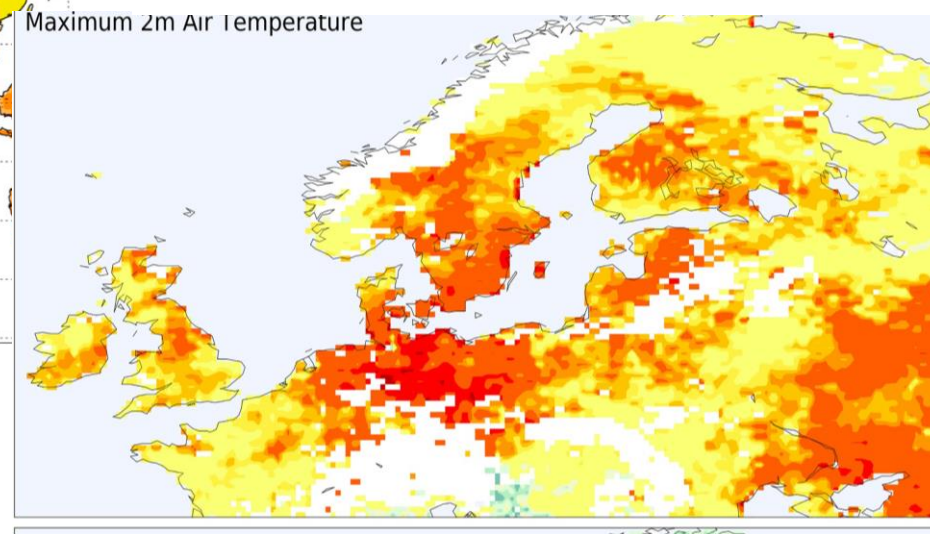
Mueller and Seneviratne 2012 PNAS

Hot-Days correlation with 3-month antecedent P deficit



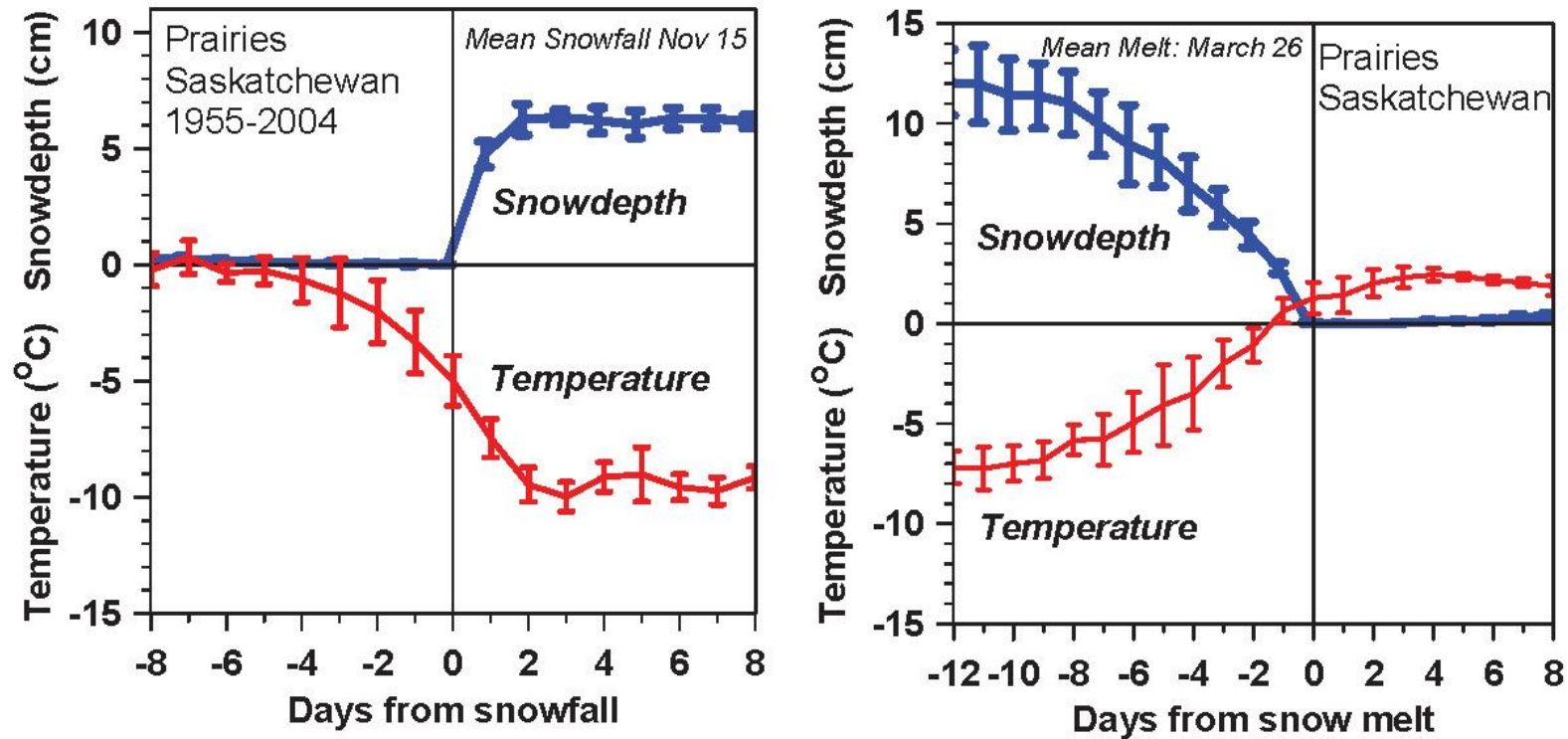
Dirmeyer et al. 2018

SM-Tmax coupling during North Europe heat wave in summer 2018



Albergel et al. 2013 JHM show dominance of significant drying trends for soil moisture in both reanalysis and satellite-based soil moisture dataset, with possibly larger areas of land surface predictability

# Earth surface role, observational evidence (snow)



***Snow reflects sunlight; shift to cold stable BL***

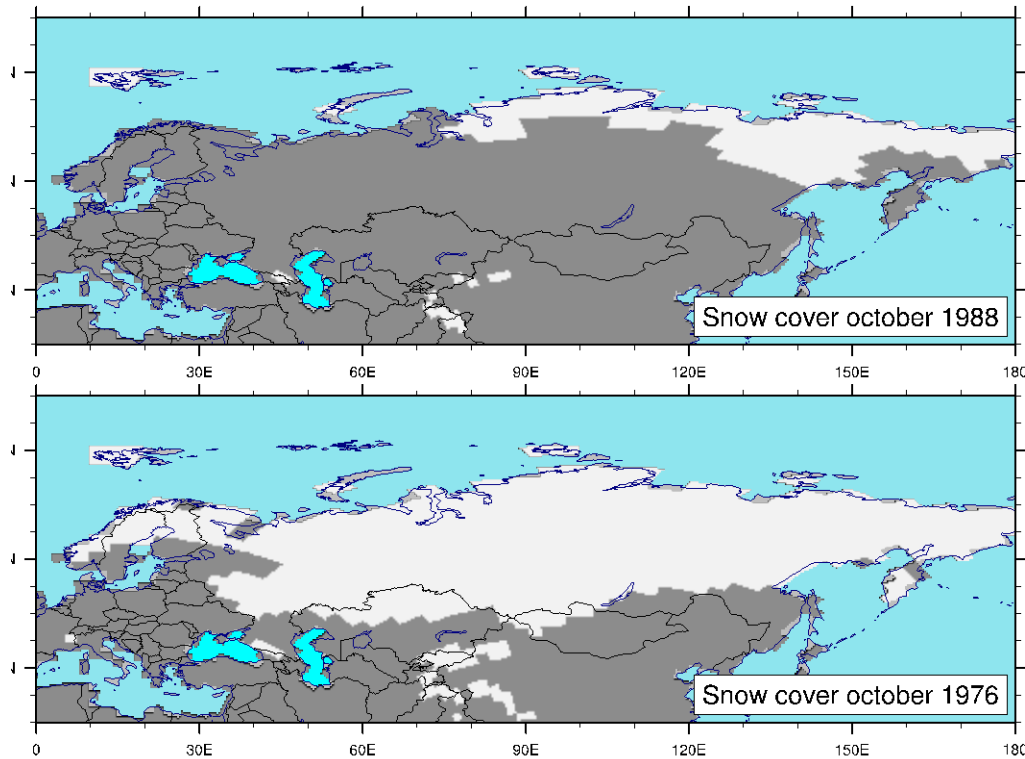
***Local climate switch between warm and cold seasons***

***Winter comes fast with snow***

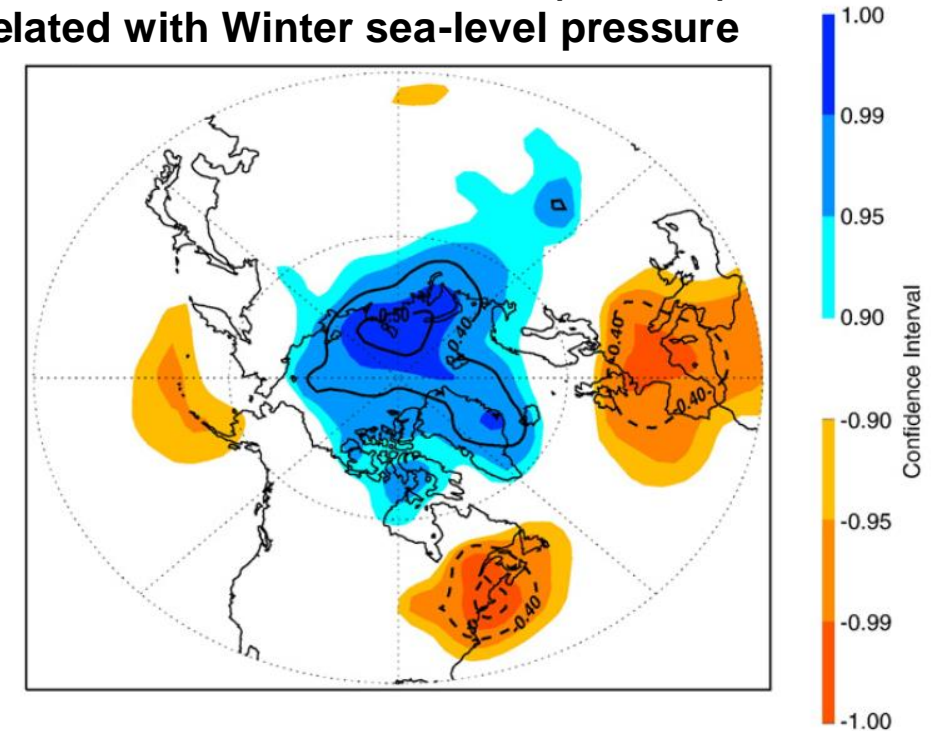
*Betts et al. 2014*

# Earth surface role, snow cover variability

Snow cover shows large interannual variability



October snow cover anomalies (Eurasia)  
correlated with Winter sea-level pressure



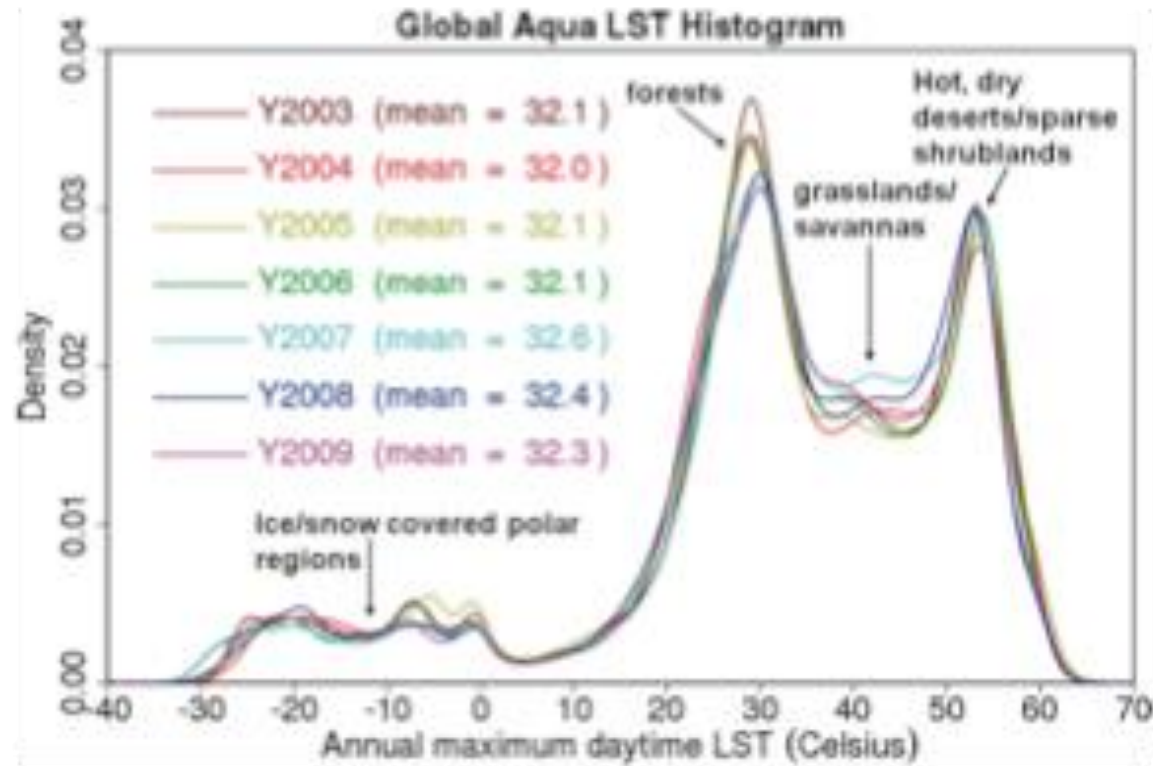
(Cohen and Saito 2003,  
Gong et al. 2007)

The spatial anomaly patterns resemble the Artic Oscillation  
pattern of variability

# Modelling surface heterogeneity and coupling with the atmosphere

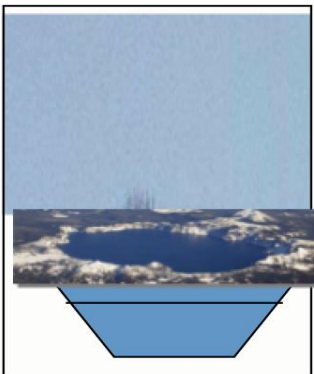
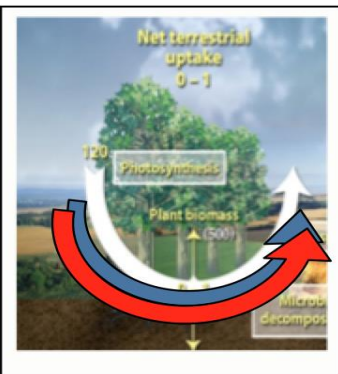
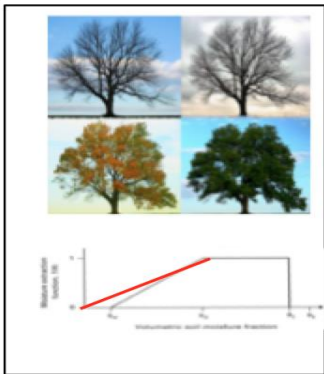
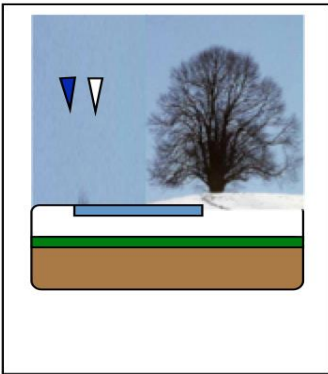
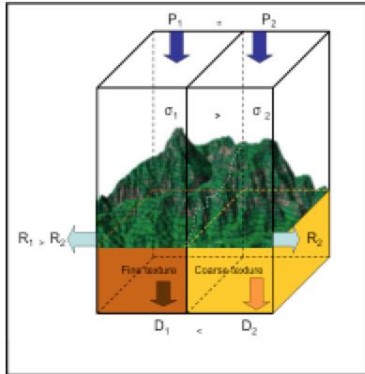
Different surfaces have different surface properties, leading to large differences in surface temperature, energy and water fluxes and their diurnal cycle

Histogram of maximum surface temperature from satellite captures the influence of land-cover types



# Earth surface modelling components @ECMWF

- Hydrology-TESSEL**  
Balsamo et al. 2009
  - tiling scheme
  - 4-layer soil heat and water budget
  - 7 soil texture/type
  - Variable infiltration and runoff
- Snow module**
  - snow multi-layer (5 layers, Arduini et al. 2019)
  - single layer (Dutra et al. 2009)
- Vegetation & LAI**  
Boussetta et al. (2013)  
New satellite-based  
Leaf-Area-Index
- Carbon**  
Integration of Carbon/Energy/Water  
Boussetta et al. 2013  
Agesti-Panareda et al. 2015
- Lake & Coastal area**  
Mironov et al (2010),  
Dutra et al. (2010),  
Balsamo et al. (2012, 2010)  
Extra tile (9) to for sub-grid lakes and ice  
LW tiling (Dutra)
- Sea-ice module**  
4-layer Ice thermodynamic model  
Fixed sea-ice depth, 1.5m



## ECMWF land-surface, "ecLand"

Land surface 1D-model: soil, snow, vegetation, lakes and coastal water. Same horizontal resolution as atmosphere.

Ocean, waves and sea-ice\* : NEMO3.4, ecWAM, LIM2

\*only sea-ice fraction

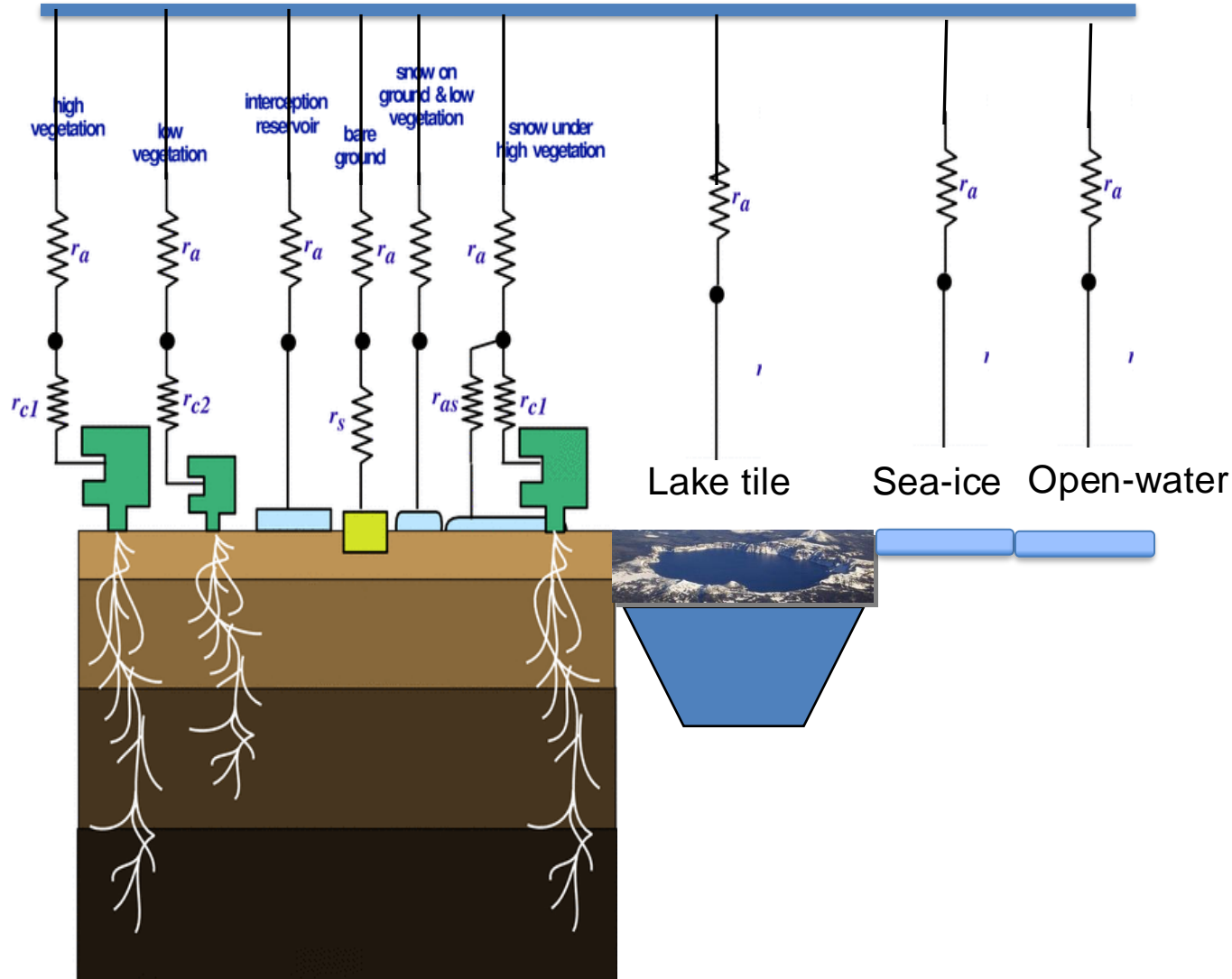


EUROPEAN CENTRE FOR MEDIUM-RANGE WEATHER FORECASTS

Atm/Land resol.	ECMWF Config, 2023
80 km	ERA1
32 km	ERA5 SEAS5
9 km	ENS
9 km	HRES

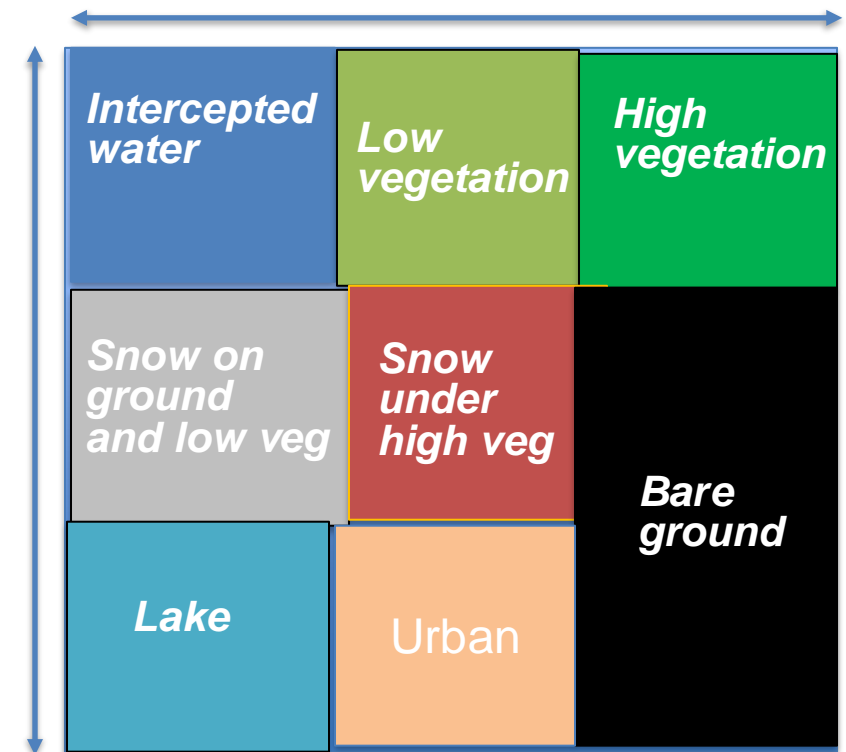
# Modelling surface heterogeneity: ecLand tiling approach

## Lowest atmospheric model level

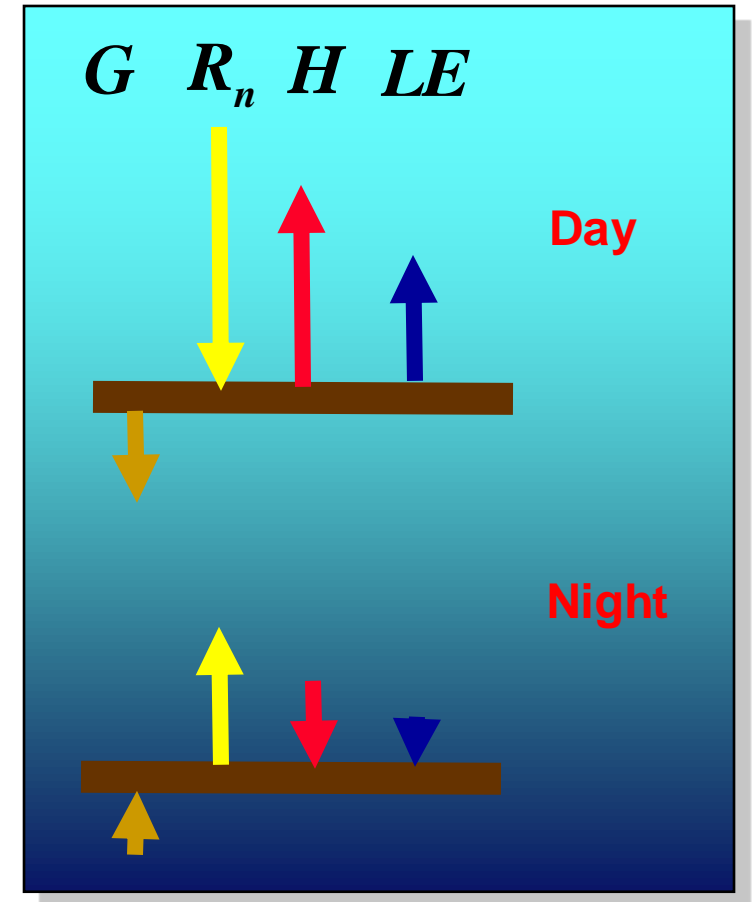
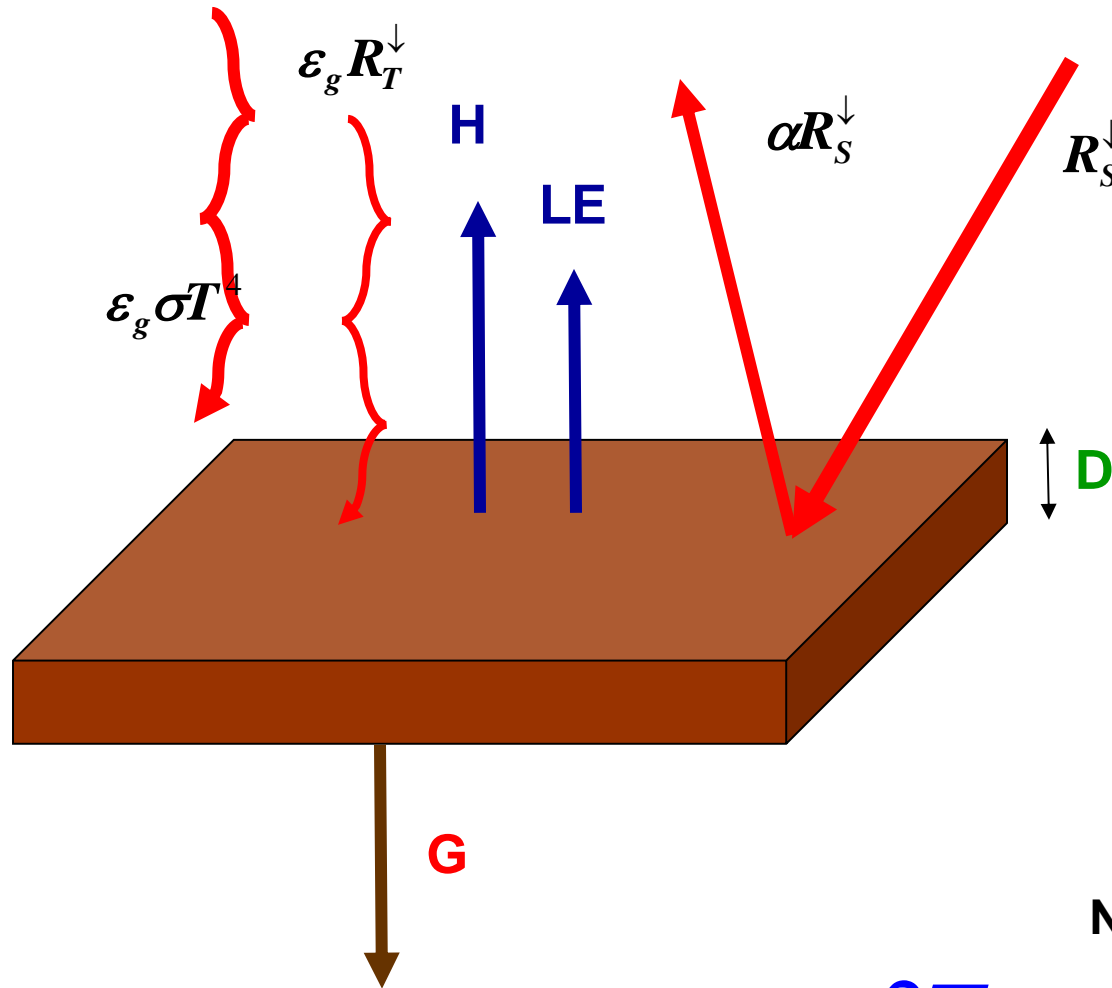


- Atmospheric resolution is too coarse to represent surface heterogeneities
- The surface energy balance and heat fluxes are computed for each tile
- Each tile communicates with the lowest model level above and snow/soil underneath

## Model grid-box over land



# Schematic for the energy balance at the surface



$$(\rho C)_g D \frac{\partial T_s}{\partial t} = R_n + H + LE + G$$

Net radiation  $\rightarrow R_n$   
 Sensible heat  $\rightarrow H$   
 Latent heat  $\rightarrow LE$   
 Conducted heat  $\rightarrow G$

# ecLand skin temperature equation, aerodynamic perspective

For each tile:

$$(1 - \alpha_i) R_S^\downarrow + \varepsilon_g R_T^\downarrow - \varepsilon_g \sigma T_{sk,i}^4 + \quad \leftarrow \text{Net radiation (Shortwave and longwave)}$$

$$\rho C_{h,i} u_L (C_p T_L + gz - C_p T_{sk,i}) + \quad \leftarrow \text{Turbulent Sensible heat flux}$$

$$\rho C_{h,i} u_L [a_{L,i} q_L - a_{s,i} q_{sat}(T_{sk,i}, p_s)] + \quad \leftarrow \text{Turbulent latent heat flux}$$

$$\Lambda_{sk,i} (T_s - T_{sk,i}) = \quad \leftarrow \text{Ground (basal) heat flux}$$

$$= 0 \quad \leftarrow \text{no storage term!}$$

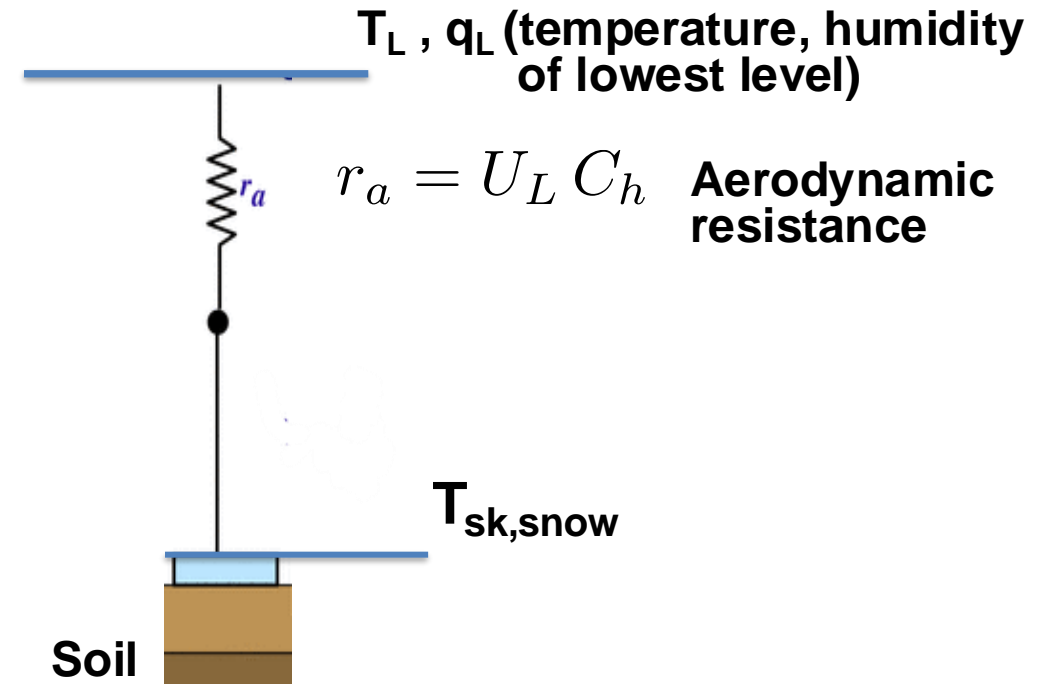
Grid-box quantities

$$H = \sum_i C_i H_i$$

$$E = \sum_i C_i E_i$$

$$T_{sk} = \sum_i C_i T_{sk,i}$$

$C_i$  Tile fraction



# ecLand heat transfer

Soil temperature equation, no phase changes

$$(\rho C)_g \frac{\partial T_s}{\partial t} = -\frac{\partial G}{\partial z} = \frac{\partial}{\partial z} \lambda_T \frac{\partial T}{\partial z}$$

$(\rho C)_g$  Soil volumetric heat capacity

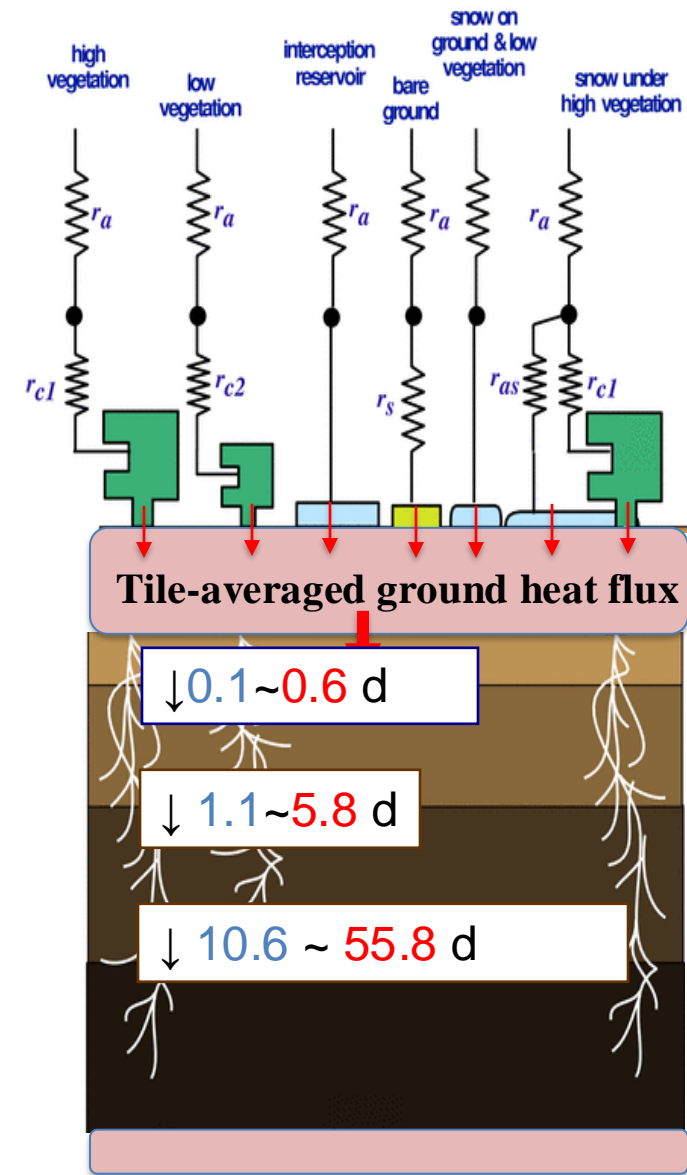
$\lambda_T$  Thermal conductivity

$k = \frac{\lambda_T}{(\rho C)_g}$  Thermal diffusivity

For an homogeneous soil,

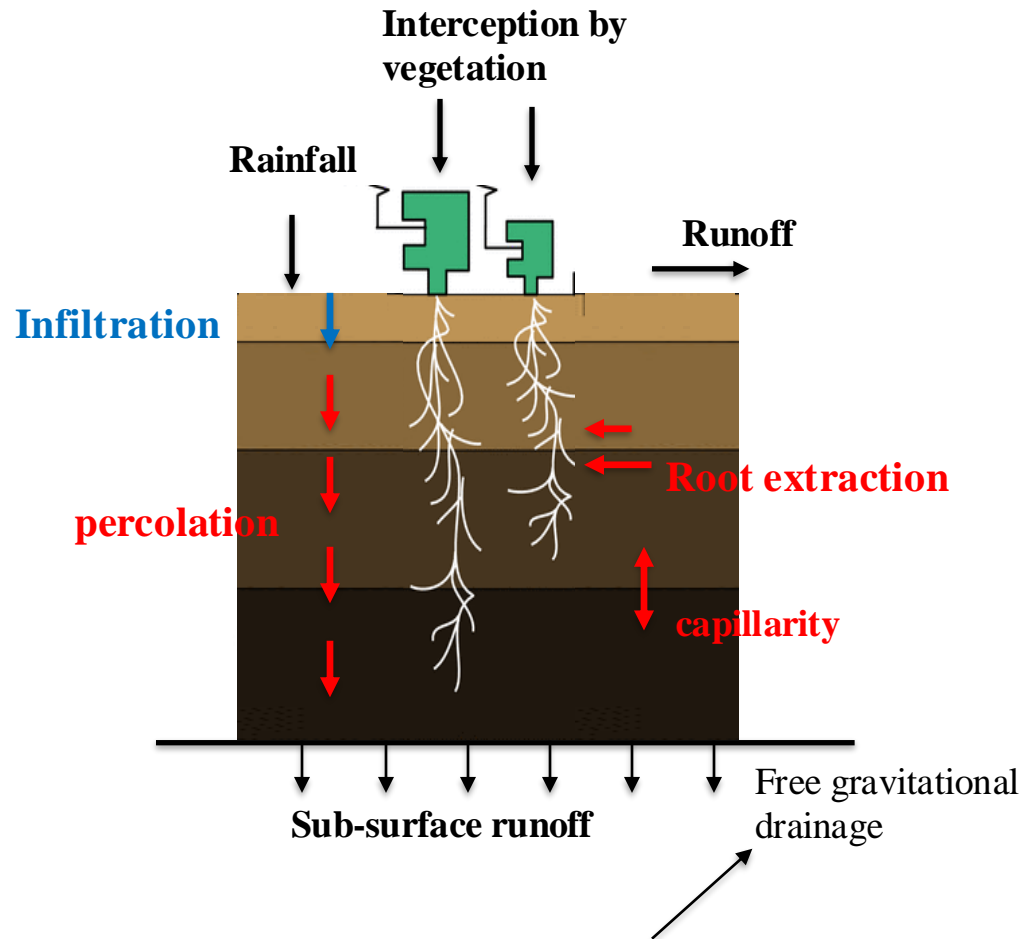
$$\frac{\partial T_s}{\partial t} = k \frac{\partial^2 T}{\partial z^2}$$

- Solution of heat transfer equation with the soil discretized in 4 layers of depths 7, 21, 72, and 189 cm.
- No-flux bottom boundary condition
- Heat conductivity dependent on soil water content
- Thermal effects of soil water phase change



Time-scale for downward heat transfers in wet/dry soil

# Soil water budget



No groundwater or bedrock representation

$$\rho_w \frac{\partial \theta}{\partial t} = -\frac{\partial F}{\partial z} + \rho_w S_\theta$$

$$\theta \text{ soil water } [ ] = m^3 m^{-3}$$

$$F \text{ Soil water flux } [ ] = kg m^{-2} s^{-1}$$

$$S_\theta \text{ Soil water source/sink, ie root extraction}$$

## Boundary conditions

Top: Infiltration = precip – interception – runoff

Bottom: Free drainage

## Root extraction

The amount of water transported from the root system up to the stomata and then available for transpiration

## Coupling with the soil temperature

In frozen soil, infiltration and percolation are minimum → most of water goes into runoff

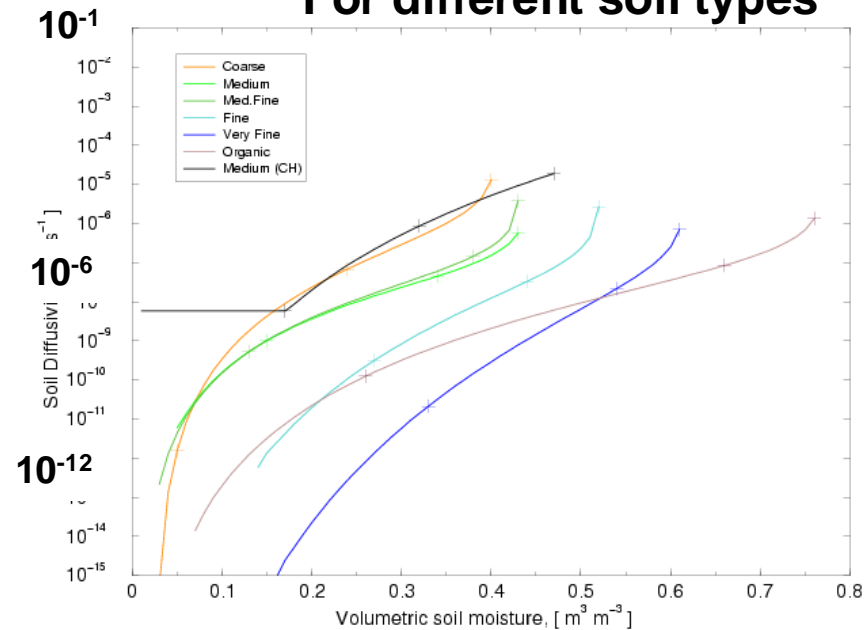
# Soil water flux

$$F = -\rho_w \left( \lambda \frac{\partial \theta}{\partial z} - \gamma \right) \quad \text{Darcy's law}$$

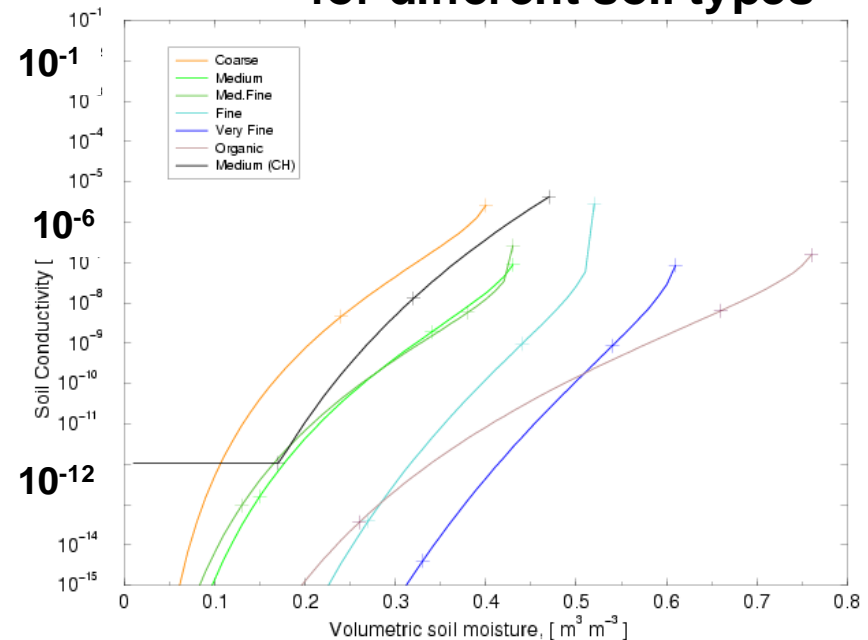
$\lambda$  hydraulic diffusivity  $[\lambda] = m^2 s^{-1}$

$\gamma$  hydraulic conductivity  $[\gamma] = m s^{-1}$

**Hydraulic Diffusivity,  
For different soil types**

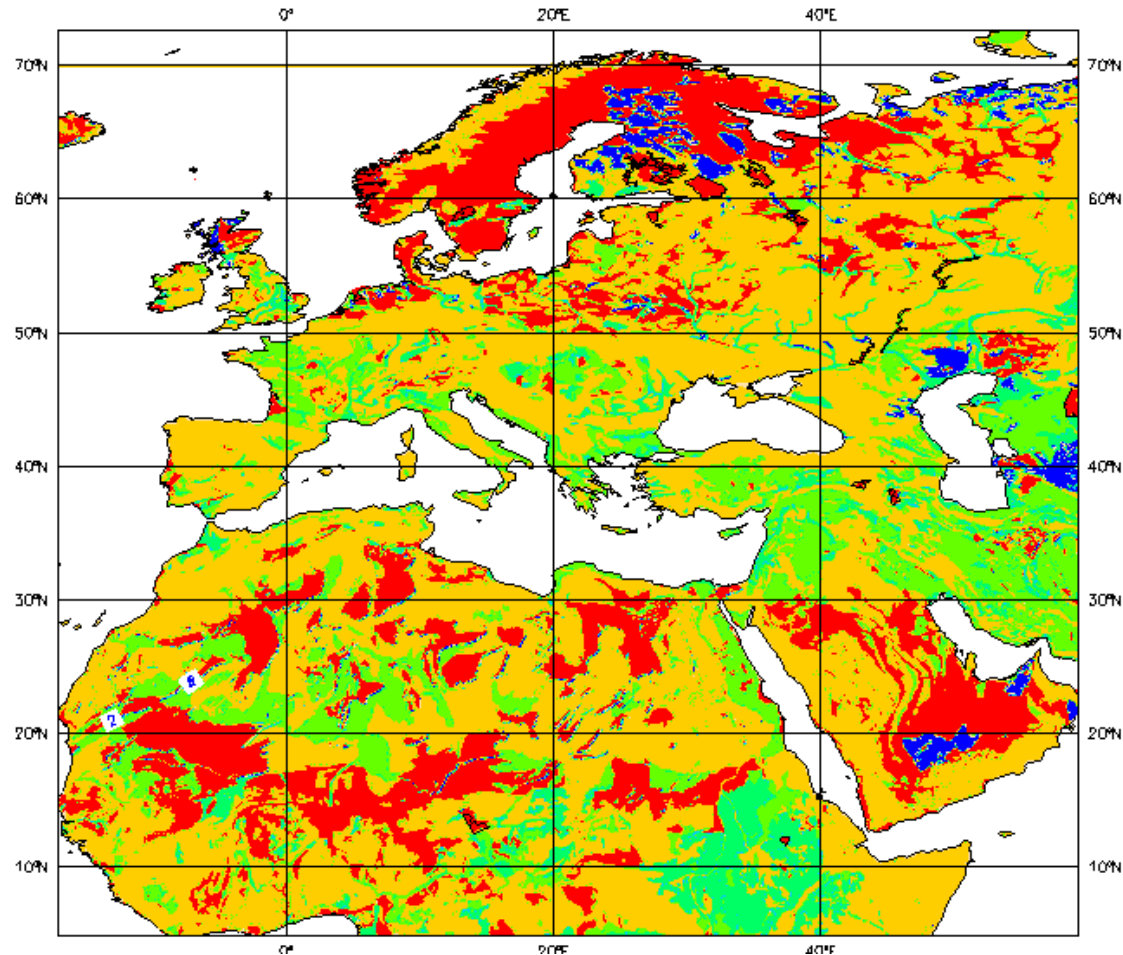


**Hydraulic conductivity,  
for different soil types**



# ecLand hydrology scheme(1)

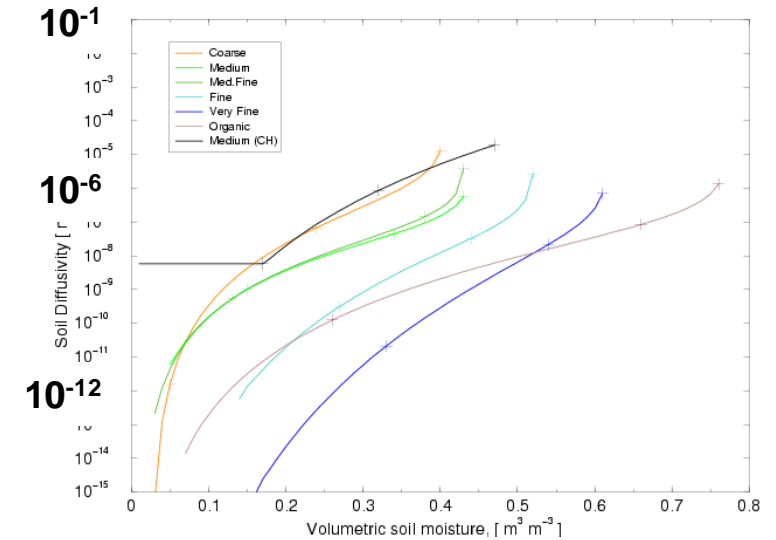
coarse medium med-fine fine very-fine organic



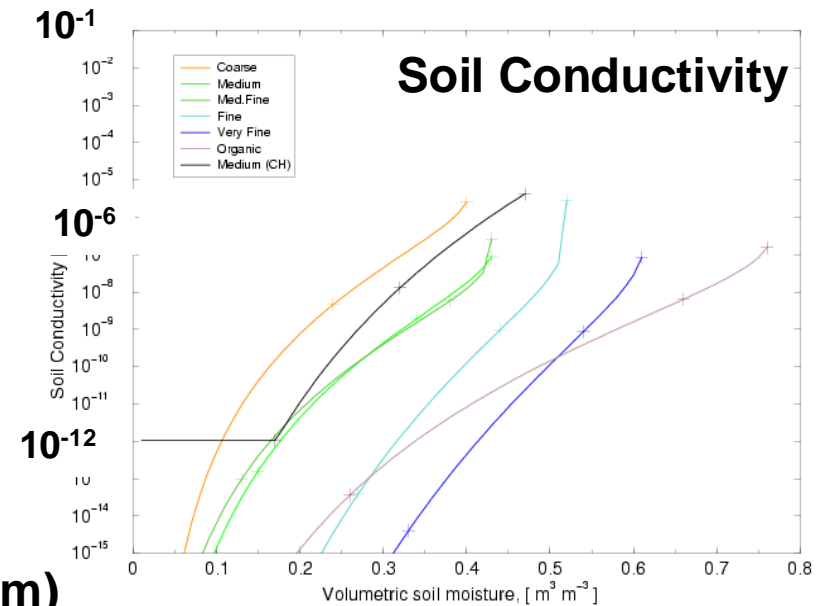
Dominant soil type from FAO2003 (at native resolution of ~ 10 km)



## Soil Diffusivity



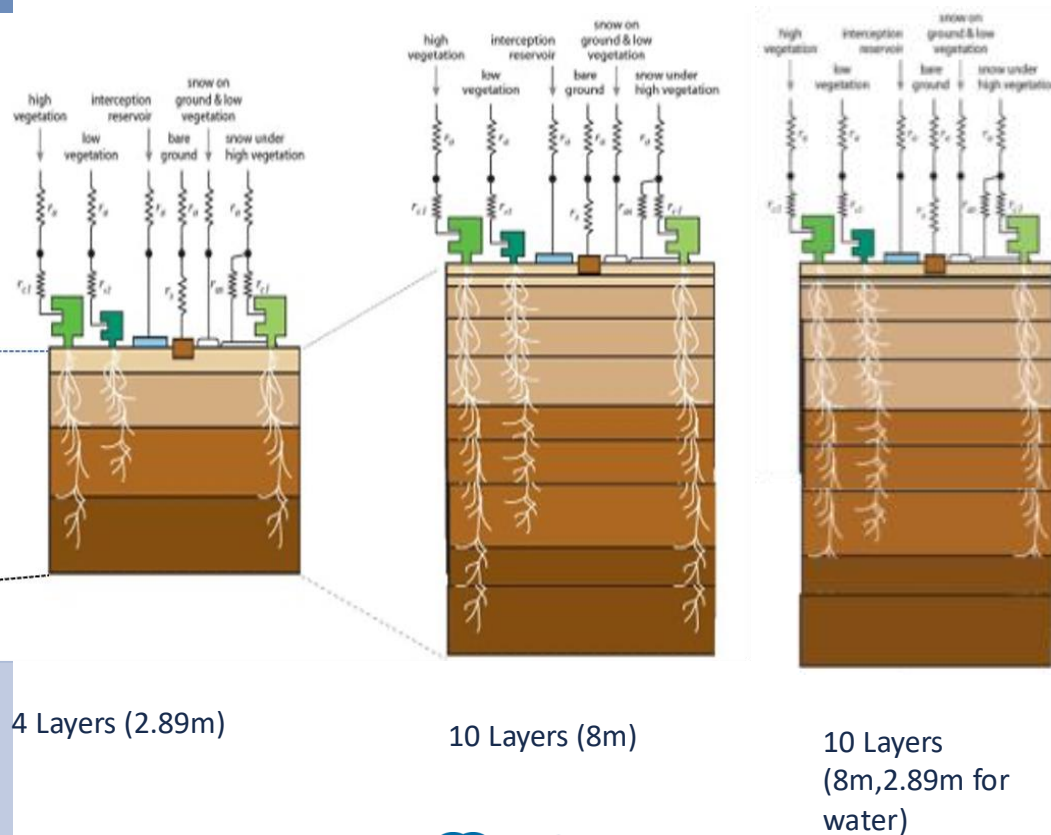
## Soil Conductivity



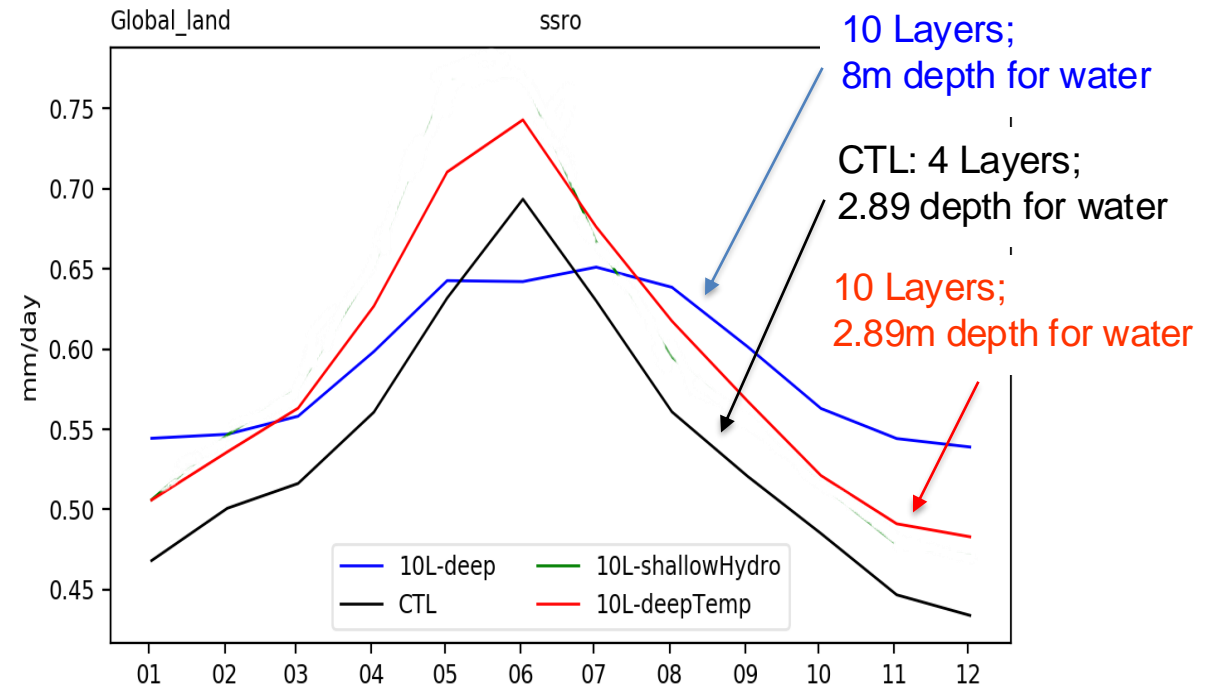
# Water transfer, considerations on vertical discretization

- The number of vertical layers and the thickness of the topmost layer is important to correctly represent water fluxes like bare soil evaporation, transpiration and runoff

*Best solution is to have a large number of layers, a deep column, and very fine layer close to the surface interface*



Global sub-surface runoff from simulation with different number of layers and total column depth



# Modelling of inland water bodies

A representation of **inland water bodies and coastal areas** in NWP models is essential to simulate large contrasts of albedo, roughness and heat storage

A lake and shallow coastal waters parametrization scheme, FLake, was introduced as part of ecLand.

**FLake (Mironov et al. 2010, BER)** is a two-layer bulk model based on a self-similar parametric representation of the evolving temperature profile within lake water and ice.

**ecLand - HTESSEL**

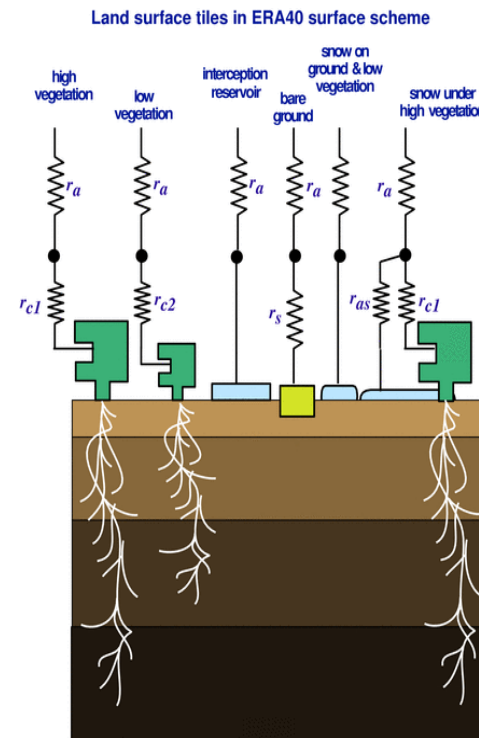
Hydrology - Tiled ECMWF

Scheme for Surface Exchanges over Land

+

**FLake**

Fresh water Lake scheme



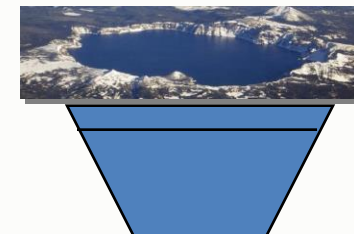
## ● Lake tile

Mironov et al (2010),

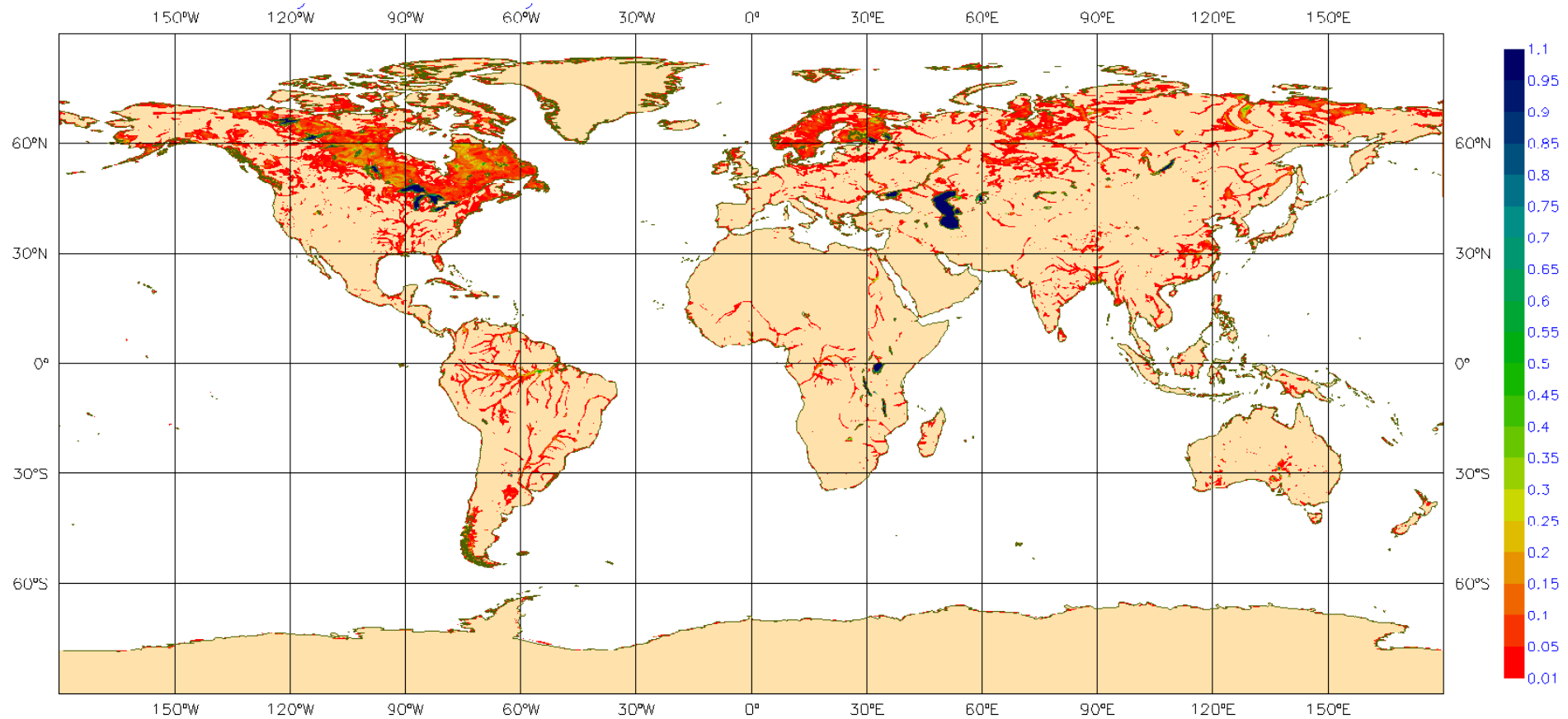
Dutra et al. (2010),

Balsamo et al. (2010, 2012,  
2013)

Extra tile (9) to account  
for sub-grid lakes



# Inland water bodies fraction



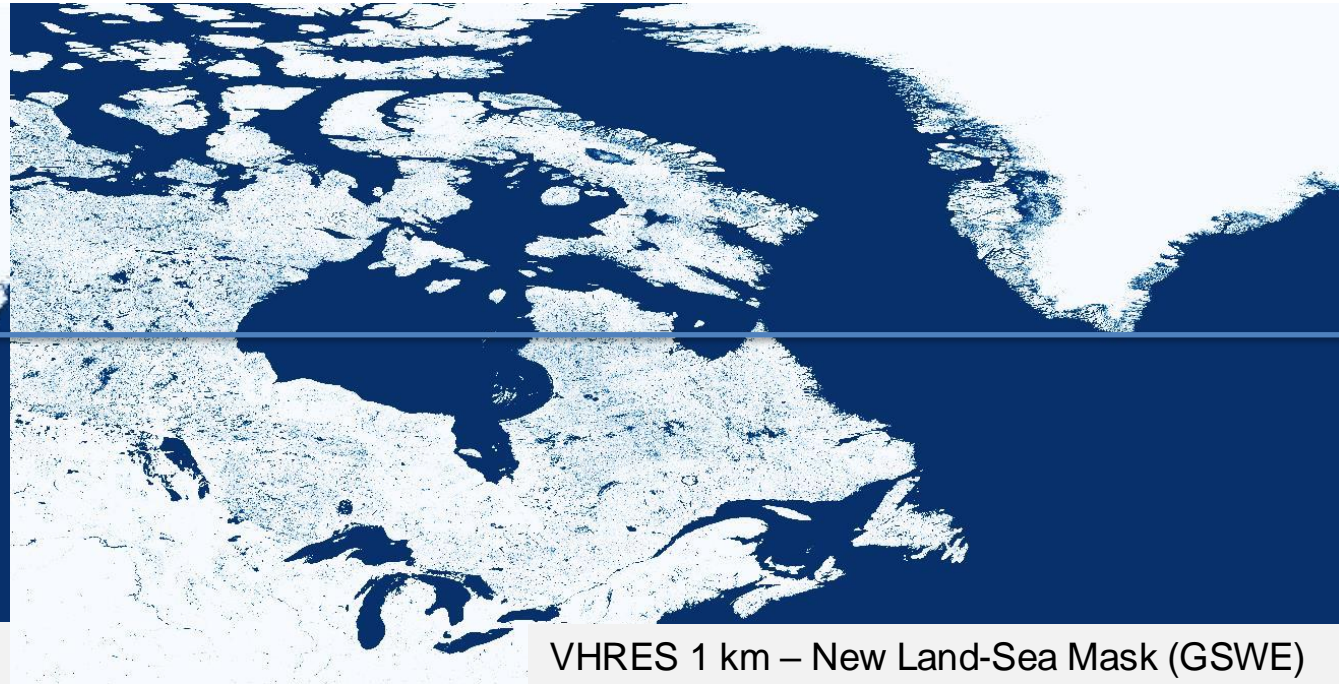
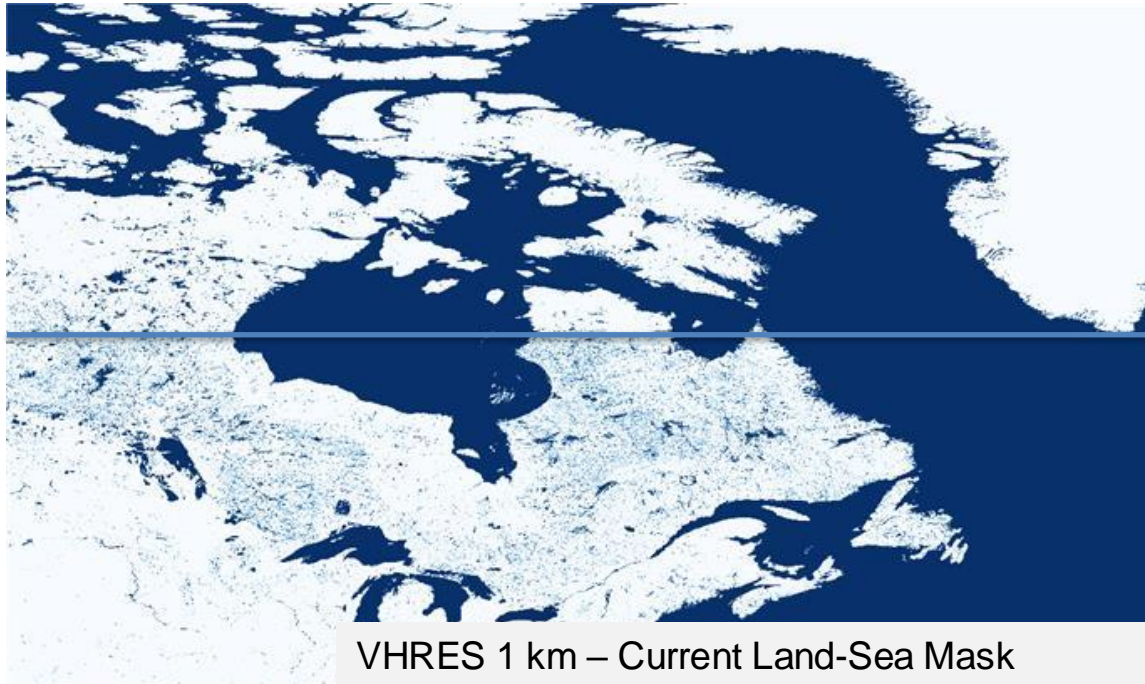
**Aggregated from GLOBCOVER 300m**

# Towards representing inland water surfaces at km-scale resolution

Margarita Choulga, Souhail Boussetta et al.

Moving towards global 1km enable resolving more of the inland water surfaces affecting the surface temperature

Mapping water surfaces correctly is essential to have an inter-consistent treatment of land surface



Choulga et al. 2021. Example of land sea mask obtained by the global 30m resolution GSWE aggregated to 1km on Google Earth Engine.

<https://hess.copernicus.org/articles/23/4051/2019/>

# Energy fluxes: diurnal cycle impact of lakes

Manrique-Suñén et al. (2013, JHM)

## Monthly diurnal cycle of energy fluxes for July

### ● Lake tile

Mironov et al (2010),

Dutra et al. (2010),

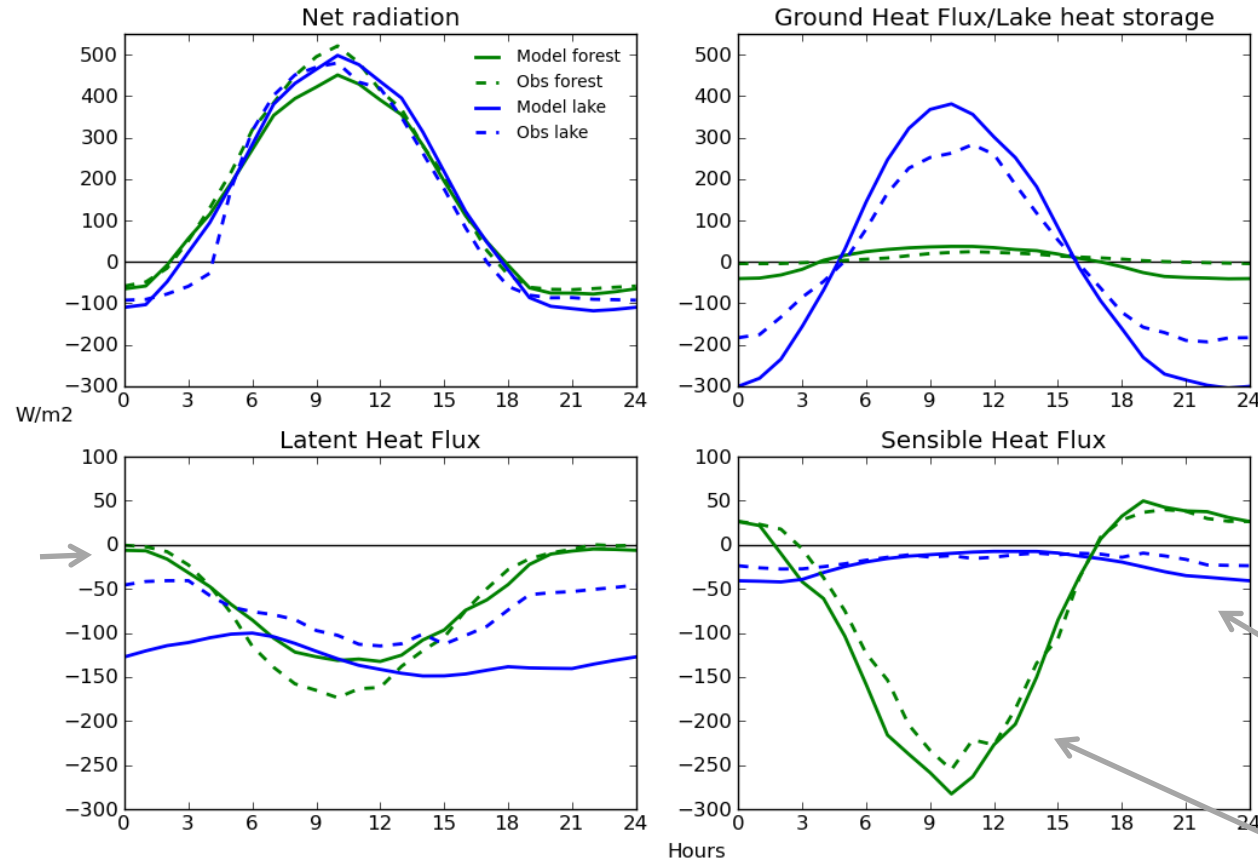
Balsamo et al. (2010, 2012, 2013)

Extra tile (9) to account for sub-grid lakes



**Forest**  
evaporation is driven by vegetation, so it is zero at night

**Lake LH**  
diurnal cycle: over-estimation in evaporation



Very good representation by the model of diurnal cycles and particularities of each surface

Lake SH maximum is at night

Forest SH maximum is at midday

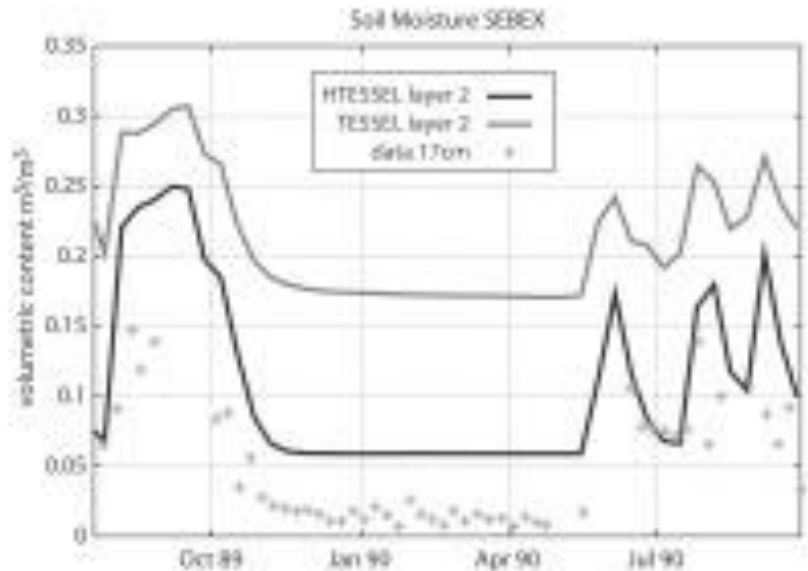
Main difference between lake & forest sites is found in energy partitioning

# Process evaluation using in-situ observations

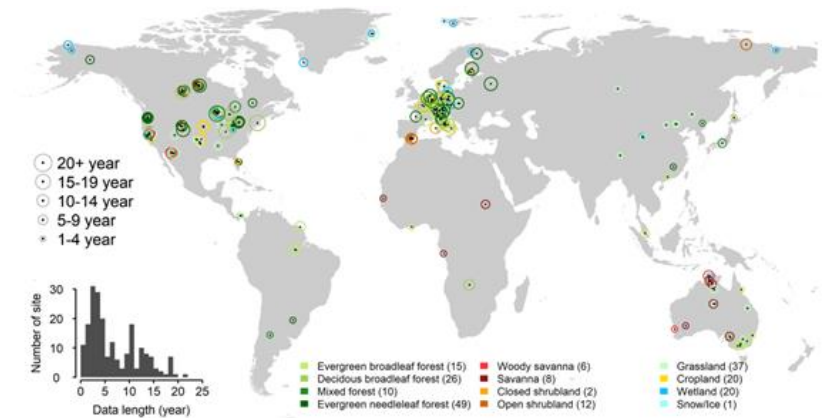
Balsamo et al 2009 JHM, Dutra et al. 2010 JHM, Arduini et al. 2019, ....

Evaluation of land-surface process improvements at instrumented sites is fundamental for assessing that forecasts are improved for the right reason

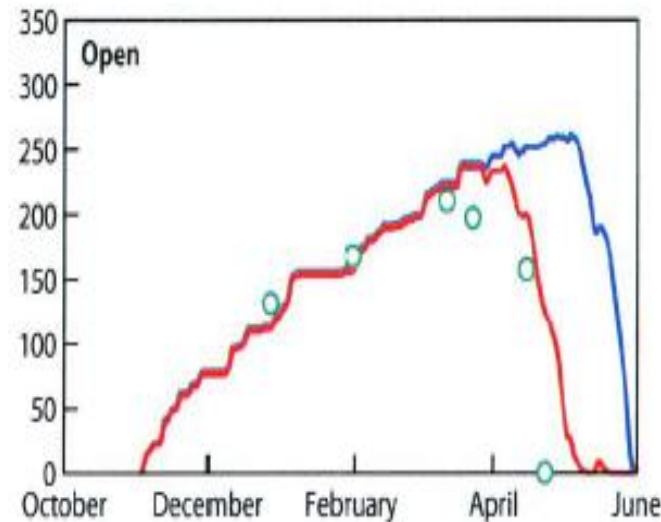
Evolution of soil moisture for a site at SEBEX site. Observations (dot), TESSEL(Grey), and HTESSEL(black) schemes.



FLUXNET 2015 sites



Evolution of snow mass at SNOWMIP 2 site in the **2010** and **old** snow scheme



Col de Porte, snow site in French Alps



# Weather forecasts impact of improved representation of soil/snow processes

- Hydrology-**TESSEL**

Balsamo et al. (2009)  
van den Hurk and Viterbo  
(2003)

Global Soil Texture (FAO)

New hydraulic properties

Variable Infiltration capacity &  
surface runoff revision

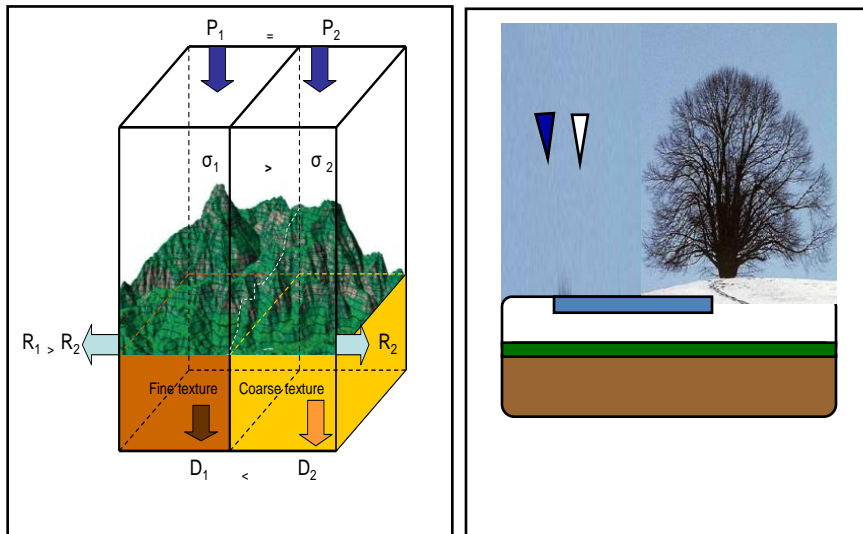
- Snow update**

Dutra et al. (2010)

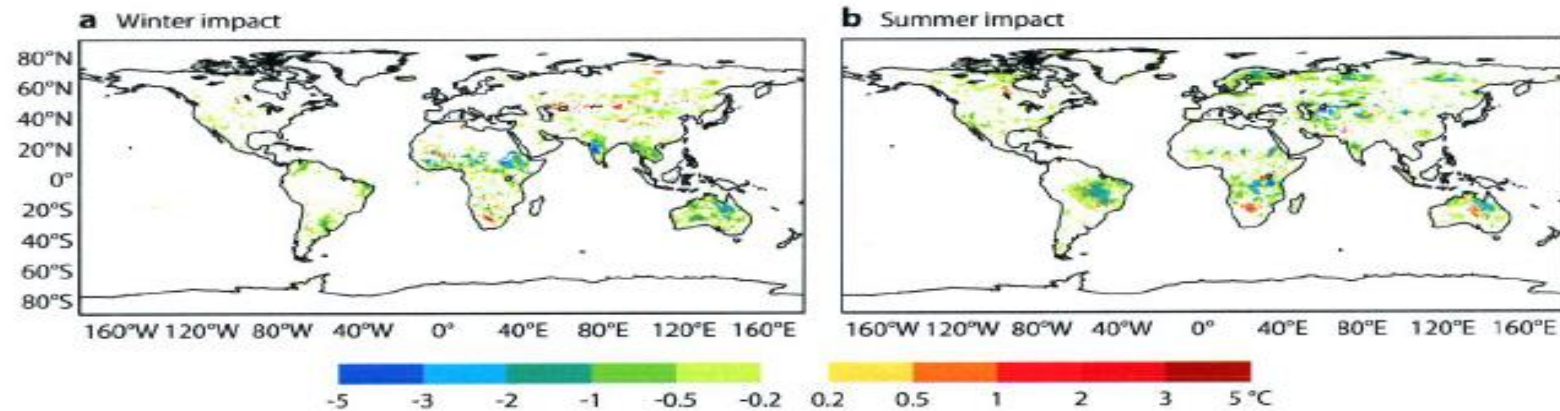
Revised snow density

Liquid water reservoir

Revision of Albedo  
and sub-grid  
snow cover



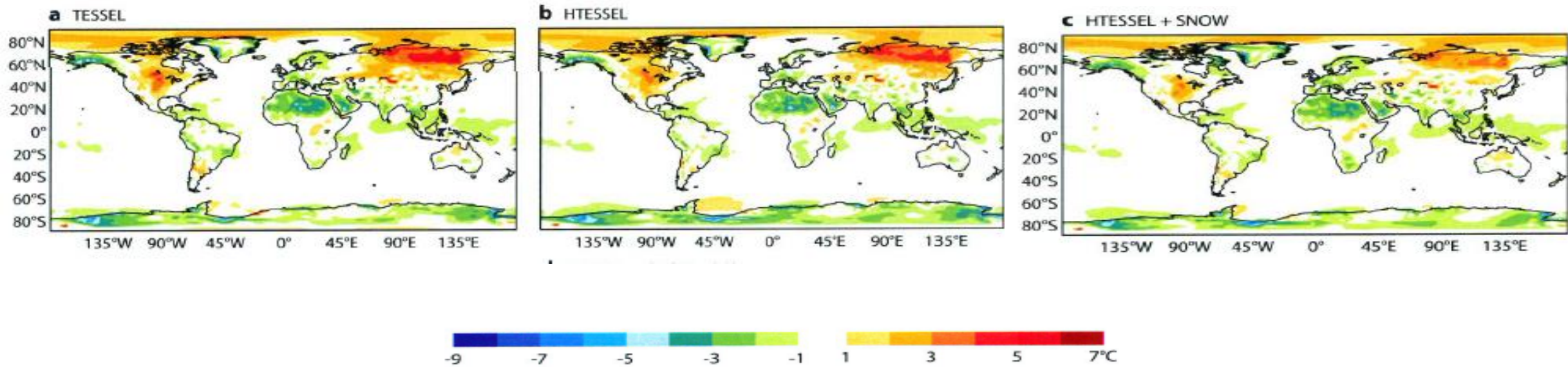
## Forecast Impact (+36-hour forecast, mean error at 2m temperature)



Improving 2m temperature

Degrade 2m temperature

# Climate improvements from land developments (soil, snow)

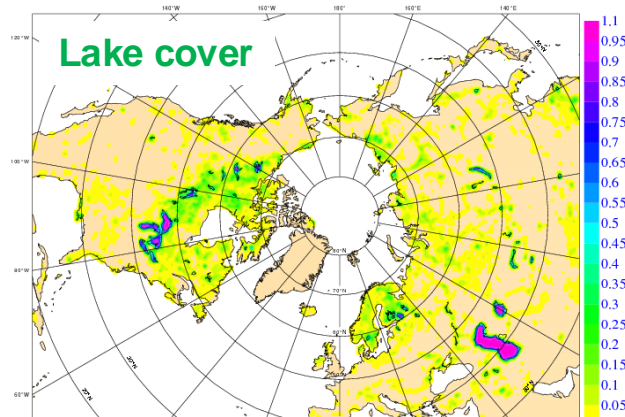


simulations colder than ERA-Interim

Warmer than ERA-Interim

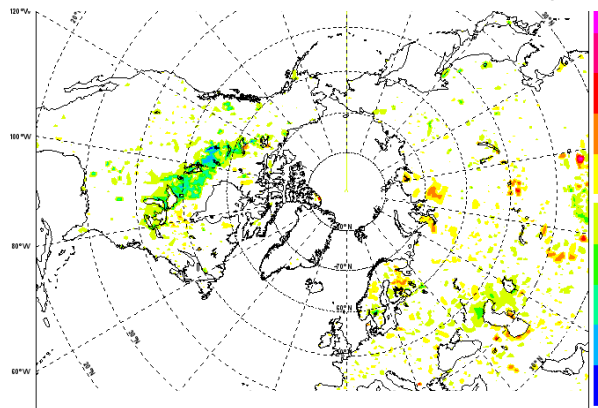
# Impact of lakes in NWP forecasts

Balsamo et al. (2012, TELLUS-A) and ECMWF TM 648



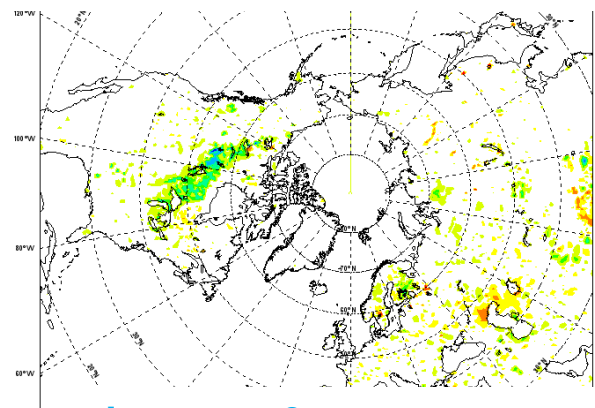
- Forecasts sensitivity and impact of lakes is shown to produce a spring-cooling on lake areas with benefit on the temperatures forecasts (day-2 (48-hour forecast) at 2m).
- The lake surface temperatures are verified with MODIS LSTs as indicative of the heat-storage accuracy of the lake model

**Forecast sensitivity**

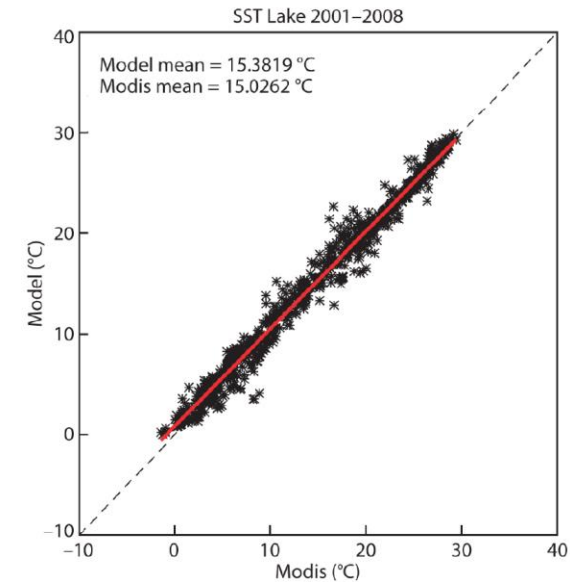


Cooling 2m temperature  
Warming 2m temperature

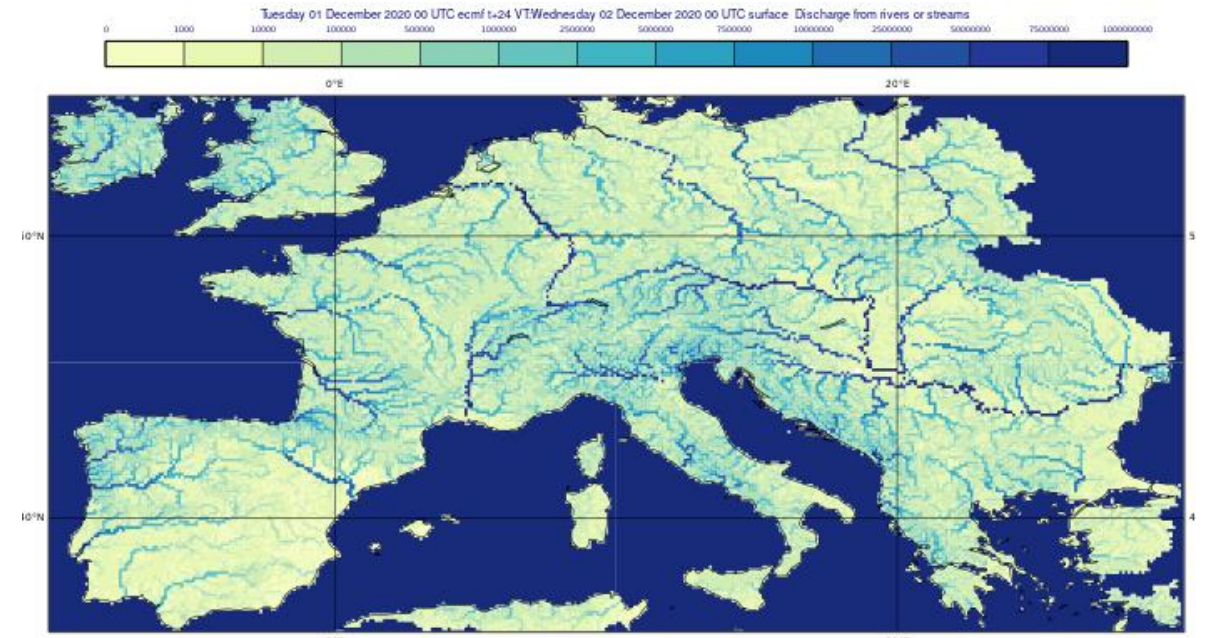
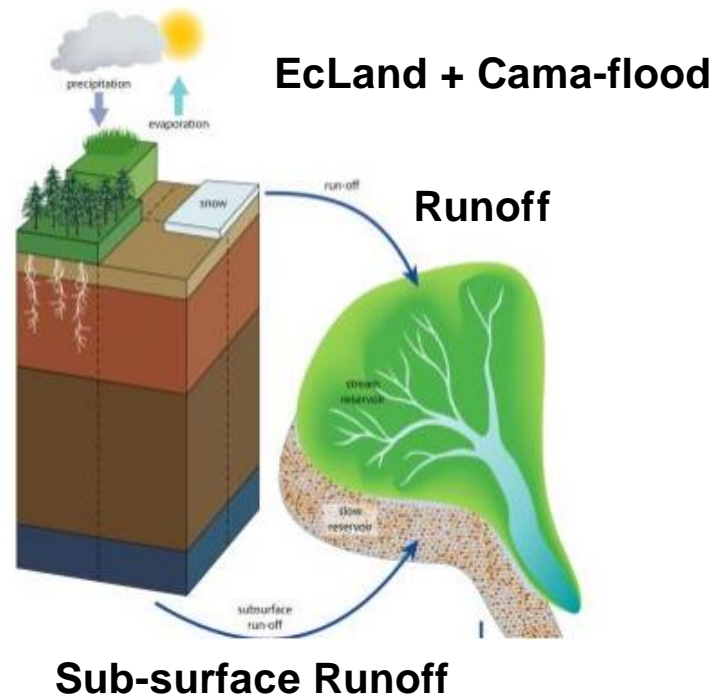
**Forecast impact**



Improves 2m temperature  
Degrades 2m temperature



# Toward simulating floods & inundations in ecLand and in the IFS

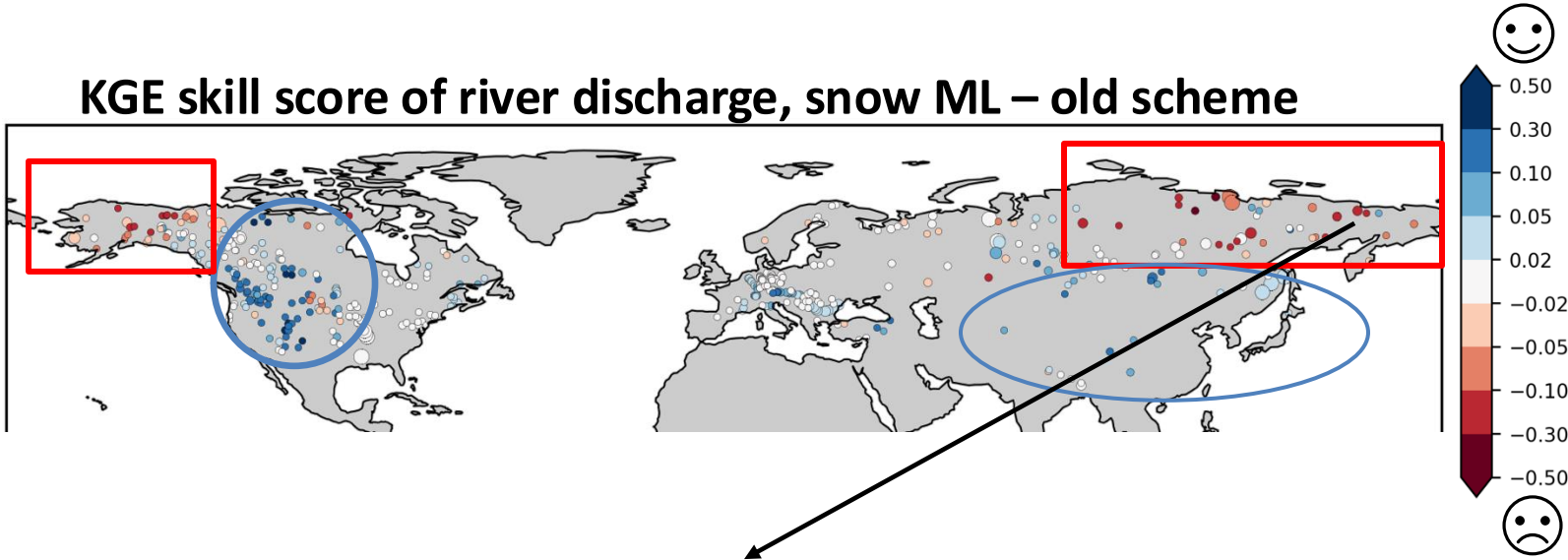


*River discharge from EcLand coupled with Cama-flood over the European domain at 3arc/min resolution (~5km)*

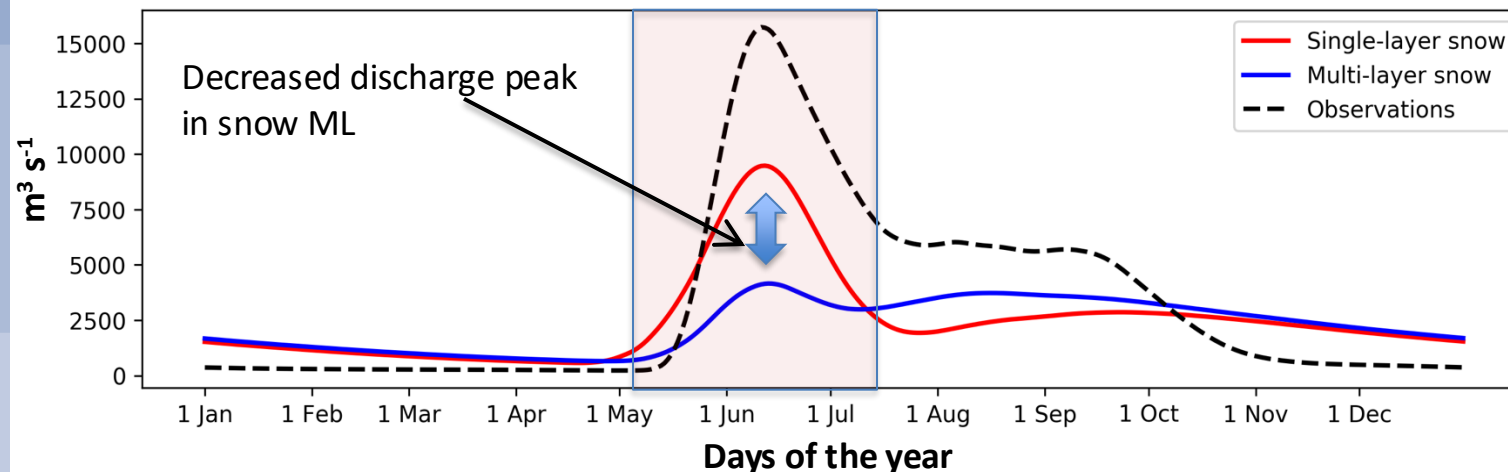
- Coupling river discharge to ecLand in future high resolution land reanalysis (running since 2023 in Destination Earth project)
- Coupling within the IFS is ongoing to permit forecasting river floods

# Hydrology to evaluate land-surface model developments – multi-layer snow example

## KGE skill score of river discharge, snow ML – old scheme

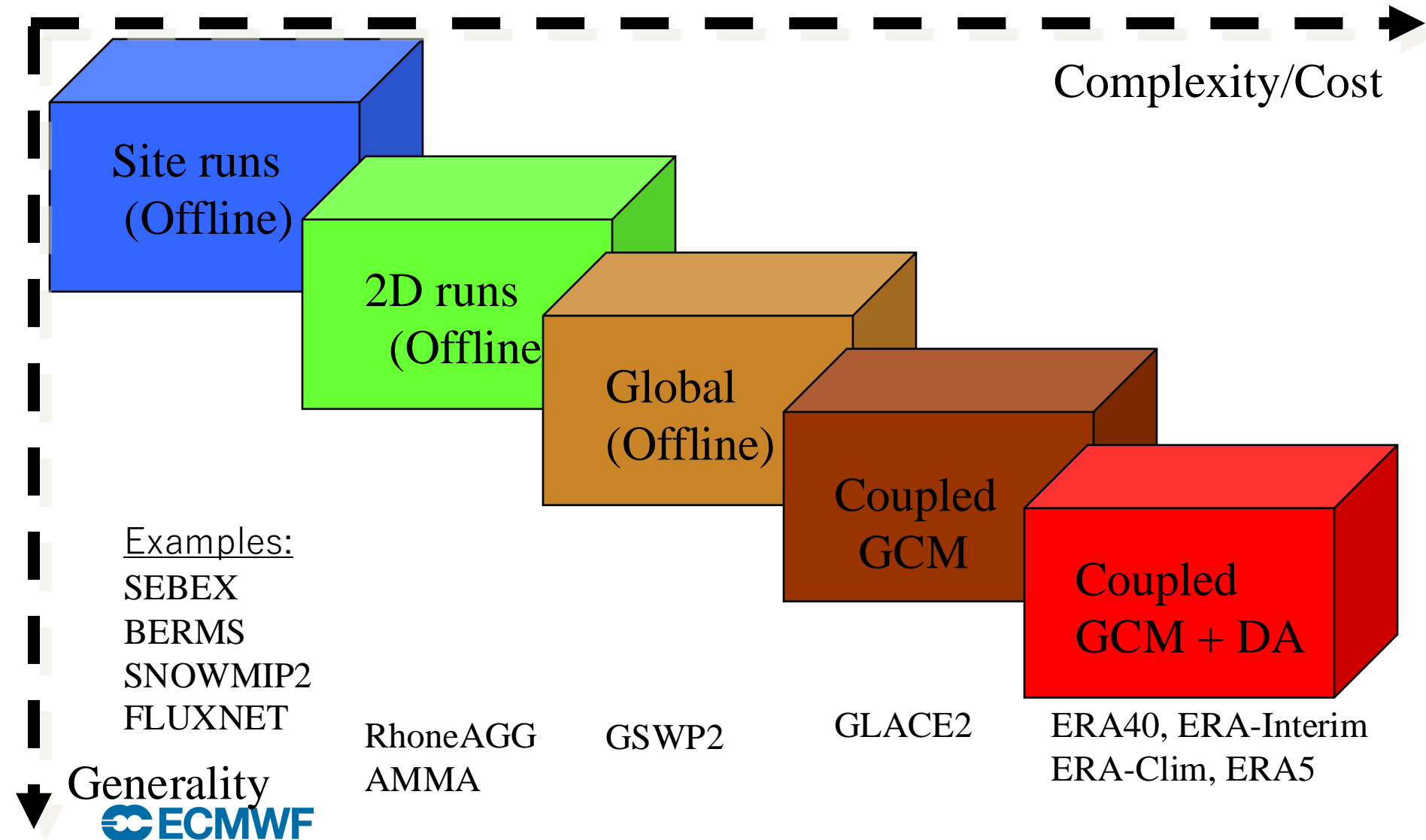


## Daily mean annual cycle of river discharge for Kolyma river, lat=68.72; lon=158.71

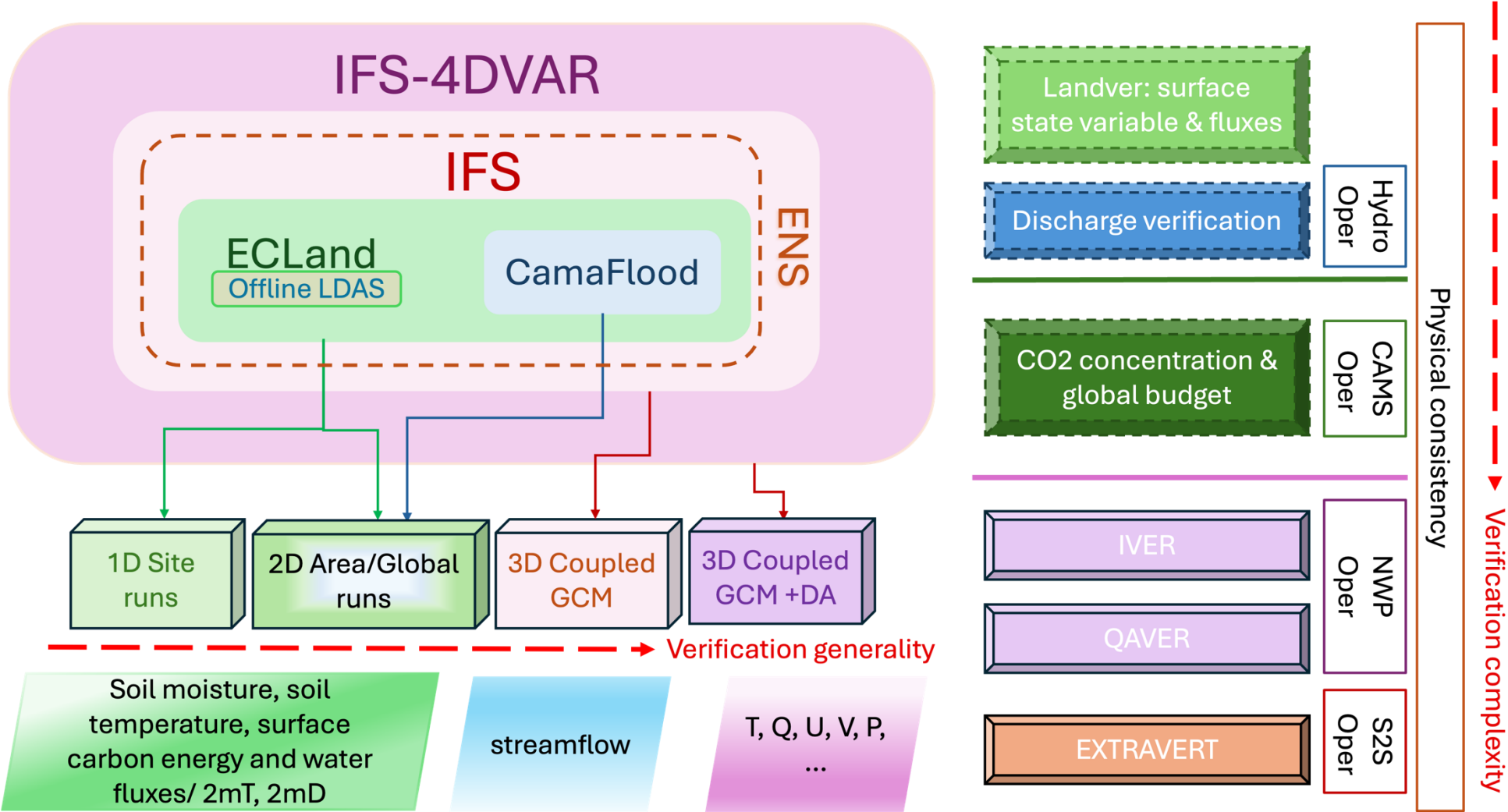


- More catchments show improvements, in particular over Rockies and mid-latitude Eurasia
- Many catchments in cold climates show lower skills (permafrost regions)
- In permafrost areas, excess of water infiltrating into the soil amplifies river discharge biases. Main causes:
  - warmer soil temperature in snowML
  - Frozen soil thawing for sub-zero temperatures

# Strategy for surface model development at ECMWF (applied)

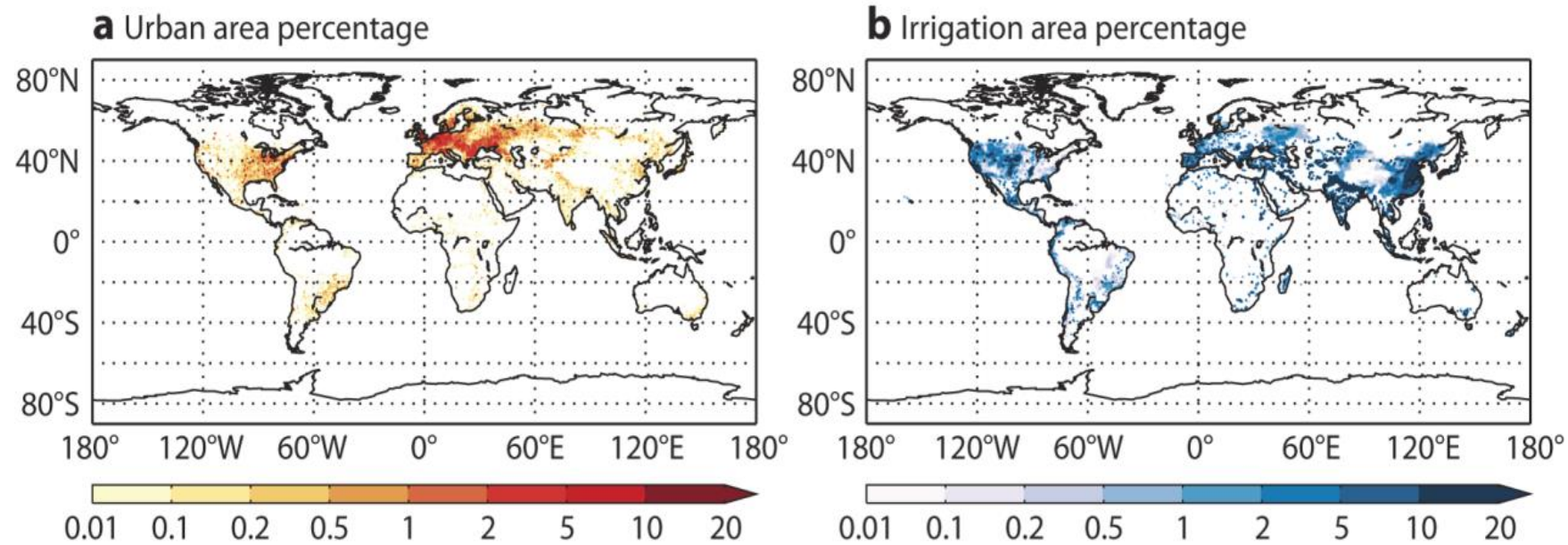


# Strategy for surface model development at ECMWF (applied 2)



# Missing surface components: An example

- Human action on the land and water use is currently neglected in most NWP models...

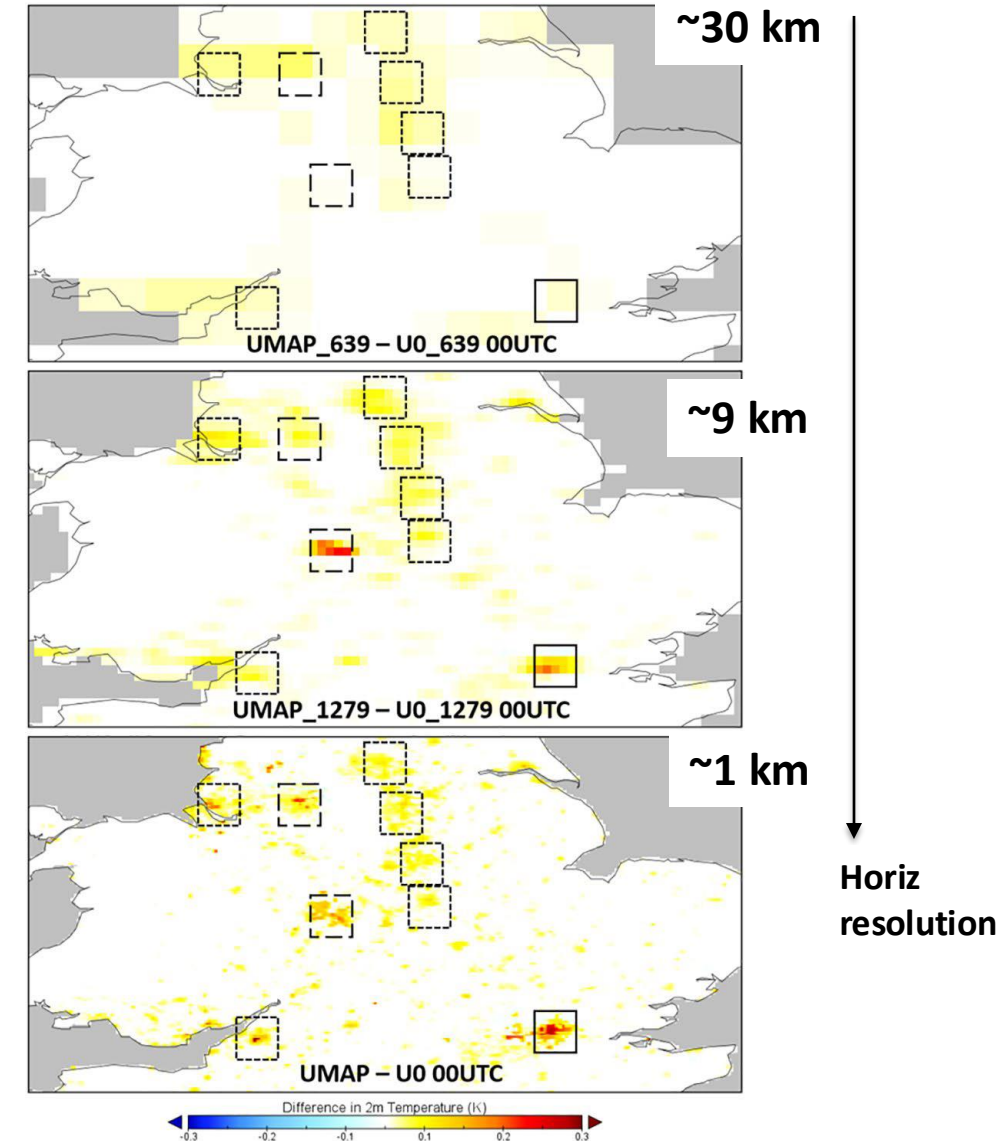
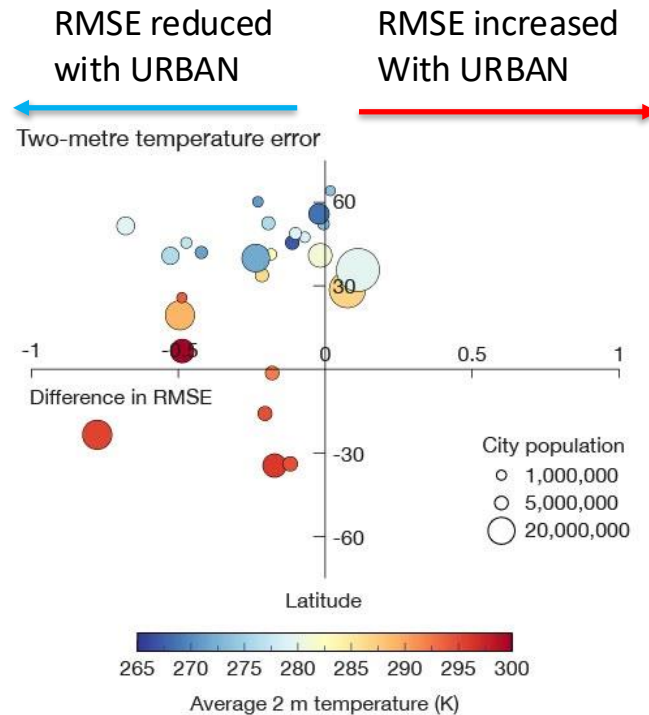
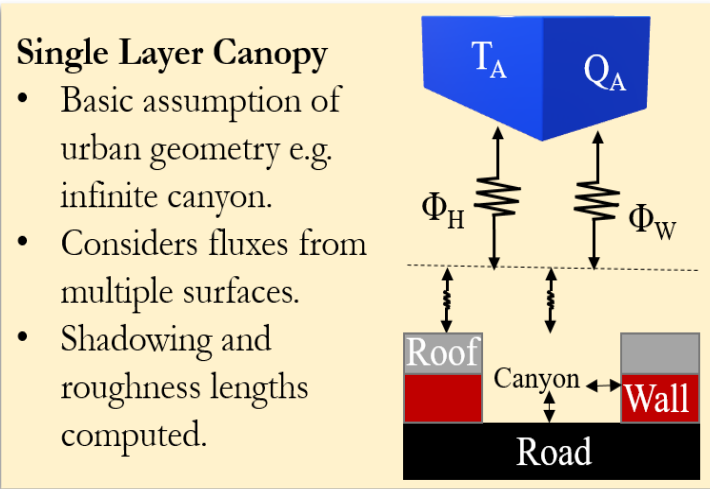


- Urban area (a, in %, from ECOCLIMAP, Masson et al., 2003) ← Operational in 2024
- Irrigated area (b, in %, from Döll and Siebert, 2002)
- Also water bodies are changing over time
- Glacier mass dynamics is missing

# Increase in horizontal resolution calls for more complexity at the surface

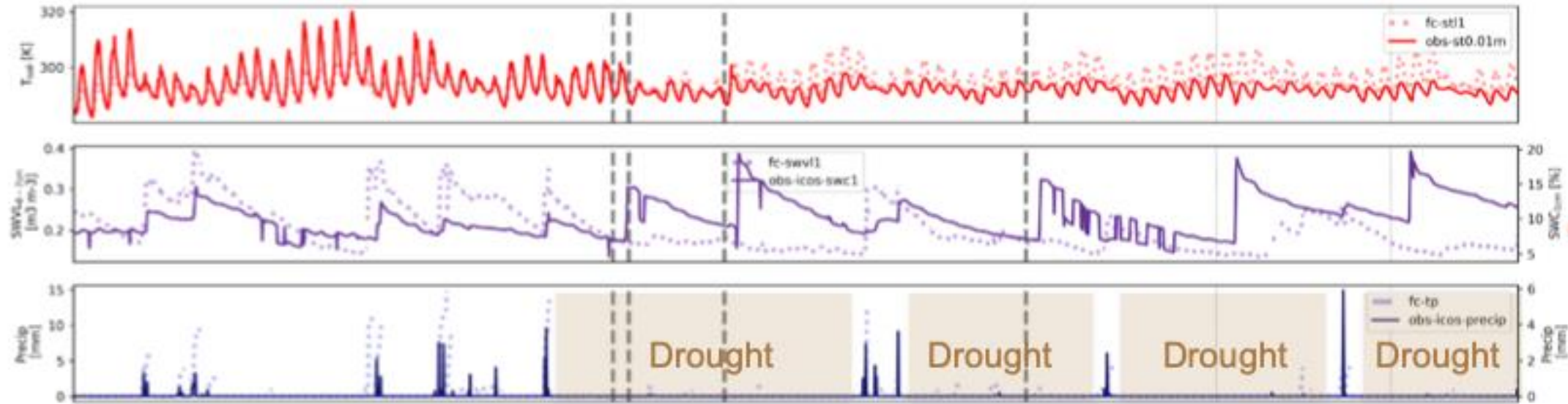
- As horizontal resolution of forecast model increases, urban environments get more and more resolved.
- Urban heat island effects can affect temperature by several Celsius
- A urban module and dedicated tile under testing for future IFS cycles

## 2-metre temperature difference between two simulation with and without an urban scheme



# Missing surface components: An example for irrigation

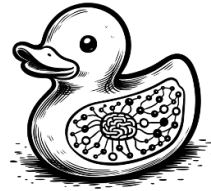
Soil temperature, Soil moisture and precipitation for an irrigated site in Germany.  
Observations compared to forecasts.



- Irrigation increases soil moisture, with a direct impact on temperature (and plant growth)
- Operationally, this is currently accounted for by the soil moisture data assimilation scheme

# Emulation of the Land Surface

ML has already successfully been used to emulate different parts of the land surface at ECMWF

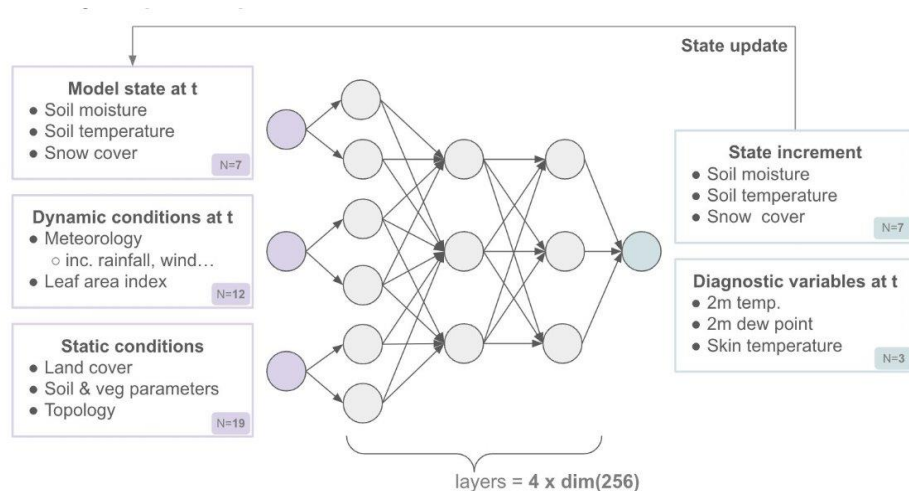


ecLand	Land & hydrology variables in the AIFS	EFAS/LISFLOOD	Probability of Fire	Orography parameters
Multilayer perceptron	Graph neural network encoder + decoder Transformer processor	Long-Short Term Memory	XGBoost	Gaussian process

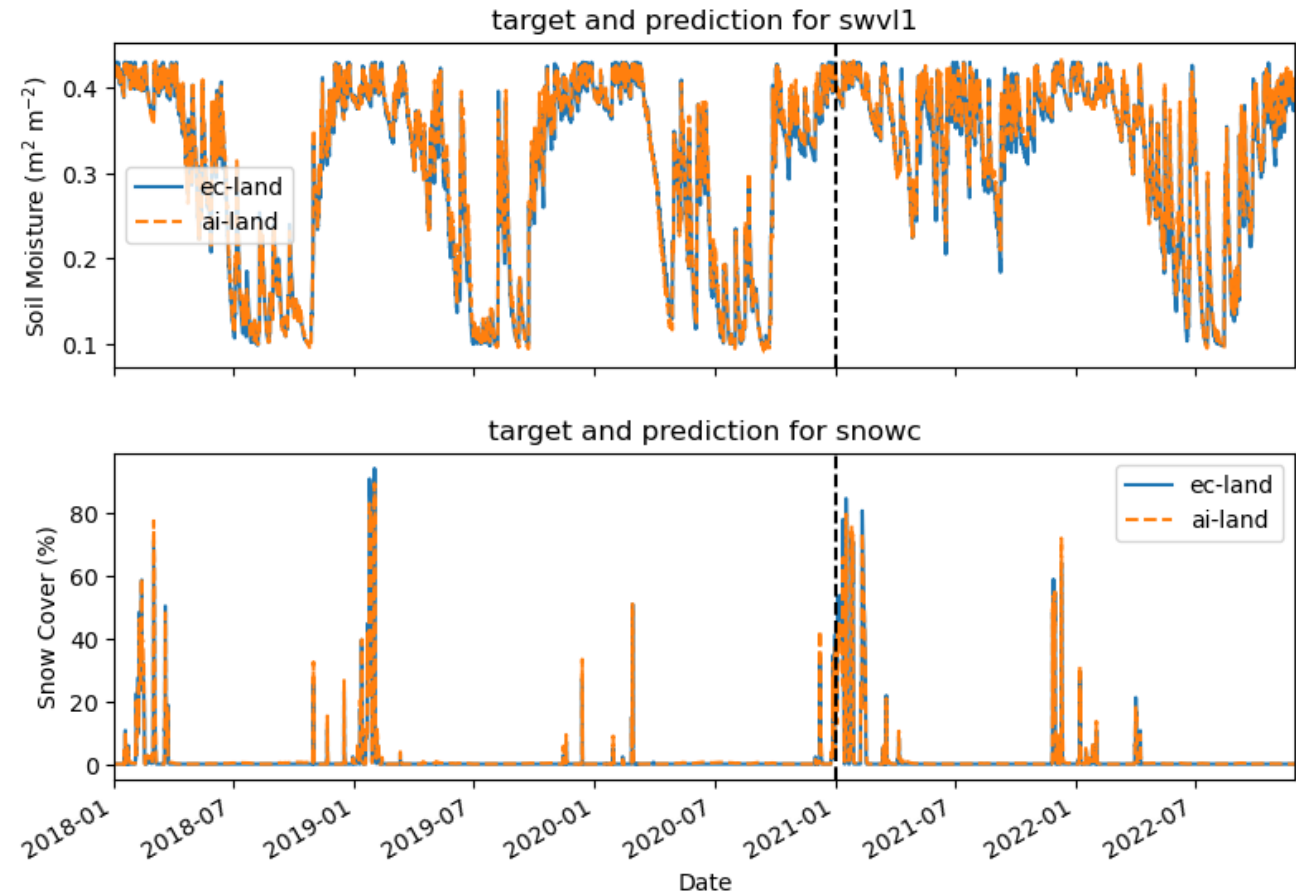
# Emulating the land surface: aiLand



- Standalone land emulator trained taking atmospheric forcing
- Simple column model MLP, stable for multiple year rollouts
- Will be used in Land Data Assimilation and Parameter Estimation routines to improve the physical model
- Looking into coupling strategies with the other Earth System components



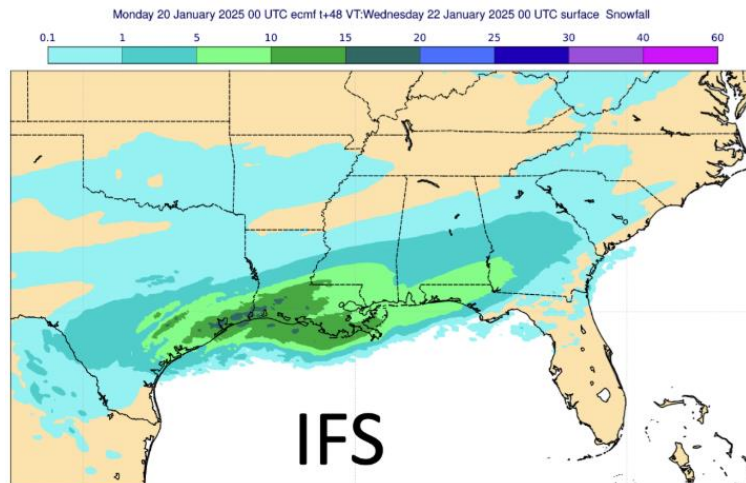
EUROPEAN CENTRE FOR MEDIUM-RANGE WEATHER FORECASTS



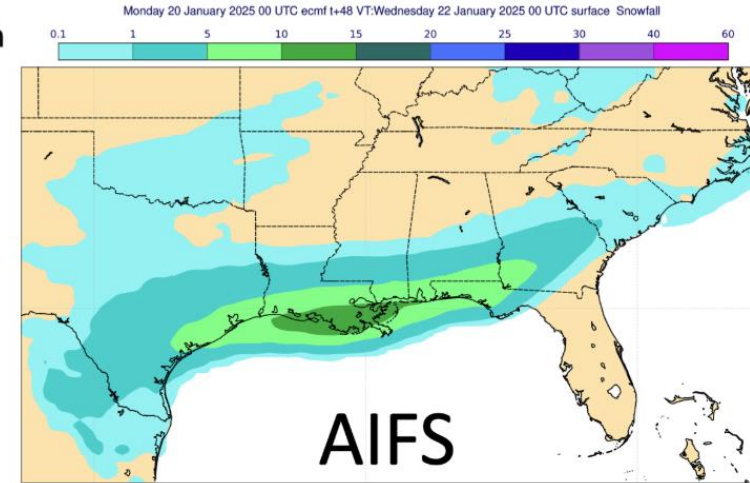
# Emulating the land surface: AIFS v1

- Variables of **soil moisture, soil temperature, runoff, solar radiations, cloud covers, 100m winds, snow fall** have been directly trained as part of the newly operational AIFS v1
- Rare snow event capture in the ML model, even though it was out of sample

## Pinnington



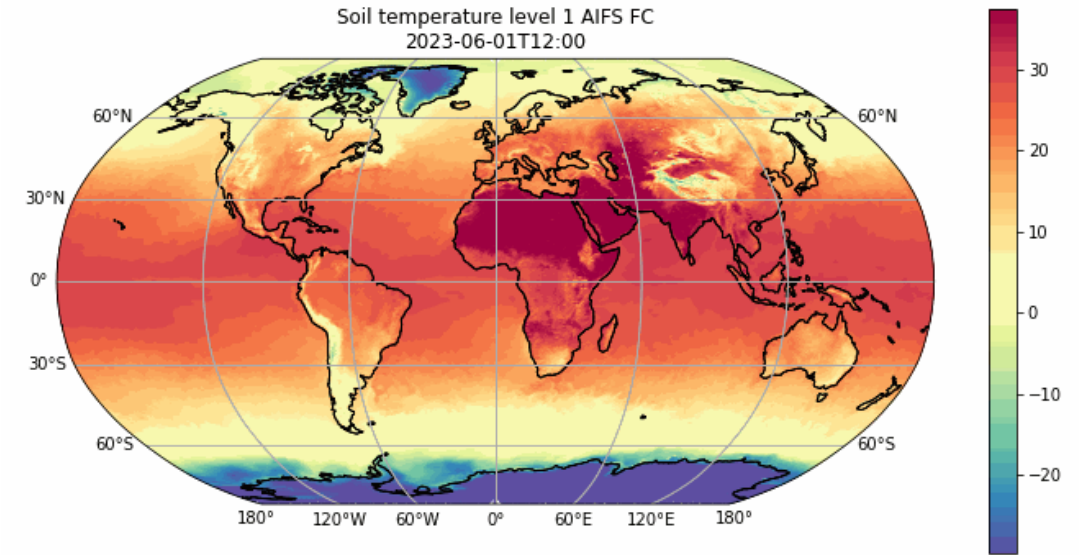
24h snowfall; T+24-48h  
VT: 21 January 2025



## Rare snow along the Gulf Coast

*"AIFS snowfall forecast looks consistent with the IFS snowfall forecasts globally and for this specific snow event near the Gulf of Mexico."*

*22/01/2025 Weather report  
Ivan & Linus*



# Perspectives

- Efforts to improve diurnal and seasonal cycles of surface state variables has transferred into weather and climate improvements and this will continue (doing things better may not sound attractive but it pays off!)
- Surface complexity is needed and permitted by the overall skill of the atmospheric processes.
- Surface representation requirements for higher resolution will not saturate at a given scale.
- Earth-Observation from Satellites provide guidance for improving processes (not only useful in the data assimilation step, but also in the model development phase) and justify complexity.
- In-situ data will provide guidance on process-level fidelity of a scheme. That cannot be expected at global scale and therefore in-situ data will always be a crucial part of verification.
- Foreseen benefits from using ML (and hybrid system) for surface modeling in different aspects (from parameters estimation and uncertainty quantification to process emulation )

# ECMWF surface model milestones

Vegetation based evaporation	1989
ML-soil (4 layers + ...)	1993 / ERA15
Initial conditions for soil water	1994
Stable BL/soil water freezing	1996
Albedo of snow forests	1996
OI increments of soil water	1999
TESSEL, new snow and sea ice	2000 / ERA40
HTESSEL, revised soil hydrology	2007
HTESSEL+SNOW, revised snow	2009
HTESSEL+SNOW+LAI, seasonal vegetation	2010
CHTESSEL (carbon-land surface)	2012
LAKETESSEL (addition of lake tile)	2015
SEAMLESS Coupling Ocean-Sea-Ice	2018
ecLand modelling platform.	2021
Multi-layer snow model	2023
Coupling with Cama Flood	2023
Urban parametrisation in IFS	2024